



UNIVERSITÀ  
degli STUDI  
di CATANIA

Dottorato di ricerca in Scienze Geologiche, Biologiche e Ambientali

**UNDERSTANDING BLOCK ROTATION ALONG  
STRIKE-SLIP FAULT ZONES IN YUNNAN (CHINA):  
PALEOMAGNETIC AND STRUCTURAL APPROACH**

**Ph.D Thesis**

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**XXXI CICLO**

**2015/2018**



## ***Introduzione e scopo del Progetto di Ricerca***

Durante gli ultimi decenni, sono stati proposti diversi modelli tettonici per spiegare la deformazione e le rotazioni di blocchi crostali, attorno ad assi verticali, in zone di taglio o *strike-slip*. Un metodo per validare questi modelli è l'analisi paleomagnetico - strutturale realizzata proprio lungo le più importanti faglie trascorrenti del mondo.

L'area di studio proposta per il progetto di ricerca di questo dottorato, è stata analizzata da un classico e noto lavoro di *Tapponnier et alii* [1982], che descrive, attraverso un modello analogico con la plastilina, l'indentazione della placca Indiana nel continente Euro-Asiatico. Ciò avrebbe prodotto, a partire da 40-50 Ma, età della collisione tra i continenti India ed Eurasia, ispessimento crostale nell'area a nord e nord-est della catena Himalayana e l'estrusione del blocco Indocinese verso sud-est. Uno dei principali punti d'interesse sulla deformazione continentale e sull'evoluzione tettonica dell'Asia [Molnar, 1988], ed ancora oggetto di acceso dibattito, è il ruolo delle grandi faglie strike-slip che, probabilmente, hanno guidato l'estrusione del blocco Indocinese.

Un primo gruppo di studiosi considera la crosta essenzialmente rigida, con la deformazione concentrata lungo i suoi margini: in questo modello, le faglie trascorrenti litosferiche hanno grandi rigetti ed grandi tassi di movimento. Un secondo gruppo lega la deformazione della crosta continentale ispessita, come quella Tibetana, allo scollamento ed alla deformazione omogenea di un sottile settore crostale, a comportamento viscoso, soggetto a rotazioni e raccorciamento. Gran parte di questa controversia è focalizzata proprio sul ruolo e sul rigetto orizzontale della faglia del Red River, lunga circa 1000 km e che si estende con orientazione NW-SE dall'angolo SE del Tibet fino al Mar Cinese Meridionale. Questa faglia, insieme alla Gaoligong e alla Chongshan, dividono strutturalmente la provincia dello Yunnan (Cina sud-occidentale) in tre blocchi crostali più piccoli caratterizzati da un'intensa attività sismica.

In questa tesi, si propone pertanto uno studio paleomagnetico e strutturale lungo queste faglie, al fine di valutare se, e con quale entità, queste siano responsabili della deformazione e delle rotazioni dei blocchi crostali che compongono la regione dello Yunnan.

## **Riassunto**

I dati raccolti durante i tre anni di attività di Dottorato di Ricerca riguardano il paleomagnetismo di rocce sedimentarie e vulcaniche affioranti lungo due importanti zone di taglio che caratterizzano la provincia dello Yunnan (Cina): la faglia Gaoligong e la faglia del Red River. Durante il primo anno, sono stati campionati 50 siti (per un totale di 503 campioni) a distanze variabili (fino a  $\sim$ 25 km) dalle miloniti affioranti lungo la zona di taglio Gaoligong. I *red beds* di età Giurassico-Cretaceo, registrano rotazioni orarie rispetto al blocco Euro-Asiatico, che raggiungono il picco massimo ( $176^\circ$ ) in prossimità della faglia e diminuiscono progressivamente spostandosi verso est. Ad ovest della faglia Gaoligong invece, i siti di età Pliocene-Olocene, affioranti nel campo vulcanico di Tengchong, non evidenziano significative rotazioni. Pertanto, i dati dimostrano che l'attività della zona di taglio Gaoligong ha prodotto rotazioni orarie significative probabilmente coeve ai principali movimenti tettonici avvenuti nell'Eocene-Miocene. Il modello di rotazione dei blocchi crostali lungo la zona di taglio Gaoligong, risponde ad un modello cinematico crostale “quasi - continuo”, con blocchi di dimensioni  $\leq$ 1 km vicino alla faglia, la cui dimensione aumenta allontanandosi verso est. I valori di rotazione e di dimensione della zona di deformazione, si traducono in un rigetto orizzontale della faglia Gaoligong di circa 230-290 km.

Al fine di investigare ulteriormente il ruolo delle faglie nella cinematica dello Yunnan, durante il secondo anno di dottorato, sono stati campionati altri 44 siti di età Triassico - Cretaceo (per un totale di 425 campioni) su entrambi i lati della zona di taglio del Red River e all'interno dei blocchi Chuandian, Lanping e Simao settentrionale. Quasi tutti i siti hanno prodotto componenti di magnetizzazione misurabili e stabili, ma la cronologia di acquisizione della magnetizzazione appare diversa nei tre blocchi. I dati evidenziano una rotazione variabile e differenziale che, anche in questo caso, non supporta un modello di rotazione rigida dei blocchi come proposto da autori precedenti, ma piuttosto suggerisce che: 1) il blocco Simao è costituito da blocchi minori (pochi km) che ruotano in senso orario, separati

da domini non rotanti di dimensioni simili; 2) una componente di magnetizzazione di alta temperatura (640-680 °C) suggerisce un comportamento rotazionale simile (blocchi secondari con rotazione oraria, antioraria e non-ruotati), anche al centro del “blocco” Lanping; al contrario, una componente di media temperatura (300-640 °C) acquisita successivamente al 28% della fase di piegamento non ha registrato alcuna rotazione. I siti vicini (entro i 25 km) al sistema di faglie della Red River, localizzati nel “blocco” Lanping producono invece grandi rotazioni dell’ordine di 180°. Anche i siti localizzati a 10-15 km di distanza dalla zona di taglio Chongshan mostrano rotazioni antiorarie di circa 90°, in accordo con il movimento sinistro della faglia.

I dati del presente lavoro, per la prima volta mostrano che le tre faglie Gaoligong, Chongshan e Red River influenzano significativamente le rotazioni di piccoli blocchi crostali nelle immediate vicinanze delle grandi faglie trascorrenti. Tuttavia, il verificarsi di ulteriori rotazioni (dell’ordine di 20°) all’interno dei blocchi Baoshan e Lanping - Simao, a distanze di oltre 20 km dalle faglie, lascia ancora aperto il quadro della deformazione e delle rotazioni di blocchi crostali che caratterizzano la provincia dello Yunnan.

## *Abstract*

Data from this study report on the paleomagnetism of sedimentary and volcanic rocks cropping out near the Gaoligong and Ailao-Shan Red River Shear Zones. Fifty paleomagnetic sites were analyzed collecting 503 samples, during the first year of Ph.D., at variable distances (up to ~25 km) from mylonites exposed along the Gaoligong fault. Jurassic-Cretaceous red bed sites yield systematic CW rotations with respect to Eurasia reaching the peak values of 176° close to the fault, and progressively decrease moving eastward, up to be virtually annulled ~20 km E of mylonite contact. West of the Gaoligong fault, Pliocene-Holocene sites from the Tengchong volcanic field do not rotate. Thus, data show that the Gaoligong Shear Zone activity yielded significant CW rotations that were likely coeval to the main Eocene-Miocene episodes of dextral fault shear. The Gaoligong zone rotation pattern conforms to a quasi-continuous crust kinematic model, and shows blocks of  $\leq 1$  km size close to the fault, which become bigger moving eastward. Rotation and width values of the rotated-deformed zone translate to a 230-290 km Gaoligong Shear Zone dextral offset, which shows that fault shear plays a significant role in Indochina CW block rotation.

During the second year of Ph.D., forty-four Triassic-Cretaceous sites (425 samples) were collected at both sides of the Ailao-Shan Red River Shear Zone (ARRSZ), within the Chuandian, Lanping and Northern Simao blocks. Nearly all sites yielded measurable and stable magnetization components, but magnetization acquisition timing was different in the three blocks. Sites from the Chuandian block show a normal polarity and were remagnetized after folding. In the northern Simao block the magnetization was acquired before folding (about 33 Ma ago), but the ubiquitous normal polarity in Jurassic-Cretaceous sites suggests a pre-folding magnetic overprint. The data show variable and different rotation that do not display evidence of a rigid block rotation, but suggest that the northern Simao block is made of small (few km size) sub-blocks rotating CW, separated by non-rotating domains of similar

size. Finally, a high-temperature (640-680°C) magnetization component suggests a similar rotational behaviour (CW-rotating and non-rotating sub-blocks) in the centre of the Lanping block. Conversely, a 300-640°C component was later acquired at 28% unfolding and subsequently underwent no rotation. The sites close (less than 25 km) to the ARRSZ yield great rotations of nearly 180°, which confirm past occurrence of significant strike-slip shear along the ARRSZ itself. Conversely, sites located at 10-15 km distance from the Chongshan Shear Zone show ca. 90° CCW rotations that imply a left-lateral shear along the fault zone, consistently with recent geological evidence. Summarizing, data from my Ph.D. study, together with previous evidence of rotations documented both near the fault zones and within the blocks themselves, show that crustal deformation of the Yunnan is extremely complex and still puzzling. The Baoshan and Lanping-Simao blocks underwent strong internal deformation and were likely fragmented in smaller independent sub-blocks whose kinematics and tectonics are still a matter of speculation.

This research is divided into four parts.

**Part I** contains six chapters. A brief introduction to the aim of the research presented in this thesis. *Chapter 1*, shows the state-of-the-art of the kinematic models of block rotation, fundamental setting to this research. The *second* and *third Chapter* include a brief overview of the matter of Paleomagnetism and the methods used for gathering paleomagnetic directions through *in situ* sampling and paleomagnetic analysis in laboratory. *Chapter 4* and *5* regard the geological setting of the studied area and the description of paleomagnetic data collected by the previous authors in adjacent areas.

**Part II**, reports the outcomes from the research along the Gaoligong Shear Zone in Yunnan (*Chapter 6*), recently published on Tectonics by *Pellegrino et al. [2018]*.

**Part III**, includes the results of the paleomagnetic analysis of the rotation pattern along the Ailaoshan Red River Shear Zone (*Chapter 7*), and of the deformation inside the Lanping, Simao and Chuandian blocks. The use of anisotropy of magnetic susceptibility (AMS) measurements (*Chapter 8*) and magnetic mineralogy analysis (*Chapter 9*) permit to infer tectonic deformation.

Finally, in **Part IV** there are the concluding remarks of this research (*Chapter 10*), followed by references and supplementary information.

## Introduction

Previous paleomagnetic data from Yunnan (China) showed a predominant post-Cretaceous clockwise (CW) rotation pattern, mostly explained invoking huge (hundreds of km wide) blocks, laterally escaping (and/or rotating) due to India-Asia collision, separated by major strike-slip shear zones. However, the matter of crustal blocks rotation pattern inside the deforming zone of strike-slip faults, remains still highly controversial. Two set of models have been proposed to explain the rotation of Indochina block: a first set of models suggests that CW rotations of Yunnan are the result of lateral escape of Tibet that rotated CW around the north-eastern corner of the Indian Plate called “Eastern Himalayan syntaxis” [e.g. *Chen et al.*, 1995; *Funahara et al.*, 1992, 1993; *Kondo et al.*, 2012; *Sato et al.*, 2001, 2007; *Tanaka et al.*, 2008; *Tong et al.*, 2013]. However, this may explain CW rotations of the northern part of Indochina, but can hardly account for the CW rotations documented further south, at 21°-22°N latitudes.

A second set of models [*Gao et al.*, 2015; *Wang et al.*, 2008; *Zhao et al.*, 2015] argues that CW rotations occur along rigid mega-blocks (hundreds of km wide) or slates bounded by strike-slip faults, conforming to the so-called “bookshelf” tectonics [*Cowan et al.*, 1986; *Garfunkel and Ron*, 1985; *Li et al.*, 2017a; *Mandl*, 1987; *McKenzie and Jackson*, 1986; *Nur et al.*, 1986; *Ron et al.*, 1984]. This model is hard to maintain-at least for the northern Yunnan-as the elongated slates located between the faults are some 1000 km long and only 200-300 km wide, so that their rotations would lay a significant space problem according to *Pellegrino et al.* [2018]. Alternatively, the land stripes between the faults might be broken into several small slates, each rotating independently, of which –however- no field evidence has been found so far. There is a third tectonic mechanism possibly yielding vertical-axis rotations that has not been investigated in Indochina so far: strike-slip fault shear. In fact, it has been observed that strike-slip zones are bounded by damage zones where large (even >90°) vertical-axis rotations may occur in response to the ductile deformation taking place in the

lower crust [Beck, 1976; Hernandez-Moreno *et al.*, 2014 and references therein; Kimura *et al.*, 2011; Piper *et al.*, 1997; Randall *et al.*, 2011; Ron *et al.*, 1984; Sonder *et al.*, 1994; Pellegrino *et al.*, 2018]. Depending upon fault displacement and locking, and crust rheology, such local rotation zone may be even 20-30 km wide on each fault side [Randall *et al.*, 2011; Sonder *et al.*, 1994]. Therefore, paleomagnetism is a powerful tool for documenting and quantifying such tectonic data and this thesis project will aid in the unravelling of the Indochina rotation pattern puzzle.

Therefore, pursued goals of this research are focused on the understanding of: (1) the vertical axis rotation pattern associated to main strike-slip fault systems in Yunnan and (2) the behavior of crustal blocks, during the rotation process in order to (3) evaluate if - and to what extent - the faults can play an important role in the Indochina block extrusion.

To reach these outcomes I investigated the distribution of vertical-axis paleomagnetic rotations inside the deforming zone of the important strike slip fault zones in Yunnan (China), which are interpreted as the extrusion guides of the Indochina block during the collision between India and Eurasia blocks.

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*Chapter*  
**I**

## 1. KINEMATIC MODELS OF BLOCK ROTATION RELATED TO STRIKE-SLIP FAULTING

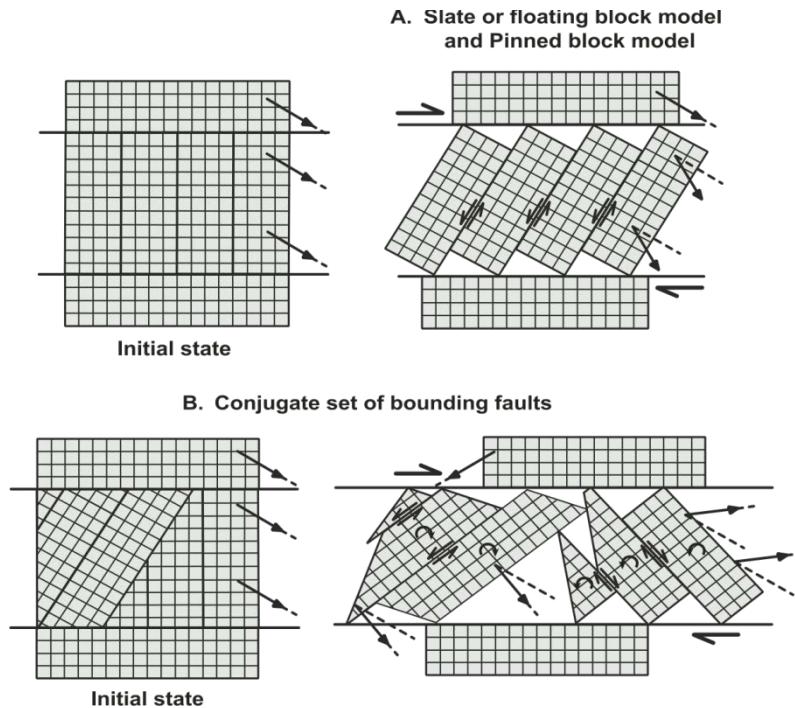
A primary component of the intraplate deformation of Earth occurs in strike-slip tectonic domains, yielding both strike-slip fault displacement and vertical-axis rotation of crustal blocks [Freund, 1974; Garfunkel, 1974; MacDonald, 1980; Lamb and Bibby, 1989; Jackson and Molnar, 1990].

In the last decades, several different geometric models have been suggested to explain the pattern of deformation and paleomagnetic rotation along strike-slip fault zones. These models consider geologic parameters, such as crust rheology, length, sense and total slip of the main shear zone, structural arrangement, deformed zone width, and deformation scale. The rotation amount can be quantified using simple mathematic relations, if some of these parameters are known with good approximation [Hernandez-Moreno, 2015].

To validate the models, paleomagnetic measurements along first-order strike-slip faults have been conducted in different parts of the world: Dead Sea transform fault [Israel, Ron *et al.*, 1984], North and Central Anatolian fault zone [Turkey, Piper *et al.*, 1997; Lucifora *et al.*, 2013], Las Vegas Valley Shear Zone [Western US, Sonder *et al.*, 1994], Alpine Fault [New Zealand, Randall *et al.*, 2011], Eneko and Tanna fault zones [Japan, Kimura *et al.*, 2004, 2011], fault systems of the Aegean Sea [central Greece, McKenzie and Jackson, 1986], Alpine Himalayan Belt [Iran, Jackson and McKenzie, 1984], San Andreas Fault System [Central California, Beck, 1986; McKenzie and Jackson, 1983; Terres and Luyendyk, 1985; Titus *et al.*, 2011], and Liquine-Ofqui fault zone [Cile, Hernandez-Moreno *et al.*, 2014, 2016].

Depending on block size and shape, and rotation pattern, crust deformation models can be subdivided into three main groups, called **discontinuous**, **continuous**, or **quasi-continuous models** [Hernandez-Moreno *et al.*, 2014].

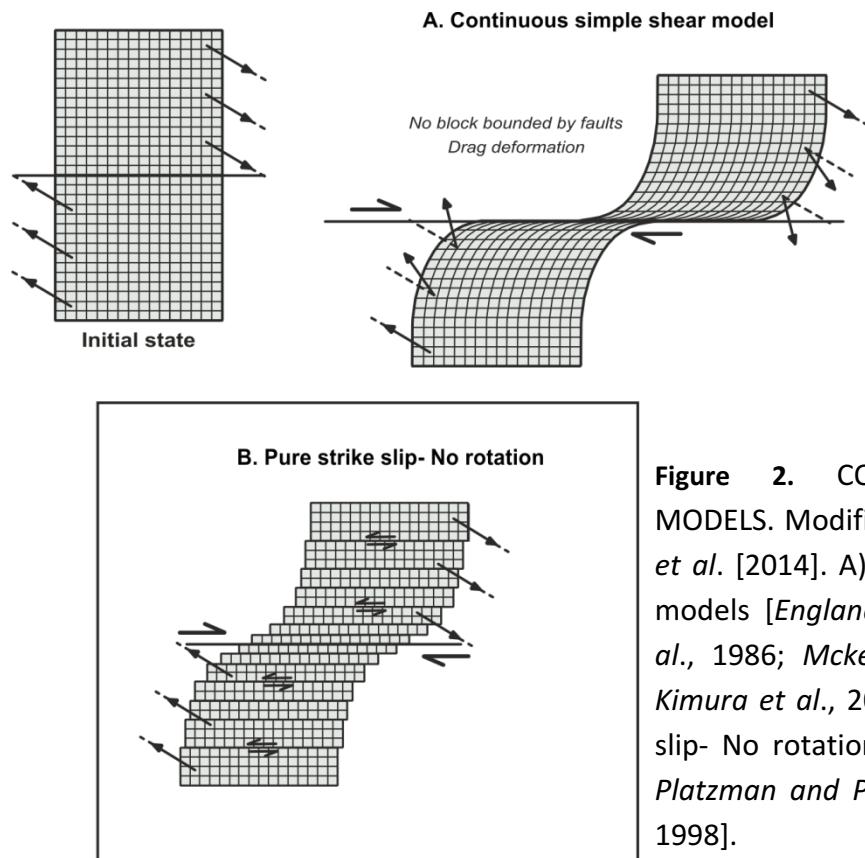
The discontinuous models consider the bookshelf tectonics [Ransome *et al.*, 1910; Morton and Black, 1975; Cowan *et al.*, 1986; Mandl, 1987], where sets of strike-slip secondary faults, within a shear zone, bound rigid or undeformed blocks with similar dimensions to the width of the shear zone that rotate as the faults move laterally [Ron *et al.*, 1984; Garfunkel and Ron, 1985; McKenzie and Jackson, 1986; Nur *et al.*, 1986] (Figure 1).



**Figure 1. DISCONTINUOUS KINEMATIC MODELS.** Modified after Hernandez-Moreno *et al.* [2014]. A) Slate or floating block model and Pinned block model [Ron *et al.*, 1984; Nur *et al.*, 1986; Garfunkel and Ron., 1985; McKenzie and Jackson, 1986]; B) Conjugate set of bounding faults [Ron *et al.*, 1984; Garfunkel and Ron, 1985].

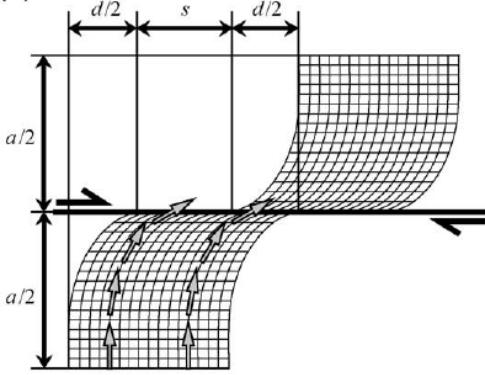
Within the same domain, rotation is spatially constant in amount and sense, cannot exceed 90°, and decreases to zero when the secondary faults become parallel to the domain boundary. In the case of sets of conjugate faults (Figure 1B), expected rotations should be of both senses: CW (CCW) at blocks bounded by sets of left-lateral (right-lateral) strike-slip faults [Freund, 1974; Garfunkel, 1974; Ron *et al.*, 1984; Garfunkel and Ron, 1985; McKenzie and Jackson, 1986; Beck *et al.*, 1993; Hernandez-Moreno, 2015 and reference therein].

The second group of models corresponds to the continuous models.



**Figure 2.** CONTINUOUS KINEMATIC MODELS. Modified by Hernandez-Moreno et al. [2014]. A) Continuous simple shear models [England et al., 1985; Sonder et al., 1986; McKenzie and Jackson, 1986; Kimura et al., 2004, 2011]; B) Pure strike slip- No rotation [Geissman et al., 1984; Platzman and Platt, 1994; Bourne et al., 1998].

McKenzie and Jackson [1986] postulated that if the scale of deformation has a large wavelength compared to the brittle layer crust thickness, the surface deformation is expected to be continuous. Differently from the discontinuous models, the shear is distributed within the deforming zone with no through-going faults over a wide uniform domain (Figure 2). Rotations are expected to be CW (CCW) in right-lateral (left-lateral) fault systems, to progressively increase towards the fault, and to be up to 90° in magnitude [Kimura et al., 2004, 2011].



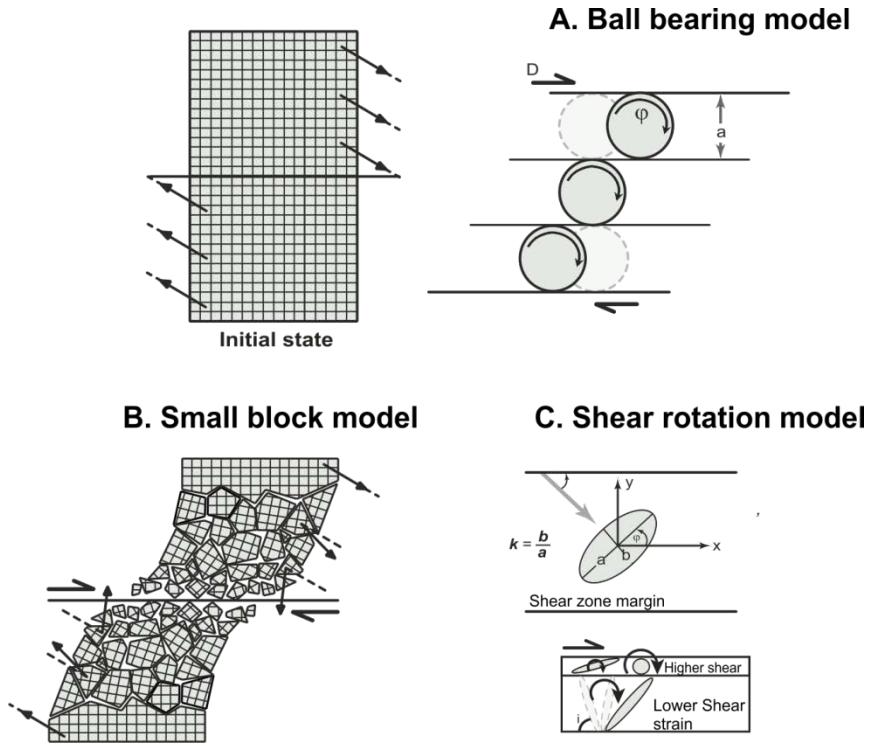
**Figure 3.** Continuous deformation model. Detail of Figure 5A. The amount of rotation at a point close to the fault trace is larger than at a point far from the fault.  $a$  width of the Shear Zone;  $d$  displacement parallel to the fault caused by strike-slip deformations,  $s$  strike-slip offset on the fault trace. Modified from *Nelson and Jones [1987]*.

Based on the assumption that the entire lithosphere behaves as a thin viscous sheet [*Bird and Piper, 1980; England and McKenzie, 1982, 1983; Sonder and England, 1986*], these models suggest that the rotation occurs in the brittle upper crust that forms a thin rigid plastic veneer upon the rest of the lithosphere. Deforming upper crust presumably follows passively the motion of the deeper parts of the lithosphere, expected to behave like a viscous medium [*England et al., 1985; Sonder et al., 1986; England and Wells, 1991*].

Several methods have been applied to quantify the total displacement of strike-slip faults: the continuous models consider necessary add the drag deformation within the shear zone ( $d$ , continuous rotational drag) to the offset on the fault trace ( $s$ , rigid deformation) (Figure 3) [*Nelson and Jones, 1987*]. Ignoring the drag deformation yields to underestimation of both individual faulting events magnitude and fault slip rates [*Salyards et al., 1992; Nagy and Sieh, 1993; Kimura et al., 2004, 2011*].

The continuous drag deformation  $d$  along a strike-slip fault and the amount of rotation at any point are mathematically expressed by *Kimura et al. [2004; 2011]* equations, but it can also be calculated based on the power law rheology model [*England et al., 1985; McKenzie and Jackson, 1983; Sonder and England, 1986; Nelson and Jones, 1987; Kimura et al., 2011*].

Finally, the third group (Figure 4), the *quasi-continuous* models, is also based on a viscous model for lithospheric deformation [*England and McKenzie, 1982; England et al., 1985; Sonder et al., 1986; Sonder and England, 1986; England and Wells, 1991*].



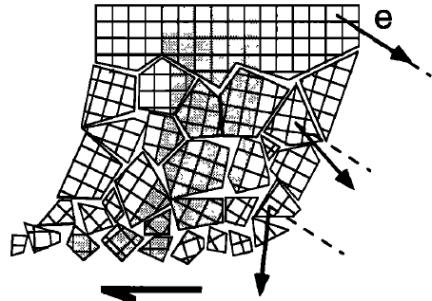
**Figure 4.** QUASI-CONTINUOUS KINEMATIC MODELS. Modified by Hernandez-Moreno *et al.*, 2014. A) Ball bearing model [Beck, 1976; Piper *et al.*, 1997]; B) Small block model [McKenzie and Jackson, 1983; Nelson and Jones, 1987; Sonder *et al.*, 1994]; C) Shear rotation model [Lamb, 1987; Randall *et al.*, 2011].

Here the thin viscous sheet represents the ductile middle and lower crusts under the uppermost brittle seismogenic crust, which unlike in the continuous models, is broken into small rigid blocks with sizes smaller than the shear zone width (Figure 4B-5). Block rotations occur in response to the angular velocity of the ductile deformation, taking place at great depth in the lower crust [Beck, 1976; McKenzie and Jackson, 1983; Lamb, 1987; Nelson and Jones, 1987; Salyards *et al.*, 1992; Sonder *et al.*, 1994; Piper *et al.*, 1997; Randall *et al.*, 2011].

The rotation magnitude will depend on fault length, lithosphere rheology, displacement amount, block aspect ratio (short/long axis) and their orientation with respect to the system-bounding fault [Lamb, 1987; Piper *et al.*, 1997; Randall *et al.*, 2011]. The rotation is CW (CCW) in regions of dextral (sinistral) shear, and increases gradually getting closer to the

fault, reaching values greater than 90° [Nelson and Jones, 1987; Sonder et al., 1994; Piper et al., 1997].

**Figure 5.** Small-block (quasi-continuous model) from Lamb [1987].

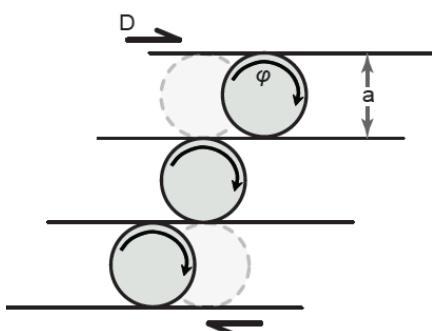


One of the most simple quasi-continuous models is the *ball bearing fashion* [Beck, 1976; Piper et al., 1997], where crustal blocks can rotate freely like balls in a bearing into narrow zones bounded by strike-slip faults parallel to the main shear zone (Figure 4A-6). As the rotation is a continue process, brittle destruction at the corners of the blocks tends to produce sub-rounded equidimensional blocks [Hernandez- Moreno, 2015].

The relationship between fault displacement  $D$  and rotation  $\varphi$  (expressed as a proportion of 360°) of a fault block of width  $a$  approximated by a freely moving circular block is:

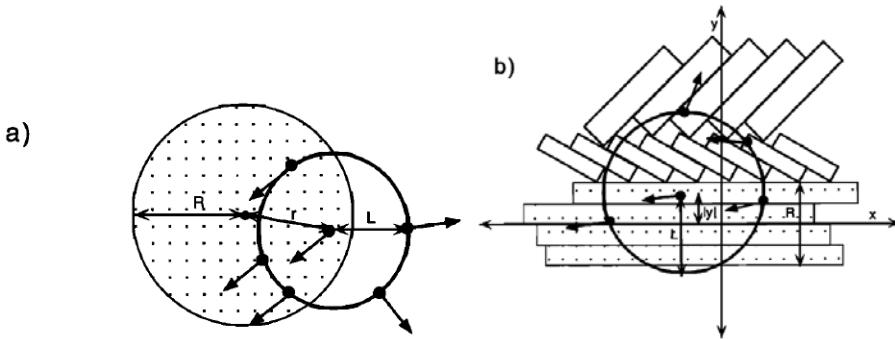
$$D = \varphi \pi a \quad (1)$$

But rotating blocks are not usually equidimensional, as might be required to accommodate very large rotations prevented by the friction between them [Nur et al., 1986].



**Figure 6.** Ball bearing fashion model from Beck [1976] and Piper et al. [1997].

Since from structural observations, may not be possible to know the block size or size distribution, Sonder *et al.* [1994] present two block geometry models to estimate the maximum characteristic block size associated with a strike-slip fault (Figure 7).



**Figure 7.** Models proposed by Sonder *et al.* [1994] to estimate block size. Arrows indicate paleomagnetic declinations at hypothetical localities. **a)** Circular block model,  $R$  is the radius of shaded rigid circular block,  $r$  the position of a locality relative to the centre of the block, and  $L$  the inter locality distance,  $\bar{p}$  (the average probability that a site a distance  $L$  from the locality at  $r$  is indistinguishable from that at  $r$ ) is related to the fraction of the circle of radius  $L$  inside the rigid block (shaded circle). **b)** Domain block model. In this case  $R$  is the width of the domain,  $|y|$  the distance from the domain axis of a locality, and  $L$  the distance of other localities from the first.

The first model (equal blocks rotating independently), considers a rigid circular block of radius  $R$  sampled at a distance  $r$  from its center (Figure 7a) [Sonder *et al.*, 1994].

The second model concerns to highly elongated blocks rotating together as domains [McKenzie and Jackson, 1983, 1986; Ron *et al.*, 1984; Taymaz *et al.*, 1991a,b] (Figure 7b).

Conversely, Lamb [1987] and Randall *et al.* [2011] in their *shear rotation model* showed ellipsoidal blocks whose rotation rate and amount are function of both their aspect ratio ( $k =$  short axis  $b$ /long axis  $a$ ) and orientation with respect to the system bounding fault through time (Figure 4C). Thus, this model predicts a marked decrease in rotation rate for elongate blocks ( $k < 1$ , constant) when they rotate into a direction more nearly parallel to the Shear Zone margins [Randall *et al.*, 2011]. Conversely, if the aspect ratio  $k$  increases by breakup during rotation ( $k \sim 1$ ), the resulting equidimensional block will continue to rotate at a constant rate [Lamb, 1987].

For a width shear zone  $W$ , and fault displacement  $D$ , the total amount of rotation at any particular time, given the initial orientation  $\phi_i$  of the long block axis  $a$ , is determined by the ratio  $D/W$  and the aspect ratio  $k$  [Lamb, 1987].

The block rotation for  $k > 5$  can be approximated within 10% accuracy to that of a passive marker line, equation 2, while for  $k \sim 1$  is defined by the equation 3 [Lamb, 1987]:

$$\theta = 90^\circ - \phi_i + \text{atan}(D/W - \tan(90^\circ - \phi_i)) \quad (2)$$

$$\theta = 0.5 D/W \quad (3)$$

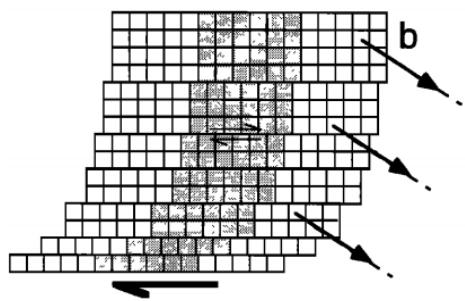
If  $k$  does not equal to 1, also there is a simplified expression for  $R$ , the instantaneous rotation rate ( $d\phi/dt$ ), by Lamb [1987]:

$$R = \frac{W}{2} \left[ \left( \frac{1 - k^2}{1 + k^2} \right) (\cos 2\phi + \tan \theta \sin 2\phi) - 1 \right] \quad (4)$$

Positive values of  $R$  indicate CCW rotations.

When  $k = 1$ ,  $R$  is constant and equals  $-W/2$  (CW to dextral shear). For  $k < 1$ , the rotation rate varies with orientation. For sufficiently small values of  $k$ ,  $R$  changes sign for certain orientations and hence the block will rotate in the opposite direction. The overall effect of all this deformation is a “straightening-out” of the major faults, where shear could be taken up by slip on the faults without any rotation of the intervening faults blocks or slow rotation.

The end-member of the quasi-continuous models would be in the case of elongated blocks trending approximately parallel to the relative plate motion across the strike-slip zone. Pure strike-slip is applied to blocks displacing parallel to the deformation zone, implying a lack of rotations [Geissman *et al.*, 1984; Platzman and Platt, 1994; Bourne *et al.*, 1998] (Figure 8).



**Figure 8.** Transcurrent deformation by displacement on parallel faults, with no rotation. Modified from Sonder *et al.* [1994].

If the frictional forces on faults walls are negligible in comparison with basal traction, the block will be in equilibrium when the net drag force on their base is zero [Bourne *et al.*, 1998]. This coincides with one of the stationary states of Lamb [1987].



*Chapter*  
***II***

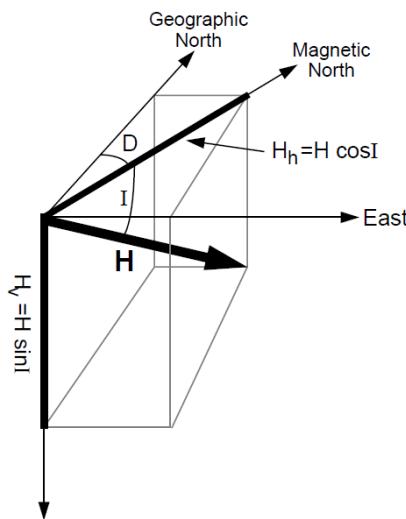
## 2. FUNDAMENTALS OF PALEOMAGNETISM

Paleomagnetism [*Tarling*, 1983; *Tauxe*, 1998; 2009; *Butler*, 1992] is a branch of Geophysics that studies the geomagnetic field behaviour of the geological past. It has widespread applications for a variety of disciplines: the study of atmosphere and biosphere interactions, the study of the early history of the Earth [e.g., *Tarduno et al.*, 2006], the physics of the Earth's interior [e.g., *Christensen and Wicht*, 2007], tectonics [e.g., *Torsvik et al.*, 2008, 2012; *Hernandez Moreno et al.*, 2014; 2016], geologic applications from magnetostratigraphy, biostratigraphy [e.g., *Opdyke and Channell*, 1996], and archaeomagnetic dating [e.g., *Lanos*, 2004; *Pavón-Carrasco et al.*, 2011].

This subject has been vastly addressed in the last forty years by both research papers and books [e.g, *Irving*, 1964; *McElhinny*, 1973; *Beck* 1976, 1980, 1984; *McElhinny*, 1976; *Jelinek*, 1977, 1978; *Hillhouse*, 1977; *Simpson and Cox*, 1977; *Beck and Burr*, 1979; *Kamerling and Luyendyk*, 1979; *Coney et al.*, 1980; *Magill and Cox*, 1980; *Bates et al.*, 1981; *Magill et al.*, 1981; *Globerman et al.*, 1982; *Magill et al.*, 1982; *Demarest*, 1983; *Merril and McElhinny*, 1983; *Tarling*, 1983; *Hillhouse and Grommé*, 1984; *Coe et al.*, 1985; *Luyendyk et al.*, 1985; *Wells and Coe*, 1985; *Beck et al.*, 1986; *Grommé et al.*, 1986; *May and Butler*, 1986; *Hagstrum et al.*, 1987; *Wells and Heller*, 1988; *Butler et al.*, 1989; *Butler*, 1992; *Dunlop and Özdemir*, 1997; *Tauxe*, 1998, 2009; *McElhinny and McFadden*, 2000].

The Earth's magnetic field is approximately a dipolar field (95% of the field component), resulting from an internal source (the dynamo in Earth's core mantle and crustal field) and external source (atmospheric field, and crustal induced field; *Merrill et al.*, 1996). The non-dipolar field is that part of internal geomagnetic field, remaining after that dipole contribution has been removed.

Assuming the validity of dipolar geocentric field, the geomagnetic field can be represented as a vector framed into a three-dimensional, orthogonal coordinate system usually with origin in a specific point on the earth surface. This vector has an  $H$  magnitude and can be broken into two components,  $H_v$  vertical and  $H_h$  horizontal (Figure 9). The direction of geomagnetic field is described by **Inclination I** (angle between the horizontal plane and the geomagnetic field vector, ranging from  $-90^\circ$  to  $90^\circ$  and positive when downward) and **declination D** (angle from geographic north to horizontal component, ranging from  $0^\circ$  to  $360^\circ$ , positive clockwise).



**Figure 9.** Geomagnetic field  $H$  components:  $H_v$  vertical component,  $H_h$  horizontal component,  $I$  inclination,  $D$  declination. Modified after McElhinny [1973].

The vertical component,  $H_v$ , of the surface geomagnetic field,  $H$ , is defined as positive downwards and is given by

$$H_v = H \sin I \quad (5)$$

where  $H$  is the magnitude of  $H$  and  $I$  is the inclination of  $H$  from horizontal, ranging from  $-90^\circ$  to  $+90^\circ$  and defined as positive downward. The horizontal component,  $H_h$ , is given by

$$H_h = H \cos I \quad (6)$$

and geographic north and east components are respectively,

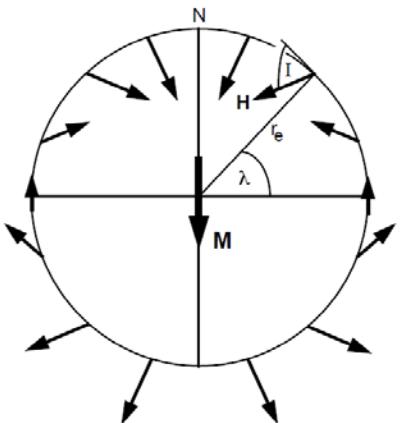
$$H_N = H \cos I \cos D \quad H_E = H \cos I \sin D \quad (7)$$

where D is declination, the angle from geographic north to horizontal component, ranging from  $0^\circ$  to  $360^\circ$ , positive clockwise [Butler *et al*, 1998].

Determination of I and D completely describes the direction of the geomagnetic field. If the components are known, the total intensity of the field is given by the equation:

$$H = \sqrt{H_N^2 + H_E^2 + H_V^2} \quad (8)$$

A central concept to many principles of paleomagnetism is the **Geocentric Axial Dipole (GAD)** model, where the magnetic field produced by a single magnetic dipole at the center of the Earth and aligned with the rotation axis is considered (Figure 10).



**Figure 10.** Modified after Butler [1998]. Geocentric axial dipole model. Magnetic dipole  $M$  is placed at the center of the Earth and aligned with the rotation axis; the geographic latitude is  $\lambda$ ; the mean Earth radius is  $r_e$ ; the magnetic field directions at the Earth's surface produced by the geocentric axial dipole are schematically shown; inclination,  $I$ , is shown for one location;  $N$  is the north geographic pole. Redrawn after McElhinny [1973].

The geographic latitude  $\lambda$  is ranging from  $-90^\circ$  at the south geographic pole to  $+90^\circ$  at the north geographic pole. The inclination of the field can be determined by

$$\tan I = \left( \frac{H_v}{H_h} \right) = \left( \frac{2 \sin \lambda}{\cos \lambda} \right) = 2 \tan \lambda \quad (9)$$

and I increases from  $-90^\circ$  at the geographic south pole to  $+90^\circ$  at the geographic north pole. This relationship between I and latitude will be essential for paleogeographic reconstructions and tectonic applications of paleomagnetism. For a GAD,  $D = 0^\circ$  everywhere.

The magnetic field in a generic point is described by the equation:

$$B = \mu_0 H + \mu_0 J \quad (10)$$

where  $B$  is magnetic induction (expressed in T),  $H$  magnetic field strength ( $A\ m^{-1}$ ),  $J$  magnetization ( $Am^{-1}$ ) or magnetic moment per unit of volume.

In vacuum,  $J$  is nil, in matter its properties depend on those of the elementary particles, according to which substances are subdivided into three categories: dia-, para- and ferromagnetic.

The **diamagnetic** response to application of a magnetic field is acquisition of a small induced magnetization,  $J_i$ , opposite to the applied field,  $H$ . The magnetization depends linearly on the applied field and reduces to zero on removal of the field. **Magnetic susceptibility values**,  $\chi$ , for a diamagnetic material is negative ( $\chi < -10^{-5}\ SI$ ) and independent of temperature (Figure 11a). An example of a diamagnetic mineral is quartz, calcite and dolomite.

For **Paramagnetic** solids, the acquisition of induced magnetization is parallel to the applied field. Magnetic susceptibility values for these materials are positive ( $10^{-5} \leq \chi \leq 10^{-3}\ SI$ ) and depend on temperature (Figure 11b). For any geologically relevant conditions,  $J_i$  is linearly dependent on  $H$ . As with diamagnetic materials, magnetization reduces to zero when the magnetizing field is removed. An example of a paramagnetic mineral are clay minerals, fayalite,  $Fe_2SiO_4$ .

**Ferromagnetic** substances "sensu latu" are magnetite, titanomagnetites, greigite. The  $\chi >> 0$  and varies in a complex way according to the variation of applied field and temperature (Figure 11c-12). The main characteristic of ferromagnetic substances is a magnetic field of their own, also in absence of an external field  $H$ .

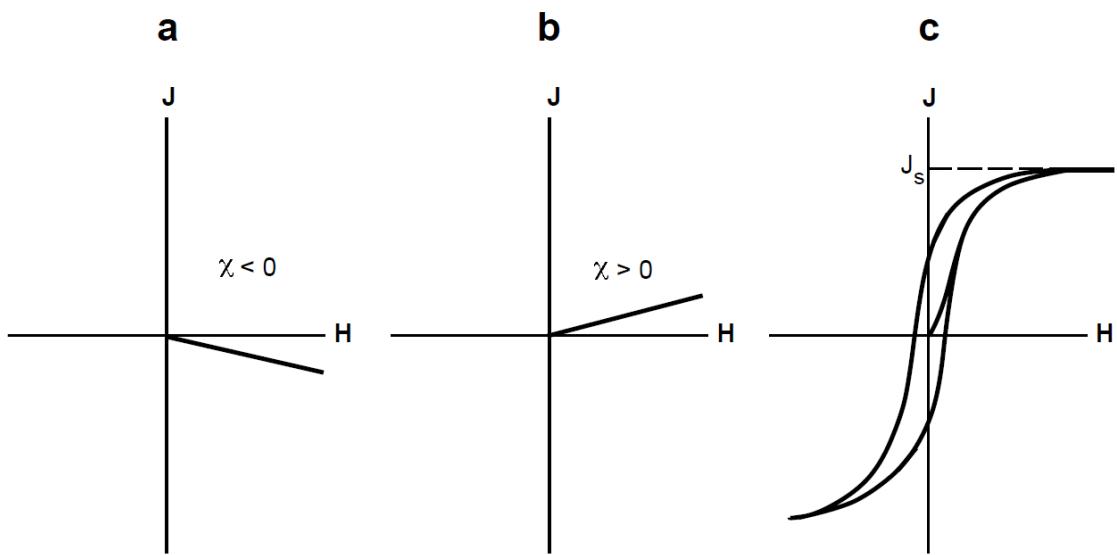
In this type of substance, very strong interaction exists at atomic level, called exchange interaction, which favour the orderly arrangement of the dipole magnetic moments. Above a

certain temperature, characteristic of each ferromagnetic substance, spontaneous magnetization material fades, and the material becomes paramagnetic. This characteristic temperature is called the ***Curie Temperature*** ( $T_C$ ), instead for anti-ferromagnetic substances (hematite) is known as ***Neel Temperature*** ( $T_N$ ). Based on the size of a single ferromagnetic granule, it is possible that its internal structure is divided into a series of microscopic volumes each characterized by a preferential direction (and versus) of magnetization  $J$ . These regions are called ***magnetic domains*** and are developed to minimize the overall magnetic energy of the ferromagnetic granule.

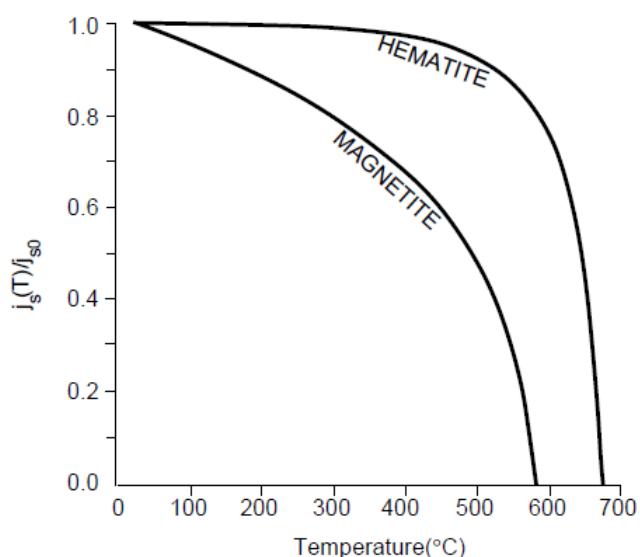
Depending on the size, granules will be:

- **Single Domain (SD)** (a granule SD of cube-shaped magnetite, has size less than 0.1  $\mu\text{m}$ );
- **Pseudo Single Domain (PSD)** (a PSD granule of cube-shaped magnetite, has dimensions between 0.1  $\mu\text{m}$  and 1  $\mu\text{m}$ ) and
- **Multi Domain (MD).**

There is another category of ferromagnetic substances called ***superparamagnetic (SP)***. A superparamagnetic granule is a ferromagnetic granule of extremely small size (of the order of 0.01-0.03  $\mu\text{m}$  for magnetite). Because of their dimensions, in order to remove an external magnetic field  $H$ , these granules are not able to maintain a remaining magnetization for a long time (paramagnetic substances loose the remaining magnetization instantly). Another feature of these substances is that they have a very high magnetic susceptibility (allowing easy identification) [Butler, 1992].



**Figure 11.** Modified after Butler [1998]. **a)** Magnetization,  $J$ , versus magnetizing field,  $H$ , for a diamagnetic substance. Magnetic susceptibility,  $c$ , is a negative constant. **(b)**  $J$  versus  $H$  for a paramagnetic substance. Magnetic susceptibility,  $c$ , is a positive constant. **(c)**  $J$  versus  $H$  for a ferromagnetic substance. The path of magnetization exhibits hysteresis (is irreversible), and magnetic susceptibility,  $c$ , is not a simple constant.



**Figure 12.**  
Modified after Pullaiah *et al.* [1975]. Normalized saturation magnetization versus temperature for magnetite and hematite.  $j_{s0}$  = saturation magnetization at room temperature; for hematite,  $j_{s0} \approx 2$  G; for magnetite,  $j_{s0} = 480$  G.

Paleomagnetism entails the assumption that the “frozen” magnetization (primary) during rock formation (consolidation, diagenesis, cooling, etc), by means of several acquisition mechanisms (detrital, thermal, chemical magnetization), is parallel to the contemporaneous geomagnetic field at the time of the formation.

The essential paleomagnetism theory was presented by the Noble prize winner Louis Néel [1949, 1955], who explained how the ancient magnetic field ( $B_{\text{anc}}$ ) might be preserved in rock’s magnetic memory.

The Néel relaxation theory [1949] defines the characteristic **relaxation time** ( $\tau$ ) by the equation, which relates  $\tau$  to frequency factor ( $\approx 10^8 \text{ s}^{-1}$ ), volume of single domain grains ( $v$ ), the anisotropy constant ( $k$ ), the absolute temperature ( $T$ ) ( $kT$  is the thermal energy), the microscopic coercitive force in single domain grains ( $h_c$ ) and the saturation magnetization of the ferromagnetic material ( $j_s$ ):

$$\tau = \frac{1}{C} \exp\left(\frac{v h_c j_s}{2kT}\right) \quad (11)$$

The Néel theory is valid for the single domain grains (SD, i.e. of  $\sim 0.03 \mu\text{m}$  with a single and stable domain). Depending on Temperature, the relaxation time can overcome the geological time or be unstable over minutes (as the  $T$  closes in the unblocking temperature of magnetic grains the relaxation time decrease exponentially).

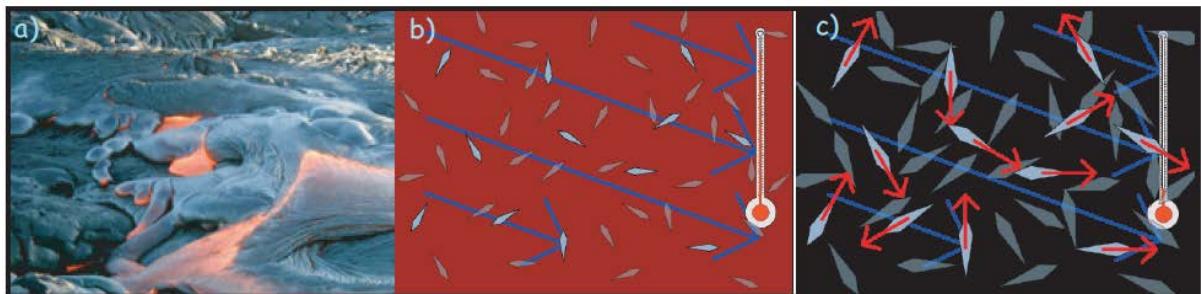
The capability of rocks to record a stable remanent magnetization depends on the relaxation time, as stated by the Néel theory [1949, 1955], depending, in turn, from temperature, and volume of magnetic grains [Butler, 1992; Tauxe, 2009].

The original magnetization can be stable along geological times, although rocks may be exposed to other magnetic field, and undergo thermal, chemical processes that can overprint or even remove the primary magnetization, then acquiring secondary remagnetization.

Secondary magnetizations can be partial or total, deleting completely the original magnetization. Let's distinguish:

**The Natural Remanent Magnetization (NRM)** is the remanent magnetization in a rock sample prior to laboratory treatment, and depends on the geomagnetic field and the geological processes occurred during rock formation. NRM is typically composed of different components; a primary component is the component acquired during the rock formation, secondary components are acquired subsequently, altering or even obscuring the primary NRM. The main forms in which the NRM can be recorded are the TRM, DRM and CRM.

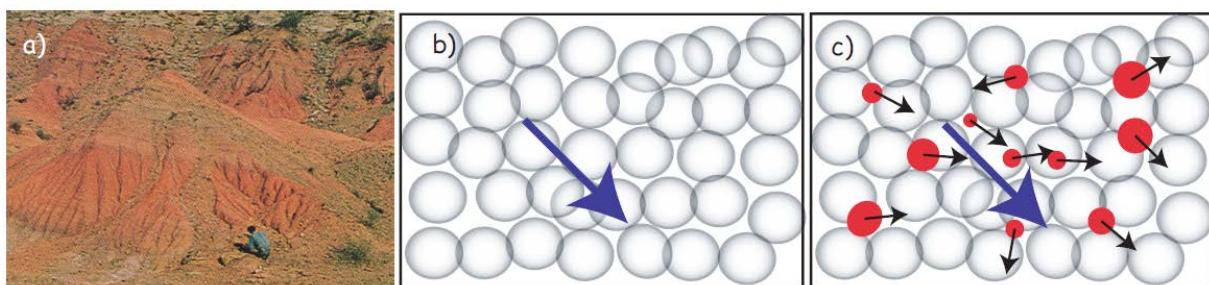
**The Thermo Remanent Magnetization (TRM)** is acquired during cooling from high temperatures to temperatures below Curie Temperatures (temperature below which the magnetic material retains a remanent magnetization (Figure 13); it changes with respect to magnetic minerals).



**Figure 13.** Modified after Tauxe, 2005; a) Picture of lava flow courtesy of Daniel Staudigel. b) While the lava is still well above the Curie temperature, crystals start to form, but are non-magnetic. c) Below the *Curie temperature* but above the blocking temperature, certain minerals become magnetic, but their moments continually flip among the easy axes with a statistical preference for the applied magnetic field. As the lava cools down, the moments become fixed, preserving a thermal remanence.

**The Detrital Remanent Magnetization (DRM)** is acquired during the accumulation of magnetic minerals during sedimentation and diagenesis.

**The Chemical Remanent Magnetization (CRM)** is acquired during the formation (precipitation of ferromagnetic minerals from a solution) of magnetic minerals within a rock, or the alteration of pre-existent magnetic minerals (Figure 14).



**Figure 14.** Modified after Tauxe [2005]; Grain growth CRM. a) Red beds of the Chinji Formation, Siwaliks, Pakistan. The red soil horizons have a CRM carried by pigmentary hematite. b) Initial state of non-magnetic matrix. c) Formation of superparamagnetic minerals with a statistical alignment with the ambient magnetic field (shown in blue).

Secondary component of remanent magnetization should be detected and rejected (as the Isothermal and Viscous Remanent Magnetizations - IRM and VRM).

For the interpretation it is critical to isolate the primary and secondary magnetizations and evaluate rocks post-depositional tilting, or folding and other geological processes overprinting the original component of remanent magnetization.

For further details, see reference books by *Butler* [1998], *Lanza and Meloni* [2006] and *Tauxe* [2009].

## 2.1. GEOLOGICAL AND TECTONIC APPLICATIONS

Since decades [e.g. *Torsvik et al.*, 2008, 2012; *Hernandez Moreno et al.*, 2014 and reference therein] the paleomagnetism has played a central role in the solution of tectonic problems from lithospheric plates scale (paleogeographic reconstructions, poles migrations, etc.) to small scale, detecting tectonic motions as rotations with respect to a reference paleomagnetic pole.

The paleomagnetism is the only tool that permits to measure the intensity H, inclination I and declination D of the magnetic field, and therefore deduce all the components of a remanent magnetization of rocks, necessary to detect vertical-axis rotation and latitudinal motion of crustal blocks.

To determine the expected paleomagnetic direction for rocks of any age at any point on Earth surface and then the motions of crustal blocks with respect to the rotation axis, we need to know reference poles. If the measured paleomagnetic declination and inclination are different with respect to those expected, in relation to the geographic position and the age of the analyzed rocks, tectonic rotations can be inferred.

There are two basic methods of analyzing vertical-axis rotations and latitudinal motions from paleomagnetic directions, the direction-space and pole-space approaches. These methods have been developed by *Beck* [1976, 1980], *Demarest* [1983], and *Beck et al.* [1986].

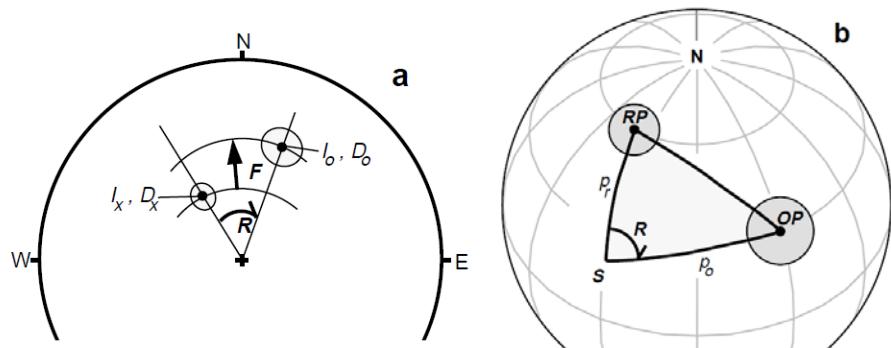
In this work, the direction-space approach is considered, where the observed paleomagnetic direction for a particular site ( $I_o, D_o$ ) is simply compared with the expected direction ( $I_x, D_x$ ) obtained on rocks of the same age from the stable part of the continent (Figure 15). The inclination flattening or latitudinal motion, F, and the rotation, R, of declination is given by, respectively:

$$F = I_x - I_o$$

$$R = D_o - D_x$$

$R$  is defined as positive when  $D_o$  is clockwise with respect to  $D_x$ .

Both the expected and observed directions have associated confidence cones ( $\alpha_{95}$ ), so  $F$  and  $R$  have 95% confidence limits  $\Delta F$  and  $\Delta R$ , respectively. Results of direction-space analyses are usually reported by listings of  $R \pm \Delta R$  and  $F \pm \Delta F$ . A significant positive flattening of inclination,  $F \pm \Delta F$ , indicates motion toward the paleomagnetic pole. To a complete explanation of the confidence limit mathematical development refer to *Demarest* [1983] and *Butler* [1992].



**Figure 15.** Modified after *Butler* [1992]. a) Equal-area projection of an observed discordant paleomagnetic direction, inclination  $I_o$  and declination  $D_o$ , compared to an expected direction, inclination  $I_x$  and declination  $D_x$ . The observed direction is shallower than the expected direction by the flattening angle  $F$ . Observed declination is clockwise from the expected declination by the rotation angle  $R$ . b) Comparison of observed and reference paleomagnetic poles. The discordant paleomagnetic pole  $OP$  (observed pole) was determined from paleomagnetic analysis of rocks at the collection location labeled  $S$ ;  $RP$  is the reference paleomagnetic pole; the spherical triangle with apices at  $S$ ,  $OP$ , and  $RP$  is shown by the heavy lines;  $p_r$  = great circle distance from  $S$  to  $RP$ ;  $p_o$  = great circle distance from  $S$  to  $OP$ ; poleward transport  $p$  =  $p_o - p_r$ ; vertical-axis rotation  $R$  = angle of spherical triangle at  $S$ .

This analytical approximation can be used when the 95% confidence cone of the paleomagnetic direction is less than  $10^\circ$  and the mean inclination is not too close to vertical [*Clark and Morrison*, 1983; *Demarest*, 1983]. For  $I > 80^\circ$  the paleomagnetic declination is poorly defined.

## 2.2. MAGNETIC ANISOTROPY

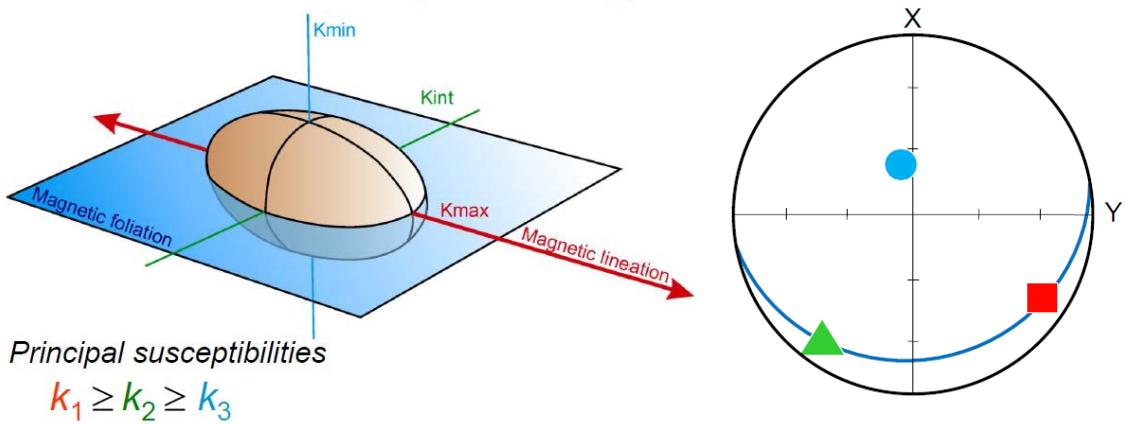
Other important structural information may also come from the **magnetic anisotropy** studies, a field also widely investigated in the last forty years [*Janák, 1967; Hrouda and Janák, 1976; Jelinek, 1977, 1978; Borradaile, 1981, 1987, 1988; Hrouda, 1982; Borradaile and Tarling, 1984; MacDonald and Ellwood, 1987; Hrouda and Jelinek, 1990; Parés et al., 1999*].

Rocks in which intensity of induced or remanent magnetization depends on direction of the applied magnetic field gain a magnetic anisotropy. One type of magnetic susceptibility (**K**, which represents the ratio between the induced magnetization **M** and the applied magnetic field **H**) is the Anisotropy of Magnetic Susceptibility (AMS), in which susceptibility is a function of direction of the applied field.

Measurement of the AMS is the technique used to determine the preferred orientation of fine-grained minerals in sediments where no mineral fabric is appreciable to the naked eye (“**magnetic fabric**”, eg. *Owens and Barnford, 1976; Goldstein, 1980; Hrouda, 1982; Borradaile, 1988; Lowrie, 1989; Pearce and Fueten, 1989; Jackson and Tauxe, 1991; Rochette et al., 1992; Sagnotti and Speranza, 1993; Tarling and Hrouda, 1993; Borradaile and Henry, 1997; Bouchez, 1997; Mattei et al., 1997; Cifelli et al., 2005; Pares et al., 2015; Caricchi et al., 2016*).

The fabric is the result of various forces acting during the formation and eventual geologic history of the rock: mainly gravity, Earth’s magnetic field, hydrodynamic forces and tectonic stress.

The AMS can be geometrically defined as a second-rank tensor, commonly expressed in terms of a **triaxial ellipsoid** with its own orientation defined by the directions of the three main axes (Figure 16):  $K_1 \geq K_2 \geq K_3$ , maximum, intermediate, and minimum susceptibility.



**Figure 16.** The magnetic susceptibility ellipsoid. M maximum (or  $\mathbf{K}_1$ - Magnetic Lineation), K intermediate (or  $\mathbf{K}_2$ ), K minimum (or  $\mathbf{K}_3$  - Magnetic Foliation). Modified after *Chadima and Jelinek* [2009].

The AMS ellipsoid principal axes usually correspond to the strain ellipsoid axes, indicating that the magnetic fabric is a strain proxy, and suggesting that magnetic fabric measurements are significant with respect to the strain history of rocks [e.g. *Goldstein and Brown*, 1988].

Several parameters are used for the quantification of the magnitude of anisotropy and for defining the shape of the AMS ellipsoid [Figure 17; *Jelinek*, 1981; *Hrouda*, 1982; *Tarling and Hrouda*, 1993; *Winkler et al.*, 1997].

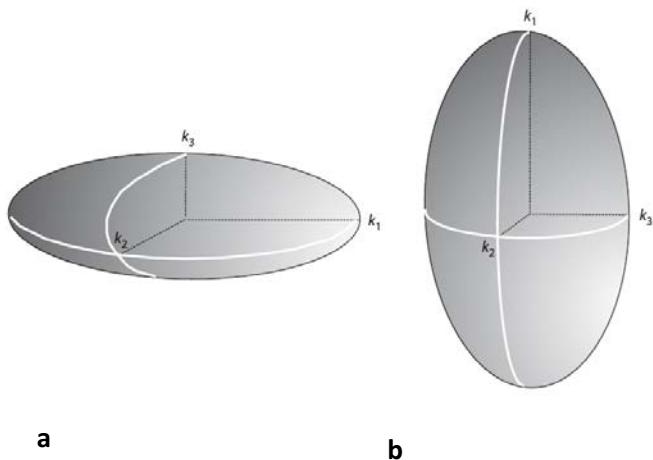
Parameter name	Parameter symbol	Parameter formula	Author
Anisotropy degree	$P$	$k_1/k_3$	Nagata (1961)
Corrected anisotropy degree	$P'$	$\exp \sqrt{2[(\eta_1 - \eta)^2 + (\eta_2 - \eta)^2 + (\eta_3 - \eta)^2]}$	Jelinek (1981)
Magnetic lineation	$L$	$k_1/k_2$	Balsley and Buddington (1960)
Magnetic foliation	$F$	$k_2/k_3$	Stacey <i>et al.</i> (1960) Balsley and Buddington (1960)
Ellipsoid shape	$T$	$2(\eta_2 - \eta_3)/(\eta_1 - \eta_3) - 1$	Jelinek (1981)
Mean susceptibility	$k_m$	$(k_1 + k_2 + k_3)/3$	Nagata (1961)

$$\eta_1 = \ln k_1, \eta_2 = \ln k_2, \eta_3 = \ln k_3, \eta = (\eta_1 + \eta_2 + \eta_3)/3.$$

**Figure 17.** Most commonly used AMS parameters (SI units). Modified after *Winkler et al.*, [1997], see references therein.

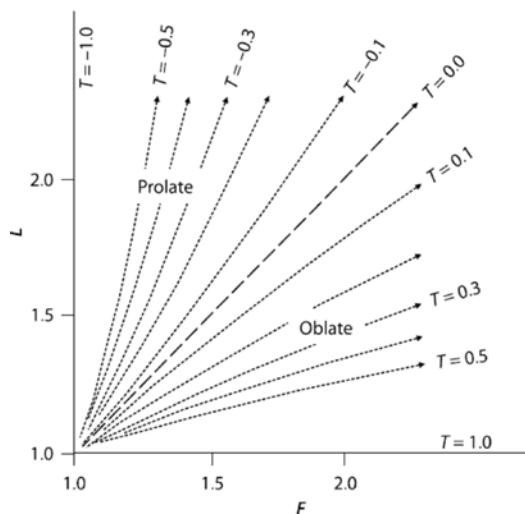
The magnetic lineation (L) and foliation (F) describe the shape of ellipsoid with a geometrical meaning: the lineation corresponds to the direction of K<sub>1</sub>, the foliation to the plane being defined by the directions of K<sub>1</sub> and k<sub>2</sub> and hence orthogonal to k<sub>3</sub>.

When the ellipsoid is prolate ( $k_1 > k_2 = k_3$ ) L prevails, and when it is oblate ( $k_1 = k_2 > k_3$ ) F prevails (Figure 18-19).



**Figure 18.** Shape of AMS ellipsoids;  
**a)** oblate (= planar fabric);  
**b)** prolate (= linear fabric).  
 Modified after *Lanza and Meloni, [2006]*

A more detailed evaluation of the shape of the ellipsoid is given by the Jelinek shape parameter, T, for which  $-1 \leq T \leq 0$  corresponds to prolate ellipsoids, and  $0 < T \leq 1$  to oblate ellipsoids [Jelinek, 1981; Figure 19].



**Figure 19.** Flinn diagram [Flinn, 1962].  
 Shape parameter T as a function of lineation (L) and foliation (F). Arrows point towards increasing degree of anisotropy (P). Modified after Tarling and Hrouda [1993].

One of the main breakthroughs in the last decade has been the wide recognition of a **specific magnetic mineralogy related to AMS**. In particular, there are important mineral sources of susceptibility that are not carriers of NRM (ferromagnetic minerals *sensu lato*) such as the diamagnetic, paramagnetic, and anti-ferromagnetic minerals, referred to as matrix minerals because they constitute the main volume fraction of common rocks [see detail in Figure 20; *Owens and Barnford, 1976; Borradaile, 1987; Rochette, 1987*]. The paramagnetic minerals, often play a major role in the magnetic susceptibility of rocks [*Rochette, 1987*], different from diamagnetic compounds like quartz, calcite or water which do not vary to a great extent, having a mean value of about  $-14 \times 10^{-6}$  SI.

<i>Mineral</i>	<i>Symmetry</i>	<i>Type</i>	<i>K<sub>m</sub></i>	<i>P</i>	<i>Reference</i>
<i>Diamagnetic</i>					
Quartz	0	...	-14.5	<1.01	1,2,13
<i>Paramagnetic</i>					
Calcite	1c	<i>N(C)</i>	-13.0	1.13	1,2
Biotite	1c	<i>N(CS)</i>	1-3	1.35	2,3,4,5
Other					
phyllosilicates	1c	<i>N(CS)</i>	0.05-1	1.2-1.35	2,3,5
Pyroxenes	4	<i>N(S)?</i>	0.5-5	1.2-1.4	2
Amphiboles	4	<i>N(S)?</i>	0.5-5	1.08-1.3	2
Riebeckite	1a	?	2.6	1.16	0
Orthoferrosilite	1a	?	5	1.21	14
Staurolite	2b	?	0.8	1.06	0
Garnet	0	...	3	1.001	0
Tourmaline	1c	<i>I(S)</i>	0.9	1.12	0
Cordierite	1c	<i>I(S)</i>	0.6	1.15-1.31	0
Siderite	2c	<i>I(C)</i>	3.8-4.2	1.7	6
Other Fe					
carbonates	2c	<i>I(C)</i>	0.05-0.7	1.08-1.45	6
<i>Ordered</i>					
Goethite	1c	<i>I(S)</i>	1.3-5	2?	7,8,9
Hematite	3	<i>N(CS)</i>	2-50	2.5-100	2,9,10
Pyrrhotite	3	<i>N(CS)</i>	50-300	>100	9,11
Magnetite MD	4	<i>N(S)</i>	$\leq 3000$	<5	2,12
Magnetite SD	3	<i>I(S)</i>	$\leq 1500$	$\infty?$	12
Magnetite SP	4	<i>N(S)?</i>	$\leq 5000$	...	12

*K<sub>m</sub>* is in  $10^{-3}$  SI, except for the diamagnetic minerals ( $10^{-6}$ ). Symmetry code: 0, isotropic; 1, uniaxial oblate; 2, uniaxial prolate; 3, triaxial oblate; 4, triaxial prolate. For uniaxial, the symmetry is indicated by the crystallographic axis of revolution. Type code: normal *N* or inverse *I* with mechanism of preferred orientation either controlled by shape *S* or by intracrystalline gliding during ductile deformation *C*. References: 0, unpublished or *Rochette [1988b]*; 1, *Rochette [1987]*; 2, *Hrouda [1982]*; 3, *Ballet [1979]*; 4, *Zapletal [1990]*; 5, *Borradaile et al. [1987]*; 6, *Rochette [1988a]*; 7, *Hedley [1971]*; 8, *Rochette and Fillion [1989]*; 9, *Dekkers [1988]*; 10, *Dunlop [1971]*; 11, *Rochette [1988a]*; 12, *Maher [1988]*; 13, *Hrouda [1986]*; 14, *Wiedenmann et al. [1986]*.

The anti-ferromagnetic matrix behaviour exemplified by goethite and hematite, leads to a linear susceptibility smaller than the paramagnetic susceptibility. According to *Neel and Pauthenet [1952]*, at this point it is worth mentioning a difficulty arising from the present definition of matrix susceptibility: hematite contributes to both the matrix component, with an isotropic anti-ferromagnetic susceptibility, and the ferromagnetic component which is alone responsible for the anisotropy.

**Figure 20.** Modified after *Rochette et al. [1992]*, see references therein. Selected AMS Data for rock-forming Minerals.

In weakly deformed (non-metamorphic) sedimentary rocks, AMS reflects the pristine fabric produced during incipient deformation at the time of, or shortly after, deposition and diagenesis of the sediment [Sintubin, 1994; Mattei *et al.*, 1995, 1997; Sagnotti *et al.*, 1998, 1999; Parés *et al.*, 1999; Coutand *et al.*, 2001; Cifelli *et al.*, 2004, 2005, 2009; Soto *et al.*, 2009].

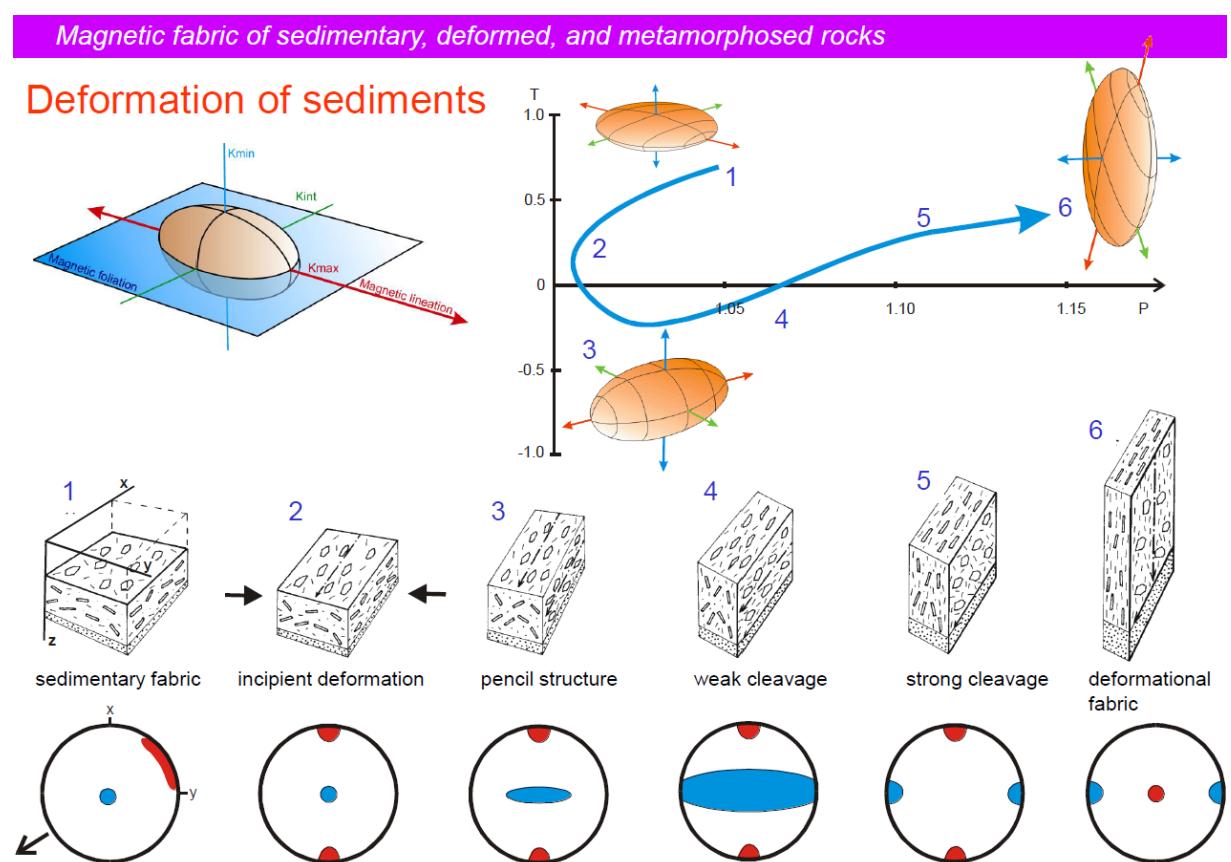
AMS analysis of weakly deformed sediments, have frequently been used in orogenic settings to document the syn-sedimentary tectonic regime [Kissel *et al.*, 1986; Mattei *et al.*, 1995, 1997, 1999; Sagnotti and Speranza, 1993; Sagnotti *et al.*, 1998; Parés *et al.*, 1999, 2015; Maffione *et al.*, 2008, 2012, 2015; Macrì *et al.*, 2014; Cifelli *et al.*, 2013, 2015; Caricchi *et al.*, 2016].

During deposition, sedimentary rocks acquire the so called '**sedimentary fabric**' characterized by the  $k_{\max}$  and  $k_{\text{int}}$  axes dispersed within a plane (magnetic foliation, the orthogonal plane to the minimum magnetic susceptibility direction) that is sub-parallel to the stratification plane. This sedimentary fabric can be partially overprinted by a '**tectonic fabric**' during incipient deformation [e.g., Parés *et al.*, 1999; Alimohammadian *et al.*, 2013]. The result of this process is the development of a magnetic lineation (the direction of maximum susceptibility in a rock) whereby  $k_{\max}$  aligns parallel to the maximum axis of stretching ( $\varepsilon_1$ ), hence perpendicular to the maximum axis of compression ( $\sigma_1$ ). This mechanism allows a direct correlation between the AMS and strain ellipsoids [e.g., Parés *et al.*, 1999, 2015].

A series investigations of correlation between the AMS fabric and structural observations, confirmed that in **compressional settings** the magnetic lineation is usually sub-horizontal and parallel to fold axes or local thrusts strikes [e.g. Borradaile and Henry, 1997; Mattei *et al.*, 1997; Sagnotti and Speranza, 1993; Mattei *et al.*, 1997; Sagnotti *et al.*, 1998; Maffione *et al.*, 2008], whereas it coincides with the stretching direction in extensional basins [e.g. Sagnotti *et al.*, 1994; Cifelli *et al.*, 2005; Oliva-Urcia *et al.*, 2009]. In **extensional settings**, the magnetic

lineation coincides with the local dip of the bedding, and is therefore perpendicular to the local normal fault planes [Sagnotti *et al.*, 1994; Mattei *et al.*, 1997, 1999; Cifelli *et al.*, 2004, 2005; Maffione *et al.*, 2012; Porreca and Mattei, 2012]. Few attempts of using AMS analyses in strike-slip tectonic settings have been done in the past [e.g., Cifelli *et al.*, 2013; Ferré *et al.*, 2002].

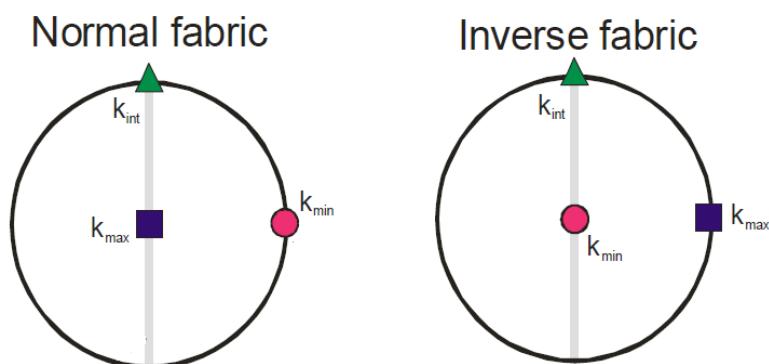
Increasing deformation progressively modifies the shape of the AMS ellipsoid from a pure sedimentary fabric (oblate ellipsoid:  $k_{\max} \approx k_{\text{int}} \gg k_{\min}$ ), to a sedimentary fabric with a marked tectonic imprint (triaxial ellipsoid:  $k_{\max} > k_{\text{int}} > k_{\min}$ ), to a tectonic fabric (prolate ellipsoid:  $k_{\max} \gg k_{\text{int}} \approx k_{\min}$ ), and eventually returning during the highest strain to an oblate ellipsoid with the magnetic foliation parallel to the cleavage/schistosity [e.g., Parés, 2004, 2015] (Figure 21).



**Figure 21.** Magnetic fabric evolution during progressive deformation. The ellipsoid is oblate (flattened) when  $K_1 \gg K_2$  but  $K_2 > K_3$ , and prolate (cigar-shaped) when  $K_1 > K_2$ ,  $K_2 \gg K_3$ . Modified after Graham [1966] Pares *et al.* [1999].

In the last stage of deformation, which corresponds to incipient metamorphism, the pristine tectonic fabric developed during the initial (syn-sedimentary) phases of deformation is completely obliterated. Conversely, the pristine tectonic fabric is not easily overprinted by small strains at low temperature [e.g., *Borradaile*, 1988; *Sagnotti et al.*, 1994, 1998; *Cifelli et al.*, 2004, 2005; *Parés*, 2004; *Soto et al.*, 2009]. This implies that, compared to classical structural geological analysis, where the definition of the age of deformation requires additional constraints (i.e., crosscutting relationships or unconformities), the finite strain determined from AMS analyses of weakly deformed rocks is a direct and powerful tool that can be used to study the deformation active at the time of sedimentation.

Sometimes, unusual relationships between structural and magnetic axes, consisting in the “exchange” of the maximum and minimum anisotropy axes (Figure 22, so called *inverse magnetic fabric*) can occur because of the presence of certain magnetic minerals, either SD magnetite, iron-bearing carbonates or various paramagnetic minerals like tourmaline, cordierite, goethite or siderite [*Rochette et al.*, 1992]. Rocks with fine-grained magnetite, are particularly prone to anomalous AMS fabric. Therefore it is strongly recommended to investigate the mineralogical source of AMS and to compare AMS with other types of anisotropy or mineralogical investigation. Accordingly, in the past, authors strongly recommended to investigate the mineralogical source of AMS and to compare AMS with other data of different types of anisotropy or mineralogical investigation.



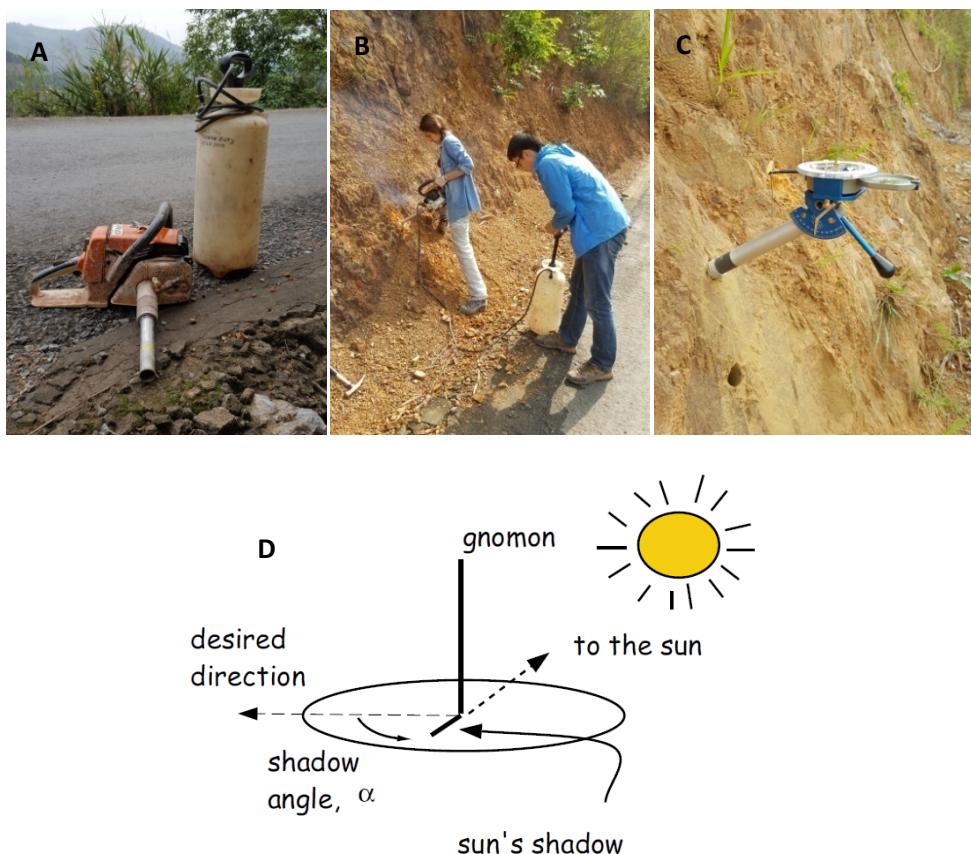
**Figure 22.**  
Example of Inverse  
Fabric projection.

*Chapter*  
***III***

### 3. SAMPLING AND LABORATORY METHODS

During this research activity, two field campaigns for sample collection were carried out on April 2016 and 2017, providing a total number of 930 core samples. The samples were collected following a classic paleomagnetic sampling method: several sites have been selected in the same rock unit and several samples in the same site (ca. 10 on average). The samples include volcanic rocks (n. 17 basalt sites) and sedimentary rocks (n. 5 whitish siltstones sites and n. 72 continental red beds).

Sampled sites were aligned along transects cutting the main strike-slip faults considering that several studies demonstrated that shear-related rotations virtually end within 10–20 km from the fault trace [Ron *et al.*, 1984; Sonder *et al.*, 1994; Piper *et al.*, 1997; Randall *et al.*, 2011; Hernandez-Moreno *et al.*, 2014, 2016].

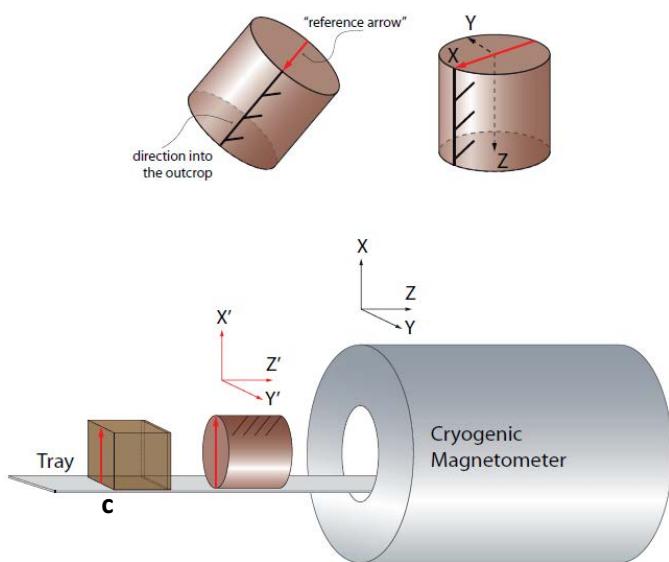
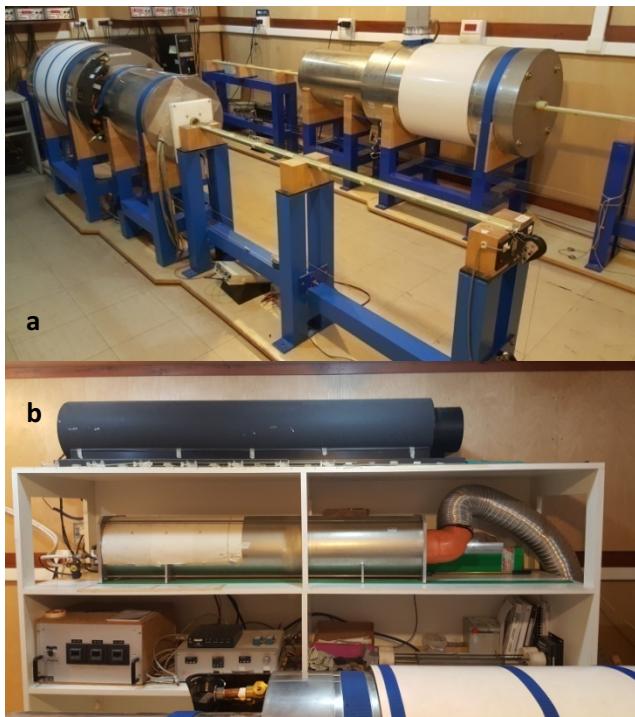


**Figure 23.** A) Portable petrol-powered drill; B) paleomagnetic sampling; C) Pomeroy orientation device in use as a sun magnetic compass; D) Schematic of the principles of sun compass orientation (by Tauxe, 2005).

The samples were drilled using a portable petrol-powered and water-cooled drill (Figure 23). Cores were oriented *in situ* before extraction using a magnetic compass, corrected for the local magnetic declination for year 2016/2017 (between 0° and 0.9° W according to NOAA's National Geophysical Data Center, <http://www.ngdc.noaa.gov/geomag/declination.shtml>) and, when possible, the sun. Afterward, the cores were cut into standard cylindrical paleomagnetic specimens of 22 mm height.

Paleomagnetic measurements were later carried out in the shielded room of the paleomagnetic laboratory of the Istituto Nazionale di Geofisica e Vulcanologia (Rome), using a 2G Enterprises DC-superconducting quantum interference device cryogenic magnetometer (Figure 24).

Neogene volcanic samples were demagnetized in 10 steps by alternating field (AF) yielded by three orthogonal coils in-line with the magnetometer up-to a maximum AF peak of 100 mT (Figure 24a). Red beds, Jurassic basalts and whitish siltstones samples were thermally demagnetized using a Pyrox shielded oven (Figure 24b) in 12 steps (20°, 200°, 300°, 350°, 400°, 440°, 480°, 520°, 560°, 600°, 640°, 680°C) of temperature up to 680°C. Demagnetization data were plotted on orthogonal vector component diagrams [Zijderveld, 1967].



**Figure 24. a-c)** View of the magnetically shielded room of the paleomagnetic laboratory (INGV, Rome) and schematic description of samples position into cryogenic magnetometer ; **b)** The Pyrox oven and its controller, to demagnetize up to 700 °C.

The magnetization components were identified by principal component analysis (PCA) [Kirschvink, 1980] using Remasoft 3.0, a freeware distributed by Agico [Chadima and Hrouda, 2007]. The site mean paleomagnetic directions were computed using Fisher [1953] statistics, and plotted on equal-angle projections. Finally, the rotation and flattening values with respect to Eurasia were evaluated according to Demarest [1983], using the reference paleopoles by Torsvik *et al.* [2012] for 170-140 Ma sites (Europe), and the most recent East Asia-focused poles by Cogné *et al.* [2013] for 130 to 10 Ma sites (Figure 25-26).

Age (Ma)			Palaeopole						
Window	Mean	(Myr)	N	N <sub>rot</sub>	N <sub>ign</sub> /N <sub>sed</sub>	Latitude (°N)	Longitude (°E)	k	A95(°)
0	1.0	20	5	0	0/5	89.0	76.6	200.7	5.4
10	12.4	20	21	5	13/8	87.5	257.3	154.8	2.6
20	18.2	20	20	7	16/4	84.7	255.9	112.5	3.1
30	23.7	40	25	10	17/8	83.9	259.6	106.7	2.8
40	41.7	40	14	7	8/6	81.9	260.8	94.8	4.2
50	55.0	40	13	2	6/7	81.6	238.8	51.6	5.9
60	65.2	40	22	5	11/11	81.9	247.5	57.4	4.1
70	77.3	40	45	5	17/28	79.7	219.5	42.6	3.3
80	79.7	40	44	4	19/25	79.5	216.9	42.2	3.4
90	87.8	40	53	5	23/30	77.9	212.8	46.2	2.9
100	105.1	20	16	6	8/8	82.2	205.4	118.1	3.5
110	109.8	20	22	6	8/14	80.8	199.5	78.9	3.5
120	121.6	20	41	9	15/26	80.6	192.9	60.4	2.9
130	126.2	20	44	14	16/28	79.9	193.5	70.0	2.6

Notes: Width: temporal window width; N: Number of data in the statistics; Nrot: number of rotated poles in the computations; Nign/Nsed: number of poles from igneous/sedimentary formations; k. A95: Fisher (1953) statistics parameters.

**Figure 25.** East Asia paleopoles from Cogné et al. [2013].

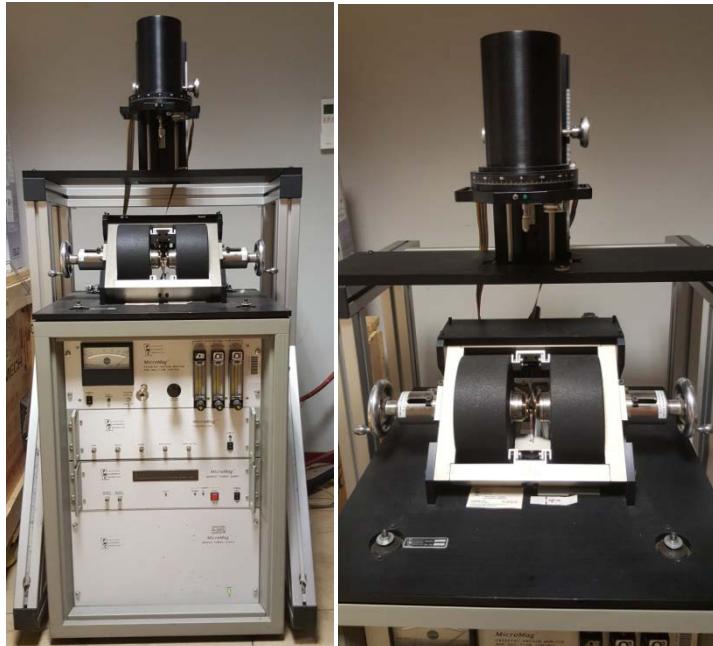
Age	N	A95	North America		Europe		India	
			Plat	Plon	Plat	Plon	Plat	Plon
0	24	1.9	-88.5	353.9	-88.5	353.9	-88.5	353.9
10	49	1.8	-86.4	342.2	-86.7	330.0	-87.2	60.4
20	31	2.6	-83.7	343.2	-84.4	332.1	-83.7	74.7
30	24	2.6	-82.1	338.7	-83.1	326.5	-79.7	101.7
40	24	2.9	-80.1	337.2	-81.1	324.3	-74.7	106.8
50	33	2.8	-76.4	354.5	-78.9	344.7	-65.1	98.4
60	44	2.1	-73.6	7.5	-78.2	352.6	-48.5	100.8
70	32	2.5	-73.5	12.6	-79.2	355.7	-36.4	100.7
80	25	2.9	-74.7	10.5	-79.7	357.9	-29.0	103.5
90	28	2.5	-76.8	4.8	-80.4	347.2	-20.9	111.4
100	14	3.3	-78.7	358.4	-80.8	332.3	-19.7	113.0
110	21	3.3	-75.2	21.7	-81.2	13.1	-11.1	115.9
120	28	2.6	-73.7	15.5	-79.0	10.1	-8.6	116.4
130	18	2.8	-70.9	5.4	-75.0	3.4	1.0	117.1
140	9	6.0	-67.9	5.7	-72.4	7.9	5.3	117.9
150	15	6.4	-72.6	339.4	-72.9	334.2	-2.9	122.3
160	19	5.1	-72.0	325.9	-70.5	323.4	-9.7	126.6
170	18	4.6	-70.1	314.8	-67.3	316.4	-10.7	125.8

**Figure 26.** Paleopoles from Torsvik et al. [2012].

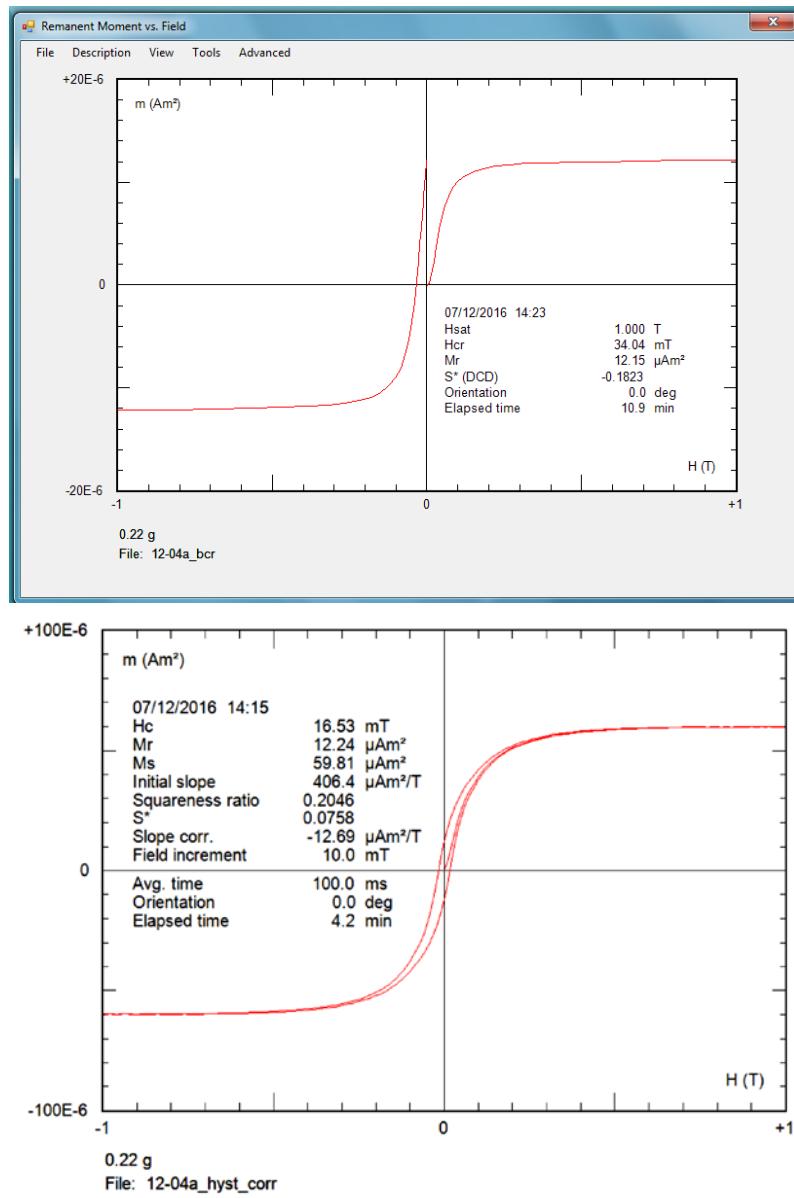
To define the sense and amount of rotation, have been always considered the smaller angle between the observed and expected declinations, thereby calculating rotation values always  $\leq 180^\circ$ . This is a conservative approach, although we are aware that in the past, some authors considered rotation values exceeding  $180^\circ$  [e.g. Hernandez-Moreno et al., 2014; Nelson and Jones, 1987; Nelson and Piper et al., 1997].

Moreover, for each sampling site, have been selected specimens for additional magnetic analyses carried out with the aim of characterizing the magnetic mineralogy.

For hysteresis measurements, the samples were crushed into powder and then placed in pharmaceutical gel caps #4 (corresponding to a filled volume of about 0.15 ml), in order to vibrate by means of a carbon fiber probe in the Princeton Measurement Corporation Micromag 3900 Vibrating Sample Magnetometer (VSM) (Figure 27). The software *MicroMagVSM* was used for hysteresis data analysis (Figure 28).



**Figure 27.** The Micromag magnetometer, for high-sensitivity hysteresis loops.



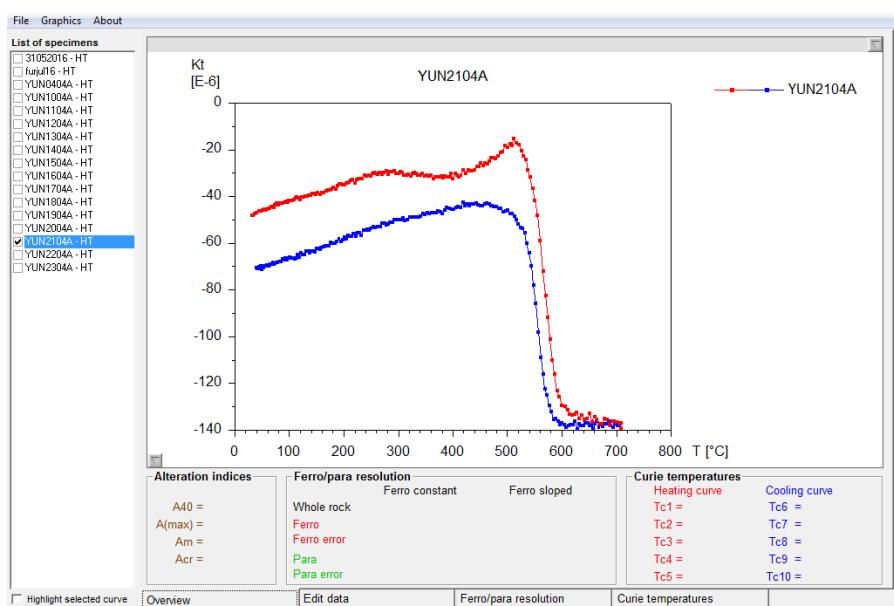
**Figure 28.** Graphic interface of *MicroMagVSM* program used for data visualization of hysteresis measurements

The coercive force (Bc), the saturation or maximum remanent magnetization (Mrs) and the saturation/maximum magnetization (Ms) were measured using the VSM under cycling in a maximum field of 1.0 T, and determined after subtracting the high field paramagnetic linear trend when saturated samples. The coercivity of remanence (Bcr) values have been extrapolated from the backfield remagnetization curves up to -1 T, following a forward magnetization in +1 T field. Bcr represents the negative field needed to remove the remanent magnetization after applying the maximum positive field. The saturation remanence to

saturation magnetization (Mrs/Ms) vs. the ratio of remanent coercive force to coercive force ( $B_{cr}/B_c$ ) has been plotted in a Day plot [Day *et al.*, 1977; Dunlop, 2002].

First order reversal curves (FORCs) have been measured using the Micromag operating software, and processed, smoothed and drawn with the FORCINEL Igor Pro routine [Harrison and Feinberg, 2008]. FORCs are a series of partial hysteresis loops made after the sample magnetization is saturated [Pike *et al.*, 1999; Roberts *et al.*, 2000]. Selected FORCs have been measured in steps of 2 mT with an averaging time of 100 ms; the maximum applied field was 1.0 T. The optimum smoothing factor was calculated by the FORCINEL software.

For one specimen from each basalt and red bed sites, have been also measured the variation of the low-field magnetic susceptibility ( $K$ ) during a heating and cooling cycle performed in air, from room temperature up to 700°C, using an AGICO KLY-3 Kappabridge coupled with a CS-3 furnace (Figure 29). The Curie/Néel points have been estimated as the temperature, or range of temperatures, at which the paramagnetic behavior starts to dominate, following the approach outlined by Petrovsky and Kapicka [2006]. Cureval8 software was used for data analysis (Figure 29).

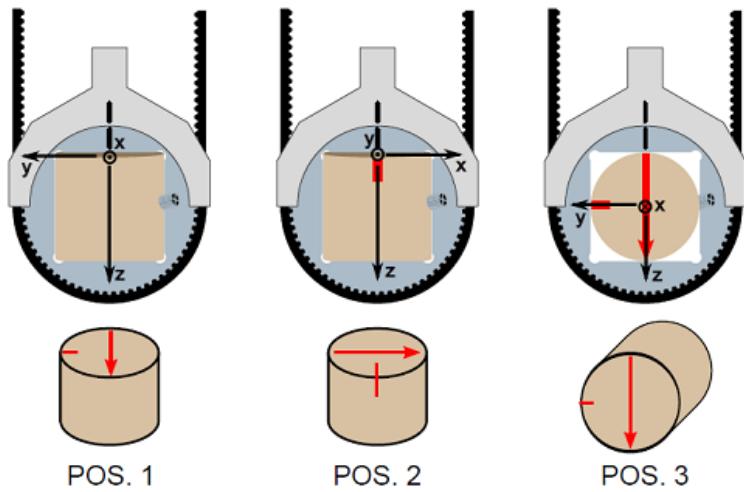


**Figure 29.**  
Graphic interface of Cureval8 program used for data visualization of thermal curves.

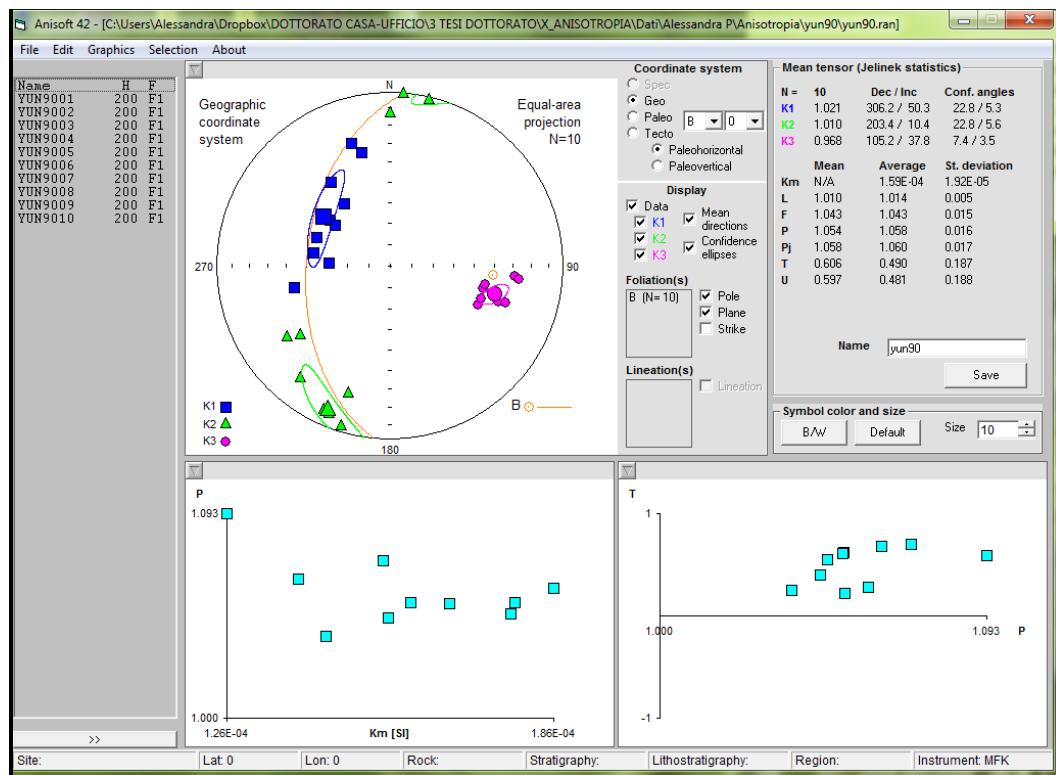
Analysis of low-field anisotropy of magnetic susceptibility (AMS) was done using a MFK1 Kappabridge (AGICO) (Figure 30-31). During measurement, the specimen slowly rotates subsequently along three perpendicular axes (total of 64 measurements are made during one spin). The AMS parameters were evaluated using Jelinek statistics [Jelinek, 1977, 1978; see details in chapter II]. Anisoft42 software was used for data analysis [Chadima and Jelinek, 2009] (Figure 32).



**Figure 30.** AGICO MFK1-FA susceptibility meter, connected to the furnace for measuring the variation of susceptibility vs. temperature.



**Figure 31.** Three specimen positions for the automatic AMS measurements using the rotator.



**Figure 32.** Graphic interface of the program used for data visualization AMS analysis [Chadima and Jelinek, 2009].

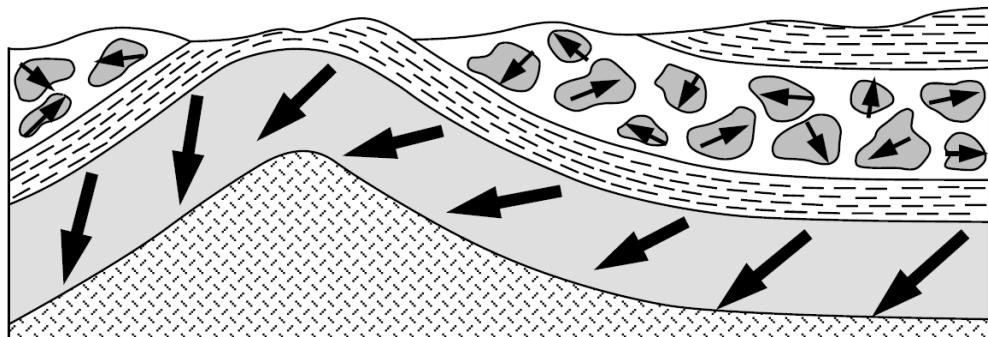
Finally, have been also used a tool to apply a pulse magnetic field up to 2.7 T, in order to impart an isothermal remanent magnetization. This method, called the "**Three Axis Method**" or the "**Lowrie Method**" [Lowrie, 1990], applies to a sample three decreasing fields along three orthogonal directions (2700 mT, 600 mT and 120 mT respectively from z, y and x axes) using a pulse magnetizer (Figure 33). Then the sample is thermally demagnetized by increasing temperature steps. From the analysis of the demagnetization curves along the three x, y and z components, it's possible to discriminate the Curie (or Néel) temperatures of distinct ferromagnetic fractions of analyzed sample and characterized by specific coercivity spectrum.



**Figure 33.**  
The 2G pulse magnetizer, the  
field can reach up to 2.7 T

### 3.1 THE FOLD TEST

With the fold test (or bedding-tilt test) can be evaluated relative timing of acquisition of a component of NRM (usually ChRM) and folding. If a ChRM was acquired prior to folding, directions of ChRM from sites on opposing limbs of a fold are dispersed when plotted in geographic coordinates (*in situ*) but converge when the structural correction is made (“restoring” the beds to horizontal). The ChRM directions are said to “pass the fold test” if clustering increases through application of the structural correction or “fail the fold test” if the ChRM directions become more scattered. The fold test can be applied either to a single fold or to several sites from widely separated localities at which different bedding tilts are observed (Figure 34). The fold test is used to understand when the magnetizations were acquired with respect to the tilting process [Butler, 1992].



**Figure 34.** Modified after Butler [1992]. Schematic illustration of the fold and conglomerate tests of paleomagnetic stability. Bold arrows are directions of ChRM in limbs of the fold and in cobbles of the conglomerate; random distribution of ChRM directions from cobble to cobble within the conglomerate indicates that ChRM was acquired prior to formation of the conglomerate; improved grouping of ChRM upon restoring the limbs of the fold to horizontal indicates ChRM formation prior to folding. Redrawn from Cox and Doell [1960].

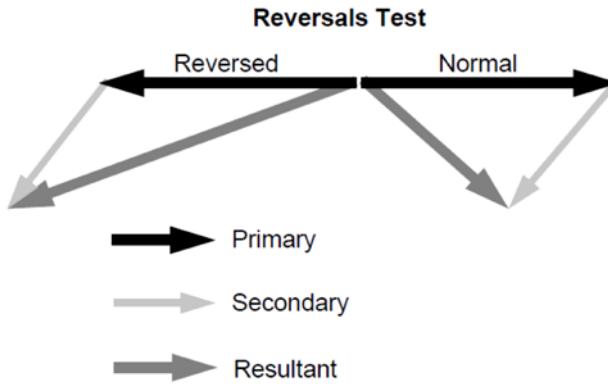
### 3.2 THE REVERSAL TEST

The time-averaged geocentric axial dipolar nature of the geomagnetic field holds during both normal- and reversed-polarity intervals.

At all locations, the time-averaged geomagnetic field directions during a normal-polarity interval and during a reversed-polarity interval differ by  $180^\circ$ . This property of the geomagnetic field is the basis for the reversals test of paleomagnetic stability. If a set of paleomagnetic sites affords adequate averaging of secular variation during both normal and reversed polarity intervals, the average direction of primary NRM for the normal polarity sites is expected to be antiparallel to the average direction of primary NRM for the reversed-polarity sites [Butler, 1992]. However, acquisition of later secondary NRM components will cause resultant NRM vectors to deviate by less than  $180^\circ$ . ChRM directions are said to “**pass the reversals test**” if the mean direction computed from the normal-polarity sites is antiparallel to the mean direction for the reversed-polarity sites. Passage of the reversals test indicates that ChRM directions are free of secondary NRM components and that the time sampling afforded by the set of paleomagnetic data has adequately averaged geomagnetic secular variation. Furthermore, if the sets of normal- and reversed-polarity sites conform to stratigraphic layering, the ChRM is probably a primary NRM.

If a paleomagnetic data set “**fails the reversals test**”, the average directions for the normal and reversed polarity sites differ by an angle that is significantly less than  $180^\circ$ . Failure of the reversals test can indicate either: (1) presence of an un-removed secondary NRM component or (2) inadequate sampling of geomagnetic secular variation during either (or both) of the polarity intervals. Because polarity reversals are characteristic of most geologic time intervals, paleomagnetic data sets often contain normal- and reversed-polarity ChRM (Figure 35).

The reversals test of paleomagnetic stability is often applicable and, unlike the conglomerate or fold test, does not require special geologic settings. Quantitative evaluation of the reversals test involves computation of the mean directions (and confidence intervals about those mean directions) for both normal and reversed-polarity groups and comparison of one mean direction with the antipode of the other mean direction [Butler, 1992].



**Figure 35.** Modified after Butler [1992]. Schematic illustration of the reversals test of paleomagnetic stability. Solid arrows indicate the expected anti-parallel configuration of the average direction of primary NRM vectors resulting from magnetization during normal- and reversed- polarity intervals of the geomagnetic field; an unremoved secondary NRM component is shown by the lightly stippled arrows; the resultant NRM directions are shown by the heavily stippled arrows. Redrawn from McElhinny [1973].



*Chapter*

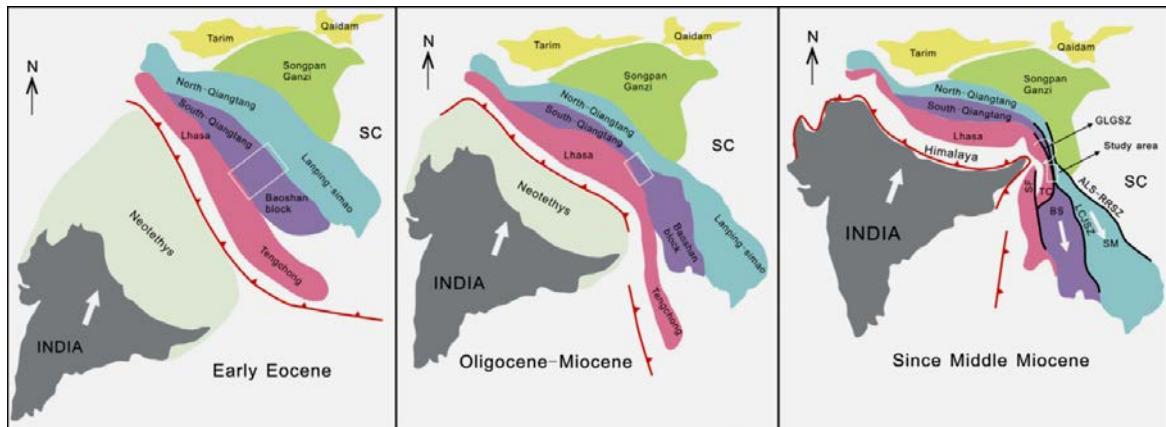
*IV*

#### 4. GEOLOGICAL SETTING

##### EASTWARD DRIFT OF TIBET AND TECTONIC DEFORMATION OF EAST ASIA

Asia formed during the Phanerozoic by the welding of several continental blocks including Siberia, Tarim, North China, South China, Indochina, India, and others small-sized micro continents. All published models acknowledge subduction, and accretion of Gondwana derived fragments drifting to the North, but in many cases the timing and nature of collision is poorly defined.

Since early Cenozoic times, the northward drift of Indian plate and its progressive indentation into Asia [Figure 36; *Tapponnier and Molnar, 1977*] is amazingly producing spreading and crustal deformation in SE Asia [*Leloup et al., 2001; Morley, 2007; Roger et al., 1995; Searle, 2006; Tapponnier et al., 1986; Yin, 2010; Zhong et al., 1990; Aitchison et al., 2007*].



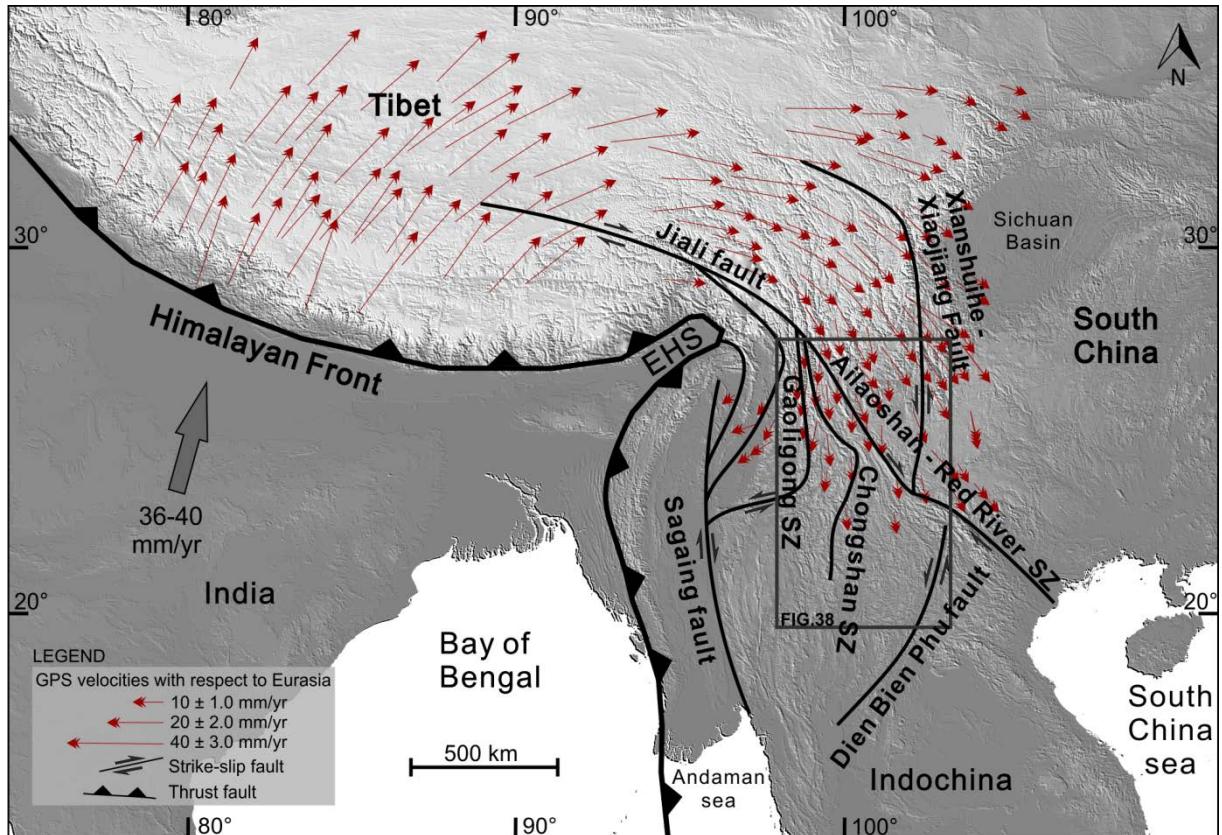
**Figure 36.** Tectonic reconstructions of the Himalayan–Tibet area since early Eocene. Modified by *Huang et al. [2015]* after *Replumaz and Tapponnier [2003]; Royden et al. [2008]; Pan et al. [2012]; Xu et al. [2013]*.

Plate collision yielded the 2000 km wide and 5000 m high Tibet Plateau, the most outstanding high elevation plateau of the Earth [*Besse et al., 1984; Najman et al., 2010; Searle et al., 2007*], edged by the Himalayan arc. Since the Oligocene/Miocene, the closure of Neotethys and consequent plate convergence and collision has been partly accommodated by the SE-ward lateral extrusion along major continental-scale-strike-slip shear zones in the South China

and Indochina blocks [*Tapponnier et al.*, 1990; *Leloup et al.*, 1995; *Wang E. et al.*, 1998; *Sato et al.*, 1999; *Wang and Burchfield*, 2000; *Molnar and Dayem*, 2010]. Geological data suggested about  $700\pm200$  km southward displacement of Indochina block during the late Paleogene to early Miocene time span [*Lacassin et al.*, 1996; *Leloup et al.*, 1995, 2001]. As clearly shown by GPS, seismological and geophysical data [*Maurin et al.*, 2010; *Liang et al.*, 2013; *Huang et al.*, 2017; *Zhu et al.*, 2017], Tibet is nowadays squeezed by the surrounding plates and is escaping towards the E-SE, likely floating above hot and ductile lower crust. The upper brittle crust of SE Tibet is, in fact, characterized by a set of curved strike-slip faults, running around the north-eastern edge of the Indian indenter (Eastern Himalayan Syntaxis, EHS; Figure 37, 38), which are inferred to concentrate the deformation between major crustal blocks [*Wang et al.*, 1998]. Identifying and determining the role and kinematic variability of the continental-scale strike-slip faults, responsible for accommodating a significant portion of the post-50 Ma India-Eurasia convergence, have great significance to understand the geologic evolution during Tibet Plateau growth [*Hall*, 2002; *Hall et al.*, 2008; *Houseman and England*, 1986, 1993; *Molnar and Dayem*, 2010; *Peltzer and Tapponnier*, 1988; *Royden et al.*, 2008; *Tapponnier et al.*, 1982, 1990; *Vilotte et al.*, 1986; *Wang and Burchfiel*, 2000; *Yin and Taylor*, 2011].

In the past, the ductile high temperature deformation of the Shear Zones, addressed by structural geology studies coupled with radiometric dating, traced an activity spanning a wide time window between 35 and 15 Ma (late Eocene to middle Miocene; *Leloup et al.*, 1993, 1995; *Searle et al.* 2006, 2007 and references therein). However, the apparent lack of post-15 Ma shear zone activity conflicts with both GPS evidence [Figure 37; *Liang et al.*, 2013], showing southward drift of Indochina at a  $\sim 2$  cm/yr rate, and significant seismic activity with mainly strike-slip focal mechanisms [*Socquet and Pubellier*, 2005].

Lower magnitude events are diffuse in the whole region, and focal mechanisms of the main events are available in the Harvard Global CMT catalog (<http://www.globalcmt.org/>) [Ekstrom et al., 2012].



**Figure 37.** Schematic tectonic map of the South Asia and surrounding areas generated by GIS- Geographic Information System (DEM source: Esri, User Community, Coordinate System & Projection: world geodetic system 1984-Web Mercator Auxiliary Sphere). The grey rectangle represents the study area (see Figure 38). Red thin arrows show present-day global positioning system velocities around the Tibetan Plateau relative to stable Eurasia [Liang et al., 2013]. EHS Eastern Himalayan Syntaxis; SZ shear zone.

There are two end-member models commonly used to interpret the deformation mechanisms of eastern and SE Tibet. The first is the ‘tectonic escape’ model: strain is localized on a limited number of major faults bounding rigid lithospheric mega-blocks, and is driven solely by stresses exerted at plate boundaries [Avouac and Tapponnier, 1993; Ji et al., 2000; Molnar and Tapponnier, 1975; Replumaz and Tapponnier, 2003; Tapponnier et al., 1982; Tapponnier

*and Molnar* 1976, 1977]; the second is the ‘crustal flow’ model: deformation is pervasive and driven, for a significant part, by buoyancy forces resulting from lateral variations of crustal thickness. Deformation of the crust is continuous, and there are no blocks separated by strike-slip faults [*Bird*, 1991; *Clark and Royden*, 2000; *England and Houseman*, 1986; *Houseman and England*, 1993; *Royden et al.*, 1997]. Additionally, *Wang and Burchfiel* [1997] were the first to notice that the extrusion was accompanied by strong internal deformation and rotation of smaller crustal fragments. *Clark et al.* [2006] and *Liu-Zheng et al.* [2008] noticed that Mesozoic-Cenozoic red beds in northern Indochina are tightly folded around NNW trending fold axes, and *Hall* [2002] also suggested a major shortening in N-NW Indochina block. Finally, *Li et al.* [2017a] proposed that this shortening is in part accommodated by large-scale block rotations.

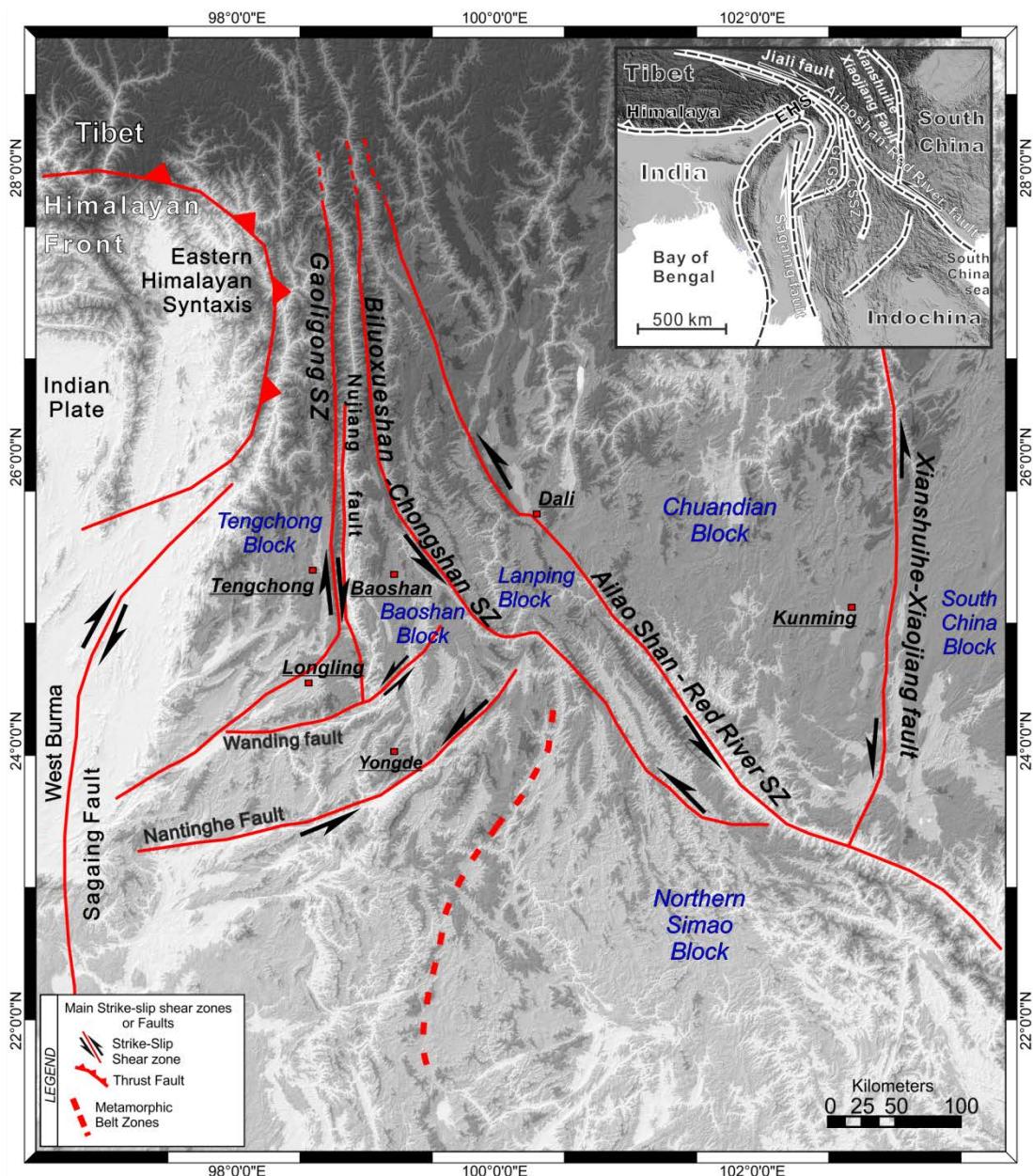
A key for evaluating the validity of the two models may lie in the investigation of block motion scale and crustal deformation style and timing, especially in the vicinity of large strike-slip faults [*Wang et al.*, 2001; *Zhang et al.*, 2004; *Gan et al.*, 2007; *Maurin et al.*, 2010; *Liang et al.*, 2013].

Paleomagnetism may represent a powerful tool for documenting and quantifying such tectonic data and this work will focus on this main goal.

#### 4.1 TECTONICS OF YUNNAN AND CHARACTERISTICS OF ITS MAJOR SHEAR ZONES

To the SE of Tibet, the Indochina block is thought to have been extruded south-eastward during the initial phases of India-Eurasia collision [Molnar and Tapponnier, 1975; Molnar *et al.*, 1988; Patriat and Achache, 1984; Replumaz and Tapponnier, 2003]. This convergence not only formed the Himalaya and the mountain system of the Central Asia, but also induced an intensive internal tectonic deformation within the Eurasia. Around the Eastern Himalayan Syntaxis (Figure 37, 38) and the area to the southeast, the large-scale strike-slip-dominated deformation zones curve from trending WNW-ESE at the eastern extremities of Tibetan plateau to trending NW-SE in northwestern Yunnan, China, and the rocks are strongly metamorphosed and deformed [Tapponnier *et al.*, 1990; Leloup *et al.*, 1993, 2007; Searle, 2006]. Within this area itself, four major ~N-S Shear Zone occur: the Sagaing, Gaoligong, Biluoqueshan-Chongshan (hereinafter referred as to Chongshan) and the Ailaoshan Red River Shear Zones [Akciz *et al.*, 2008; Leloup *et al.*, 1995; Lin *et al.*, 2009]. These faults separate different tectono-metamorphic mega-terrane (Figure 38), focus of this thesis, called Tengchong, Baoshan, Lanping–Simao and Chuandian blocks.

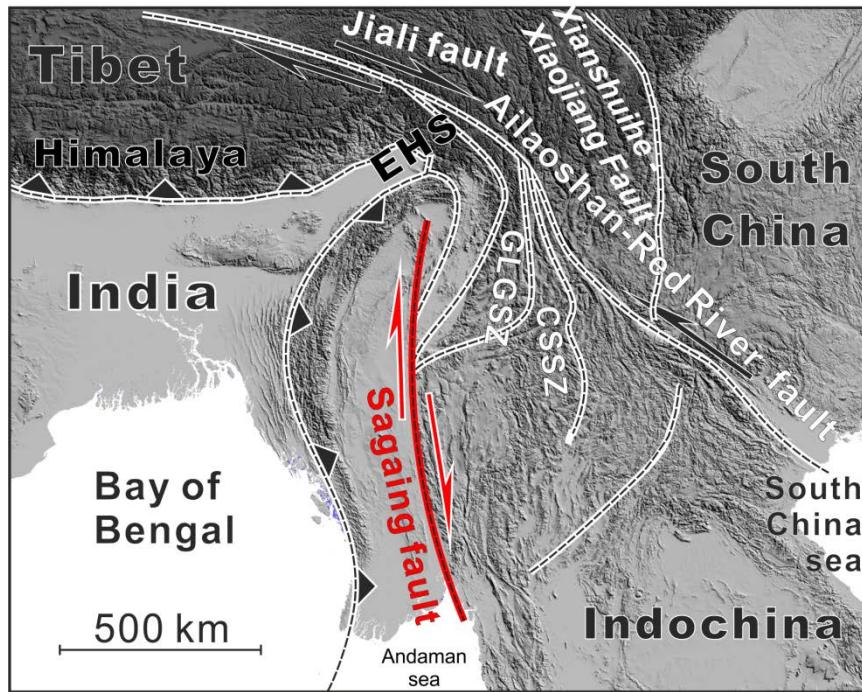
They are characterized by different lithologies, protolith ages, intracontinental deformation, paleogeography, and orogenic history, with additional internal deformation and variable vertical axis rotations [Kondo *et al.*, 2012; Li *et al.*, 2017; Sato *et al.*, 2007; Sato *et al.* 2001; Tanaka *et al.*, 2008; Tong *et al.*, 2013; Wang and Burchfiel, 1997; Metcalfe, 2006; BGMRYP, 1990; BGMRXZR, 1993].



**Figure 38.** Schematic tectonic map of the Yunnan area (China). Detail of the grey rectangle showed in Figure 37. Here are shown the different blocks (in blue) and the major strike-slip faults (in red).

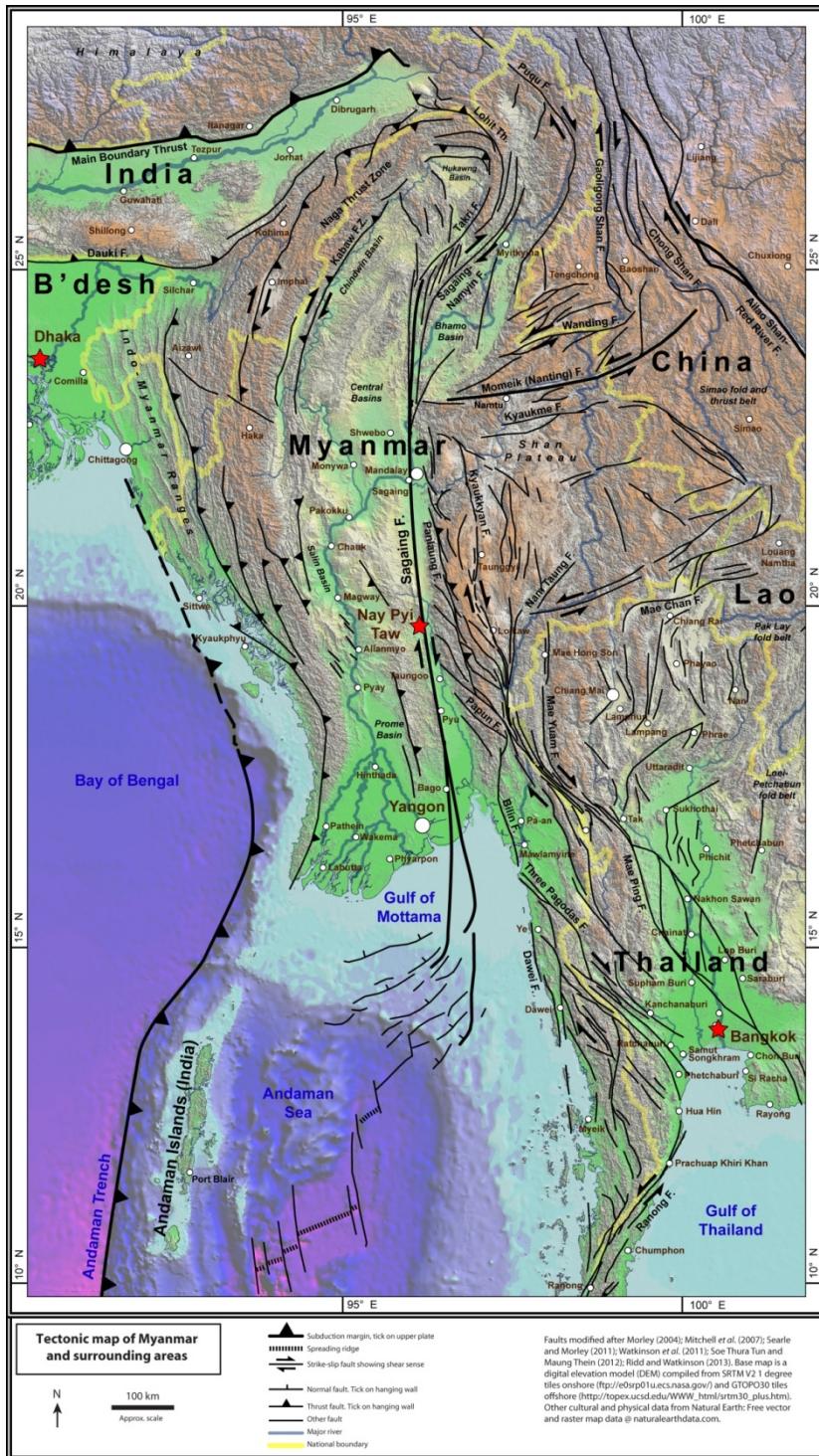
A brief description of the most important faults that characterize the Yunnan area, two of which are the subject of this study (Gaoligong and Ailaoshan Red River) follows.

## 4.2 THE SAGAING FAULT



**Figure 39.** Schematic tectonic map of South Tibet. In red The Sagaing Fault.

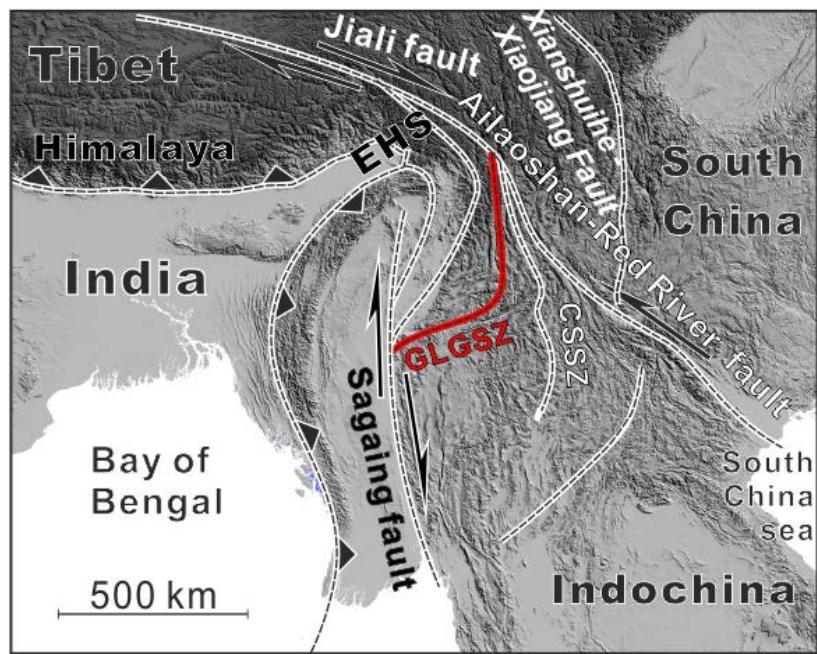
The Sagaing fault is interpreted as a recent dextral continental strike-slip fault which has long and straight traces for 1000 km along the entire length of Myanmar and separates the Indochina continental fragment from the Burma block [Figure 38-39; Akciz *et al.*, 2008; Lin *et al.*, 2009; Metcalfe, 2013; Replumaz and Tapponnier, 2003; Wang *et al.*, 2006]. The Sagaing fault is linked with Central Andaman spreading center to the south [Curray *et al.*, 1979]. The Andaman Sea has been formed by seafloor spreading along short ENE-striking spreading center that is offset by NNW striking transform faults [Figure 39, 40; Curray *et al.*, 1979]. Extension and rifting in the Central Andaman Basin began around 11 Ma ago and extension and sea floor spreading has been ongoing since 4-5 Ma [Khan and Chakraborty, 2005]. Spreading in a  $335^\circ$  ( $N25^\circ W$ ) direction, relative to present N, is at an average rate of 30 mm/yr the northward component is 27 mm/yr [Curray, 2005]. The average trend of the Sagaing Fault is  $351^\circ(N9^\circ W)$  [Soquet *et al.*, 2006].



Since its formation about 22 to 15 million years ago, during the Miocene epoch [Curray, 2005; Socquet and Pubellier, 2005; Searle *et al.*, 2007], the Sagaing Fault has accommodated about 330 km [Curray, 2005] to 460 [Mitchell, 1993] of dextral displacement between its eastern and western sides. It is characterized at present by several right-lateral focal mechanisms, particularly along its northern end [Ratschbacher *et al.*, 1996; Wang *et al.*, 1998; Xu *et al.*, 2015].

**Figure 40.** Tectonic features of Myanmar and surrounding areas (<http://www.sagaingfault.info/index.html>).  
Faults modified after Morley [2004].

#### 4.3 THE GAOLIGONG SHEAR ZONE



**Figure 41.** Schematic tectonic map of South Tibet. In red the Gaoligong Shear Zone (GLGSZ).

The Gaoligong Shear Zone extends eastward and southward from the Eastern Himalayan Syntaxis forming the boundary between the Tengchong and Baoshan blocks [Figure 38, 41; *Li et al.*, 2004; *Mitchell*, 1992, 1993; *Xu et al.*, 2012]. It is exposed completely along the Gaoligong Mountains west of the Nujiang river valley [*Wang and Burchfiel*, 1997; *Wang et al.*, 2008], where dextral ductile shear indicators are widely preserved.

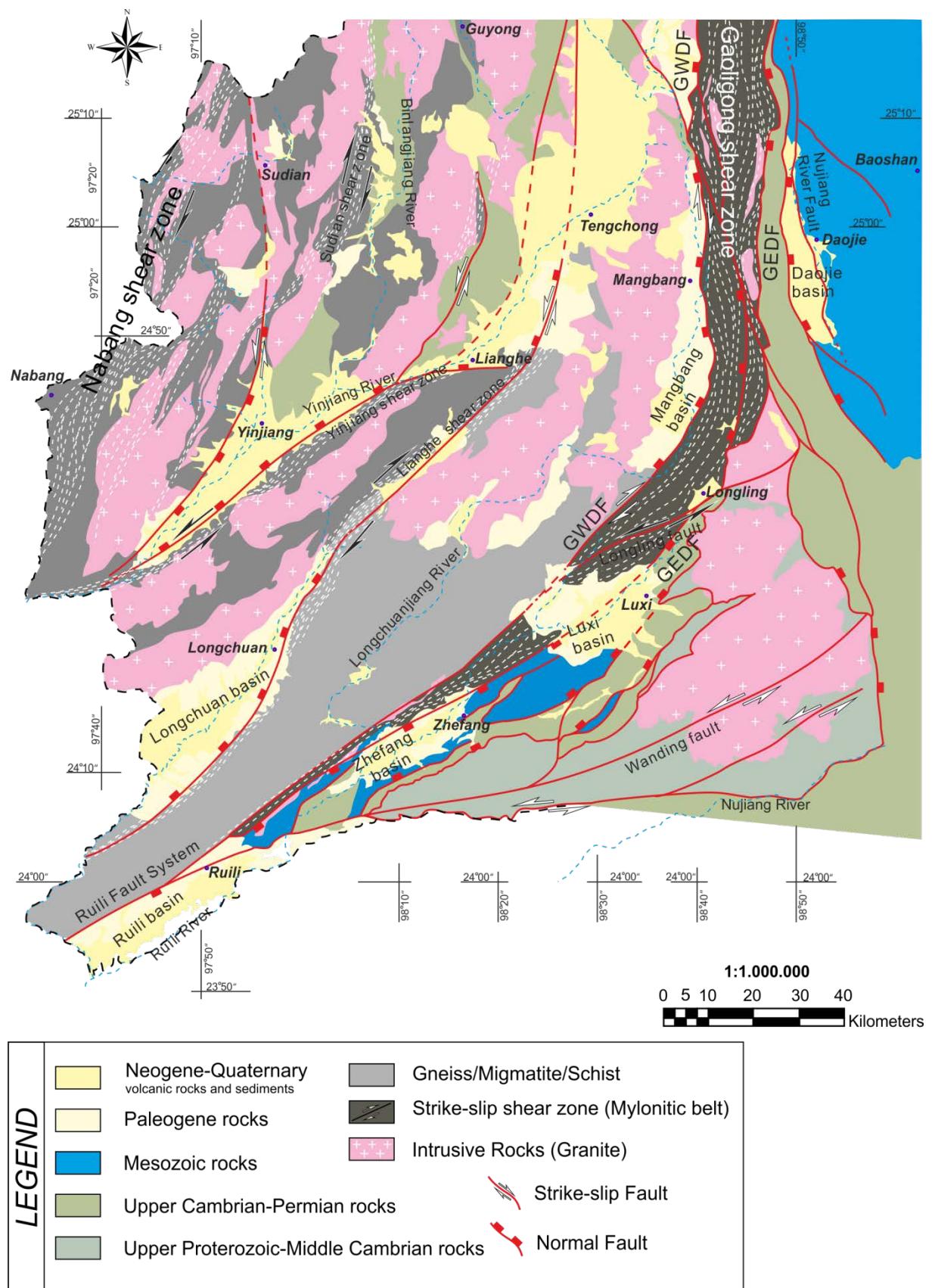
Field observations along the Gaoligong Mountains showed that the 500 m to 6 km wide Gaoligong Shear Zone contains mylonitic granites, gneisses, quartz schists and marble lens, and is characterized by steep foliation and nearly N-S trending sub-horizontal stretching lineations [*Wang et al.*, 2008]. The banded mylonitic rocks were tightly folded with the hinges parallel to the lineations [*Zhang et al.*, 2012a]. The S-C fabrics in shear bands and asymmetric boudins of quartz veins, confirm the dextral movement [*Xu et al.*, 2015].

The Shear Zone is truncated by brittle normal faults with dextral shear component along both sides of the mylonitic belt (GEDF and GWDF in Figure 40). Moving southwards, shear zone

width increases from 10 to 20 km, and the mountains decrease in elevation from 3500 m to 2500 m. Near Longling, the mylonitic rocks of the Gaoligong Shear Zone curve to the southwest, are exposed at the southwest tip of the mountain range, and finally crop out close to the Zhefang and Ruili basins (Figure 42 from *Pellegrino et al.*, 2018).

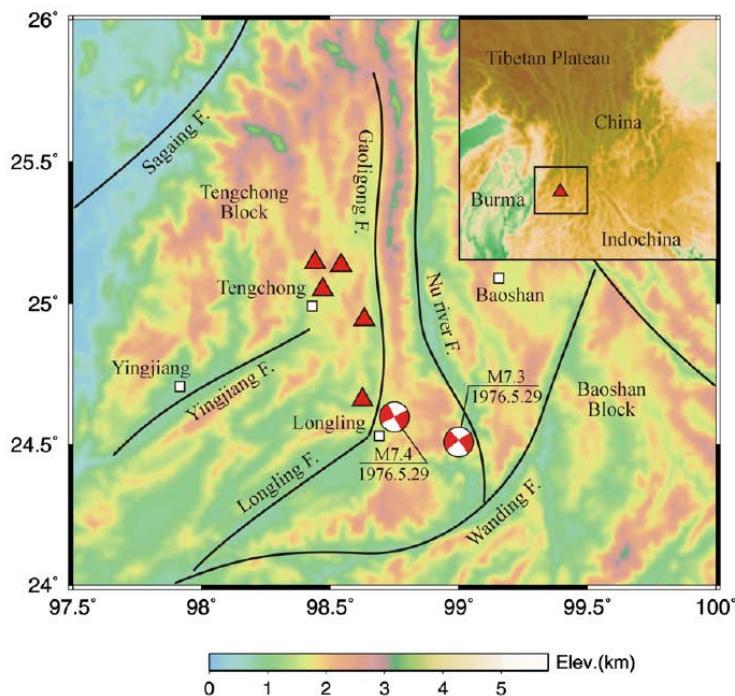
The age of activity and kinematics along the shear zone was constrained by several Ar/Ar ages and structural studies on ductile deformation, which consistently suggest a dextral shear sense. *Lin et al.* [2009] and *Zhang et al.* [2012b] provided 13-18 and 10-16 Ma activity windows (respectively) relying on Ar/Ar mica ages, whereas an older 27-32 Ma activity age was radiometrically documented by *Wang et al.* [2006]. Ductile right-lateral deformation indicators are widespread all along the shear zone. No younger activity is documented in its central-northern N-S branch.

To the south, the Gaoligong Shear Zone is reactivated and truncated with left-lateral movement by the active NE-trending Longling and Wanding brittle faults [*Wang et al.*, 2008; *Xu et al.*, 2012; Figures 42] yielding the M 7.4 Longling earthquake in 1976. Apatite fission track data by *Wang et al.* [2008] constrain the onset of left-lateral brittle fault activity between 8.4 and 0.9 Ma. More recent (and possibly on-going) dextral strike-slip activity of the Gaoligong Shear Zone was proposed by *Socquet and Pubellier* [2005] relying on satellite image analysis. In agreement with the latter work, *Morley* [2007] suggested that the shear zone is a late Neogene transpressional deformation zone.



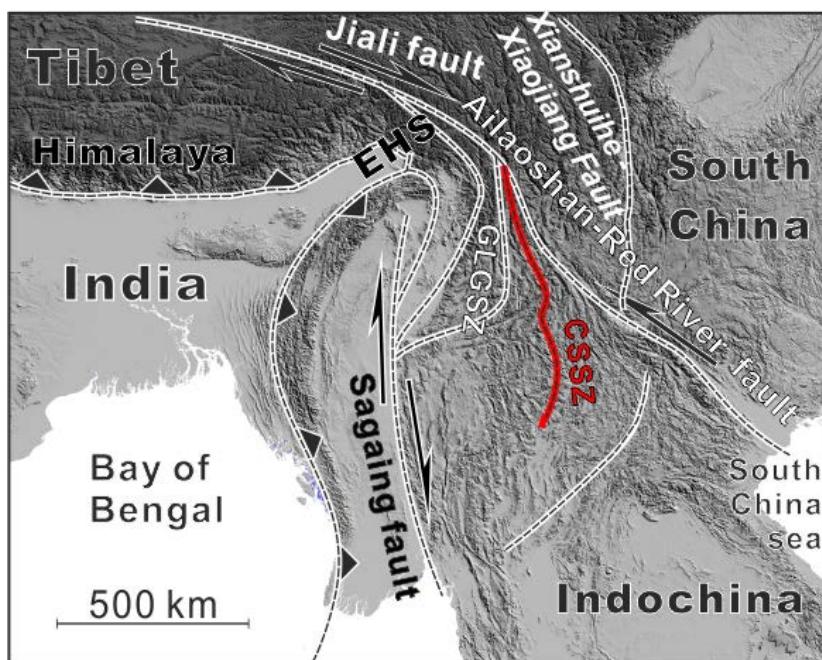
**Figure 42.** Modified after Pellegrino *et al.* [2018]. Simplified geological and tectonic map of the Gaoligong Shear Zone and adjacent areas. GEDF = Gaoligong east detachment fault; GWDF = Gaoligong west detachment fault.

Seismic activity has been recorded along the southern NE-trending part of the Gaoligong Shear Zone (also called Ruili fault system, Figure 42). On May 29, 1976, two large magnitude earthquakes ( $M_s > 7$ ) occurred in the Longling area south of Tengchong [Wang *et al.*, 2006; Zhang *et al.*, 1994]. The first  $M_s = 7.4$  event was located close to the NE-trending Gaoligong Shear Zone branch [Holt *et al.*, 1991]. The second  $M_s = 7.3$  event occurred close to the Wanding fault [Figure 42, 43; Xu *et al.*, 2012]. Both of them are strike-slip events with steep fault planes and ~N-S P axes, and have been interpreted as related to the recent left-lateral shear along NE-trending Longling and Wanding faults [Figure 42, 43; Wang *et al.*, 2006; Xu *et al.*, 2012]. It is noteworthy that the present-day sinistral kinematics along NE faults entering and cross-cutting the southern branch of the Gaoligong Shear Zone is opposite to the dextral kinematics reported for Oligo-Miocene times.



**Figure 43.** Tectonic outline from [Xu *et al.*, 2012]. Red triangles are main volcanic locations. Focal mechanisms are two  $M > 7.0$  earthquakes on May 29, 1976 (derived from the Harvard Global CMT Catalog). Their locations and magnitudes refer to the China National Earthquake Bulletins. Black solid lines are major active faults. The color scale bar denotes topography of the study area. Inset sub-figure is a large scale map showing the location of the study area.

#### 4.4 THE CHONGSHAN SHEAR ZONE



**Figure 44.** Schematic tectonic map of South Tibet. In red the Chongshan Shear Zone (CSSZ).

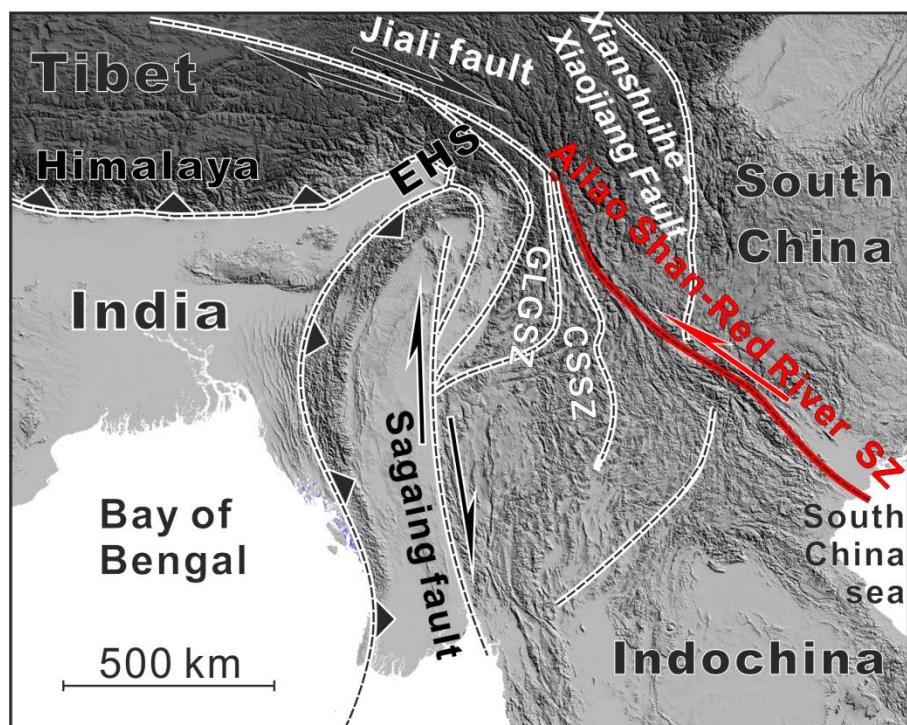
The Chongshan Shear Zone extends for hundreds of kilometers along the Biluoxueshan-Chongshan Mountains in western Yunnan with a width of ca. 10 km (Figure 38, 44). The shear zone is inconsistent along its southern and northern branches, in fact is divided into two belts (N-S trending) with contrasting metamorphic facies: an eastern low-grade schist belt and a western high-grade gneiss belt, considered to be a paired-metamorphic zone [Wang *et al.*, 2008; Zhang *et al.*, 1993; Zhang *et al.*, 2010, 2011, 2012b]. The shear zone formed during the Cenozoic, and exhumation (from a depth of approximately 20 km) occurred from late Paleocene (ca. 57 Ma) to Miocene times [Akciz *et al.*, 2008].

The northern segment of the Chongshan Shear Zone show a dextral displacement exceeding 100 km, related to the ductile deformation, spanning along 57-16 Ma time window. Subsequently a dextral and sinistral strike-slip displacement reported for the 32-14 Ma time windows, was documented along its northern and southern segment, respectively [Zhang *et al.*, 2010; 2012a]. Sediments predominantly older than strike-slip fault activity are widely

exposed in the western sector (Lanping block, see description later), and are made almost entirely of a 7.5 km thick succession of continental red beds of Jurassic-early Cenozoic age.

The ductile sinistral Chongshan and dextral Gaoligong Shear Zone were formed in Oligocene and are thought to be conjugate shear zones [Wang *et al.*, 2006; Akciz *et al.*, 2008]. They probably deformed contemporaneously with the Ailaoshan Red River Shear Zone, and all of their motions accommodated the southeastward extrusion of Southeast Asia.

#### 4.5 THE AILAOSHAN-RED RIVER SHEAR ZONE



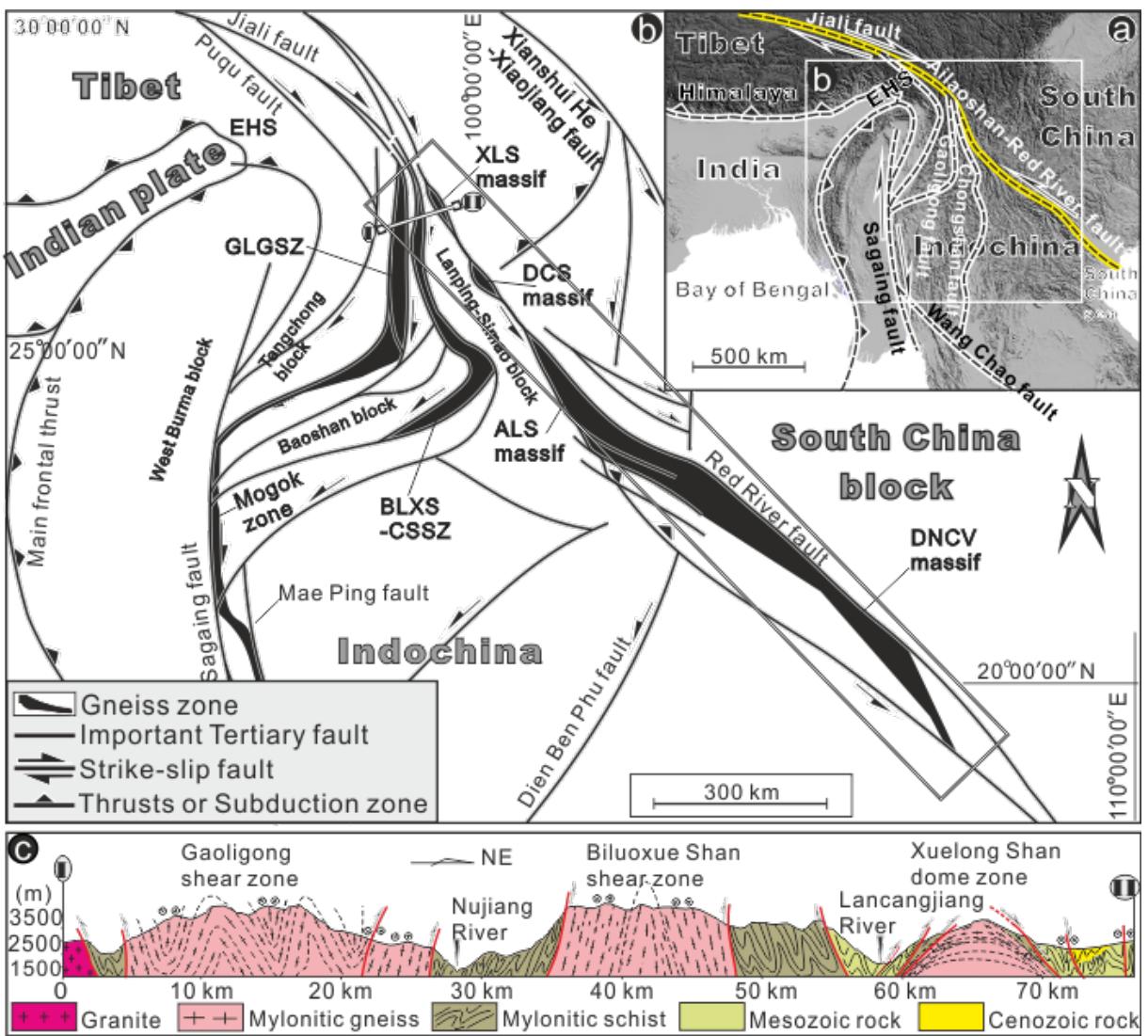
**Figure 45.** Schematic tectonic map of South Tibet. In red the Ailaoshan-Red River Shear Zone.

The Ailaoshan-Red River Shear Zone (ARRSZ) is a first-order tectonic structure that may be followed from the east margin of the Tibetan Plateau to the South China Sea and plays an important role in the distributed shortening model, with Tibet rotating around the eastern

Himalayan Syntaxis by introduction of a weak crustal layer through lower crustal flow southeastward [Figure 45; *Royden et al.*, 1997; *Clark et al.*, 2005a,b].

The Ailaoshan Red River Shear Zone is the major geological discontinuity that separates the Chuandian and South China blocks to the north-east from the extruding Indochina block to the south-east [Figure 38; *Leloup et al.*, 1995; *Tapponnier et al.*, 1990; *Zhang et al.*, 2006; *Zhong et al.*, 1990]. The Shear Zone extends over a length of more than 1000 km between the South China Sea to the south and the Tibet Plateau to the north, and stands out as the most striking discontinuity in the morphology and geology of Yunnan. It is regarded as the eastern boundary of the southeastward extruded Indochina block, in agreement with the observed left-lateral ductile shear [*Leloup et al.*, 1995; *Tapponnier et al.*, 1982]. The onshore segment of the Red River Fault is exposed in four ranges about 10–20 km wide underlain by mylonitic gneiss, metamorphosed to amphibolite grade: Xuelong Shan (XLS), Diancang Shan (DCS), Ailaoshan (ALS) and Day Nui Con VOI (DNCV) [Figure 46; *Leloup et al.*, 1995; *Zhang et al.*, 2017]. In the central sector, the foliation is usually steep, and the lineation is usually nearly horizontal, both being almost parallel to the trend of the mylonitization in the gneiss massifs [e.g., *Leloup et al.*, 1993, 2001; *Jolivet et al.*, 2001; *Cao et al.*, 2010; *Zhang et al.*, 2017].

The timing and the amount of displacement along the shear zone have been strongly debated over the last decades. Its reflects at least two successive deformation phases [*Zhu et al.*, 2009], but reversal of slip on the ARRSZ from 17 to 5 Ma is not well documented. Pliocene–Quaternary right-lateral movement of the Red River Fault is documented onshore by sharp geomorphic fault traces, large river offsets, and well-preserved cumulative scarps. Dextral motion on the Red River Fault likely started around 8–5 Ma, and the total slip has been estimated to be  $\leq 40$  km [*Allen et al.*, 1984; 2016; *Replumaz et al.*, 2001; *Burchfield and Wang*, 2002; *Schoenbohm et al.*, 2006].



**Figure 46.** Simplified regional tectonic framework in southeastern Asia (modified after Zhang et al., [2017] after Tapponnier and Molnar [1977], Tapponnier et al. [2001], Leloup et al. [1995, 2001], Wang and Burchfiel [1997], Ding et al. [2005], Searle [2006], Liu et al. [2007], and Xu et al. [2015]). **(a and b)** The Cenozoic major fault systems, the Jiali-Ailaoshan-Red River fault, and the Gaoligong-Sagaing- Wang Chao fault curving around the EHS region extending to southeast. **(c)** A regional profile I-II across the Gaoligong, Biluoqeshan, and Xuelong Shan massifs [Zhang et al., 2010, 2014]. Xuelong Shan (XLS), Diancang Shan (DCS), Ailaoshan (ALS), Gaoligong Shan Shear Zone (GLGSZ), Biluoqeshan-Chong Shan Shear Zone (BLXS-CSSZ), Day Nui Con Voi (DNCV), Eastern Himalaya Syntaxis (EHS).

In addition, the Red River Fault was characterized by left-lateral movement about 34–17 Ma (Oligo-Miocene), with widely varying estimation of total slip of 100–1400 km specifically between  $700 \pm 200$  km in the NW and ~250 km in the SE [Tappognier *et al.*, 1990; Zhong *et al.*, 1990; Harrison *et al.*, 1992, 1996; Leloup and Kienast, 1993; Scharer *et al.*, 1994; Leloup *et al.*, 1995; Chung *et al.*, 1997; Wang E. *et al.*, 1998; Zhang and Scharer, 1999; Wang *et al.*, 2000; Leloup *et al.*, 2001; 2007; Gilley *et al.*, 2003; Searle, 2006; Zhang *et al.*, 2017; Li *et al.*, 2017a]. The latter offset of the ARRSZ has been strongly debated.

Following the pioneering work of Tappognier *et al.* [1990], many geologists have posited- mainly relying on displaced geological markers- that the ductile, left-lateral deformation of the Red River Fault absorbed hundreds of kilometers of slip and a substantial part of Asian shortening during the Cenozoic India–Asia collision, and led to the opening of the South China Sea [e.g. Scharer *et al.*, 1994; Leloup *et al.*, 2001]. However, some evidence is incompatible with the continental extrusion model [Schoenbohm *et al.*, 2006 and reference therein], because this would imply that the ARRSZ was a proper transform fault and that Indochina behaved as a microplate during Oligo-Miocene times. On the other hand, Searle [2006; 2007] recognized the ARRSZ as an Oligo-Miocene crustal strike slip fault reactivating an older exhumed metamorphic core complex, and stressed that the left-lateral offset cannot be constrained.

After the description of the major shear zones, below is a brief geological description of the adjacent blocks, known as Tengchong, Baoshan, Lanping-Simao and Chuandian blocks. They are characterized by tectono-metamorphic settings showing different lithologies, intracontinental deformation and orogenic history [Lee *et al.*, 2003; Leloup *et al.*, 1995; Wang and Burchfiel, 1997; Wang *et al.*, 2006].

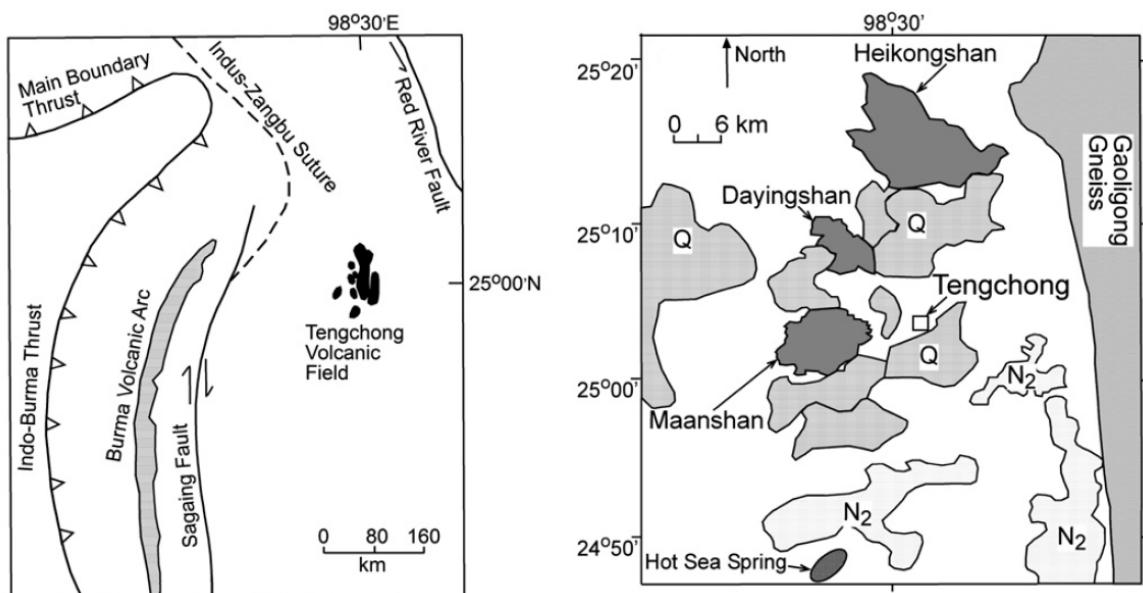
#### 4.6 THE TENGCHONG BLOCK

The Tengchong block, located in southwestern China (Yunnan Province) and part of the SE Himalayan tectonic belt [Zhu *et al.*, 1983], underwent ca. E–W extension in the Late Cenozoic along ca. N-trending fault-bounded grabens [e.g. Wang *et al.*, 2008]. The basement consists of Neoproterozoic to Paleozoic gneisses, migmatites, granites, and small lenses of mafic rocks [e.g. Chen *et al.*, 2002], covered by mostly Late Paleozoic sedimentary rocks and intruded by volcanic rocks of late Miocene to Holocene (sporadic Late Mesozoic to Early Cenozoic granites, locally foliated) age which constitute the well-known “Tengchong Volcanic field” [Wang *et al.*, 2007; Zhu *et al.*, 1983; Zou *et al.*, 2010; Kornfeld *et al.*, 2014; Xu *et al.*, 2015].

The magmatism was active since Mesozoic times [Dong *et al.*, 2006; Li *et al.*, 2011]. Widely exposed Pliocene-Holocene volcanics of the Tengchong field are mainly distributed along NE–NNE trending basins, alternated with Mesozoic granites and sedimentary strata. The distribution of upper Cenozoic volcanic rocks covers an area of 792 km<sup>2</sup>. Pliocene-Holocene volcanic activity of the Tengchong field can be grouped into at least three different types [Jiang, 1998; Li *et al.*, 1999]: (1) Middle-Upper Pliocene to Upper Pleistocene basalts, (2) Lower Pleistocene silicic pyroclastic rocks and (3) Holocene basaltic andesites [Li *et al.*, 1999; Mu *et al.*, 1987; Wang *et al.*, 1999; Zou *et al.*, 2014]. Further south, the Mangbang basin was interpreted as a half-graben developed along the Gaoligong Shear Zone. It is filled with upper Miocene-Pliocene sandstones and mudstone intercalated with tuffs and basalts [Figure 42; Wang *et al.*, 2008].

Volcanic centers are located within a 90 km long and 50 km wide N-S-trending graben. There are 68 late Neogene–Quaternary volcanic centers, 145 hot springs and 25 preserved volcanic craters and cones [Jiang, 1998; Jiang *et al.*, 2003].

The three youngest of the 68 Quaternary volcanoes in Tengchong Volcanic Province [Jiang, 1998; Huangpu and Jiang, 2000; Turker *et al.*, 2013] are thought to be of Holocene age. The centers are named Dayingshan (Shooting Eagle Mountain), Maanshan (Saddle Peak), and Heikongshan (Black Emptied Peak) (Figure 47). These are active volcanoes, identified by their current geothermal activity [Shangguan, 2005], geophysical surveys that indicate the presence of magma, and dating of young volcanic rocks [Li *et al.*, 1999, 2000; Wang *et al.*, 1999]. Dayingshan Volcano (E 98°; N 25°) is the oldest and largest of the three, with a base to the cone of 2 x 1 km and cone height of 120 m and it is made up of lavas erupted in several episodes and a scoria cone.



**Figure 47.** Location map of the Tengchong volcanic field in SE China (left) and distribution and relative ages of the three most recently active volcanoes and surrounding rocks. Modified after Turker *et al.* [2013].

To the west of the Gaoligong strike-slip Shear Zone, new mapping identified widespread mylonitic ortho- and paragneisses with gently dipping foliation and nearly horizontal stretching lineation defined by long axes of amphibole and sillimanite grains and boudinage of felsic veins. Despite the varying attitude of the foliation, the stretching lineation has a consistent NE-azimuth, parallel to the strike of the Gaoligong Shear Zone. These strongly sheared rocks can be traced for more than 100 km from the northern Longling area, the

western Mangshi area, to the northern Ruili area, forming a gently folded, low-angle ductile Shear Zone, the Donghe detachment [Xu *et al.*, 2015].

#### 4.7 THE BAOSHAN BLOCK

The Baoshan Block appears to pinch out where the GSSZ and CSSZ merge [Akciz, 2004, 2008]. The Baoshan Block accreted to Asia in the Late Mesozoic shortened internally along N-trending folds and thrusts (present-day orientation) before or during the Eocene–Oligocene, possibly associated with the strike-slip movements along the GSSZ and CSSZ [Wang and Burchfiel 1997; Wang *et al.*, 2008]. The Baoshan Block exposes thick Upper Proterozoic to Triassic marine strata consisting of carbonates, sandstones, siltstones and shales deposited in a platform environment, unconformably covered by Middle Jurassic marine sedimentary rocks and continental deposits made of Jurassic-lower Cenozoic red beds [Huang *et al.*, 2015; Wang and Burchfiel, 1997; Kornfeld *et al.*, 2014a].

The Paleozoic strata are mainly carbonatic and clastic and locally intercalated with mainly rift-related tholeiitic basalts and basaltic volcanoclastic rocks that erupted when the Baoshan Block separated from Gondwana [e.g. Wopfner, 1996; Wang *et al.*, 2008; Kornfeld *et al.*, 2014a; Xu Y. *et al.*, 2015]. In the central part of the BST, between Baoshan and Yongde city (Figure 38), several approximately N-S trending thrust faults develop, which were offset by a series of E-W trending strike-slip faults. The Eocene-Oligocene formations (conglomerate and sandstone resting unconformably on older rocks) have undergone approximately east-west compression forming folds with approximately N-S trending fold axes [e.g. Wang and Burchfiel, 1997; Tong *et al.*, 2016 and reference therein]. Pliocene volcanic and sedimentary rocks and Quaternary strata remain mostly undeformed [Wang and Burchfiel, 1997; Kornfeld *et al.*, 2014a].

Cenozoic rocks were involved in intensive shortening along a number of N-S-trending folds and thrust faults, which narrow and merge with each other at the northern end of the Baoshan block. To the east of the Gaoligong fault, just to the south, the Baoshan block is cut by three sinistral strike-slip shear zone NNE-SSO oriented called Longling, Wanding and Nujiang fault, interrupted by normal faults running parallel to the Gaoligong east detachment fault. Further south, the Baoshan block is bounded by the active left-lateral Nantinghe fault with the same orientation as the three just mentioned faults [Figure 38; *Socquet and Pubellier, 2005*]. Seismicity and stream-offset analysis indicate that the Nantinghe fault may accommodate at present a 1 mm/yr deformation rate [*Lacassin et al., 1998*].

#### **4.8 LANPING SIMAO BLOCK**

The Lanping-Simao block contains the Proterozoic basement and the Palaeozoic marine strata [BGMRY, 1990]. The Lanping-Simao Basin in this terrane consists of the Mesozoic (Upper Triassic to Eocene) continental red beds, which unconformably overlie the Pre-Mesozoic strata. These Mesozoic red beds are generally affected by the NNW-SSE trending folds and thrusts, but an arcuate trend is maintained in the Simao Basin [*Leloup, 1995*]. Timing of fold and thrust formation remains controversial. Although, fairly recent fold and trust growth has been claimed by *Leloup [1995]* on the basis of its study on ramp anticlines, other researchers [*Wang and Burchfiel, 2000; Socquet and Pubellier, 2005; Akciz, 2008; Kondo et al., 2012*] have suggested an older age (between the Paleocene and Oligocene) for compressive tectonics.

The Cretaceous sequence of the Simao fold belt is of terrestrial origin and has been subdivided into four different formations, they are: the Lower Cretaceous Jinxing Formation, the Middle Cretaceous Nanxin Formation, the Upper Cretaceous Hutousi Formation and the

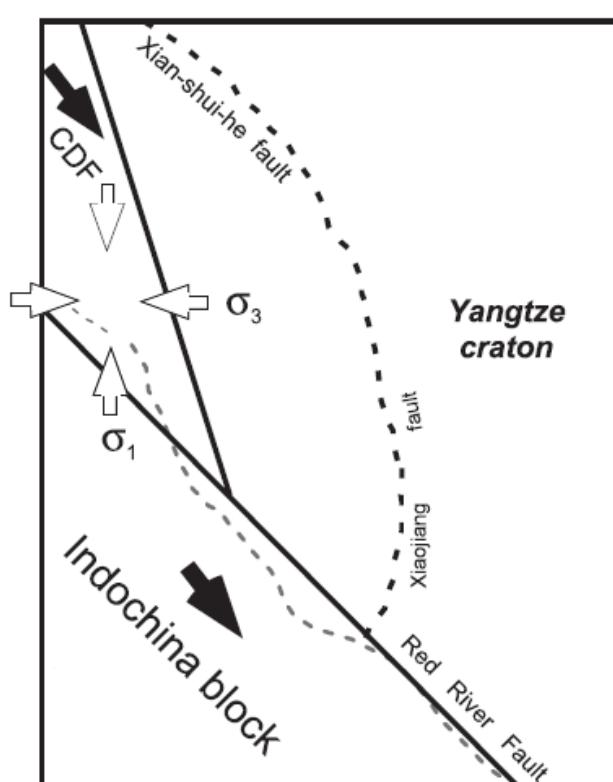
Upper Cretaceous Mankuanhe Formation [BGMRY 1990; *Leloup* 1995; *Kondo et al.*, 2012]. The Jinxing Formation is mainly composed of greyish sandstone intercalated with purplish-red, greyish and green mudstone. The presence of rich Lamellibranchiate indicates the Early Cretaceous age for this formation. The Middle Cretaceous Nanxin Formation mainly consists of purplish-red sandstone, which conformably overlies the Early Cretaceous Jinxing Formation. This formation is overlain by the Upper Cretaceous Hutousi Formation. The occurrences of Estheria, Cypridea, Gastropods, and Sporopollen indicate an age of Middle Cretaceous for the Nanxin Formation. The Upper Cretaceous Mankuanhe Formation is distributed only in the Pu'er-Mengla area [*Kondo et al.*, 2012]. The Nanxin Formation, around Jingdong area, forms an E-W trending anticlinal structure; more southern, in the Zhengyuan area, this formation is separate by several N-S trending anticlines and, to the west others monoclinal stuctural. Also in Mengla locality, a monoclinal structure with N-S trending axis and a dip of 2°-63° always in Nanxin Formation (from northeastward to eastward) has been observed by *Tanaka et al.* [2008].

#### 4.9 CHUANDIAN BLOCK

The Chuandian block is delimited by the Ailaoshan Red River Shear Zone and the Xianshuihe-Xiaojiang fault, which is considered to be a boundary fault to accommodate the different rotations around the eastern Himalayan Syntaxis and the total offset is of 78–100 km [*Wang et al.*, 1998]. The Xianshuihe-Xiaojiang fault developed only in the last 10 Ma, with a brittle sinistral offset of ca. 60 km [*Wang et al.*, 1998; *Zhang et al.*, 2004; *Meade*, 2007; *Wang et al.*, 2016a,b]. GPS measurements and geological observations along the Xianshuihe-Xiaojiang fault system suggest a southward displacement of the Chuandian Fragment (CDF) during the last 12 Ma [*Roger et al.*, 1995; *Wang et al.*, 1998, 2001; *Chen et al.*, 2000]. In the Chuxiong basin, Upper Triassic gray sandstones and marls grade upwards into Jurassic and

Cretaceous red sandstones. The lower red sandstones display decametric cross-bedding [Leloup , 1995]. The youngest red beds, cropping out to the north of Chuxiong (N25°02', E101°31'), are of Eocene age [BGMRY, 1983]. To the west, the maximum thickness of the Chuxiong Mesozoic-Tertiary continental sequence is about 6 km [Gao *et al.*, 2017].

Previous studies reported from the area suggest that this fragment experienced more than 300 km of southward displacement after Cretaceous times. Neogene southward displacement of this fragment is also supported by geological observation and GPS measurements [Tamai *et al.*, 2004 and reference therein]. The indentation of India into Asia brought about some internal deformation to the Chuandian Fragment during its southward translation; this deformation can be seen in significant amount of rotation deduced from Cretaceous paleomagnetic data of this fragment [Funahara *et al.*, 1992; Avouac and Tapponnier, 1993; Huang and Opdyke, 1992; Lacassin *et al.*, 1997; Otofuji *et al.*, 1998; Yoshioka *et al.*, 2003; Tamai *et al.*, 2004; Figure 48]. As suggested by these data, the southward movement of the Chuandian fragment has been continuing since the time of India-Asia collision.



**Figure 48.** A deformation model for the CDF from Yoshioka *et al.* [2003]. The paleo-shape of the CDF was originally apex. An initially linear Xianshuihe-Xiaojiang fault system (solid line) changed into present curvilinear shape (dashed line) as a result of deformation in the CDF. During its southward displacement (closed arrows) in the Cenozoic era, the CDF tectonically interacted with the Yangtze and Indochina blocks. Roughly, east–west extension ( $\sigma_1$ ) and north–south compression ( $\sigma_3$ ) are raised up (open arrows) in the southern part of the CDF. Due to these compressive stresses, an eastward bulge has been formed in the Xianshuihe-Xiaojiang fault system accompanied by a differential tectonic rotation in the Chuxiong basin.

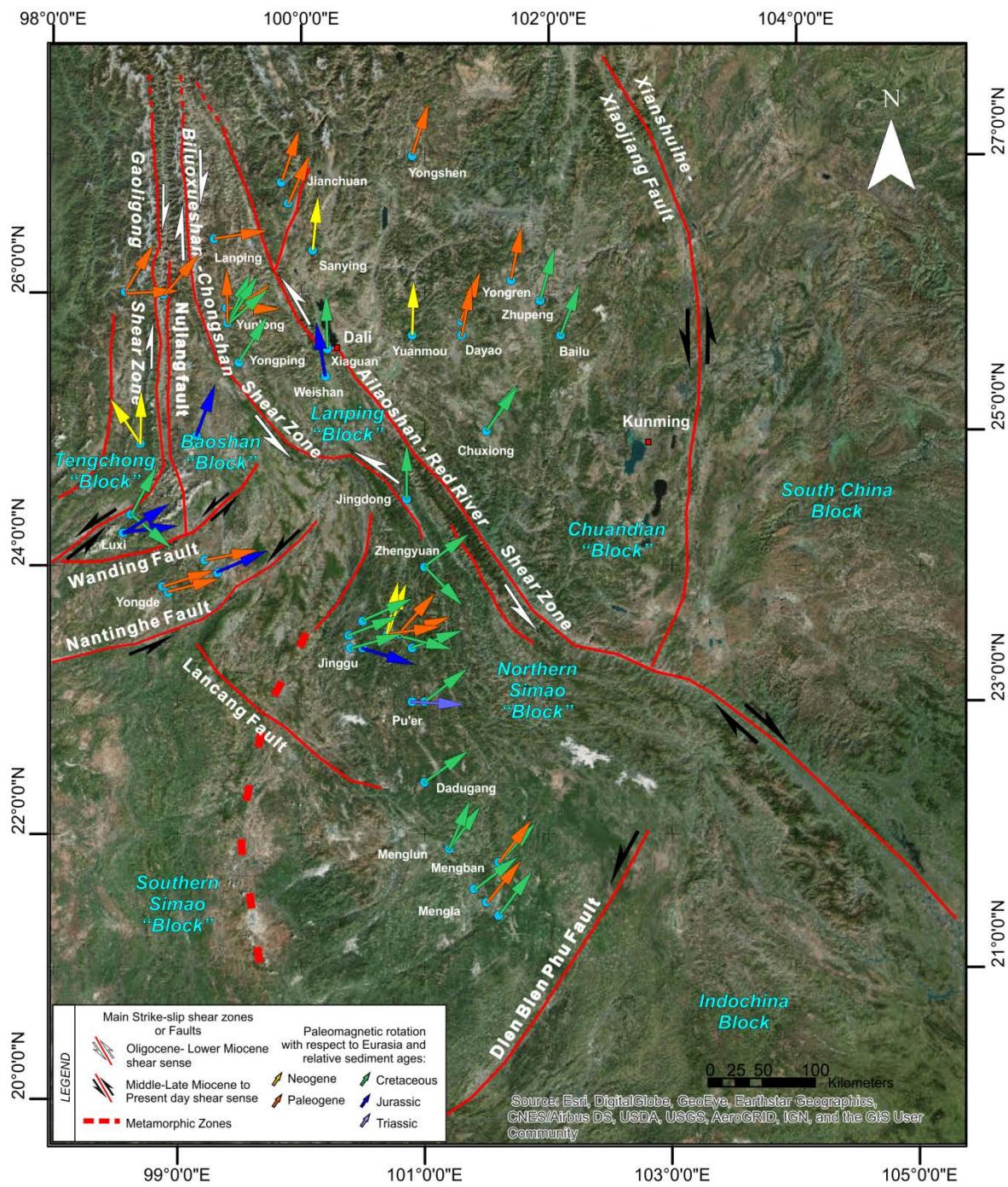




*Chapter*  
*V*

## 5. PREVIOUS PALEOMAGNETIC DATA

A wealth of paleomagnetic data has been gathered during the last decades in Yunnan area, straddling the northern Indochina block and the southeastern Tibetan Plateau (Figure 49 and Table 1). In the present work, all the previously obtained rotations were recalculated using updated 170-140 Ma Eurasian and 130-10 Ma East Asia reference poles from *Torsvik et al.* [2012] and *Cogné et al.* [2013], respectively (Chapter 3, Figure 25, 26). Data are reported, for each block, moving from west to east.



**Figure 49.** Map of western Yunnan and southeastern Eastern Himalayan Syntaxis, main tectonic features, and synthesis of the main paleomagnetic data. Base map source: GIS (Esri), WGS 1984-Web Mercator Auxiliary Sphere. Colored arrows (see legend) represent paleomagnetic rotations with respect to Eurasia, recalculated from previous paleomagnetic directions (See Table 1), according to Demarest [1983] and using updated Eurasia (from 170 to 140 Ma) and East Asia (from 130 to 10 Ma) poles by Torsvik *et al.* [2012] and Cogné *et al.* [2013], respectively.

Locality	Geographic Coordinates		Age	Age (Ma)	N	Considered paleopoles, $\lambda_{gc}$	Observed direction				Recalculated by Us		Reference
	Latitude °N	Longitude °E					In Situ	Tilt corrected	Rotation	Flattening	R ± ΔR	F ± ΔF	
<b>Tengchong Block</b>													
Pianma-Nujiang	26.0	98.7	Oligocene	30	6	30	38.5	20.0	33.3	41.4	24.5	31.2 ± 26.3	-5.0 ± 19.4
	26.0	98.7	Eocene	40	6	40	84.5	14.9	89.8	35.1	11.8	87.2 ± 11.8	-1.6 ± 10.5
Tengchong v.f.	25.1	98.4	Pleistocene-Holocene	2.5-0	16	0	352.5	25.4	—	—	44.9	-7.5 ± 40.3	-5.7 ± 14.1
Mangbang b.	24.8	98.6	Upper Miocene-Pliocene	14-2.5	16	10	327.3	48.7	—	—	14.1	-33.7 ± 17.3	-9.1 ± 11.3
<b>Baoshan Block</b>													
Luxi	24.3	98.4	Middle Jurassic	163-174	6	170	110.0	-2.9	99.7	35.2	11.3	81.3 ± 11.8	24.9 ± 9.3
Baoshan	26.0	98.8	Oligocene	30	16	30	32.6	51.8	42.2	47.0	7.8	40.0 ± 9.2	-10.6 ± 6.9
Yongde	24.4	99.3	Eocene-Oligocene	56-23	22	20	81.7	35.1	—	—	3.8	79.4 ± 4.4	0.2 ± 4.7
	24.2	99.0	Paleocene	66-56	5	60	—	—	76.9	17.8	3.8	72.5 ± 9.1	14.0 ± 9.5
<b>Lanping Block</b>													
Yunlong	25.8	99.4	Middle Cretaceous	100	20	100	50.9	29.8	40.2	49.9	3.9	32.0 ± 5.6	-9.0 ± 4.8
			Lower Cretaceous	145-100.5	23	120	103.2	61.6	59.7	41.0	11.9	49.3 ± 12.6	1.8 ± 9.7
			Upper Cretaceous	100.5-66	9	80	52.4	62.1	34.0	52.4	7.3	24.0 ± 9.8	-15.4 ± 6.9
			Paleocene	66-56	11	60	49.2	7.6	50.2	31.1	13.2	45.7 ± 12.5	3.2 ± 11.4
Lanping area	26.4	99.3	Eocene	38-29	9	30	274.0	-35.6	263.8	-42.1	14.1	81.6 ± 15.1	-5.1 ± 11.4
Yunlong area	25.9	99.4	Eocene	38-29	5	30	271.8	-19.8	273.0	-31.8	11.1	90.8 ± 10.5	4.4 ± 9.2
Weishan	25.4	100.2	Upper Jurassic	163-145	5	150	356.5	34.5	7.3	25.3	10.4	-9.5 ± 10.8	28.6 ± 9.6
Xiaoguan	25.6	100.2	Middle Cretaceous	100	9	100	8.9	59.2	6.9	47.7	8.6	-1.3 ± 10.4	-6.9 ± 7.6
Yongping	25.5	99.5	Lower Cretaceous	145-100.5	12	120	48.9	50.8	42.0	51.1	15.7	31.7 ± 20.0	-8.6 ± 12.6
<b>Northern Simao Block</b>													
Jinggu	23.4	100.9	Middle Cretaceous	100	8	100	73.2	67.7	79.4	43.3	9.1	71.3 ± 10.2	-5.4 ± 8.1
Jinggu	23.6	100.5	Middle Jurassic	174.1-163.5	10	170	75.3	31.4	83.3	36.8	5.4	65.8 ± 7.1	23.3 ± 5.2
Mengla	21.6	101.4	Middle Cretaceous	100	100	100	288.7	19.3	60.8	37.8	7.6	52.7 ± 8.0	-2.4 ± 7.2
Jinggu (Jingxing Fm.)	23.4	100.4	Lower Cretaceous	145-100.5	3	120	78.4	58.0	84.4	39.6	17.8	74.2 ± 18.4	0.3 ± 14.2
Jinggu (Mangang Fm.)	23.4	100.5	Lower Cretaceous	145-100.5	3	120	294.2	23.1	297.5	-33.3	7.1	107.3 ± 7.0	6.6 ± 6.4
	23.5	100.8	Lower Cretaceous	145-100.5	4	120	309.3	-6.2	294.4	-38.1	12.1	104.2 ± 12.3	2.0 ± 9.9
Jinggu (Mengyujing Fm.)	23.5	100.8	Paleocene	66-56	9	10*	42.8	26.1	23.9	51.6	7.9	41.7 ± 7.2	11.7 ± 6.8
Jinggu (Mengla Group)	23.5	100.7	Eocene-Oligocene	56-23	7	40	106.4	49.1	84.7	38.9	7.6	81.8 ± 8.3	-9.3 ± 8.0
Jinggu (Sanhaogou Fm.)	23.5	100.7	Miocene	23-5.3	6	10	34.6	20.2	21.1	35.5	7.5	20.0 ± 7.5	2.3 ± 6.5
Jinggu	23.5	100.8	Eocene	56-33.9	6	40	—	—	73.1	39.9	11.8	70.2 ± 12.5	-10.3 ± 10.6
Jingdong	24.5	100.8	Cretaceous	145-66	15	100	20.7	55.9	8.3	48.8	7.7	0.1 ± 9.6	-9.4 ± 7.1
Zhengyuan	24.0	101.0	Cretaceous	145-66	14	100	94.7	88.2	61.8	46.1	8.1	53.6 ± 9.6	-7.4 ± 7.4
West Zhengyuan	24.0	101.0	Cretaceous	145-66	5	100	315.6	-8.2	324.2	-49.4	6.4	136.0 ± 8.2	-10.7 ± 6.3
South Mengla	21.4	101.6	Cretaceous	145-66	19	100	353.6	50.3	51.2	46.4	5.6	43.2 ± 7.0	-11.3 ± 6.0
Pu'er	23.0	101.0	Cretaceous	145-66	31	100	82.4	49.5	59.4	45.2	5.1	51.8 ± 6.4	-7.9 ± 5.6
Zhengwan	22.8	100.9	Upper Cretaceous	100-66	11	80	267.5	57.4	51.8	47.9	6.9	41.9 ± 8.5	-15.0 ± 6.8
Dadugang	22.4	101.0	Upper Cretaceous	100-66	12	80	78.1	69.0	64.1	48.1	7.3	54.2 ± 9.0	-15.8 ± 7.0
Mengban	21.8	101.6	Paleocene-Eocene	66-33.9	6	50	12.9	45.8	43.5	23.0	13.4	37.6 ± 12.3	6.1 ± 12.9
Mengla	21.5	101.5	Eocene-Oligocene	56-23	17	40	32.3	34.0	41.8	23.8	5.8	38.8 ± 6.0	2.5 ± 7.2
	21.4	101.6	Lower Cretaceous	145-100.5	14	120	10.7	40.1	46.9	42.2	7.7	36.8 ± 8.5	-4.8 ± 6.8
Menglun	21.9	101.2	Upper Cretaceous	100-66	6	80	44.7	58.3	33.2	30.9	8.2	23.3 ± 8.0	0.6 ± 7.7
	21.9	101.2	Lower Cretaceous	145-100.5	19	120	66.7	72.6	46.2	45.9	11.0	36.1 ± 12.6	-7.9 ± 9.2
Mengban	21.8	101.6	Cretaceous	145-66	4	100	27.5	31.4	50.5	31.0	6.4	42.4 ± 6.5	4.7 ± 6.4
Pu'er	23.0	100.9	Middle Triassic	237-247.2	18	30**	95.4	28.4	98.2	22.8	6.2	93.1 ± 5.9	3.5 ± 5.9
Jinggu	23.5	100.4	Middle Cretaceous	100	47	100	73.4	42.8	77.0	43.0	2.9	68.9 ± 4.3	-5.1 ± 4.5
	23.5	100.7	Miocene	23-5.3	38	10	32.3	26.9	13.7	36.0	3.3	12.6 ± 3.8	1.8 ± 3.9
Jinggu	23.5	100.7	Miocene	23-5.3	10	10	16.4	15.4	13.5	32.2	5.8	12.4 ± 5.8	5.6 ± 5.4
			Paleocene	66-56	12	20**	43.0	41.7	36.1	31.5	8.4	40.6 ± 9.2	-7.6 ± 7.5
			Eocene	56-33.9	12	20**	27.8	32.1	35.2	35.7	6.5	25.4 ± 6.5	1.9 ± 6.3
			Eocene	56-33.9	18	20**	37.0	32.8	53.0	33.6	4.3	34.6 ± 4.7	1.2 ± 5.0
			Eocene-Oligocene	56-23	11	20**	34.9	33.2	38.4	37.3	9.7	32.5 ± 9.4	0.8 ± 8.4
			Eocene-Oligocene	56-23	17	20**	35.2	20.2	38.6	33.0	5.7	32.8 ± 5.4	13.8 ± 5.8
			Eocene-Oligocene	56-23	14	20**	45.0	22.8	50.8	31.8	5.8	42.6 ± 5.5	11.2 ± 5.8
<b>Chuandian Block</b>													
Chuxiong	25.0	101.5	Lower Cretaceous	145-100.5	11	120	28.9	37.1	43.7	36.2	17.1	33.4 ± 16.8	6.0 ± 13.6
			Upper Cretaceous	100-66	10	80	58.6	31.7	45.6	46.6	14.3	35.5 ± 16.7	-10.3 ± 11.8
Bailu	25.7	102.1	Upper Cretaceous	100-66	6	80	359.9	35.5	30.7	38.9	10.6	20.5 ± 11.0	-1.5 ± 9.1
Zhipeng	25.9	101.8	Upper Cretaceous	100-66	15	80	30.8	34.4	25.4	34.3	3.4	15.2 ± 4.3	3.3 ± 4.6
Dayao	25.8	101.3	Paleocene-Eocene	66-33.9	18	50	28.7	27.8	27.8	33.1	4.3	21.8 ± 6.3	2.2 ± 7.7
	25.7	101.3	Paleocene-Eocene	66-33.9	10	50	10.7	31.3	16.5	31.1	4.8	10.5 ± 6.5	4.0 ± 7.9
Yongren	26.1	101.7	Paleocene-Eocene	66-33.9	16	50	11.5	27.7	17.2	26.6	5.8	11.1 ± 7.0	9.2 ± 8.2
Yuannou	25.7	100.9	Pliocene	~3.0	17	0	1.4	27.8	—	—	1.1	1.4 ± 1.2	13.8 ± 1.1
Sanying	26.3	100.1	Upper Miocene-Pliocene	~5.0	0	—	—	—	—	—	—	5.6 ± 1.7	13.8 ± 1.5
Jianchuan (SJ)	26.6	99.3	Oligocene	33.9-23	8	30	213.6	-46.5	200.9	-31.3	7.7	18.7 ± 7.4	6.0 ± 8.8
Jianchuan (JL)	26.5	99.8	Lower/Middle Eocene	~56-40	16	50	28.8	35.0	29.7	32.0	5.6	23.8 ± 7.1	4.1 ± 8.0
Yongshen	27.0	100.9	Upper Eocene-Lower Oligocene	56-23	20	30	19.0	44.1	19.7	40.0	2.9	17.3 ± 3.7	-2.0 ± 3.9

**Table 1.** Paleomagnetic site-mean directions obtained from previous studies in Triassic-Neogene rocks from Tengchong, Baoshan, Lanping, Northern Simao and Chuandian blocks, and updated rotation and flattening values with respect to Eurasia calculated using paleopoles by Torsvik et al. [2012] and Cognè et al. [2013]. Age in Ma is from the geologic timescale of Cohen et al. [2013]. D and I are site mean declination and inclination calculated before and after tectonic correction;  $a_{95}$  is statistical parameter after Fisher [1953]; N is number of sites; site mean rotation R and flattening F values, and relative errors  $\Delta R$  and  $\Delta F$  (according to Demarest, 1983) are relative to coeval D and I values expected at the sampling area from East Asia paleopoles by Cognè et al. [2013], except 150 and 170 Ma from European poles by Torsvik et al. [2012]. \*\*Post-tilting remagnetized sites, R and F values are calculated using the in-situ D and I values.

**In the Tengchong block,** *Kornfeld et al.* [2014b] found 87° CW rotation in 40 Ma dykes from the Tengchong block (Table 1). In addition, *Kornfeld et al.* [2014c] reported counterclockwise (CCW) paleomagnetic rotations from lower Pliocene to Holocene volcanics of the Mangbang basin, and interpreted them as apparent rotations due to tilting around horizontal axes along listric normal faults cutting the Tengchong volcanic field.

Data from Upper Carboniferous and Early Permian basalts in the Baoshan block respectively reported by *Huang and Opdyke* [1991] and *Ali et al.* [2013], reveal their former position within the Gondwana supercontinent. However, in this thesis we will only deal with data from Triassic to Neogene age.

**In the Baoshan block,** *Huang and Opdyke* [1993] documented a strong CW rotation in Jurassic rocks from Luxi area. *Kornfeld et al.* [2014a] showed a 40° CW rotation in 30 Ma basalts from the Baoshan block. Subsequently, *Tong et al.* [2016] studied Paleocene-Oligocene red beds located at the centre of the Baoshan block, NE of Yongde (Figure 49). A primary magnetization was isolated in the Paleocene sediments, while the Eocene-Oligocene red beds were magnetically overprinted during the Miocene. In any case, both localities and sedimentary sequences yield significant CW rotations, ranging from 70° to 80° (Table 1).

**In the Lanping block,** the rotation pattern from Jurassic and Cretaceous red beds, revealed that the Xiaguan and Weishan areas did not rotate with respect to stable Eurasia since Late Cretaceous times [*Huang and Opdyke*, 1993]. In contrast, in the Yunlong and Lanping area, about 50 km west of the Ailaoshan-Red River Shear Zone, *Sato et al.* [1999, 2001, 2007] and *Yang et al.* [2001a,b], documented strong CW rotations and inferred that Indochina was squeezed out of the Asian continent due to collision of India, accompanied by CW rotations (Figure 50).

Similar evidence was found from Lower Cretaceous sites near the city of Yongping, where data show a CW rotation of  $25^\circ \pm 16^\circ$ , associated with rotation of the Yongping area [Funahara *et al.*, 1993].

Additionally, a series of works were conducted between 1993 and 2009 on the Jurassic-upper Cretaceous red beds exposed mostly in the Lanping-Simao block, between the Ailaoshan-Red River and Chongshan Shear Zones. The data show variable predominant CW rotations up to  $132^\circ$  with respect to Eurasia.



**Figure 50.** From Sato *et al.* [2007]. Declinations of the Cretaceous (black) and Tertiary (gray) paleomagnetic directions in the Lanping-Simao and Chuxiong blocks and its surrounding regions, are shown by arrows.

In the northern Simao block, Jurassic and Cretaceous red beds have been paleomagnetically studied mostly in the northern sector [Huang and Opdyke, 1993] and the rotation pattern revealed that the Jinggu-Mengla area rotated CW by  $46^\circ$ - $65^\circ$ . Chen *et al.* [1995] documented CW rotations  $>90^\circ$ ,  $77^\circ$ , and  $15^\circ$  respectively in Lower Cretaceous, Eo-Oligocene, and Miocene sediments from the Jinggu area. The available Cretaceous paleomagnetic data from the Jinggu area and neighboring regions demonstrate variable rotations in different sampling localities, possibly bounded by strike-slip faults.

This would imply that the Yunnan region and surrounding areas were not deformed as a coherent block in response to the collision of India with Asia [Chen *et al.*, 1995].

Sato *et al.* [2007] investigated Cretaceous red sandstones and siltstones in Pu'er, finding a  $46^\circ \pm 9^\circ$  CW rotation, which is consistent with the large rotations in the Mengla and Jinggu areas ( $47^\circ$  and  $66^\circ$ ). The CW rotation of the northern part of the Simao block is larger by more than  $20^\circ$  compared to the Yunlong and Yongping areas belonging to the Lanping block [Sato *et al.*, 2001]. Geological, geochronological and GPS data, obtained from the Lanping-Simao block, seem to demonstrate the close relationship between the Lanping-Simao arcuate structural zone and differential tectonic rotations [Kondo *et al.*, 2012; Tanaka *et al.*, 2008; Tong *et al.*, 2013].

Recently, Li *et al.* [2017b] focused on Paleocene-Oligocene red beds and mid-Miocene silts from the Jinggu Basin, northern Simao block. Although a pre-folding magnetization was proved in red beds, exclusively normal polarity sites were documented (consistently with previous evidence by Chen *et al.*, 1995; Gao *et al.*, 2015; Huang and Opdyke, 1993). Thus Li *et al.* [2017b] concluded that either the studied strata were deposited during the long normal Cretaceous superchron, or the red beds underwent a widespread pre-folding Tertiary overprint. In any case, red beds yield greatly variable CW rotations (ranging from  $0^\circ$  to  $>90^\circ$ ) that Li *et al.* [2017b] average to a  $30^\circ$  value, while mid-Miocene sediments do not rotate.

Li *et al.* [2017a] explained rotation with a model of semi-rigid blocks rotation. Other papers (listed below) observe that there are internal deformation and variable rotations in the Simao block [Kondo *et al.*, 2012; Tong *et al.*, 2013; Gao *et al.*, 2015 even consider the possibility of internal oroclinal deformation] rising doubts on rigid “microplate” rotations in the Yunnan.

Recently, Li *et al.* [2018], suggest that the northern Sibumasu region (area between Nantinghe fault and Dien Bien Phu fault; Figure 49), cut by a series of NE–SW strike-slip faults, experienced a much larger clockwise rotation than the remaining areas of southeast margin of

Tibetan Plateau. They ascribe this large rotation to a sum of rotations acquired after the middle Miocene and associated with Indochina extrusion and rotation and the south-eastward extrusion of Tibetan crust along the Xianshuihe-Xiaojiang fault (Figure 49-50). If correct, this indicates that the eastward motion of Tibetan crust crossed the Ailaoshan-Red River fault and was transferred through the southwestern Yunnan toward Myanmar. Furthermore, their paleomagnetic results, together with Jurassic/Cretaceous paleomagnetic data, suggest that the northern Sibumasu changed direction from a pre-collisional strike of ~N 60° W in eastern Tibet to N 10° W in southern Sibumasu.

Regarding the **Chuandian Block** (CDB, also called Chuandian Fragment, CDF), many authors support the hypothesis that it was extruded along the left-lateral Xianshuihe-Xiaojiang fault system under the tectonic influence of India–Asia collision [*Avouac and Tapponnier, 1993; Lacassin et al., 1997*], but its southward movement was eventually obstructed by the rigid Yangtze and Indochina blocks [*Wang et al., 2001; Yoshioka et al., 2003; Figure 50-51*]. GPS measurements and geological observations along the Xianshuihe-Xiaojiang fault system suggest a southward displacement of the Chuandian Fragment during the last 12 Ma [*Roger et al., 1995; Wang et al., 1998, 2001; Chen et al., 2000*].

Previous paleomagnetic studies reported from the area suggest that this fragment experienced more than 300 km of southward displacement after Cretaceous times [*Tamai et al., 2004*]; the declination values increase in latitude from north to south [*Funahara et al., 1992; Huang and Opdyke, 1992; Otofuji et al., 1998; Tamai et al., 2004*] and the magnitude of rotational motion significantly increases in the southern part of this fragment.

The Chuandian domain yields a clear pattern of ca. 20° CW rotations for Cretaceous-Oligocene red beds [*Funahara et al., 1992; Otofuji et al., 1998; Yang et al., 2001a; Yoshioka et al., 2003; Zhu et al., 2008; Li et al., 2013; 2015; Tong et al., 2015; Wang et al., 2016a; Gao et al., 2017*].

Paleomagnetic samples of Paleocene–Eocene red sandstones were investigated by *Yoshioka et al.* [2003] around the Yongren ( $26.1^{\circ}\text{N}$ ,  $101.7^{\circ}\text{E}$ ) and Dayao areas ( $25.7^{\circ}\text{N}$ ,  $101.3^{\circ}\text{E}$ ), located in the Chuxiong basin (Figure 49 and Table 1).

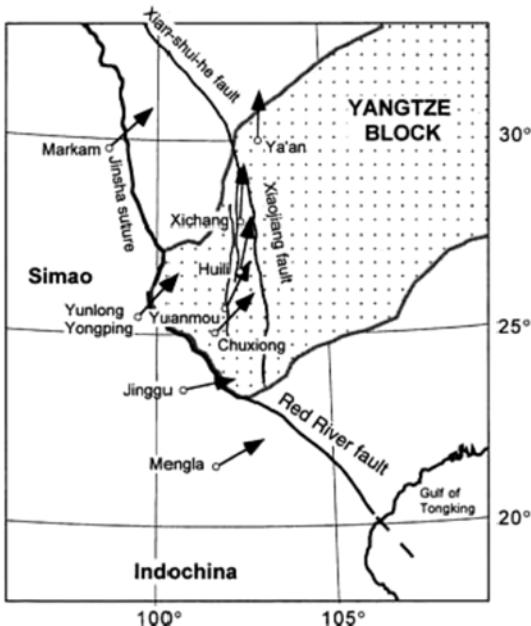
Their data of Easterly deflected declinations are consistent with those reported from other areas of the Chuxiong basin supporting a CW rotation of about  $30^{\circ}$ , indicating its wide presence in the Cretaceous–Eocene formations of the said basin.

Comparison with declination values expected from the Cretaceous–Eocene Paleopoles of Eurasia indicates that the magnitude of clockwise rotation systematically increases toward the southeast within the Chuxiong basin as well as in the Chuandian Fragment. This trend of the differential tectonic rotation in the Chuandian Fragment is consistent with curvature of the Xianshuihe-Xiaojiang fault system.

They interpret the deformation of the Chuxiong basin connected with the formation of eastward bulge in the southern part of the CDF. During southward displacement, the Chuandian Fragment was probably subjected to tectonic stresses as a result of interaction with the Yangtze and Indochina blocks, which resulted into east–west extension and north–south shortening [Figure 48-51, *Yoshioka et al.*, 2003].

Cretaceous red sandstones in Xichang were studied also by *Tamai et al.* [2004].

Combined with earlier reported paleomagnetic data from this block, they estimate a significant southward displacement of  $6.7^{\circ} \pm 3.5^{\circ}$  in latitude for the whole fragment with respect to the Sichuan Basin since the Late Cretaceous. Extrusion dynamics in the Asian continent due to its collision with India brought about the southward displacement of the CDB. In agreement with previous authors, their declination data indicate that the southern part of the block was subjected to clockwise sense rotation of up to  $45^{\circ}$  and this probably occurred during extrusion of this fragment from the north.



**Figure 51.** Horizontal projections of paleomagnetic remanent directions synthetized by Tamai *et al.* [2004] for Lanping-Simao block, Chuandian and Yangtze blocks; Ya'an [Otofuji *et al.*, 1990; Enkin *et al.*, 1991], Xichang [Tamai *et al.*, 2004], Huili [Huang and Opdyke, 1992], Yuanmou [Otofuji *et al.*, 1998], Chuxiong [Funahara *et al.*, 1992], Markam [Otofuji *et al.*, 1990; Huang and Opdyke, 1993], Yunlong [Sato *et al.*, 1999], Yongping [Funahara *et al.*, 1993], Jinggu [Huang and Opdyke, 1993; Haihong *et al.*, 1995] and Mengla [Huang and Opdyke, 1993].

More recently, late Eocene-early Oligocene red beds were studied in the Yongsheng area by Gao *et al.* [2017]. In their samples, the higher temperature components (HTC) were isolated by stepwise thermal demagnetization between  $\sim 300$ – $680^\circ\text{C}$  and they passed the fold test. However, all the HTCs are of normal polarity, which appears to conflict with the frequent occurrence of reversed polarities during the late Eocene-early Oligocene, and therefore the possibility of remagnetization needs to be considered. Widespread secondary hematite was detected in the red beds that further indicates the remagnetization of samples. From the events analysis of the sampling area, they suggest an early Oligocene ( $\sim 35$  Ma) remagnetization event. Their comparison of the pole calculated from the remagnetized remanent directions with the  $\sim 35$  Ma paleopole for Eurasia indicate that the degree of clockwise rotation in the Yongsheng area is  $17.0 \pm 4.1^\circ$  relative to stable Eurasia. The rotation value is consistent at the 95% confidence level with results obtained from Paleogene and Cretaceous strata in other areas of the CDB. Their paleomagnetic data indicate that a consistent clockwise rotation of  $20.6 \pm 6.3^\circ$  occurred in different areas in the CDB: at Yongsheng, Zhupeng, Bailu, Dayao, Chuxiong west and Jianchuan (Figure 49).

That rotation process occurred into two discrete time intervals: approximately 11° of quasi-rigid clockwise rotation occurred between ~35 and 12.7 Ma compared to stable Eurasia. Subsequently, a clockwise rotation of about 10° of the CDB with respect to stable Eurasia is inferred from the paleomagnetic results since 12.7 Ma, which is consistent with the change in the geometry of the Ailaoshan–Red River Shear Zone around the western Chuxiong Basin. The later CW rotation was related to shearing movement of the Xianshuihe-Xiaojiang fault. They also ascribe that these faults movements and their related block rotation have a close relationship with the northward subduction of the Indian lithosphere beneath the Tibetan Plateau. Thus, according to other authors, the CDB was extruded along the left-lateral Xianshuihe-Xiaojiang fault system under the tectonic influence of India–Asia collision [Avouac and Tapponnier, 1993; Lacassin *et al.*, 1997], but its southward movement was eventually obstructed by the rigid Yangtze and Indochina blocks [Wang *et al.*, 2001; Gao *et al.*, 2017].

A compilation of paleomagnetic data set available from 112 localities around the Himalayan Syntaxis is well synthesized by Otofugi *et al.* [2010] and covers a period between the Early Cretaceous and Eocene (Figure 49 and Table 1). Other recent paleomagnetic works from Yunnan, mostly realized from Cretaceous continental red beds, showed a predominant clockwise (CW) rotation pattern determining about 1000 km in the 22°-25°N latitude range [Li *et al.*, 2017a; Tong *et al.*, 2016 and references therein] and indicate that rotation of crustal fragments in this region was not homogeneous, likely due to crustal heterogeneity and spatial and temporal stress field variations [Li *et al.*, 2017a; Metcalfe, 2002, 2011; Yin and Harrison, 2000].

Although this paleomagnetic data set from Yunnan is undoubtedly robust, the tectonic explanations for such rotational pattern remain is still not completely convincing. This research study will aim at clarifying it.



*Chapter*  
*VI*

## 6. PALEOMAGNETIC RESULTS AND MAGNETIC OVERPRINT EVALUATION:

### • THE CASE OF GAOLIGONG SHEAR ZONE

*Part of the following chapter was recently published on Tectonics by myself and my colleagues [Pellegrino et al., 2018].*

## Tectonics and Paleomagnetic Rotation Pattern of Yunnan (24°N–25°N, China): Gaoligong Fault Shear Versus Megablock Drift

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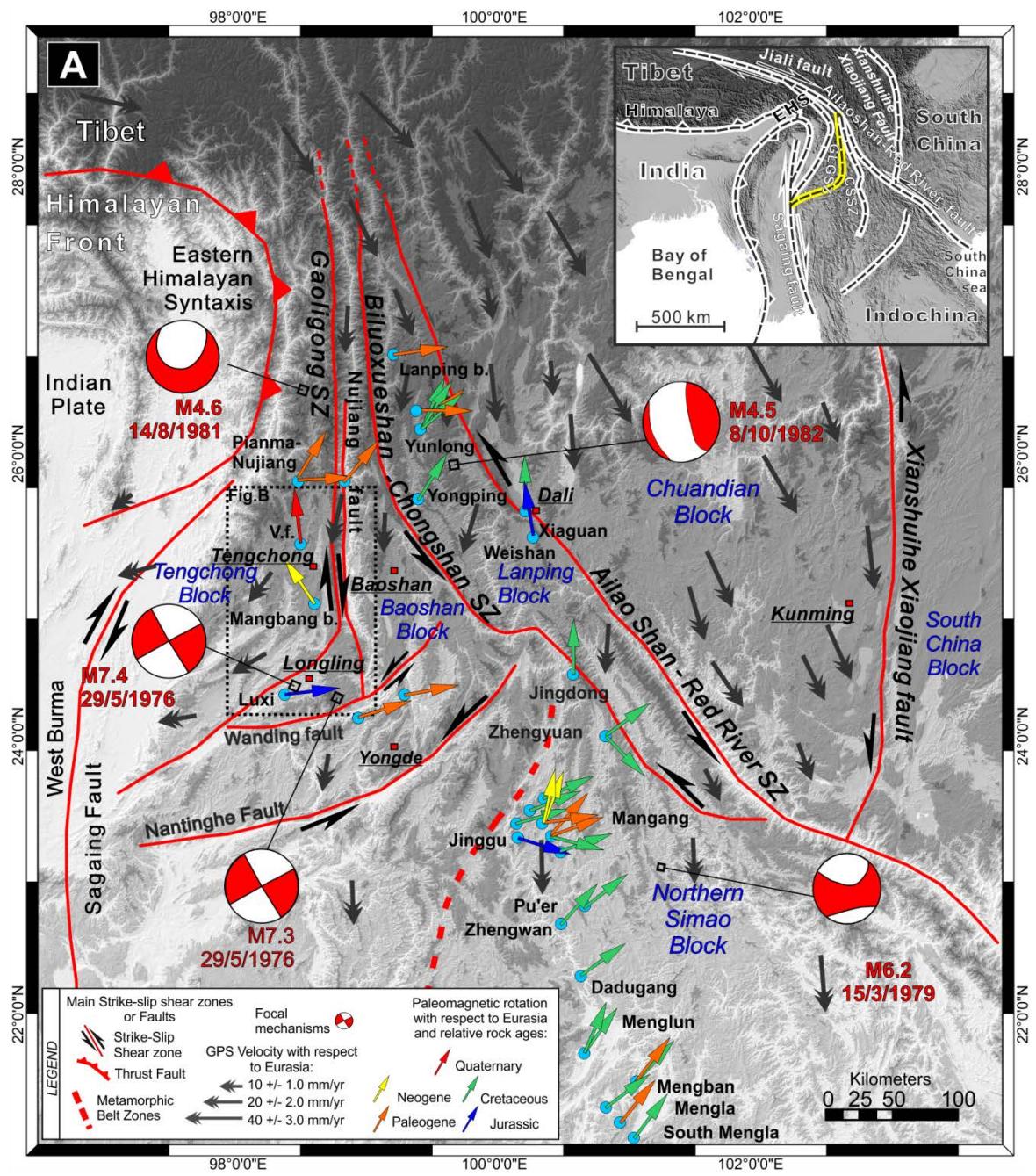
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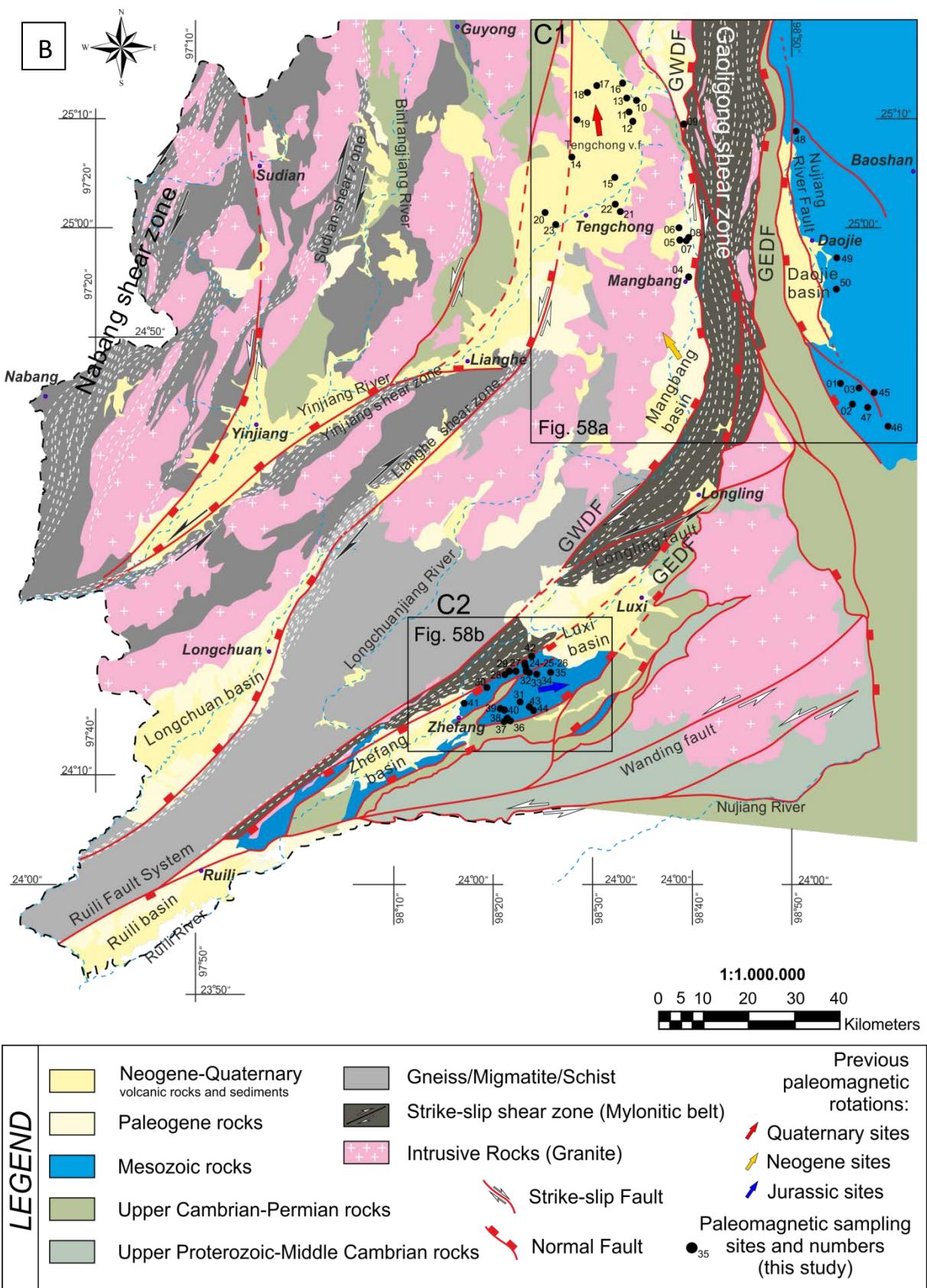
To assess the validity of models previously exposed and to better understand the structural architecture and crustal deformation processes in this region, field-based paleomagnetic and structural analyses along main strike-slip shear zone were performed. The first field work was conducted on April 2016 along the Gaoligong Shear Zone (99°E, 25°N) in Yunnan province (Figure 52A, B; 53). *Pellegrino et al. [2018]* sampled 50 paleomagnetic sites at variable distances (up to ~25 km) from the contact with mylonites and metamorphic rocks exposed along the shear zone (Figure 52).

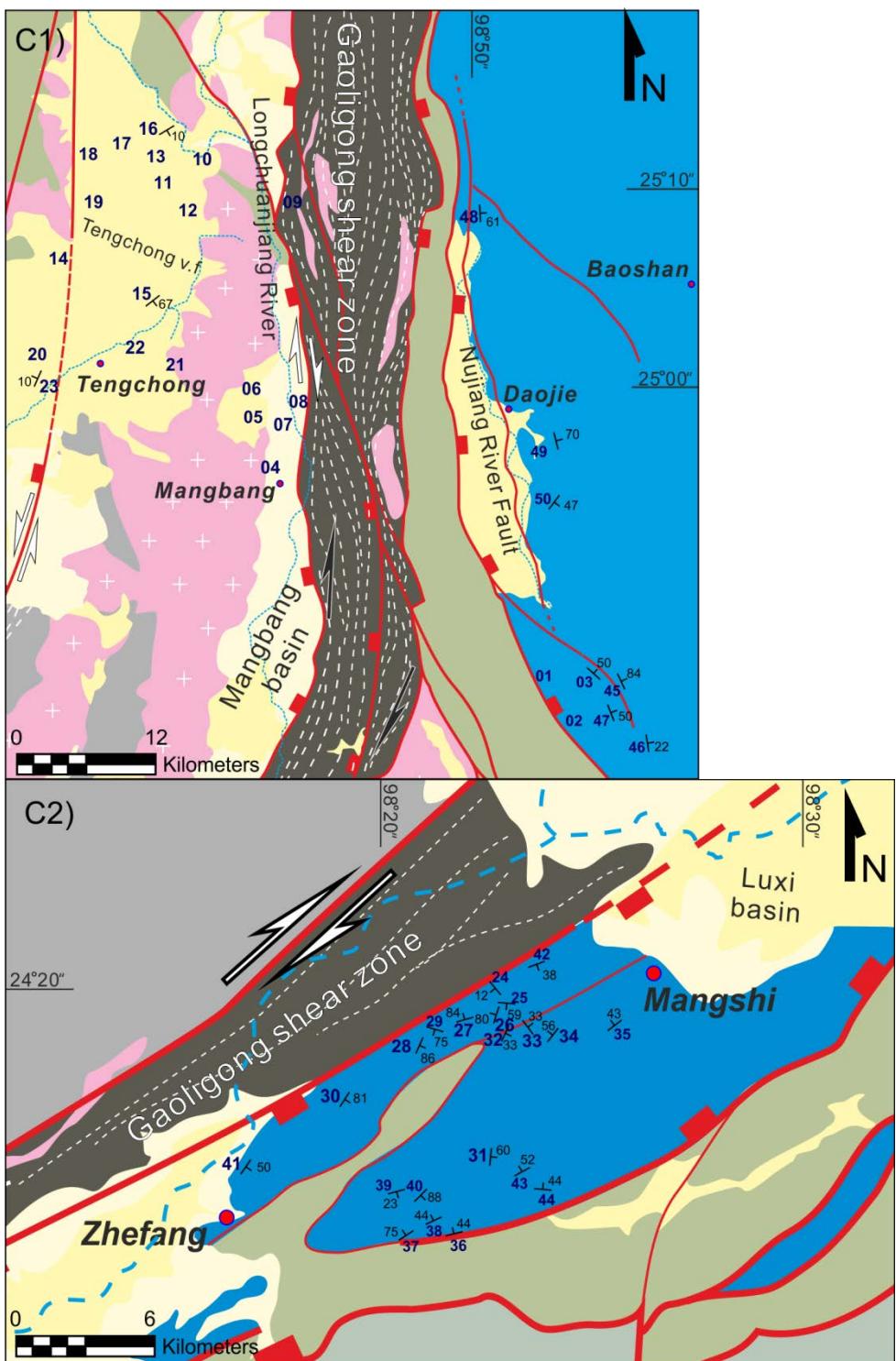
Sites were selected from three main areas to the west, east, and south of the Gaoligong Shear Zone, in the Tengchong, Baoshan, and Mangshi areas, respectively (Figures 52A, B, C1 and C2). From west to east, 15 paleomagnetic sites in Pliocene-Holocene lavas from the Tengchong volcanic field (Figure 52a,b,c1), 5 sites in Pliocene whitish lacustrine siltstones in Mangbang basin, 2 sites in Jurassic basalt and 28 sites in Jurassic-Cretaceous continental red

beds in the Baoshan block, have been gathered. The continental red beds are typically made of red fine-grained sandstone or siltstone. In the southern transect, the red beds are exposed in the Luxi Basin (Figure 52c2), delimited by two normal faults. As suggested by structural data, the Mesozoic rocks, in the basin outcropped, could represent the core of a monoclinal folded structure [*Huang and Opdyke, 1993*].

Bedding of Jurassic basalts was inferred considering the attitude of neighbor sediments, forming continuous volcano-sedimentary sequences according to available geological maps. In volcanic sites, several samples were gathered in the same flow, while samples from sedimentary sites were spread as much as possible on a given outcrop to try to average out the paleosecular variation (PSV) of the geomagnetic field. Each volcanic site is expected to record an almost instantaneous geomagnetic direction, so that PSV is surely not averaged out within the 17 sites from Pliocene-Holocene and Jurassic basalts.







**Figure 52.** Modified after Pellegrino et al. [2018]. **A)** Study area along the Gaoligong Shear Zone. Pink circles indicates the sampling locations. Focal mechanisms from Harvard Global CMT Catalog (see chapter 4.3) **B)** Simplified geological and tectonic map of the Gaoligong Shear Zone and adjacent areas. Boxes indicate the two areas of paleomagnetic investigation (see Figure 58a,b). GEDF = Gaoligong east detachment fault; GWDF = Gaoligong west detachment fault. **C1 and C2)** Structural details of the sites sampled. Sampled sites in blue.

**Table 2.** From Pellegrino et al. [2018]. Paleomagnetic site-mean directions from the Gaoligong Shear Zone (Yunnan, China)

Site	Geographic Coordinates		Rock Type	Series/ Epoch	Numerical Age (Ma)	Considered palpates age (Ma)	Bedding (deg)	n	ChRM/High-Temperature Component			k	$a_0$ (deg)	R (deg)	$\Delta R$ (deg)	F (deg)	$\Delta F$ (deg)												
									In Situ																				
	Latitude °N	Longitude °E							D (deg)	I (deg)	D (deg)																		
YUN01*	24.7766824	98.9282752	Basalts	MJ	163.174	170	80/20	10/10	52.6	-51.0	31.8	-67.1	39.3	7.8	-166.5	16.6	-6.44	6.8											
YUN02*	24.7437355	98.9482315	Basalts	MJ	163.174	170	80/20	5/10	27.9	106	29.1	-1.9	298.5	4.4	-169.2	5.9	58.7	4.6											
YUN03	24.7694685	98.9593802	Red Beds	MJ	163.174	170	42/50	7/10	346.9	81.4	33.5	34.8	40.3	9.6	15.2	10.3	25.8	8.1											
YUN04	24.9399358	98.6699462	Basalts	MP	3.6	—	—	8.8	196.3	-37.9	—	—	164.8	4.3	16.3	-6.1	5.1	4.3											
YUN05	24.9971886	98.6553678	Whitish Siltstones	LP	3.6-5.3	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN06	25.015892	98.6535897	Whitish Siltstones	LP	3.6-5.3	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN07	24.9967763	98.6655843	Whitish Siltstones	LP	3.6-5.3	—	—	6/10	5.8	31.4	—	—	13.7	18.8	5.8	22.1	11.6	18.8											
YUN08	24.9977035	98.6675056	Whitish Siltstones	LP	3.6-5.3	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN09	25.1686668	98.668388	Whitish Siltstones	LP	3.6-5.3	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN10	25.2115518	98.5821019	Basalts	MP	3.6	—	—	10/10	4.6	200	—	—	163.3	3.8	4.6	4.1	23.0	3.8											
YUN11	25.19456	98.5693205	Basalts	LPS	1.5-2.5	—	—	9/10	343.5	23.5	—	—	53.5	7.1	-16.5	7.7	20.5	7.1											
YUN12	25.1794776	98.5758987	Basalts	LPS	1.5-2.5	—	—	10/10	12.2	47.1	—	—	40.3	7.7	12.2	11.3	-4.1	7.7											
YUN13	25.213207	98.5749002	Basalts	MP	3.6	—	—	10/10	15.5	35.1	—	—	184.4	3.6	15.5	4.4	7.9	3.6											
YUN14	25.1249219	98.4719548	Basalts	H	Present- 0.01/17	—	—	9/11	2.9	24.4	—	—	261.3	3.2	2.9	3.5	18.6	3.2											
YUN15	25.0935636	98.5443339	Basalts	LPS	1.5-2.5	—	—	14/267	10/10	324.7	20.2	—	20.8	10.8	-35.3	11.6	22.8	10.8											
YUN16	25.238754	98.5577044	Basalts	H	Present- 0.01/17	—	—	149/10	10/10	6.8	34.7	12.0	42.3	176.6	3.6	6.8	4.4	8.3	3.6										
YUN17	25.234284	98.5142608	Basalts	H	Present- 0.01/17	—	—	9/11	11.5	35.0	—	—	89.9	5.5	11.5	6.8	8.0	5.5											
YUN18	25.2238856	98.4980389	Basalts	H	Present- 0.01/17	—	—	10/11	8.3	33.7	—	—	46.9	7.1	8.3	8.6	9.3	7.1											
YUN19	25.1826527	98.4799035	Basalts	H	Present- 0.01/17	—	—	9/10	356.0	16.9	—	—	84.3	5.6	-4.0	5.9	26.1	5.6											
YUN20	25.0394636	98.4258614	Basalts	H	Present- 0.01/17	—	—	9/10	3.7	28.5	—	—	297.0	3.0	3.7	3.4	14.5	3.0											
YUN21	25.0406369	98.5337038	Basalts	LPS	1.5-2.5	—	—	10/10	179.1	-42.8	—	—	63.7	6.1	-0.9	8.4	0.2	6.1											
YUN22	25.0521541	98.5450541	Basalts	LPS	1.5-2.5	—	—	10/11	230.0	-47.3	—	—	50.5	6.9	50.0	10.3	-4.3	6.9											
YUN23	25.0208672	98.4448747	Basalts	H	Present- 0.01/17	—	—	294/10	10/10	5.2	29.2	0.3	25.5	92.5	5.1	5.2	5.9	13.8	5.1										
YUN24	24.3443875	98.3921232	Red Beds	K	66-145	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN25	24.3405986	98.3929158	Red Beds	J-K	66-201	160	185/59	7/12	165.5	-72.7	13.4	-47.1	10.1	20.0	176.2	24.0	9.4	16.0											
YUN26	24.3347198	98.394079	Red Beds	K	66-145	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN27	24.331493	98.3774904	Red Beds	K	66-145	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN28	24.32577171	98.35766892	Red Beds	K	66-145	—	—	—	—	—	—	—	—	—	—	—	—	—											
YUN29	24.3314523	98.36594508	Red Beds	K	66-145	—	—	164/75	9/11	302.9	-21.1	294.8	35.9	10.7	16.5	-73.7	16.3	12.6	13.2										
YUN30	24.3064965	98.3273209	Red Beds	K	66-145	—	—	—	—	—	—	—	—	—	—	—	—	—											

Table 2. Paleomagnetic site-mean directions from the Gaoligong Shear Zone (Yunnan, China)

Site	Geographic Coordinates		Rock Type	Series/ Epoch	Numerical Age (Ma)	Considered paleopoles age (Ma)	Bedding (deg)	CHRM/High-Temperature Component			k	$\alpha_{95}$ (deg)	R (deg)	$\Delta R$ (deg)	F (deg)	$\Delta F$ (deg)											
	Latitude °N							In Situ																			
								D (deg)	I (deg)	D (deg)																	
YUN31	24.2846044	98.3841427	Red Beds	MJ	163-174	—	—	—	—	—	—	—	—	—	—	—											
YUN32	24.3316463	98.393859	Red Beds	K	66-145	—	—	—	—	—	—	—	—	—	—	—											
YUN33	24.3289722	98.39939239	Red Beds	MJ	163-174	—	—	—	—	—	—	—	—	—	—	—											
YUN34	24.326614	98.41226722	Red Beds	MJ	163-174	—	—	—	—	—	—	—	—	—	—	—											
YUN35	24.3294147	98.4357429	Red Beds	J-K	66-201	160	323/43	5/10	351.3	18.2	351.6	-20.0	86.8	8.3	154.4	8.5											
YUN36	24.255107	98.3670267	Red Beds	MJ	163-174	160	346/44	9/10	108.3	25.1	80.0	40.0	70.2	6.2	62.7	8.0											
YUN37	24.2530733	98.3559025	Red Beds	MJ	163-174	160	327/75	9/10	92.8	7.6	58.6	36.5	106.7	5.0	41.2	6.9											
YUN38	24.2588832	98.361492	Red Beds	MJ	163-174	160	335/44	6/10	118.6	-4.6	111.9	30.0	14.8	18.0	94.7	17.0											
YUN39	24.2737808	98.3497701	Red Beds	MJ	163-174	160	165/23	8/8	70.7	32.3	85.1	30.9	17.0	13.8	67.9	13.5											
YUN40	24.2716601	98.3572075	Red Beds	MJ	163-174	160	135/88	8/10	21.5	25.5	71.8	22.0	14.3	15.1	54.6	13.7											
YUN41	24.2826614	98.28916	Red Beds	K	66-145	—	—	—	—	—	—	—	—	—	—	—											
YUN42	24.3547169	98.40384259	Red Beds	K	66-145	100	160/38	5/5	14.0	38.1	53.1	62.6	51.7	10.7	45.1	18.8											
YUN43	24.2767445	98.3995939	Red Beds	MJ	163-174	160	331/52	7/10	87.8	29.0	50.8	37.5	41.1	9.5	33.6	10.6											
YUN44	24.2710038	98.4064307	Red Beds	MJ	163-174	160	5/44	10/10	110.1	30.2	82.3	31.2	69.3	5.8	65.1	7.3											
YUN45	24.7620875	98.985785	Red Beds	MJ	163-174	160	65/84	9/10	260.8	52.7	52.3	41.5	15.7	13.4	34.6	14.9											
YUN46	24.7103859	99.0088921	Red Beds	MJ	163-174	160	81/22	9/10	6.9	44.2	23.8	35.0	19.7	11.9	6.6	12.4											
YUN47	24.739466	98.9747164	Red Beds	MJ	163-174	160	69/50	10/10	319.5	70.3	42.9	43.7	45.8	7.2	25.7	9.3											
YUN48	25.1643668	98.8513117	Red Beds	MJ	163-174	—	—	—	—	—	—	—	—	—	—	—											
YUN49	24.9697763	98.9220494	Red Beds	MJ	163-174	160	76/70	10/10	310.8	58.1	44.2	35.2	29.0	9.1	27.0	10.1											
YUN50	24.9209456	98.9212281	Red Beds	MJ	163-174	160	127/47	10/10	341.1	63.1	94.3	61.9	32.7	8.6	77.1	15.3											

Ages: J=Jurassic, MJ=Middle Jurassic, K=Cretaceous, LP=Lower Pliocene, MP=Middle Pliocene, LPS=Lower Pleistocene, H=Holocene. The geographic coordinates are referred to WG84 datum. Age in Ma is from the geologic timescale of Coen et al. [2013]. Bedding is expressed in dip azimuth/dip values. D and I are site mean declination and inclination calculated before and after tectonic correction; k and  $\alpha_{95}$  are statistical parameters after Fisher [1953]; n/N is number of samples giving reliable results at a site. Site mean rotation R and flattening F values, and relative errors  $\Delta R$  and  $\Delta F$  (according to Demarest [1983]) are relative to coeval D and Eurasian values expected at the sampling area considering Eurasian paleopoles. From Torsvik et al. [2012] and Cogné et al. [2013] (in bold). R and F for Holocene sites were evaluated comparing the paleomagnetic directions with local geocentric axial dipole field direction.\* Disarded site (see text).

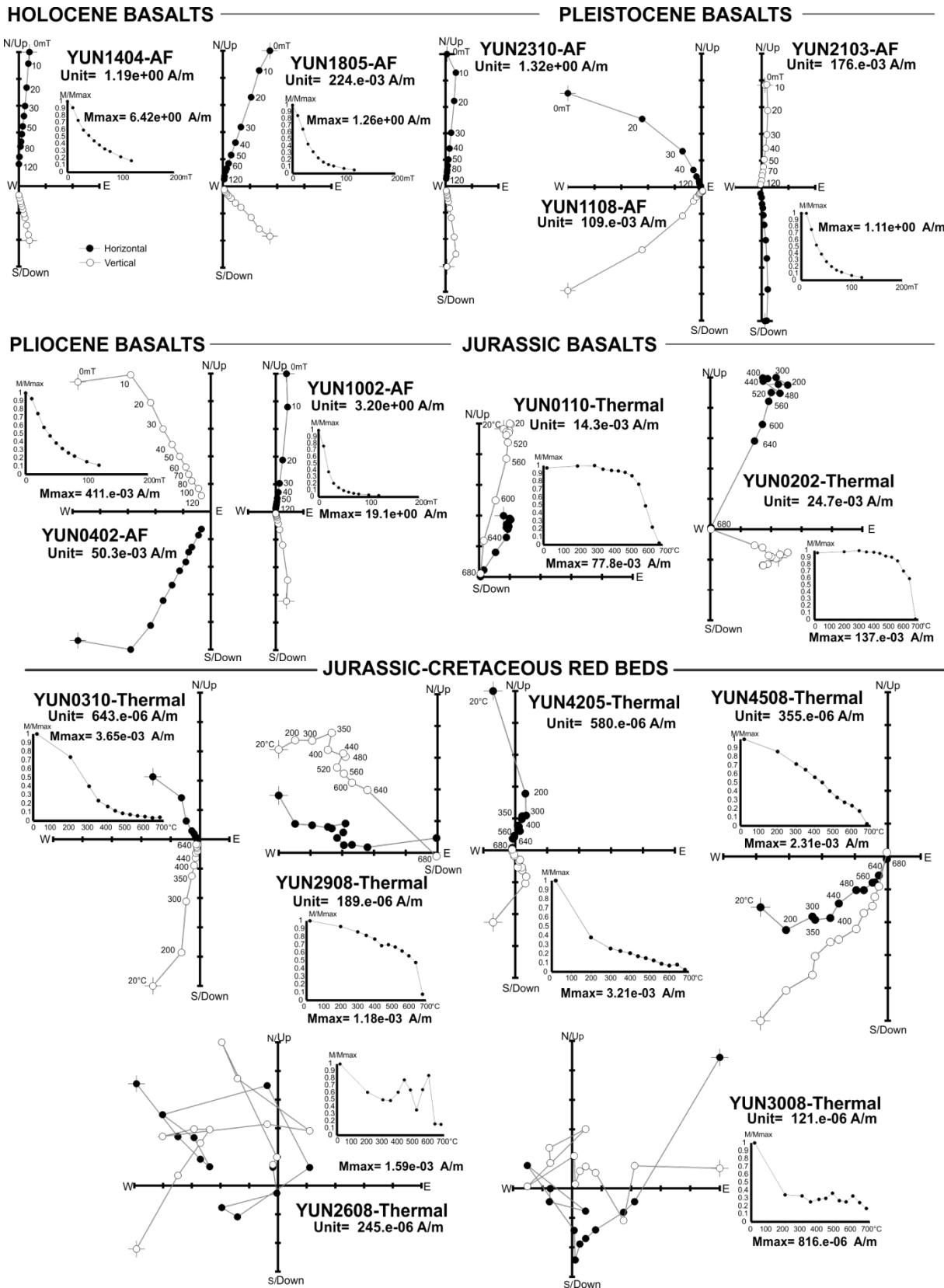
Thirty-five (18 Jurassic-Cretaceous red beds, 2 Jurassic basalts, 15 Pliocene-Holocene basalts, 1 Lower Pliocene whitish siltstone) out of 50 sites gave reliable paleomagnetic results, the remaining 14 sites yielded erratic demagnetization diagrams and were discarded from further consideration (Table 2). All Pliocene lacustrine siltstone samples, except site Yun07, gave not-interpretable diagrams.



**Figure 53.** Field pictures of some paleomagnetic sites. a) Yun03; b) Yun08; c) Yun10; d) Yun16; e) Yun13; f) Yun26; g) Yun49. See figure 52 for location and Table 2 for details.

In the AF demagnetized volcanic rocks, a viscous component was removed between 10 and 30 mT, and a characteristic magnetization (ChRM) was isolated between 30 and 120 mT

(Figure 54), confirming the occurrence of magnetite and titanomagnetite. In the thermally demagnetized Jurassic basalts, a ChRM was isolated between 480 and 680 °C.



**Figure 54.** From *Pellegrino et al.* [2018]. Orthogonal vector diagrams of typical demagnetization data (*in situ* coordinates) showing representative samples carrying characteristic magnetization components (ChRMs), low/high temperature components (Yun2908, Yun4508), and scattered magnetization (Yun2608, Yun3008). Solid (open) symbols represent projection onto the horizontal (vertical) plane. Demagnetization step values are in milli Tesla (mT) and °C for the alternating field and thermally demagnetized samples, respectively. (See supplementary information to examine all diagrams).

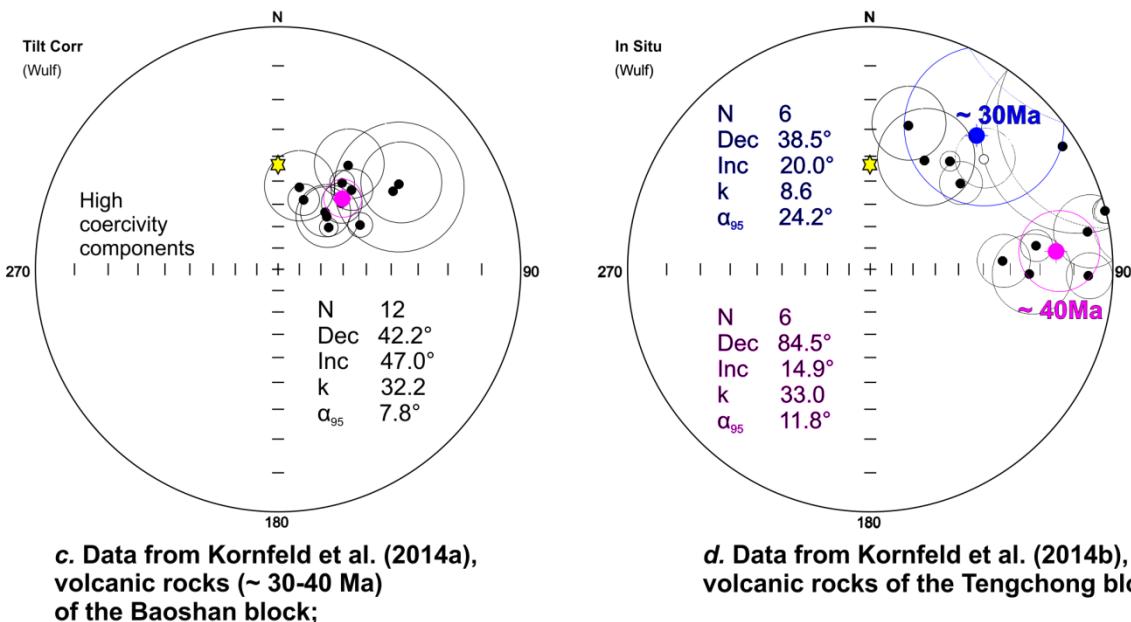
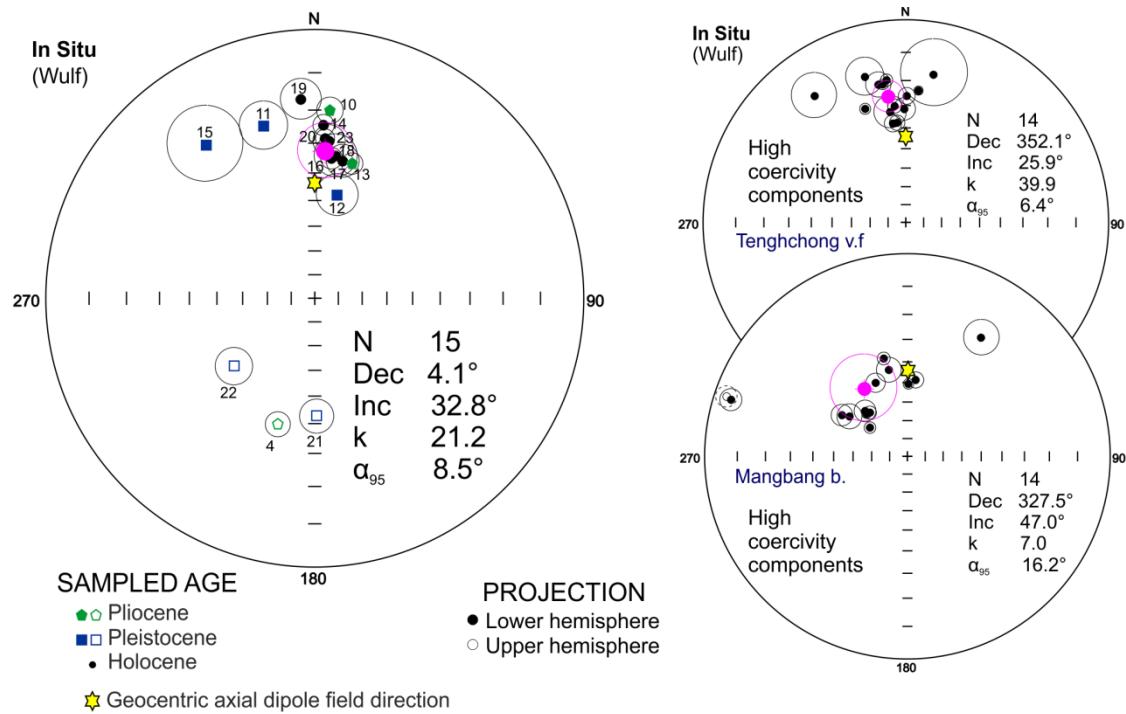
In the red beds samples, after the elimination of a viscous component up to 200°C, a ChRM was isolated in about half of the samples between 200° and 680°C. In the remaining samples, a low- and high-temperature (HT) components were defined in the 200°-400°C and 400°-680° intervals, respectively (sample Yun2908 in Figure 54). The high unblocking temperature spectra and Curie temperature of 680°C combined with high coercivity spectra show that hematite is the main magnetic carrier in the red beds (confirming evidence from previous paleomagnetic studies from Yunnan, e.g. *Sato et al.*, 2007; *Tong et al.*, 2013, 2016). Several studies carried out over nearly the last fifty years have showed that hematite in red beds may have both a detrital and a chemical origin (see review by *Jiang et al.*, 2015).

By combining ChRMs and HT components, well-defined site-mean directions ( $3.0^\circ \leq \alpha_{95} \leq 20.0^\circ$ ,  $9.1^\circ$  on average) were obtained from 15 Pliocene-Holocene volcanic sites, 1 Pliocene lacustrine site, 18 Jurassic-Cretaceous red bed sites, and 2 Jurassic volcanic sites (Table 2).

In Figure 55, are shown the paleomagnetic site mean directions from Pliocene-Holocene basalts obtained by *Pellegrino et al.* [2018] (Figure 55a) and Eocene-Pliocene volcanics sampled in neighbour regions by *Kornfeld et al.* [2014a; 2014b; 2014c; Figures 55b-55d].

*Pellegrino et al.* [2018] and *Kornfeld et al.* [2014c] data, come from the same rocks from the Tengchong volcanic field, so that they should compare. The results are effectively the same (considering respective confidence cones), and show mostly northward declinations in the Tengchong volcanic field. Data by *Pellegrino et al.* [2018] reveal higher inclination with respect to data by *Kornfeld et al.* [2014c]. Moreover, three sites (two Pleistocene and one

Pliocene in age) yield reverse polarities and inclinations as high as  $-37.9^\circ$  - $47.3^\circ$ . Therefore data by Pellegrino *et al.* [2018] demonstrate that the Tengchong volcanic field is older than 0.78 Ma (age of the Brunhes-Matuyama transition; geomagnetic polarity scale from Ogg, 2012), confirming radiometric data by Wang *et al.* [2008, and references therein].

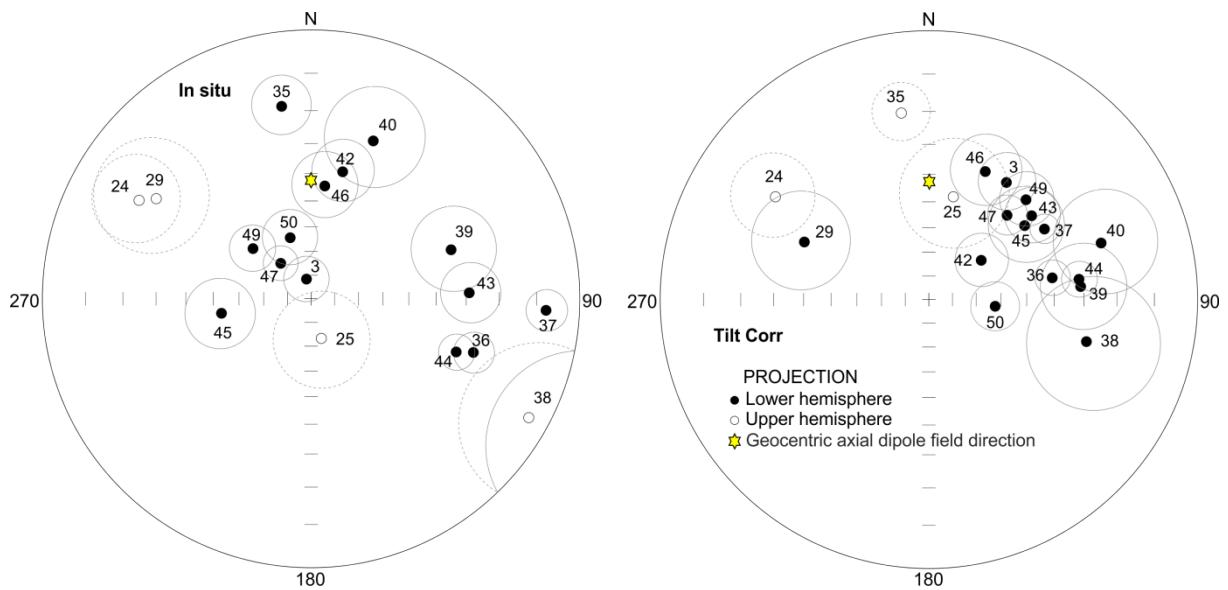


**Figure 55.** Equal-angle projection of the site-mean paleomagnetic directions gathered by *Pellegrino et al.* [2018] from Pliocene-Holocene basalts exposed west of the Gaoligong Shear Zone (a), and by *Kornfeld et al.* [2014c]; *Kornfeld et al.* [2014a]; *Kornfeld et al.* [2014b] (b–d). N = number of sites; Dec = declination; Inc = inclination; k = precision parameter;  $a_{95}$  = 95% confidence angle. The yellow stars represent the normal polarity geocentric axial dipole field direction ( $D = 0^\circ$ ;  $I = 43^\circ$ ) expected at the study area. Solid (open) symbols represent projection onto the lower (upper) hemisphere, and open circles represent confidence cones. Pink points and open circles represent the mean paleomagnetic directions and the relative 95% confidence cones, respectively.

The lower inclination with respect to the geocentric axial dipole (GAD) field inclination expected at the study area may be due to the incomplete averaging of the PSV of the geomagnetic field, as they got samples from only fifteen lava flows. Thus, *Pellegrino et al.* [2018] put forward an alternative explanation for the low normal polarity inclinations from the Tengchong volcanic field than the regional tilting proposed by *Kornfeld et al.* [2014c].

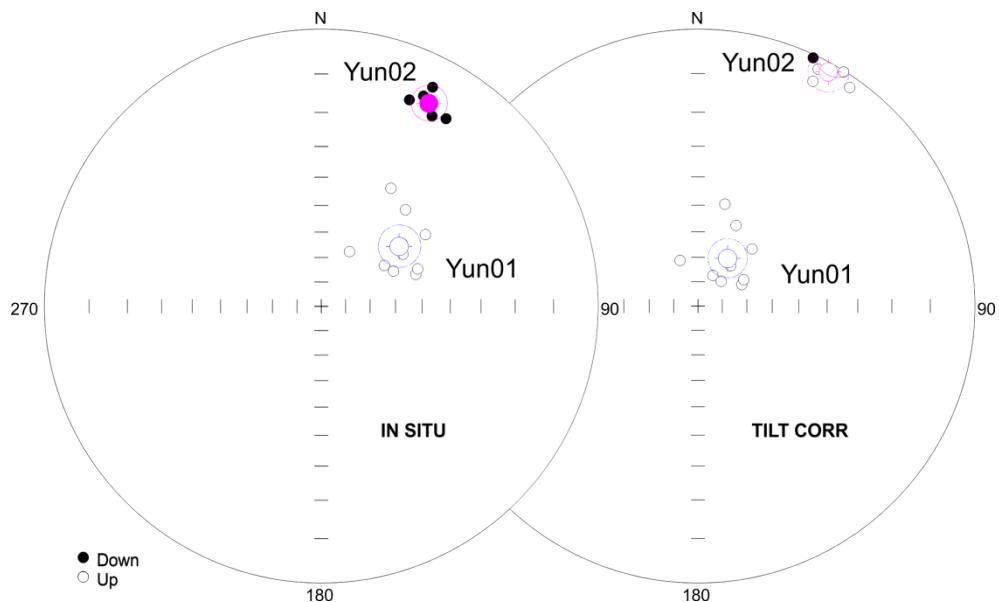
In Figure 55b, are also shown data from *Kornfeld et al.* [2014c] obtained from upper Miocene-Pliocene volcanic rocks of the Mangbang Basin. They explained the CCW rotations as apparent rotations due to tilting along the half graben hosting the basin itself. *Kornfeld et al.* [2014a] also documented a  $40^\circ$  CW rotation in 30 Ma basalts from the Baoshan block (Figure 55c) and a CW rotation of  $87^\circ$  in 40 Ma dykes from the Tengchong block (Figure 55d).

Paleomagnetic directions from Jurassic-Cretaceous red beds sampled east of the Gaoligong Shear Zone are shown in Figure 56.



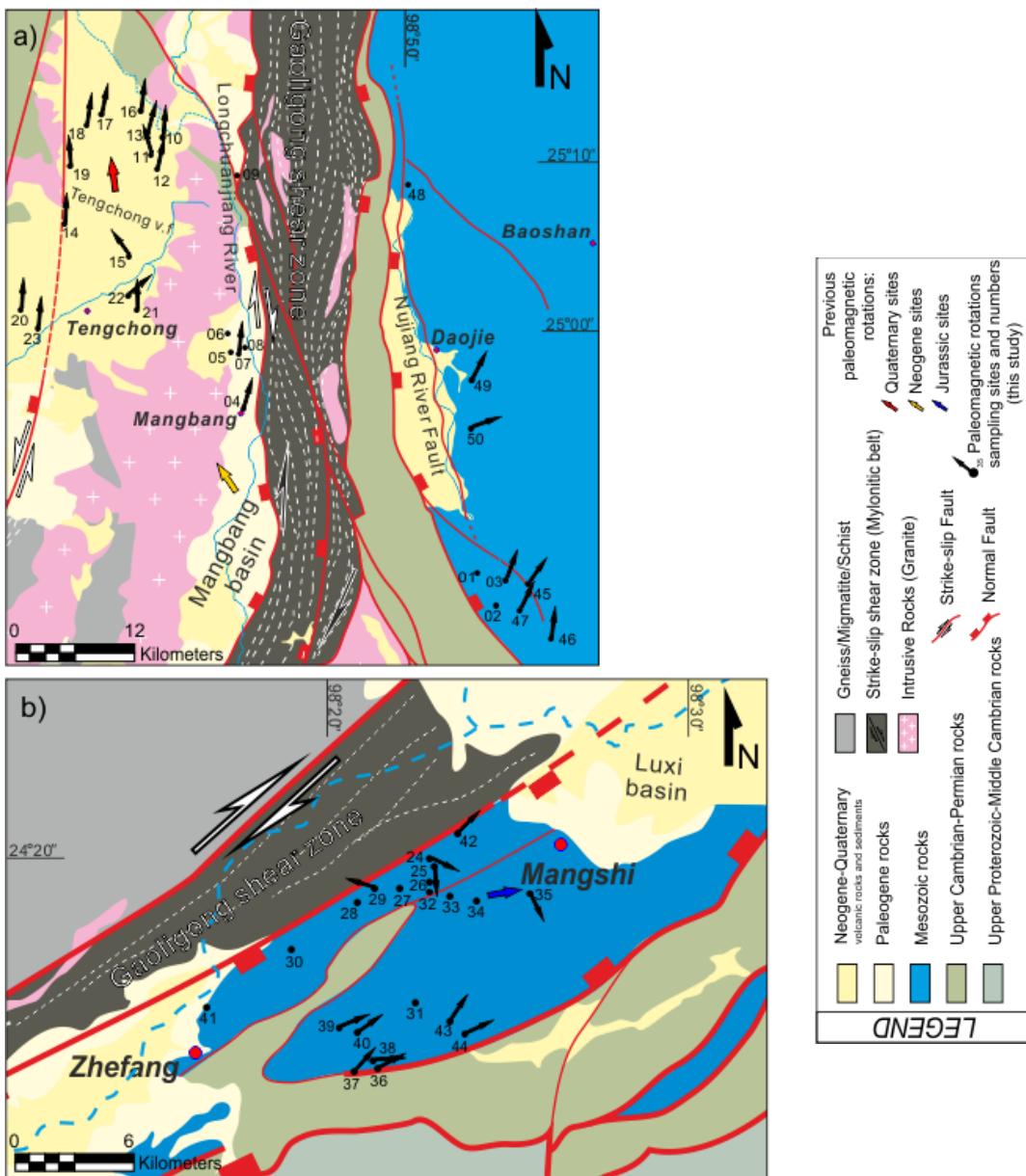
**Figure 56.** Modified after *Pellegrino et al.* [2018]. Equal-angle projection of the paleomagnetic directions obtained from Jurassic-Cretaceous red beds exposed in the Baoshan and Mangshi areas, east of the Gaoligong Shear Zone (see Figure 52-58a-b).

Site-mean directions were evaluated by averaging out ChRMs and HT components from the individual specimens. Both the in-situ and tilt-corrected directions are far from the GAD field direction expected at the sampling localities. Most of the tilt corrected directions (15 out of 18 sites) are of normal polarity and show variable amounts of westward declinations (exceeding  $100^\circ$  for site Yun38). Directions from Jurassic volcanics (sites Yun01-02) were discarded, as they were completely different from nearby red bed site directions (Figure 58 and Table 2). Moreover, site Yun02 yielded a sub-horizontal tilt-corrected direction that is incompatible with the expected Jurassic inclinations ( $42^\circ$ - $61^\circ$ , e.g. *Torsvik et al.*, 2012).



**Figure 57.** Equal-angle projection of the paleomagnetic directions obtained from Jurassic basalt exposed in the Baoshan areas, east of the Gaoligong Shear Zone. Pink points and open circles represent the mean paleomagnetic directions and the relative 95% confidence cones, respectively (See Table 2 for mean values).

The paleomagnetic data from red beds support a positive fold test at the 99% significance level according to *McFadden* [1990] ( $N=18$ ;  $S_{cos}^{insitu}=11.9$ ;  $S_{cos}^{tilt\ corrected}=1.1$ , critical  $S_{cos}$  at the 99% significance level =6.9). Maximum k and minimum  $S_{cos}$  values are obtained at 93% and 95% of complete unfolding, respectively [Pellegrino *et al.*, 2018]. This suggests a pre-folding (i.e. pre-Eo-Oligocene) [Chen *et al.*, 1995] magnetization. Furthermore, site Yun38 showed dual polarity samples (and consistent polarities in the same beds), indicating that the sampled strata encompass different polarity chronos (the reversal test on individual ChRM's according to *McFadden and McElhinny*, 1990 is indeterminate). This proves that, at least for site Yun38, the magnetization is of primary origin. The normal and reverse polarity site-mean directions are not antipodal (negative fold test according to *McFadden*, 1990), probably due to local block rotations [Pellegrino *et al.*, 2018].



**Figure 58.** Modified after Pellegrino *et al.* [2018]. Paleomagnetic sampling areas in the northern (a, Baoshan and Tengchong area) and southern (b, Mangshi basin) segment of the Gaoligong Shear Zone. Black arrows indicate paleomagnetic rotations from Jurassic to Holocene sediments and basalts (see Table 2). Rotation values were calculated with the same method and Eurasian poles as in Figure 49 and 52A.

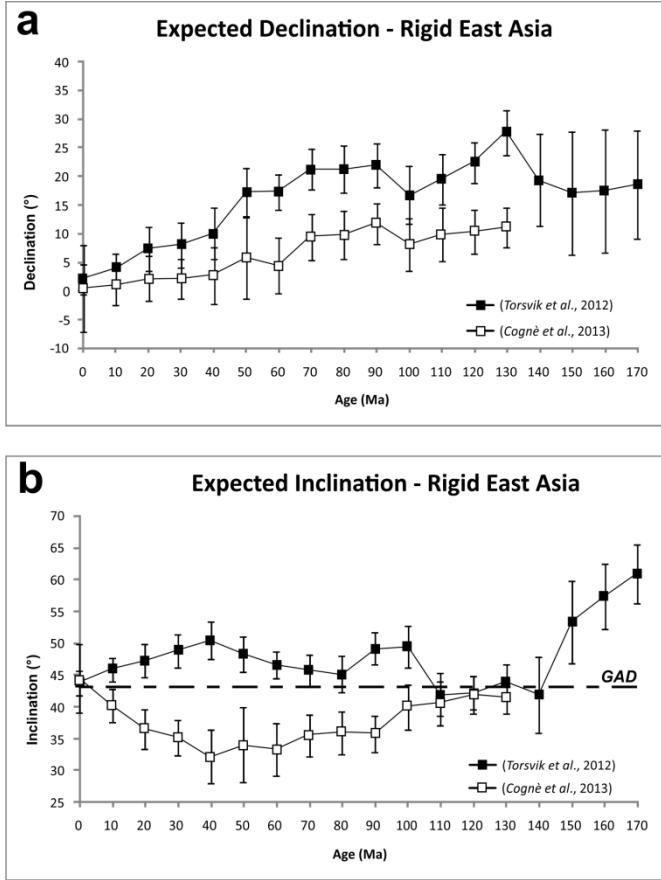
It is noteworthy that three sites (Yun24, 25 and 35) from our red bed data set yield a reverse polarity (and Yun38 dual polarity samples), while previous studies of red beds from the Lanping and Simao blocks systematically reported a normal polarity, for which a magnetic overprint has been suggested [Chen *et al.*, 1995; Gao *et al.*, 2015; Huang and Opdyke, 1993; Li *et al.*, 2017b]. Red beds sites Yun24, 25 and 35 support themselves a positive fold test at

the 95% significance level according to *McFadden* [1990] ( $S_{cos\text{insitu}} = 2.6$ ;  $S_{cos\text{tiltcorr}} = 1.8$ , critical  $S_{cos}$  at the 95% significance level = 2.1, maximum K and minimum  $S_{cos}$  values arise at 100% of complete unfolding). *Pellegrino et al.* [2018] data support for sure a pre-folding (i.e. pre-Eo-Oligocene) magnetization of the whole data set, and the reverse polarity sites suggest that magnetization is possibly of primary origin. In this case, most of the studied sites would have been deposited during the long normal Cretaceous superchron. Detailed age evaluation is often problematic in continental sediments, so that the issue of age determination and possible lower Tertiary pre-folding remagnetization of most of the sites remains open at present (see also similar conclusions by *Li et al.*, 2017b for the Simao block).

Although sites Yun24, 25, 29 and 42 are located at less than 2 km from the contact with the mylonites of the Gaoligong fault, a post-rotation (i.e. post-Oligo-Miocene) magnetic overprint is excluded (with the possible exception of site Yun42) by the evidence that both the *in situ* and tilt-corrected directions are far from the GAD field direction (Figure 56 and Table 2). The apparently puzzling lack of a late Tertiary magnetic overprint for sites adjacent to mylonites is explained considering that the high-grade rocks were exhumed along an extensional detachment that presumably elided the whole upper crust, juxtaposing high-grade metamorphic rocks and unmetamorphosed sediments [*Zhang et al.*, 2011].

## 6.1 ROTATION PATTERN ALONG THE GAOLIGONG SHEAR ZONE

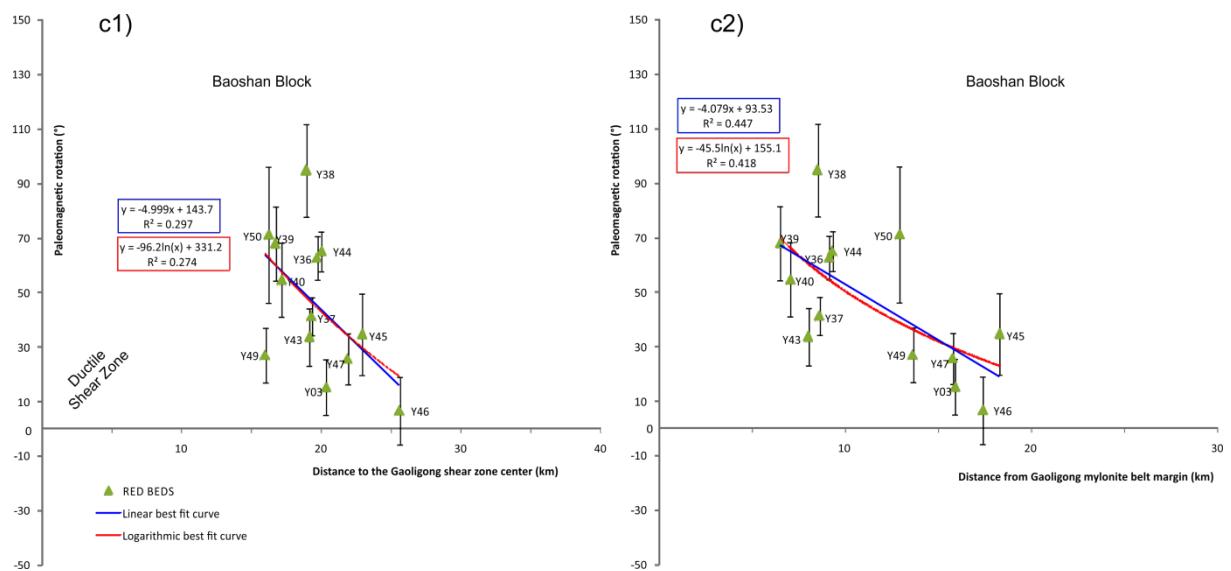
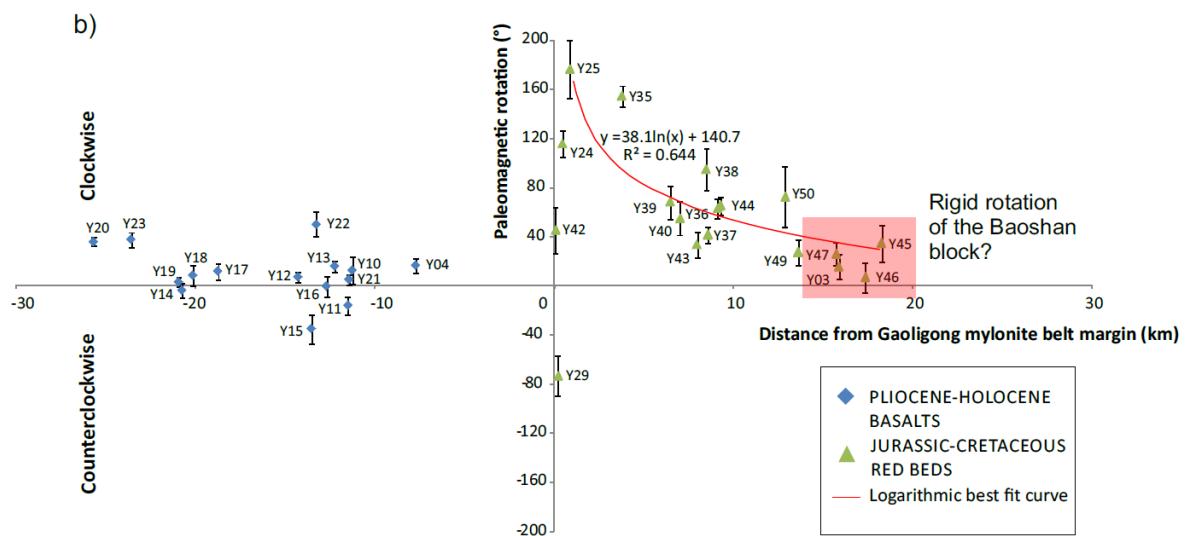
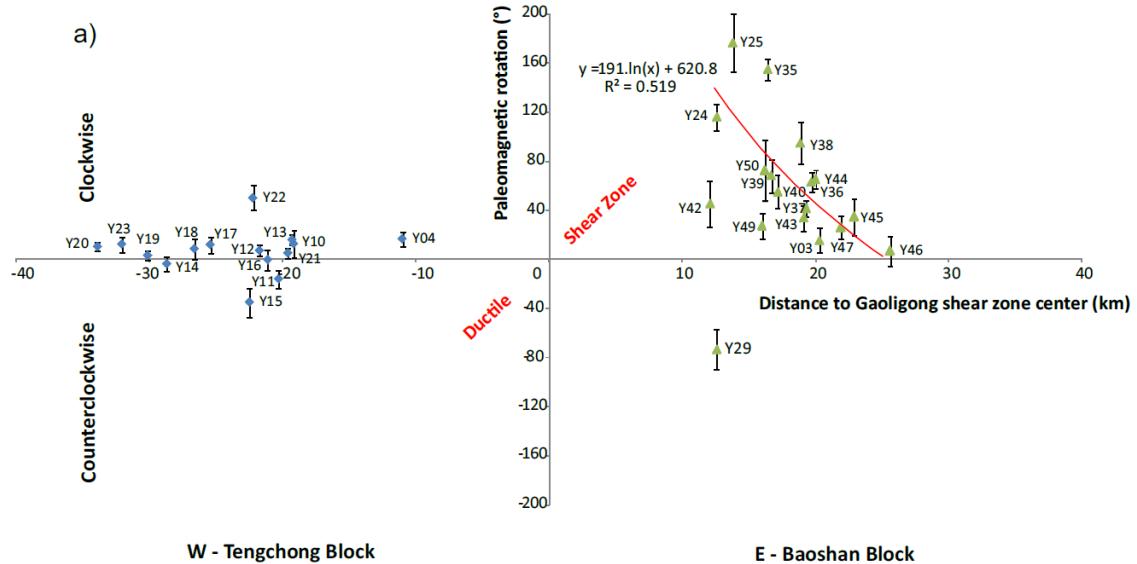
To evaluate tectonic rotations with respect to Eurasia, the paleomagnetic directions were compared to coeval Eurasian paleopoles from *Torsvik et al.* [2012] and *Cogné et al.* [2013] (see details in chapter 3 Figure 25-26). Sites younger than 5 Ma (i.e. Pliocene to Holocene in age) were compared to the local GAD field direction, assuming that no paleomagnetically significant East Asia rotation or drift occurred after 5 Ma. The expected Eurasian declinations and inclinations at Yunnan are shown in Figure 59.



**Figure 59.** From Pellegrino *et al.* [2018]. Expected declination (a) and inclination (b) values (and relative error bars) in the study area, considering Eurasian (from 170 to 140 Ma) and East Asia (from 130 to 10 Ma) paleopoles by Torsvik *et al.* [2012] and Cogné *et al.* [2013], respectively. Reference point is at  $25.21^{\circ}\text{N}$ ,  $98.58^{\circ}\text{E}$ . The black line indicates the geocentric axial dipole field inclination ( $43^{\circ}$ ) for the study area.

Inclination data show rapid changes in the 180-150 Ma time window, followed by contrasting paths when the two paleopoles lists are considered. Expected declinations are always positive, although they are significant (around  $20^{\circ}$ ) only at 180-150 Ma, between 140 and 80 Ma they are around  $10^{\circ}$ - $20^{\circ}$ , and are progressively annulled afterwards.

Site rotations with respect to their distance to the Gaoligong Shear Zone are shown in Figure 56. Site distances are calculated with respect to both the Shear Zone center (Figure 60a), and the contact with the fault mylonites (Figure 60b).



**Figure 60.** From *Pellegrino et al.* [2018]. Plot of rotations with respect to Eurasia versus site distance to both Gaoligong Shear Zone center (a) and mylonite- sediment contact (b). See Figure 52-58a-b for site location. Red lines are logarithmic best fit curves for the red bed sites located east of the Gaoligong Shear Zone (disregarding sites Yun24, Yun29, and Yun42 located adjacent to fault mylonites). (c) Detail of rotations with respect to Eurasia versus site distance to both Gaoligong Shear Zone center (c1) and mylonite- sediment contact (c2) disregarding site Yun24, 25, 29, 35 and 42. Red and blue lines are logarithmic and linear best fit curves. Sense and amount of rotation are defined by the smaller angle between observed and expected Eurasian declinations, thus rotation values are  $\leq |180^\circ|$  by definition.

To the west of the fault, Pliocene-Holocene basalts overall show null rotations, consistently with results from the Pliocene whitish silt site Yun07. As discussed above, PSV from basalt sites is likely not averaged out, and we may expect a geomagnetic declination scatter that mostly ranges within  $\pm 20^\circ$  at such latitudes [*Hernandez-Moreno et al.*, 2016, and discussion therein]. Thus a bias can exist on the average null rotation obtained from volcanic data, but such error is expected to be much smaller than the large CW rotations (up to  $176^\circ$ ) observed in Mesozoic red beds. *Pellegrino et al.* [2018] conclude that data from basalts -although probably scattered by PSV effects–do prove that rotations ceased in Yunnan by Early Pliocene (5 Ma) times. Red beds east of the fault yield always CW rotations (apart from site Yun29), reaching values as high as  $176^\circ$  for site Yun25. Apart for scattered rotations adjacent to the fault (sites Yun24, 29, 42), in both plots there is a clear trend of decreasing rotation values moving away from the fault. Rotations are virtually annulled at ca. 20 km from mylonite-sediment contact (Figure 60b). Sites with number of directions less than 8/10 samples (eg. Yun25 and 35; Figure c) have been excluded, however, as clearly shown in figures (c1) and (c2), the rotation pattern is still the same in both cases (distance to the Gaoligong Shear Zone center and from Gaoligong mylonite belt margin). Linear and logarithmic best fit curves fit well and are in agreement with the interpretation proposed by *Pellegrino et al.* [2018]. There is not clear evidence, in the field, of the width of the rotating blocks. By scrutinizing rotations of sites Yun24, 25, 29, and 42, located adjacent to the

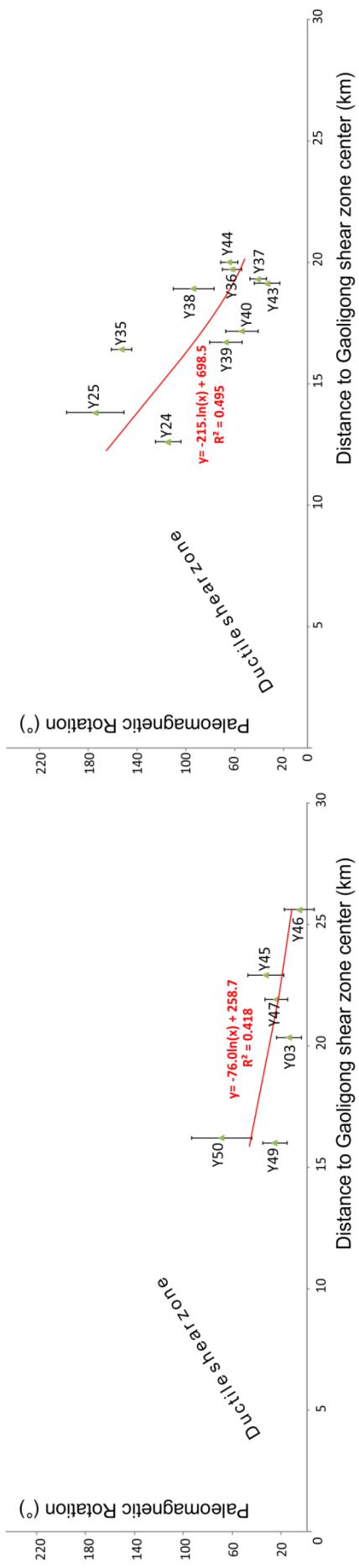
Gaoligong Shear Zone contact (Figure 60), rotations change abruptly from close sites, suggesting that the crust along the Gaoligong damage zone is broken into small blocks, width in the order of 1 km, or smaller (see sites Yun24 and Yun25). Obviously, block size seems to increase moving away from the fault, and in fact sites located in the southern boundary of the Mangshi basin (Figure 58b-60) show similar rotations over a distance of 6-7 km.

Rotations of sites with respect to the distance of the Gaoligong Shear Zone center and mylonite belt margin, in northern and southern transect, have been shown in Figure 61. There is a clear trend of decreasing rotation values moving away from the fault. All these tests, which gave consistent results, are useful for further confirming the proposed interpretations (Figures 60 c1,c2 and 61).

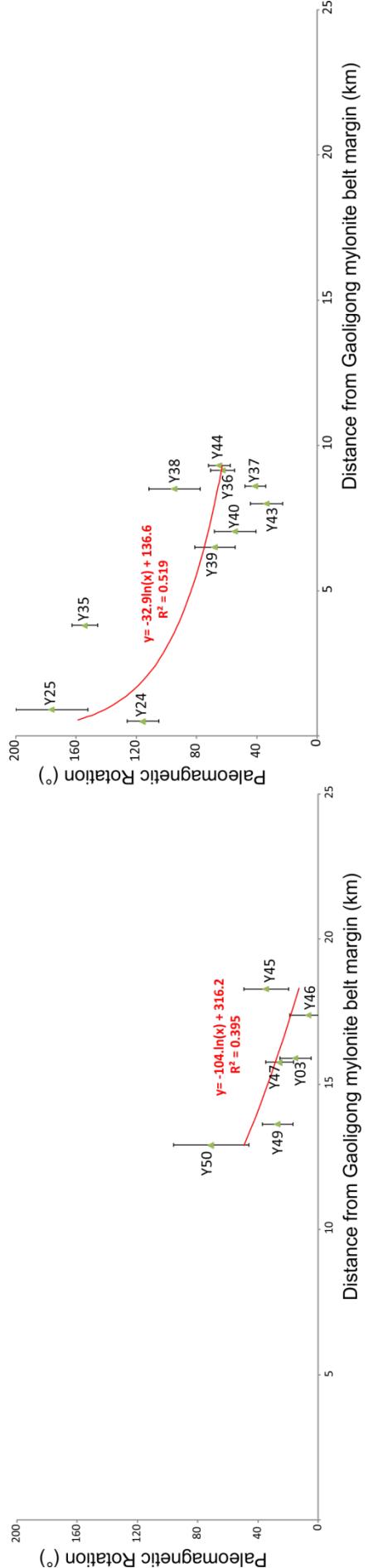
**Figure 61.** Plot of rotations with respect to Eurasia versus site distance to both Gaoligong Shear Zone center and mylonite-sediment contact in Northern (a) and southern (b) transects. See Figure 52-58a-b for site location. Red lines are logarithmic best fit curves for the red bed sites located east of the Gaoligong Shear Zone (disregarding sites Yun24, 29, and 42 located adjacent to fault mylonites).

### Gaoligong Fault rotational pattern

a. Northern transect



b. Southern transect



The described CW paleomagnetic data seem to support a “*quasi-continuous*” crust deformation model of strike-slip fault zones, where the upper brittle crust is broken into small rigid blocks with sizes smaller than the shear zone width, the rotation is CW (CCW) in regions of dextral (sinistral) shear and gradually increases getting closer to the fault, reaching values  $> 90^\circ$  [Figure 62c; *Hernandez-Moreno et al.*, 2014; 2016; *McKenzie and Jackson*, 1983; *Nelson and Jones*, 1987; *Randall et al.*, 2011; *Sonder et al.*, 1994].

According to such model, the rotation is not directly driven by fault shear, but is a consequence of the angular velocity of the ductile deformation, taking place in the ductile lower crust.

*Pellegrino et al.* [2018] note that this CW rotation pattern is not consistent with 1) the continuous crust deformation models (where even the brittle upper crust behaves as a viscous fluid and rotations are  $\leq 90^\circ$  [*Bird and Piper*, 1980; *England and McKenzie*, 1982; *England and Wells*, 1991; *England et al.*, 1985; *Kimura et al.*, 2004, 2011; *Sonder and England*, 1986; *Sonder et al.*, 1986], or 2) the discontinuous models, where set of secondary faults inside the deforming zone bound large rigid domains that rotate uniformly CW and CCW  $\leq 90^\circ$  [*Garfunkel and Ron*, 1985; *McKenzie and Jackson*, 1986; *Nur et al.*, 1986; *Ron et al.*, 1984].

In quasi-continuous deformation models, the rotation magnitude of the blocks has been assumed to depend upon fault length, displacement amount, fault wall locking, and lithosphere rheology [*Piper et al.*, 1997; *Randall et al.*, 2011; *Sonder et al.*, 1994]. The simpler equation by *Lamb* [1987], which relates fault zone deformation-rotation width (W) and displacement (D) with the maximum rotation value of equidimensional blocks ( $\Theta$ ) was recently successfully used to infer the displacement along the 1000 km-long Liquiñe-Ofqui fault from southern Chile [*Hernandez-Moreno et al.*, 2014; 2016]:

$$\Theta = 0.5 D/W$$

Other formulas by *Lamb* [1987; see previous chapter] consider also the occurrence of elongated blocks, and their aspect ratio. However, *Pellegrino et al.* [2018] have no kind of evidence of the shape of the rotating blocks east of the Gaoligong Shear Zone, so that we assume for simplicity the occurrence of equidimensional blocks.

By considering  $165^\circ$  as maximum CW rotation value (average rotation of sites Yun25 and Yun35, Figure 60, 61), and 50 and 40 km as total deformation zone width W (although we have no red bed data west of the fault) with respect to the shear zone center (Figure 60a) and the contact with mylonites (Figure 60b), respectively, the formula by *Lamb* [1987] gives displacement of 290 and 230 km. Such values would be too large if an intra-continental strike-slip fault is considered [*Cao et al.*, 2011a, 2011b; *Otofuji et al.*, 2010; *Tanaka et al.*, 2008; *Tappognier et al.*, 1990; *Zhong et al.*, 1991]. However, by taking in mind that the Gaoligong Shear Zone would have bounded rigid blocks - or “microplates”- escaping laterally from the India-Eurasia collision (thus being a sort of transform fault), the 230-290 km offset values given by paleomagnetism become realistic. *Pellegrino et al.* [2018] remind that an even greater total offset of ca. 700 km was proposed for the Ailaoshan-Red River Shear Zone [*Chung et al.*, 1997; *Leloup et al.*, 1995, 2001], although such estimate was considered unrealistically high by other authors [*Allen et al.*, 1984; *Replumaz et al.*, 2001; *Schoenbohm et al.*, 2006a; *Wang E. et al.*, 1998].

*Pellegrino et al.* [2018] note that sites Yun24, Yun42 and Yun29 yield smaller CW ( $116^\circ$ ,  $45^\circ$ ) and  $74^\circ$  CCW rotations (respectively), even if they were sampled adjacent to the fault (Figures 58b and 60). Therefore, it is possible that such sites rotated in fact CW more than  $180^\circ$ .

*Pellegrino et al.* [2018] conclude that small blocks located adjacent to the fault underwent variable and/or very large CW rotations exceeding largely  $180^\circ$ , but their behavior is not representative of deformation occurring in the 20 km-wide crust slices located at both fault

edges. Probably a sort of “rotation channel” exists adjacent to the fault where small blocks rotate freely by large amounts reminding the “ball-bearing model” of Beck [1976].

Average flattening (F) values for volcanics and red beds are  $12.7^{\circ}\pm9.4^{\circ}$  and  $17.4^{\circ}\pm14.0^{\circ}$ , respectively. This confirms that Pliocene-Holocene volcanics have smaller inclinations than those predicted by a GAD field model, likely due to the incomplete averaging of the paleosecular variation of the geomagnetic field. On the other hand, the positive F for red beds may be due to the so-called “*inclination shallowing*”, a phenomenon frequently observed in sediments and related to compaction-diagenetic effects on rock fabric and magnetic grains [e.g. *Arason and Levi*, 1990; *Deamer and Kodama*, 1990; *Tauxe and Kent*, 2004]. By the way, the positive F value would testify that no remagnetization process occurred after sediment diagenesis, thus it may represent an additional proof for the primary nature of the ChRM and HT components isolated by us in the red beds.

## 6.2 DISCUSSION

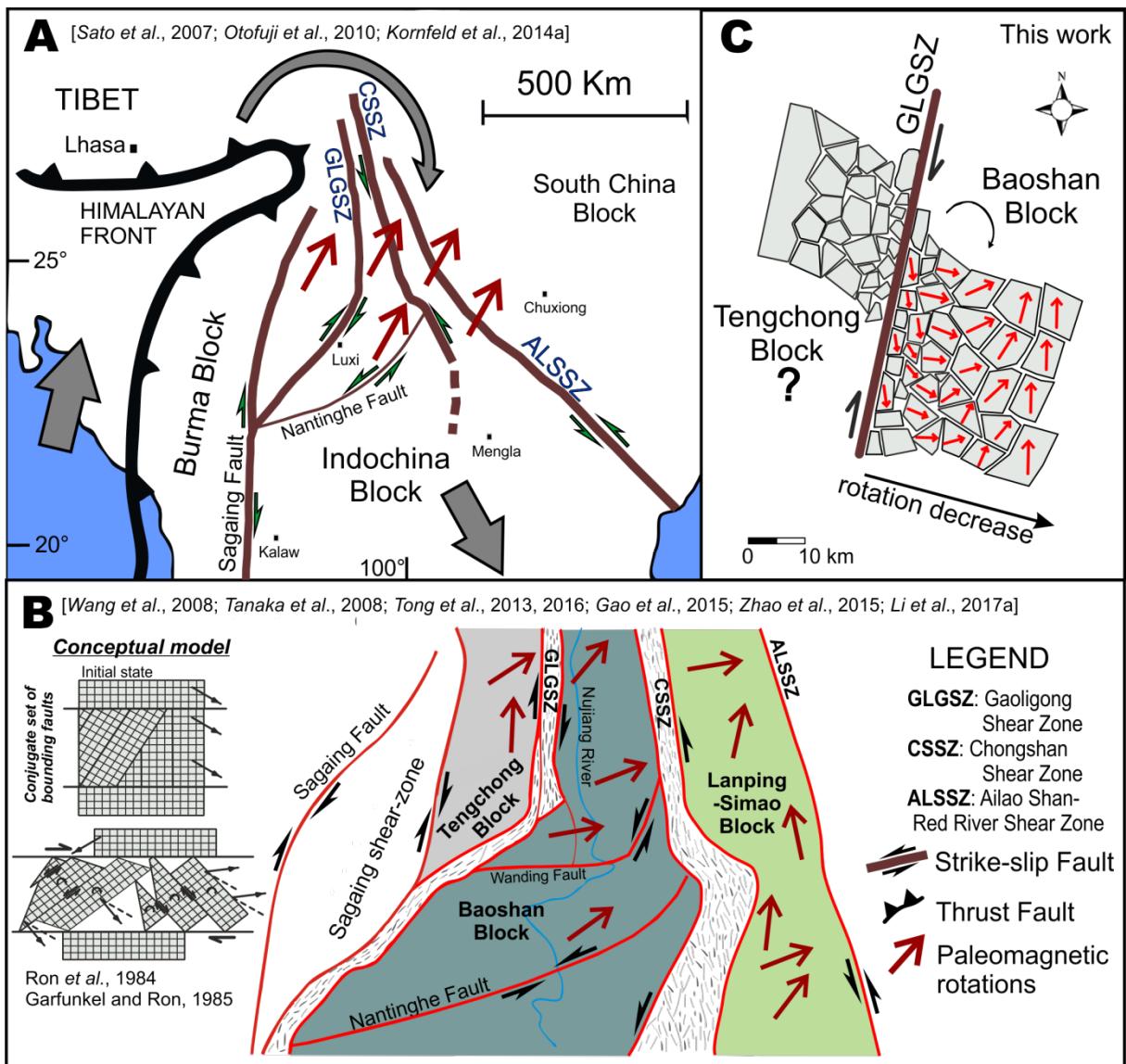
*Pellegrino et al.* [2018] paleomagnetic transects across the Gaoligong Shear Zone show that the dextral shear of the fault caused in the Baoshan block CW rotations of the Jurassic-Cretaceous red beds, whose magnitude decreases progressively moving away from the fault (Figure 60). Sites located adjacent to the fault underwent scattered rotations, and likely rotated CW by more than  $180^{\circ}$ . To the west of the fault, Pliocene-Holocene volcanics do not record significant rotations. Rotation of red beds blocks (whose size is likely  $\leq 1$  km adjacent to the fault) occurred presumably during the period of main Shear Zone activity, which is Oligo-Miocene [*Lin et al.*, 2009; *Wang et al.*, 2006; 2008; *Zhang et al.*, 2012a,b].

The paleomagnetic data show that the dextral shear along the Gaoligong Shear Zone ended 5 Ma ago (at least) being later reactivated as sinistral, consistently with the focal mechanism of

the 1976 M=7.4 event, which yields for the southern part of the Gaoligong zone a sinistral strike-slip fault kinematics (Figure 52a). Moreover, the CW rotation shown today by GPS data in SE Tibet and northern Yunnan [Liang *et al.*, 2013; Figure 48a] is certainly not related to the paleomagnetic rotation that ended by 5 Ma.

The paleomagnetism of the Gaoligong Shear Zone thus suggests that in the Yunnan area the extrusion of blocks bounded by strike-slip shear zones occurred in Oligo-Miocene times. Such tectonics ceased around 10 Ma (considering the younger Ar/Ar ages provided by Zhang *et al.*, 2012b), and was followed by an almost continuous southward drift of Yunnan driven by ductile lower crust squeezed out from Tibet. The present-day southern drift of Yunnan and its collision with Indochina is consistent with both GPS data [Liang *et al.*, 2013; Meade, 2007; Vergnolle *et al.*, 2007], and focal mechanisms of the major earthquakes (Figure 52a), yielding ~N-S sub-horizontal P axes, thus ~N-S horizontal shortening.

In the past, the predominantly CW rotation pattern of Yunnan (Figure 62) has been related to (1) the lateral extrusion of Tibet describing CW rotation trajectories (Figure 62; Allmendinger *et al.*, 2007; Copley, 2008; Gan *et al.*, 2007; Wang *et al.*, 2001; Zhang *et al.*, 2004), and 2) large semi-rigid block rotation driven by strike-slip fault activity following the “broken slate” model (Figure 62a; Li *et al.*, 2017a; Tanaka *et al.*, 2008; Yang and Besse, 1993; Yang *et al.*, 2001b).



**Figure 62.** From Pellegrino et al. [2018]. Kinematic models of block rotations proposed for the Yunnan region. (a) Eastward extrusion of the Tibet plateau, yielding clockwise crust rotation around the East Himalayan Syntaxis, and inducing constant-magnitude clockwise rotations in Yunnan [Kornfeld et al., 2014a; Otofuji et al., 2010; Sato et al., 2007]; (b) rigid mega-block rotations induced by shear along the Gaoligong and adjacent Shear Zones of Yunnan (similar to “broken slate” model by Ron et al., 1984; Garfunkel and Ron, 1985; Nur et al., 1986). Each block rotates rigidly and is bounded by a major Shear Zone [Gao et al., 2015; Li et al., 2017; Tanaka et al., 2008; Tong et al., 2013; Wang et al., 2008; Zhao et al., 2015]; c) quasi-continuous block rotation model [e.g., Lamb, 1987; Randall et al., 2011; Sonder et al., 1994; see also Hernandez-Moreno et al., 2014, 2016] for the Tengchong and Baoshan blocks at both edges of the Gaoligong Shear Zone, based on paleomagnetic rotations reported by Pellegrino et al. [2018]. Rotations of small blocks of Jurassic-Cretaceous red beds east of the Gaoligong Shear Zone (Baoshan block) are caused uniquely by dextral shear along the Gaoligong Shear Zone. Rotations reach maximum values ( $176^\circ$ ) close to the fault and diminish moving east of it being virtually annulled at ca. 20 km from the Shear Zone contact. The dimension and shape of individual crustal blocks are speculative, although— relying on

rotation value scatter—blocks adjacent to the fault are  $\leq$ 1 km wide (see Figure 52-58a-b). Pliocene-Holocene sites located west of the Gaoligong Shear Zone (Tengchong block) do not rotate, likely because dextral fault activity and its related rotations are older than Pliocene times.

*Pellegrino et al. [2018]* data show for the first time that CW rotations in Yunnan are also due to strike-slip fault activity itself, and occurred in small (few km of width) blocks inside the 40-50 km wide damage zone of the Gaoligong Shear Zone (total width at both sides of the fault, although we lack red bed data west of the fault). The rotation pattern and block dimensions are consistent with a quasi-continuous crust deformation model (Figure 62c).

However, it is clear that the rotation kinematics of Yunnan is complex and cannot be solely related to the rotation zones bounding the major strike-slip faults. In fact in *Pellegrino et al. [2018]* working zone, CW rotations of  $42^\circ$ - $100^\circ$  were obtained from 170-30 Ma rocks at distances that are generally greater than 20 km from both the Gaoligong and Chongshan Shear Zones, roughly within the centre of the Baoshan and Tengchong blocks (Table 1, *Huang and Opdyke, 1993; Kornfeld et al., 2014b; Tong et al., 2016*). Thus, such rotations cannot be definitely caused by Gaoligong or Chongshan lateral shear alone. Further E and S of our study area, the recent overview by *Li et al. [2017a]* similarly suggests a rigid CW rotation for both the Lanping and Simao blocks.

Relying on *Pellegrino et al. [2018]* and literature data, the rotational behavior of the Baoshan block itself is far from being clear. *Tong et al. [2016]* reported  $70^\circ$ - $80^\circ$  CW rotations from two localities (NE of Yongde, Figure 52 and Table 1) that are 60-80 km far from the Gaoligong Shear Zone, thus unquestionably not influenced by fault kinematics. Yet, *Pellegrino et al. [2018]* data suggest that at  $\sim$ 25 km from the contact with fault mylonites, the Baoshan block rotation is virtually annulled (Figure 62b). It is also true –however – that a statistically significant rotation of  $20.5^\circ \pm 12.2^\circ$  can be evaluated by considering the four farthest sites from fault contact (Yun03, 45, 46, 47; Figure 60b). A  $20^\circ$  rigid rotation of the

Baoshan block would imply that the maximum rotation value due to fault shear is reduced to  $145^\circ$  (instead of  $165^\circ$ ), and this would translate –by using the formula by *Lamb* [1987] – into a smaller 200-250 km paleomagnetically-evaluated Gaoligong fault offset.

These results of *Pellegrino et al.* [2008] study put forward two hypotheses to explain the complexity of the Yunnan rotational pattern:

- 1) Rotations of the whole Tengchong, Baoshan, and Lanping mega-blocks coexist with rotations of small blocks (few km in size) inside the fault damage zones (Figure 62). Their relative magnitude and kinematic relation –however–have not been clarified yet;

However, it is difficult to imagine that such elongated rock stripes (400 km long and only 20-30 km wide) located between strike-slip faults underwent large rotation, as a big problem of space would arise (Figure 62b). Thus, it is possible that such rotations occurred (at least partly) before that the strike-slip activity of the huge N-S shear zones started in Eocene times.

- 2) Other unrecognized strike-slip faults exist within the Tengchong and Baoshan blocks, and their activity caused the rotations documented in apparently “undeformed” block centers. Rotating folds and thrust sheets might also contribute to the rotation pattern, as suggested for the Simao block by *Kondo et al.* [2012], *Tong et al.* [2013], and *Gao et al.* [2015].

Therefore, during the second year of Ph.D. project, to try unraveling the complexity of the Yunnan rotational pattern, three transect orthogonal to ARRSZ and within Simao, Lanping and Chuandian blocks have been paleomagnetically sampled.

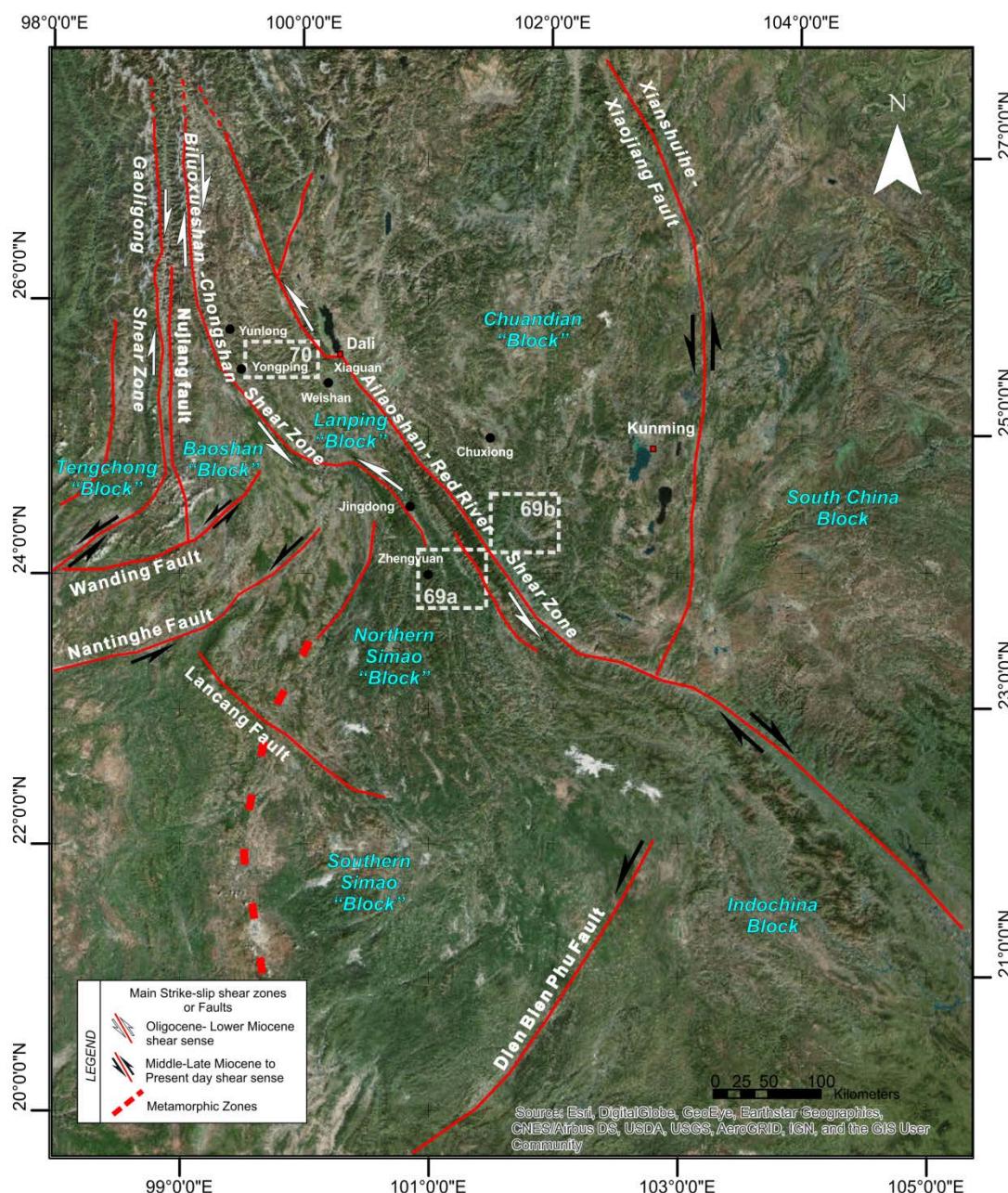


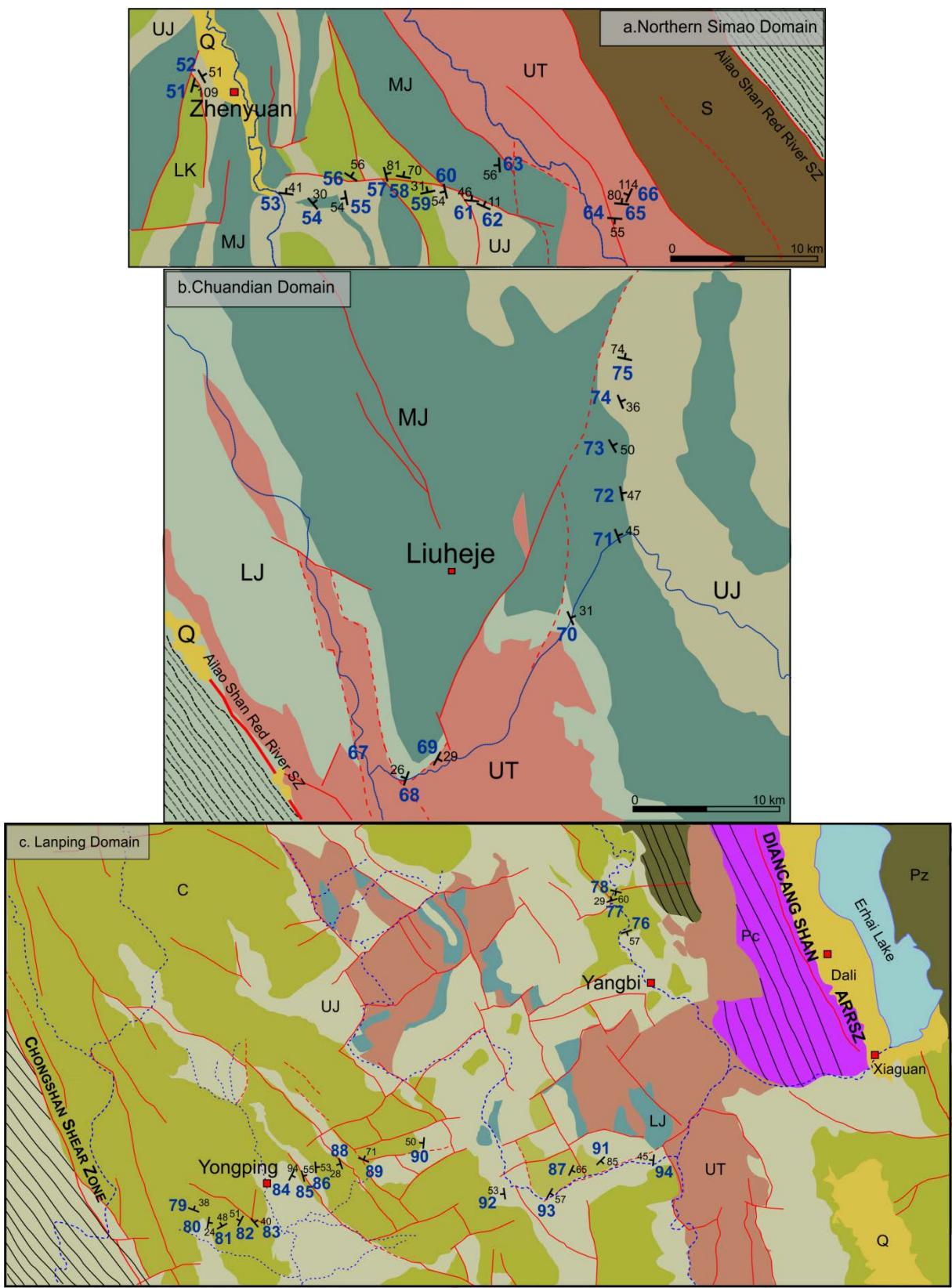
*Chapter*  
**VII**

## 7. PALEOMAGNETIC RESULTS AND MAGNETIC OVERPRINT EVALUATION:

- THE CASE OF AILAOSHAN RED RIVER SHEAR ZONE

All sites sampled along the Ailaoshan Red River Shear Zone into the Lanping, Simao and Chuandian “Block” (or Domain), during the second field campaign on April 2017, yielded a measurable magnetization.





**Figure 63.** Schematic and structural map of Yunnan region modified after *Bureau of Geology and Mineral Resources of Yunnan Province* [1990]. The white boxes indicate the three areas of paleomagnetic investigation with structural details of the sites sampled [a= Figure 69a, b= Figure 69b and c= Figure 70] in Northern Simao, Chuandian and Lanping Domain. Sampled sites in blue.

Sampled sites are 44 (tot. 425 samples, 6 Upper Triassic, 20 Jurassic and 18 Lower Cretaceous red beds sites), located at both sides of the Ailaoshan Red River Shear Zone moving along transect perpendicular to the fault: 16 sites in the southern area within the northern Simao block (south-west of the fault), 9 sites within the Chuandian block (south-east of the fault) and 19 within Simao block (North-west of the fault) between the Ailaoshan Red River Shear Zone and the Chongshan Shear Zone (Table 3, Figure 63, 69a,b and 70).

Differently from the Gaoligong case, we chose sites within a maximum distance of 40 km from the contact with the shear zone, because the estimated offset of the Red River fault is greater than other faults. Furthermore, investigation took place not only in the area adjacent to the fault but also inside the blocks bounding the main structure. Continental red beds, Triassic to Cretaceous in age, were sampled. Sampling methods and laboratory analysis have been described in previous chapters.

The sampled area in the northern Simao Domain, between Zhengyuan city and the Ailao Shan Red River shear zone (Figure 63a), presents several anticlines N-S trending and, to the westernmost others monoclinal stuctural [Tanaka *et al.*, 2008; Bureau of Geology and Mineral Resources of Yunnan Province, 1990]. Sites located into the Chuandian Domain (Figure 63b), have been sampled in the southern part of the so called Chuxiong Basin. In this basin, Upper Triassic gray sandstones and marls grade upwards into Jurassic and Cretaceous red sandstones. The youngest redbeds of this basin are placed in Eocene age [Yoshioka *et al.*, 2003; Gao *et al.*, 2017 and reference therein]. The Upper Cretaceous–Eocene redbeds in the Chuxiong basin is folded into NW to NNW trending anticlines [Leloup *et al.*, 1995]. The folding took place in the Upper Eocene–Oligocene, as evidenced by an unconformity between the Lower Eocene and the Upper Eocene–Oligocene [Bureau of Geology and Mineral Resources of Yunnan Province, 1990; Yoshioka *et al.*, 2003]. Finally, the red beds sampled in the Lanping Domain, between Yongping and Dali city (Figure 63c), are folded, often with steep NNW-striking axial clivage. The folding took place in the Paleogene, resulting in north-

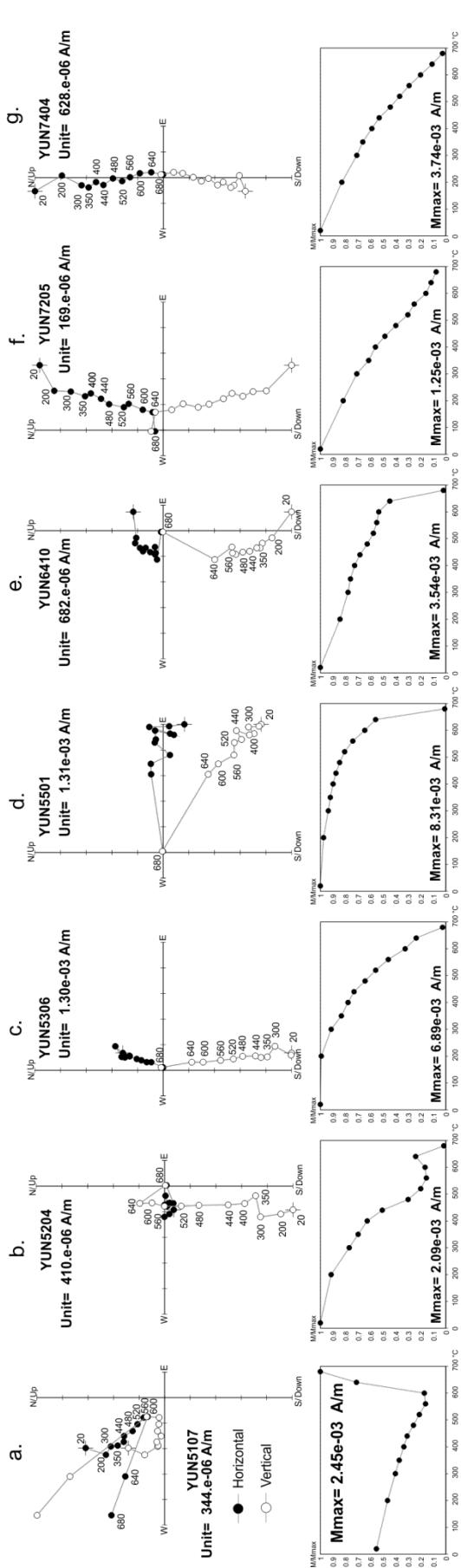
south trending fold axes in the Mesozoic system of Yunnan Province [Tapponnier *et al.*, 1990; Sharer *et al.*, 1994; Funahara *et al.*, 1993].

Red beds samples were thermally demagnetized using a Pyrox shielded oven in 12 steps of temperature up to 680°C. Demagnetization data were plotted on orthogonal vector component diagrams [Zijderveld, 1967]. In most of the sites, after the elimination of a viscous component up to 200°C, a characteristic component (ChRM) was isolated in the 300°-680°C interval or 300°-640°C (called MT component) (Figure 64). Only in 12 samples was isolated also an high-temperature (HT) component defined in the 640°-680°C interval (Figure 64 and Table 3). In agreement with results cropping out from sampled red beds along the Gaoligong fault (see previous chapter), the high unblocking temperature spectra and the Curie temperature of 680°C combined with high coercivity spectra, seem to confirm that hematite is the main magnetic carrier in the red beds (confirming evidence from previous paleomagnetic studies from Yunnan, e.g., Sato *et al.*, 2007; Tong *et al.*, 2013, 2016). Furthermore, several studies carried out over nearly the last 50 years have showed that hematite in red beds may have both a detrital and a chemical origin (see review by Jiang *et al.*, 2015, 2017 and our results of magnetic analysis in next chapter). By combining ChRM-MT and HT components, well-defined site-mean direction ( $1.8 \leq \alpha_{95} \leq 20.5^\circ$ ,  $11.0^\circ$  on average) were obtained from 6 Upper Triassic sites, 1 Lower Jurassic site, 6 Middle Jurassic sites, 13 Upper Jurassic sites and 18 Lower Cretaceous sites (Table 3). See supplementary information to all diagrams.

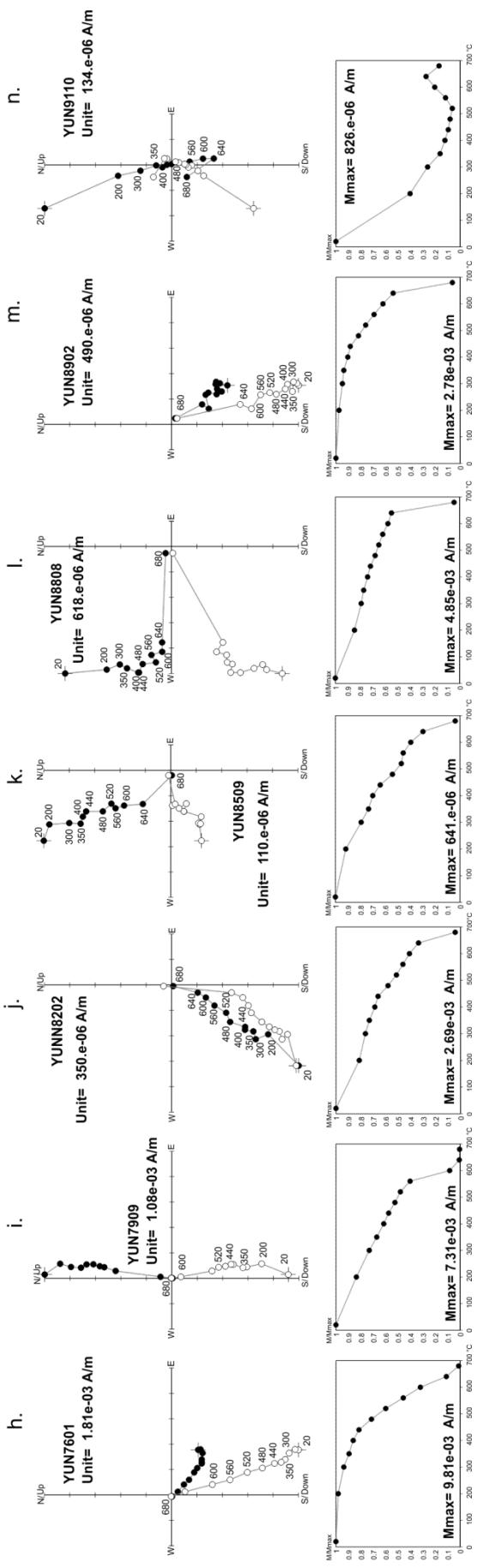
**Figure 64.** Orthogonal vector diagrams of typical demagnetization data (in situ coordinates) showing representative samples carrying characteristic magnetization (ChRM), medium temperature (MT) or high temperature (HT) components (See Table 3 for more details). Solid (open) symbols represent projection onto the horizontal (vertical) plane. Demagnetization step values are in degrees Celsius.

## Northern Simao Domain

## Chuandian Domain



## Laping Domain



**Table 3.** Paleomagnetic site-mean directions from the northern Simao, Chuandian and Lanping domains (Northern Indochina).

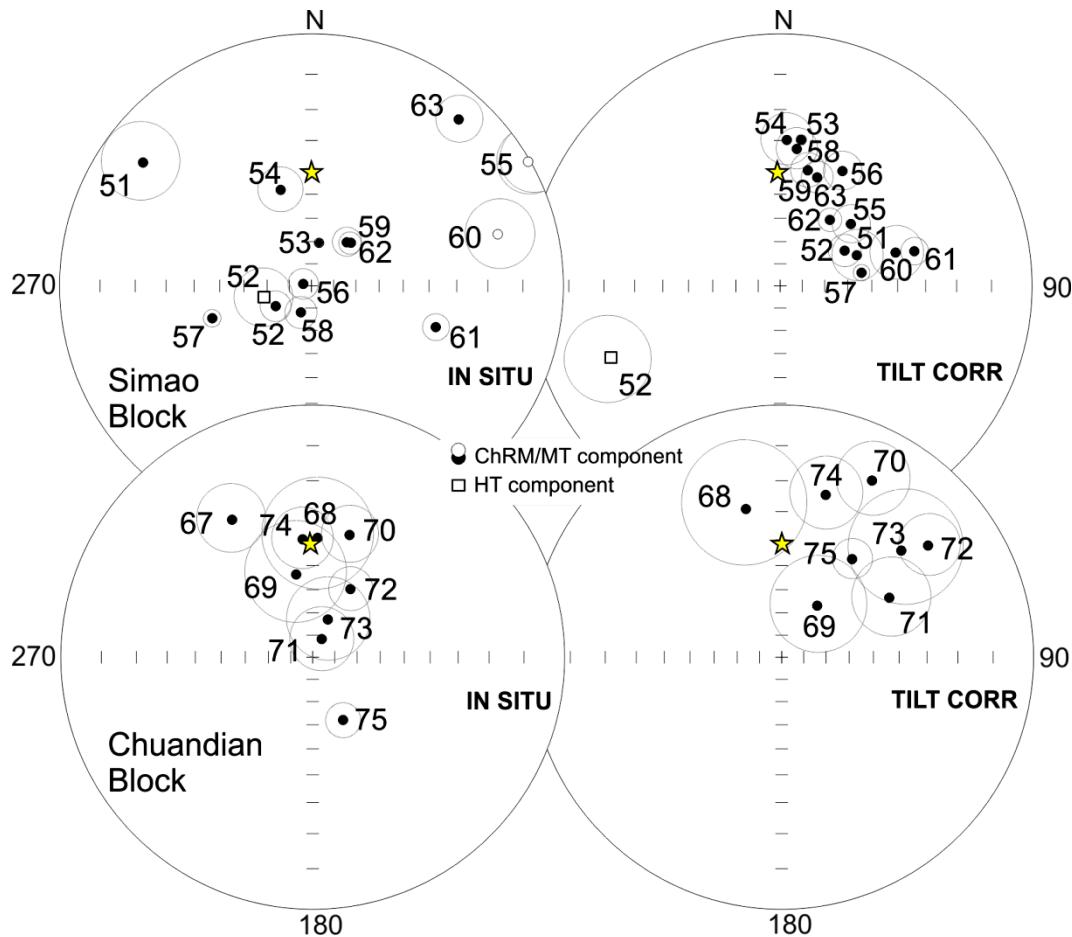
Site	Geographic Coordinates		Age	Age (Ma)	Considered paleopoles age (Ma)	Bedding (deg)	n/N	In Situ		Tilt Corrected		k	$a_{95}$ (deg)	R (deg)	$\Delta R$ (deg)	F (deg)	$\Delta F$ (deg)	
	Latitude °N	Longitude °E						D (deg)	I (deg)	D (deg)	I (deg)							
<b>Northern Simao Block</b>																		
YUN51	24°0'36.8"	101°04'47.8"	LK	100°145	30	ChRM	105/109	7/10	306.2	10.7	68.2	53.7	34.2	10.5	65.9	14.2	-20.2	8.9
YUN52	24°0'39.6"	101°05'12.6"	UJ	145°163	30	MT	61/51	9/10	240.5	71.5	61.3	57.5	58.7	6.8	58.9	10.2	-24.0	6.3
YUN53	23°56'26.8"	101°08'43.3"	UJ	145°163	30	ChRM	7/41	10/10	256.8	-68.1	247.1	-17.8	17.0	12.8	50.7	12.1	35.0	11.3
YUN54	23°55'57.0"	101°09'59.6"	MJ	163°174	30	ChRM	50/30	8/10	342.2	46.4	2.6	29.8	40.9	8.8	0.2	8.2	3.6	7.6
YUN55	23°56'22.3"	101°11'17.3"	UJ	145°163	30	ChRM	260/54	9/10	60.2	-0.5	48.8	49.1	46.3	7.6	46.4	9.4	-15.7	6.8
YUN56	23°56'56.3"	101°13'13.9"	LK	100°145	30	ChRM	33/56	10/10	283.7	86.1	28.5	35.2	49.4	6.9	26.1	7.0	-1.8	6.4
YUN57	23°57'06.9"	101°12'41.4"	LK	100°145	30	ChRM	76/81	9/10	252.0	45.0	80.8	53.8	224.6	3.4	78.4	5.0	-20.4	4.3
YUN58	23°56'55.6"	101°13'31.7"	LK	100°145	30	ChRM	10/70	9/10	201.8	77.1	6.9	32.6	52.1	7.2	4.5	7.1	0.8	6.6
YUN59	23°56'33.2"	101°14'43.46	LK	100°145	30	ChRM	348/31	9/10	38.8	64.7	13.4	39.5	66.7	6.4	11.0	6.9	-6.1	6.0
YUN60	23°56'25.8"	101°14'57.7"	UJ	145°163	30	MT	257/54	10/10	74.5	-15.0	73.9	39.0	24.9	9.9	71.5	10.2	-5.6	8.4
YUN61	23°56'00.4"	101°16'22.8"	UJ	145°163	30	ChRM	5/46	10/10	108.3	35.1	75.5	32.4	102.8	4.8	73.1	5.0	1.0	5.0
YUN62	23°55'53.2"	101°16'38.7"	UJ	145°163	30	ChRM	24/11	9/10	41.9	64.3	37.0	53.7	118.3	4.8	34.6	6.7	-20.3	5.0
YUN63	23°57'15.3"	101°17'35.5"	MJ	163°174	30	ChRM	266/56	9/10	41.5	7.2	18.8	41.1	75.3	6.0	16.4	6.6	-8.2	5.8
YUN64*	23°55'27.5"	101°22'21.3"	UT	201°237	30	HT	184/55	9/10	357.0	71.2	187.8	53.6	17.3	12.7	-5.4	33.6	-37.8	10.5
YUN65*	23°56'09.7"	101°22'33.8"	UT	201°237	30	HT	5/80	10/10	2.1	49.8	2.9	-30.1	27.7	9.3	-0.3	11.5	-16.4	8.0
YUN66*	23°56'15.9"	101°22'44.1"	UT	201°237	30	ChRM	294/114	6/10	18.7	51.3	336.0	-21.7	30.6	12.3	16.3	15.7	-17.9	10.2
<b>Chuandian Block</b>																		
YUN67*	24°1'508.5"	101°23'44.3"	UT	201°237	20	ChRM	—	10/11	329.6	25.4	—	—	19.8	11.1	-32.8	9.9	9.8	
YUN68*	24°1'333.6"	101°30'47.7"	UT	201°237	20	ChRM	290/26	6/10	2.3	39.3	346.4	27.7	11.7	20.4	-0.1	21.0	-4.1	16.3
YUN69*	24°1'426.2"	101°31'53.9"	UT	201°237	20	ChRM	118/29	6/10	348.7	53.0	34.6	62.1	11.6	20.5	-13.7	27.9	-17.8	16.4
YUN70*	24°1'910.4"	101°32'20.9"	LJ	174/201	20	ChRM	66/31	8/10	16.8	36.3	27.1	13.6	29.8	10.3	14.4	10.3	-1.0	8.8
YUN71*	24°2'159.7"	101°38'47.1"	MJ	163°174	20	ChRM	69/45	7/10	26.3	80.8	61.1	38.0	18.6	14.4	23.9	14.2	-45.4	14.4
YUN72*	24°2'408.8"	101°39'42.9"	MJ	163°174	20	MT	80/47	9/10	29.0	55.7	52.7	17.7	33.1	9.1	26.6	13.0	-26.2	8.0
YUN73*	24°2'612.2"	101°39'24.7"	MJ	163°174	20	MT	61/50	7/10	21.9	71.7	48.3	25.1	11.8	18.3	19.5	70.2	-36.2	14.7
YUN74*	24°2'719.3"	101°39'34.2"	UJ	145°163	20	ChRM	68/36	7/10	355.1	39.7	15.2	22.6	28.7	11.4	-7.3	11.9	-4.2	9.6
YUN75*	24°2'915.0"	101°39'49.1"	UJ	145°163	20	MT	127/4	7/10	154.2	59.1	35.7	38.8	67.0	7.4	151.8	11.6	-23.5	6.8

**Table 2.** Paleomagnetic directions from red beds from the northern Simao, Chuandian, and Lanping blocks (Northern Indochina)

Site	Geographic Coordinates		Age (Ma)	Considered paleopole age (Ma)	Dipangle/azimuth	Bedding (deg)	n/N	In Situ			Tilt Corrected			$k$	$a_{is}$ (deg)	R (deg)	$\Delta R$ (deg)	F (deg)	$\Delta F$ (deg)
	Latitude °N	Longitude °E						D	I	D	D	I	I						
<b>Laminc Block</b>																			
YUN76	25°43'20.7"	099°55'31.2"	LK	100-145	120	ChRM	162/57	10/10	123.0	71.4	149.8	18.0	212.3	3.3	139.4	3.7	24.8	3.9	
YUN77	25°45'05.9"	099°54'44.2"	LK	100-145	120	ChRM	342/29	9/10	317.7	6.3	316.2	-20.1	13.6	14.5	125.8	12.3	22.8	11.7	
YUN78	25°45'43.0"	099°54'41.6"	LK	100-145	120	ChRM	193/60	8/10	322.6	58.9	227.7	45.5	19.8	12.8	-142.6	14.6	-3.1	10.4	
YUN79**	25°26'14.1"	099°26'31.1"	LK	100-145	20	ChRM	21/38	8/10	3.7	40.1	7.8	3.3	811.9	1.9	5.5	3.0	33.6	3.8	
YUN80	25°25'44.7"	099°27'06.5"	LK	100-145	120	ChRM	101/24	7/10	290.1	34.6	295.3	58.1	19.4	14.0	-75.0	21.4	-15.7	11.3	
YUN81	25°25'28.5"	099°28'24.0"	LK	100-145	120	ChRM	331/48	8/10	64.5	76.8	348.6	41.4	69.6	6.7	-21.7	7.4	0.9	6.0	
YUN82	25°25'35.9"	099°30'01.3"	LK	100-145	120	MT	297/51	8/10	205.1	30.4	230.5	19.9	15.8	14.4	-139.8	12.2	22.4	11.6	
YUN83	25°25'23.1"	099°30'57.6"	LK	100-145	120	HT	50/40	9/10	333.0	-33.5	305.4	-32.9	25.3	10.4	115.1	10.0	9.4	8.6	
YUN84**	25°27'56.0"	099°31'17.6"	UJ	145-163	20	MT	117/94	9/10	319.3	12.7	(325.1)	(38.1)	14.9	13.8	-37.2	14.0	-1.2	11.3	
YUN85**	25°27'49.3"	099°33'41.6"	UJ	145-163	20	MT	75/55	9/10	355.9	51.3	(13.2)	(45.9)	10.1	17.0	10.9	19.5	-9.0	13.7	
YUN86	25°28'16.2"	099°34'42.2"	UJ	145-163	150	HT	86/53	9/10	324.3	11.2	345.5	31.9	55.1	7.0	-31.4	8.8	21.9	7.5	
YUN87**	25°28'11.2"	099°52'09.6"	LK	100-145	20	MT	115/65	9/10	343.9	27.7	(384.1)	(39.0)	18.4	12.3	-8.2	12.7	-2.1	10.2	
YUN88**	25°29'21.7"	099°36'26.9"	LK	100-145	20	MT	252/28	8/10	0.6	29.4	(356.0)	(31.6)	10.4	18.0	-6.3	16.8	5.3	14.5	
YUN89	25°29'24.4"	099°38'09.6"	LK	100-145	120	HT	7/10	283.1	16.2	282.1	-8.0	12.0	18.1	-22.3	18.7	20.5	15.5		
YUN90	25°30'52.2"	099°41'15.1"	UJ	145-163	150	HT	22/71	10/10	143.6	64.6	47.1	30.5	80.3	5.4	36.8	5.5	11.9	5.2	
YUN91**	25°28'36.2"	099°54'08.4"	LK	100-145	20	MT	135/85	9/11	8.1	19.9	(18.9)	(33.2)	8.0	19.4	16.6	18.4	3.7	15.5	
YUN92**	25°27'12.7"	099°47'08.9"	UJ	145-163	20	MT	262/53	9/10	9.2	44.5	(352.8)	(47.1)	40.4	8.2	-9.5	9.8	-10.2	7.3	
YUN93**	25°27'09.4"	099°50'24.9"	LK	100-163	20	MT	120/57	8/10	341.9	16.4	(346.9)	(28.6)	35.8	9.4	-15.4	8.8	8.3	8.1	
YUN94**	25°29'03.7"	099°57'02.7"	MJ	163-174	20	ChRM	276/45	8/10	354.6	41.1	(344.4)	(37.3)	37.0	9.2	-17.9	9.4	-0.3	8.0	

Ages: U1, Upper Triassic; L1, Lower Jurassic; MJ, Middle Jurassic; UJ, Upper Jurassic; LK, Lower Cretaceous. Demagnetization component: ChRM, characteristic remanent magnetization (300-680°C); MT, medium temperature (300-640°C) component; HT, high temperature component (640-680°C). The geographic coordinates are referred to WGS84 datum. Ages in Ma is from the geologic timescale of Chen et al. [2013]. Bedding is expressed in dip azimuth/dip values. D and I are site mean declination and inclination calculated before and after tectonic correction;  $\Delta$  and  $\alpha$  are statistical parameters after Fisher [1953]; n/N is number of samples giving reliable results; number of samples at site. Site mean rotation R and flattening F values, and relative errors AR and AF according to Demarest [1983] are relative to coeval D and F values. \* syn-tectonic remagnetized sites, R and F values are calculated using the in-situ D and I values. \*\* post-tectonic remagnetized sites, the R and F values are obtained using 150 Ma. \* post-tectonic remagnetized sites, the R and F values are calculated using the 150 Ma. \* post-tectonic remagnetized sites, the R and F values are obtained using the 150% unfolding (in parenthesis).

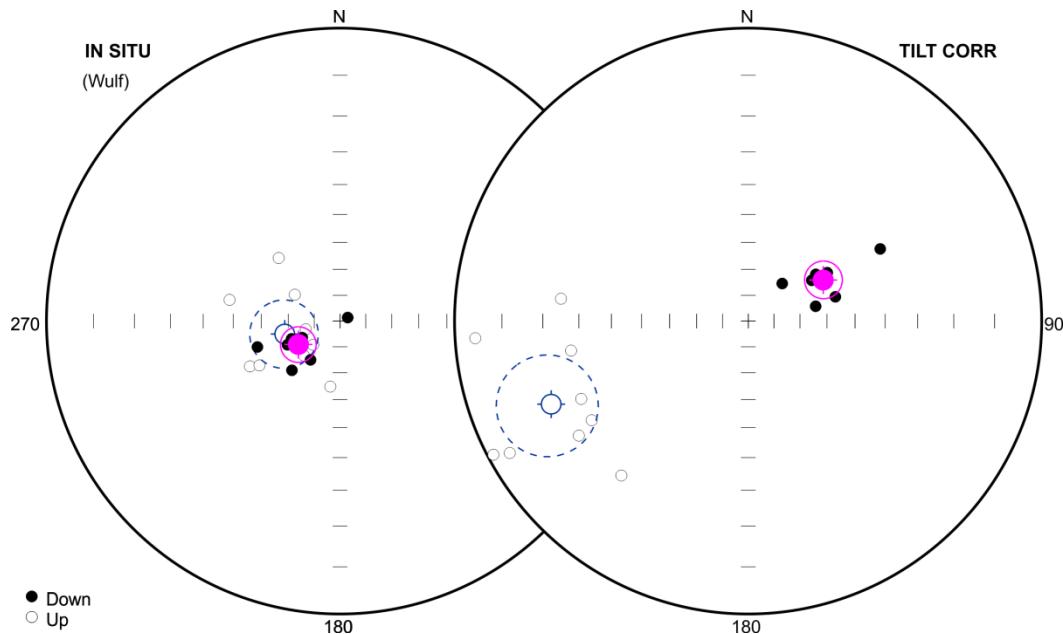
All paleomagnetic direction are shown in Figure 65 for Northern Simao and Chuandian domains and in Figure 68 for Lanping domain. Site-mean directions were evaluated by averaging out ChRM, MT and HT components from the individual specimens. Data are shown in Table 3, and a fold test was applied separately to the different zones and group sites (see Table 4).



**Figure 65.** Equal-angle projection of the paleomagnetic directions obtained from Triassic-Cretaceous red beds exposed in the Simao and Chuandian domains. The three Triassic sites (Yun64, 65, 66) from the Simao domain are shown in Figure 67. The yellow stars represent the normal polarity geocentric axial dipole field direction ( $D=0^\circ$ ;  $I=42^\circ$ ) expected at the study area.

As shown in Figure 65, both the *in situ* and tilt-corrected directions, for Simao and Chuandian domains, are far from the GAD field direction expected at the sampling localities (represented by the yellow star). All tilt-corrected directions from Simao domain (sites from Yun51 to

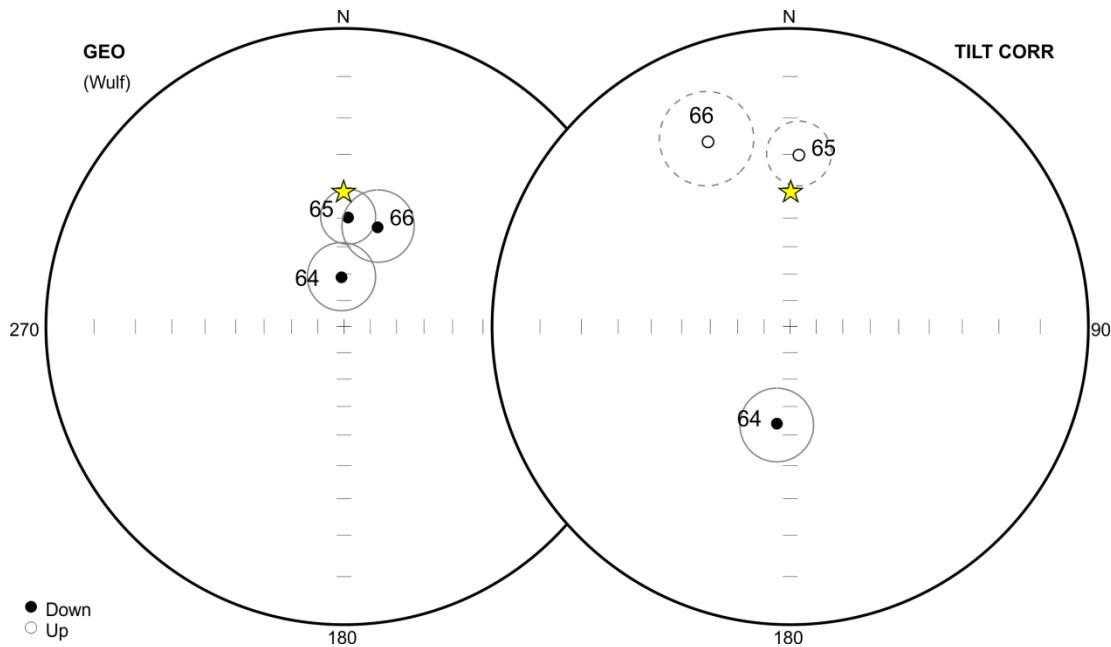
Yun63; Yun64, Yun65, Yun66 are Triassic red beds) are of normal polarity, excepted for site Yun52, where a reverse-polarity HT of possible primary origin was observed. Two different components (of both normal and reverse polarity) was isolated in the 350-600°C (MT) and 640°C (HT) interval, respectively (Figure 66). The mean values are shown in Table 3.



**Figure 66.** Different components and polarity of the Yun52. Blue (pink) open (solid) circles represent the mean paleomagnetic directions and relative 95% confidence cone of HT and MT components, respectively (see Table 3 for mean values). The yellow stars represent the normal polarity geocentric axial dipole field direction ( $D=0^\circ$ ;  $I=44^\circ$ ) expected at the study area.

The paleomagnetic data from Jurassic-Cretaceous sites from Northern Simao domain (from Yun51 to Yun63) support a positive fold test at 99% significance level according to McFadden [1990;  $N=13$ ;  $\text{Scos}_{\text{insitu}}= 8.5$ ;  $\text{Scos}_{\text{tilt corrected}}= 1.8$ ; critical Scos at the 95% significance level = 4.2; critical Scos at the 99% significance level = 5.9; see table 4]. Maximum  $\kappa$  and minimum Scos values arise at 92% of complete unfolding. This suggest a pre-folding [ $\sim 30$  Ma according to Yang *et al.*, 2001b; Sato *et al.*, 2001; Gao *et al.*, 2017 and reference therein].

In contrast, Upper Triassic red beds sites Yun64, Yun65 and Yun66 support a negative fold test at 99% significance level [N=3;  $\text{Scos}_{\text{insitu}} = 1.4$ ;  $\text{Scos}_{\text{tilt corrected}} = 2.7$ ; critical Scos at the 95% significance level = 2.1; critical Scos at the 99% significance level = 2.7; see Table 4]. Maximum  $\kappa$  and minimum Scos values arise at 0% of complete unfolding. This suggest a post-folding magnetization (post- 30 Ma).



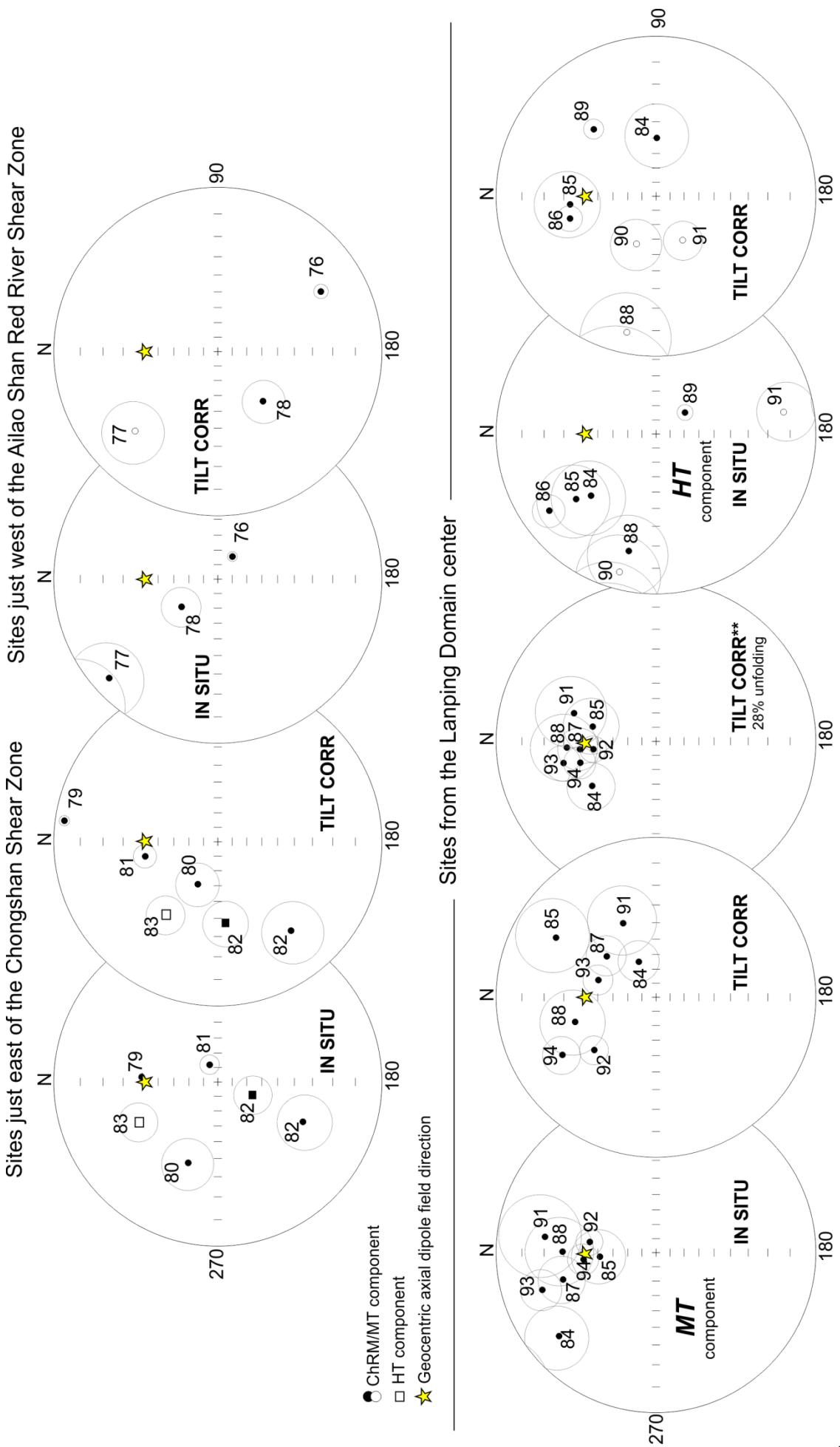
**Figure 67.** Equal-angle projection of the paleomagnetic direction obtained from three Triassic red beds sites exposed in the Simao block. The yellow stars represent the normal polarity geocentric axial dipole field direction ( $D=0^\circ$ ;  $I=42^\circ$ ) expected at the study area.

All components of Chuandian block sites are also of normal polarity and far from the GAD field direction expected at the sampling localities (Figure 65). The paleomagnetic data from this group of samples (sites from Yun67 to Yun75) support a negative fold test at 99% significance level [N=9;  $\text{Scos}_{\text{insitu}} = 0.6$ ;  $\text{Scos}_{\text{tilt corrected}} = 3.9$ ; critical Scos at the 95% significance level = 3.5; critical Scos at the 99% significance level = 4.8; see Table 4]. Maximum  $\kappa$  values arise at 32% of complete unfolding and minimum Scos values arise at 0% of complete unfolding. Also in this case, these suggest a post-folding magnetization (post-Late Eocene-Early Oligocene).

Different behavior is observed at sites close to the faults ARRSZ and Chongshan fault (CF), and at the centre of the Lanping block. In the Lanping block, our three Lower Cretaceous sites (Yun76, Yun77, Yun78), located just west of the Ailaoshan Red River fault (near Yangby city, Figure 68) support a positive fold test at 99% significance level suggesting a pre-folding magnetization [ $N=3$ ;  $Scos_{insitu}= 2.5$ ;  $Scos_{tilt\ corrected}= 0.1$ ; critical  $Scos$  at the 95% significance level = 2.1; critical  $Scos$  at the 99% significance level = 2.7; see Table 4]. Maximum  $\kappa$  values arise at 100% of complete unfolding and minimum  $Scos$  values arise at 96% of complete unfolding.

Sites located near the Chongshan fault (from site Yun80 to Yun83) support an indeterminate fold test at 99% significance level [ $N=4$ ;  $Scos_{insitu}= 0.4$ ;  $Scos_{tilt\ corrected}= 0.8$ ; critical  $Scos$  at the 95% significance level = 2.3; critical  $Scos$  at the 99% significance level = 3.2; see Table 4]. Maximum  $\kappa$  values arise at 8% of complete unfolding and minimum  $Scos$  values arise at 29% of complete unfolding. No evidence of remagnetization is apparent here.

**Figure 68.** Equal-angle projection of the paleomagnetic directions obtained from Jurassic-Cretaceous red beds exposed in the Lanping domain. The yellow stars represent the normal polarity geocentric axial dipole field direction ( $D=0^\circ$ ;  $I=44^\circ$ ) expected at the study area.



Sites with MT component (Yun84, 85, 86, 88, 89, 90, 91) support a positive fold test at 99% significance level suggesting a syn-folding magnetization [N=8; Scos<sub>insitu</sub>= 5.3; Scos<sub>tilt corrected</sub>= 6.9; critical Scos at the 95% significance level = 3.3; critical Scos at the 99% significance level = 4.6; see Table 4]. Maximum  $\kappa$  values arise at 32% of complete unfolding and minimum Scos values arise at 28% of complete unfolding.

At the end, sites with HT component support a positive fold test fold test at 99% significance level suggesting a pre-folding magnetization [N=7; Scos<sub>insitu</sub>= 6.1; Scos<sub>tilt corrected</sub>= 0.3; critical Scos at the 95% significance level = 3.1; critical Scos at the 99% significance level = 4.2; see Table 4]. Maximum  $\kappa$  values arise at 100% of complete unfolding and minimum Scos values arise at 98% of complete unfolding.

**Table 3.** Results of the fold test [according to McFadden, 1990] for sites from different domains.

Sites (Yun)	N	95% cr. va	99% cr. va	In Situ Statistics				Unfolded Statistics				Minimum SCOS Statistics				Test result				
				D (deg)	I (deg)	k (deg)	$\alpha_{95}$ (deg)	D (deg)	I (deg)	k (deg)	$\alpha_{95}$ (deg)	D (deg)	I (deg)	k (deg)	$\alpha_{95}$ (deg)	SCOS	% unfolding			
<b>Northern Simao Domain</b>																				
51,52,53,54,55,56, 57,58,59,60,61,62,63	13	4,200	5,860	31.1	66.3	2.6	32.2	8,494	37.4	45.7	11.9	12.5	1,756	37.3	47.6	12.3	0.087	92%	P	
64,65,66	3	2,076	2,662	7.6	57.7	36.9	20.6	1,355	339.8	-3.2	1.1	180.0	2,701	7.6	57.7	36.9	20.6	1,355	0%	N
<b>Chuandian Domain</b>																				
67,68,69,70,71, 72,73,74,75	9	3,497	4,849	4.2	58.2	7.2	20.5	0,606	24.4	31.0	6.0	22.8	3,914	4.2	58.2	7.2	20.5	0,606	0%	N
<b>Laping Domain</b>																				
76,77,78	3	2,076	2,662	128.7	64.8	1.8	180.0	2,479	162.8	33.5	3.8	74.7	0,112	162.4	34.7	3.8	75	0.009	96%	P
80,81,82,83	4	2,335	3,180	205.9	64.0	2.7	69.9	0,435	259.5	74.2	2.0	92.3	0,787	217.2	69.0	2.7	70.7	0.002	29%	I
84,85,87,88,91,92,93,94 (ChrM/MT-HT)	8	3,298	4,562	350.7	31.3	16.3	14.1	5,321	12.8	45.2	5.8	25.2	6,865	353.6	38.3	30.1	10.3	0.502	28%	S
84,85,86,88,89,90,91 (HT)	7	3,086	4,253	6.4	42.6	1.5	79.6	6,095	54.6	48.5	4.0	34.5	0,268	54.0	48.9	4.0	34.5	0.031	98%	P

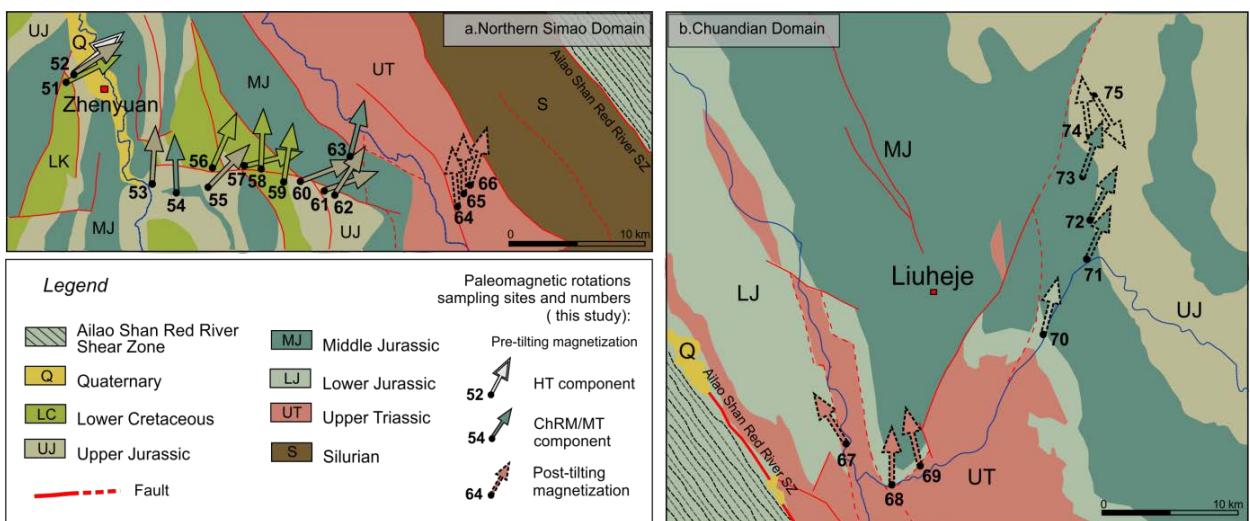
N, number of sites; D and I are mean declination and inclination, respectively; k and  $\alpha_{95}$  are statistical parameters after Fisher [1953]; SCOS, 95% cr.va and 99% cr.va are statistical parameters and 95-99% critical values, respectively [see McFadden, 1990]. Test results: P positive, S syn-folding, N negative, I indeterminate.

**Table 4.** Results of the fold test [according to McFadden, 1990] for sites from different domains.

## 7.1 ROTATION PATTERN

To evaluate tectonic rotations with respect to Eurasia, the paleomagnetic directions were compared to coeval Eurasian paleopoles from *Torsvik et al.* [2012] and *Cogné et al.* [2013]. Paleomagnetic rotations (R) and flattening (F) with relative errors, according to *Demarest* [1983], are reported in Table 3.

As discussed above, the paleomagnetic results revealed different structural behavior in the analyzed domains. **In the Northern Simao block** (west to the fault, Figure 69) clockwise rotation between  $\sim 30^\circ$ - $70^\circ$  CW have been recorded. Some sites show a null rotation. At site Yun52, located at  $\sim 30$  km from the western boundary of the Ailaoshan Red River Shear Zone, rotations from MT and HT components are similar, confirming the reliability of the normal polarity data set.

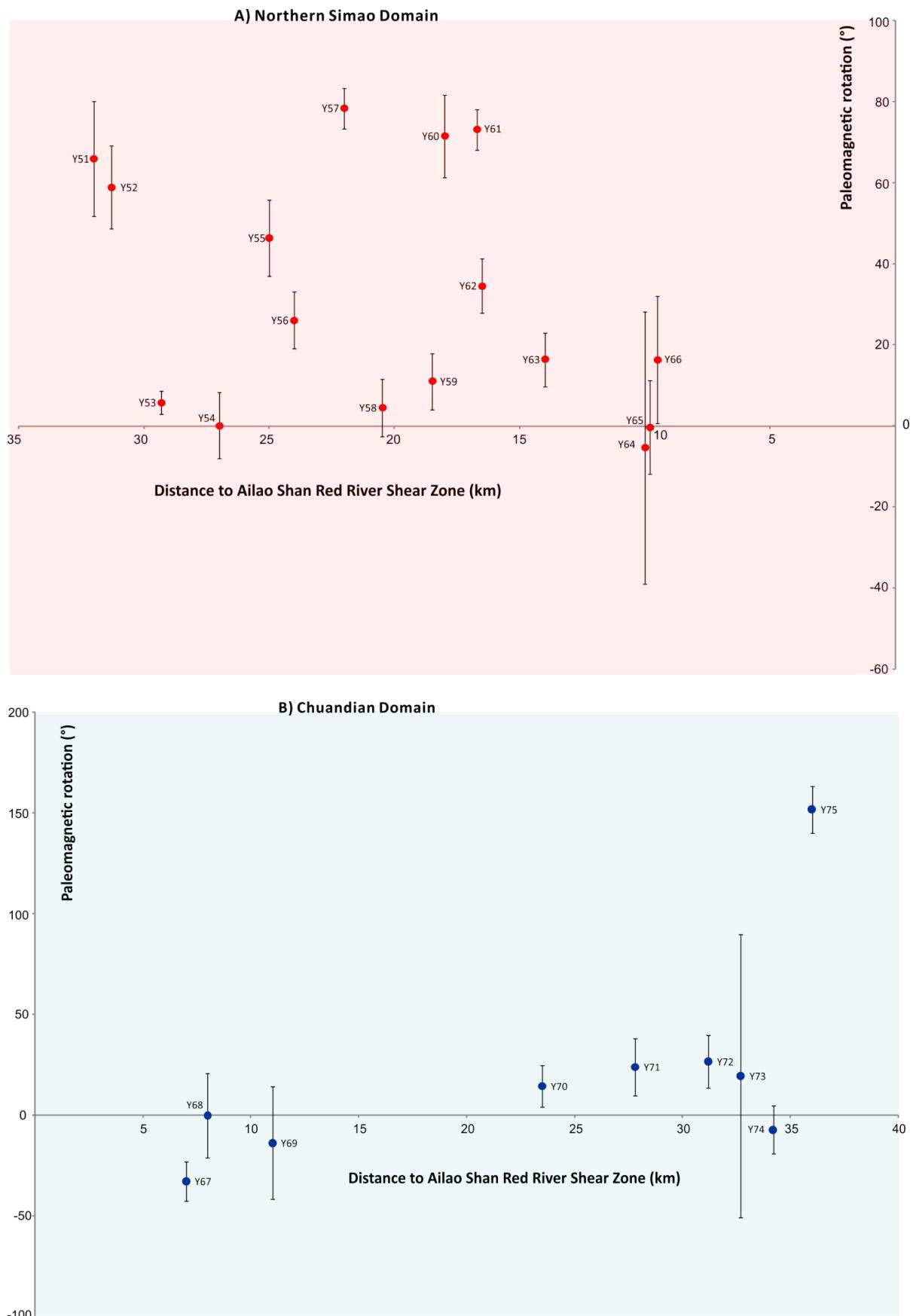


**Figure 69.** Paleomagnetic rotations with respect to Asia from the (a) Northern Simao and (b) Chuandian domains, along the Ailaoshan Red River Shear Zone (geological map modified after *Bureau of Geology and Mineral Resources of Yunnan Province*, 1990). Rotation values are reported in Table 3. Solid contour arrows show the rotation values with respect to tilt-corrected characteristic/medium temperature (ChRM/MT) paleomagnetic directions and pre-tilting normal-polarity magnetic overprint (30 Ma Asia paleopole, see text). The rotation of reverse-polarity high temperature (HT) component from site Yun52 was also calculated with respect to 150 Ma paleopole, coeval to sediment age. Dashed arrows show rotation values relative to in-situ paleomagnetic directions (post-tilting remagnetization).

From W to E (Figure 69a), we report three blocks undergoing  $57^\circ \pm 9^\circ$ ,  $48^\circ \pm 26^\circ$  and  $70^\circ \pm 1^\circ$  CW rotations, separated by two unrotated blocks (sites Yun53-54 and Yun58-59). Finally the two easternmost sites Yun62-63 define a  $23^\circ \pm 13^\circ$  CW rotation.

Directions of the three Triassic sites located at about 10 km from the western boundary of the ARRSZ (Yun64, 65, 66) and sites from the **Chuandian domain** (Yun67 to Yun75) post-date tilting, evaluated using the 20 Ma Asia pole and in-situ paleomagnetic directions. Rotation values are small (Figure 69a,b), except for easternmost site Yun75 located at ~35 km from the fault, which is possibly an outlier (CW rotation more than  $150^\circ$  and bedding was completely different than in nearby sites, Table 2). The post-tilting rotations of the three Triassic sites in the Northern Simao domain ( $4^\circ \pm 11^\circ$  on average), the three sites located just east the ARRSZ ( $-15^\circ \pm 16^\circ$ ) and the five sites located farther east in the Chuandian domain ( $15^\circ \pm 14^\circ$ ) seem to be remagnetized after folding, according to the fold test result. These are not or barely significant, implying an overall null rotation of the ARRSZ and the Chuandian domain occurring after magnetic overprint (in turn post-dating 25-15 Ma folding). Thus our paleomagnetic data from the ARRSZ and the Chuandian domain are not relevant to constrain their rotation history, as we cannot exclude a recent overprint (i.e. within the last 0.78 Ma of the Brunhes polarity chron).

Differently from the Gaoligong case, the distribution of the rotation values with respect to the distance to both sides of the Ailaoshan Red River Shear Zone (km), do not show a significant logarithmic trend. West of the fault, in the Northern Simao block, sites show different crustal portion (size < 5 km) undergoing independent variable rotations (Figure 70a). East of the fault, in the Chuandian domain, the low and post-tilting rotations suggest a recent overprint, probably after 25-15 Ma (Figure 70b).



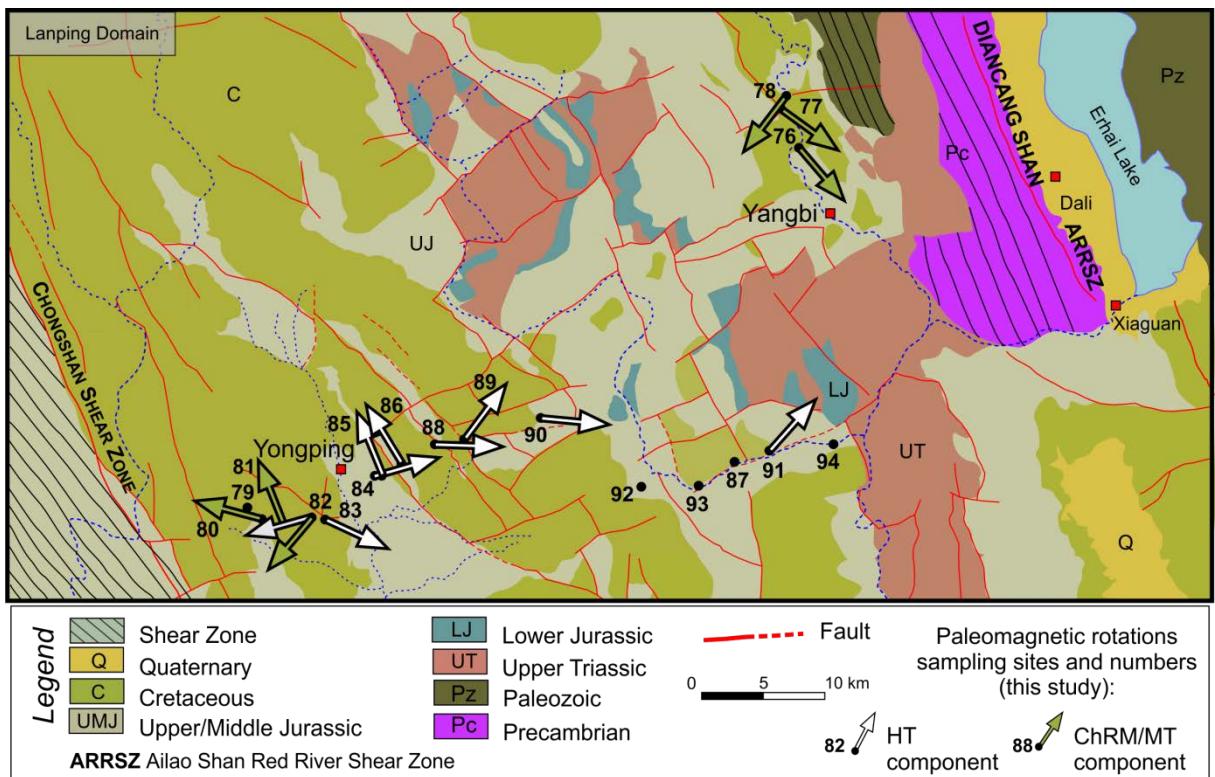
**Figure 70.** Plot of rotation with respect to Eurasia versus site distance to both Ailaoshan Red River Shear Zone in Northern Simao (a) and Chuandian (b) domains. See figures 69 an 70 for site location. Sense and amount of rotation are defined by the smaller angel between observed and expected Eurasian declinations, thus rotation values are  $\leq 180^\circ$  | by definition.

In the northern transect adjacent to Lanping block (Figure 71), we identified three different rotation patterns. Sites sampled close to the Chongshan Shear Zone (Yun80 to Yun83- site Yun79 is remagnetized post-folding) yield a  $22^\circ$ - $140^\circ$  CCW rotation (the MT and HT components from site Yun82 give a  $106^\circ \pm 13^\circ$  and  $140^\circ \pm 12^\circ$  CCW rotation, respectively). Sites from the block centre (Yun84 to Yun94) showed two different magnetization components: HT component ( $640$ - $680^\circ\text{C}$ ) yield variable  $37^\circ$ - $115^\circ$  CW rotation interrupted by two sites just E of Yongping (Yun85 to Yun86) where a slightly  $27^\circ \pm 6^\circ$  CCW rotation is recorded. MT components ( $300$ - $640^\circ\text{C}$ ) from the same sites are yield remagnetization later acquired at 28% unfolding, and a subsequent lack of rotation (Figure 68).

Finally, three sites located adjacent to ARRSZ, yields a strong CW (CCW) rotation between  $126^\circ$  and  $143^\circ$  in Yun76 and Yun77 (Yun78), respectively. The sense is unclear because they approach  $180^\circ$ , so cannot exclude that this site rotated CW more than  $180^\circ$ .

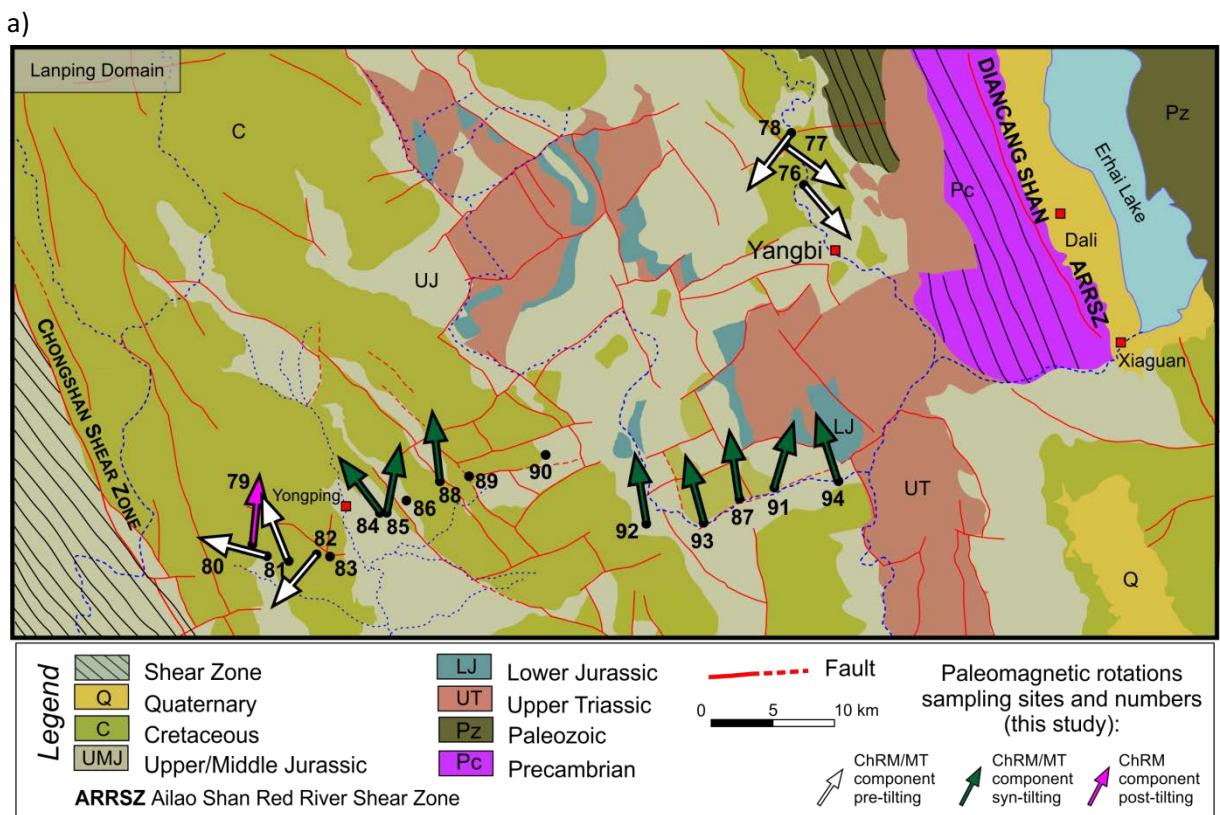
Unfortunately, sites located adjacent to the shear zone are not enough to evaluate the pattern of rotation approaching major structures.

In Figure 72a,b report separately the direction of ChRM/MT and HT component pre-, syn- and post-tilting. The ChRM/MT give a congruous rotation values of  $8^\circ \pm 16^\circ$  on average, scenario that could be confirm related magnetic overprint due to fluid migration.

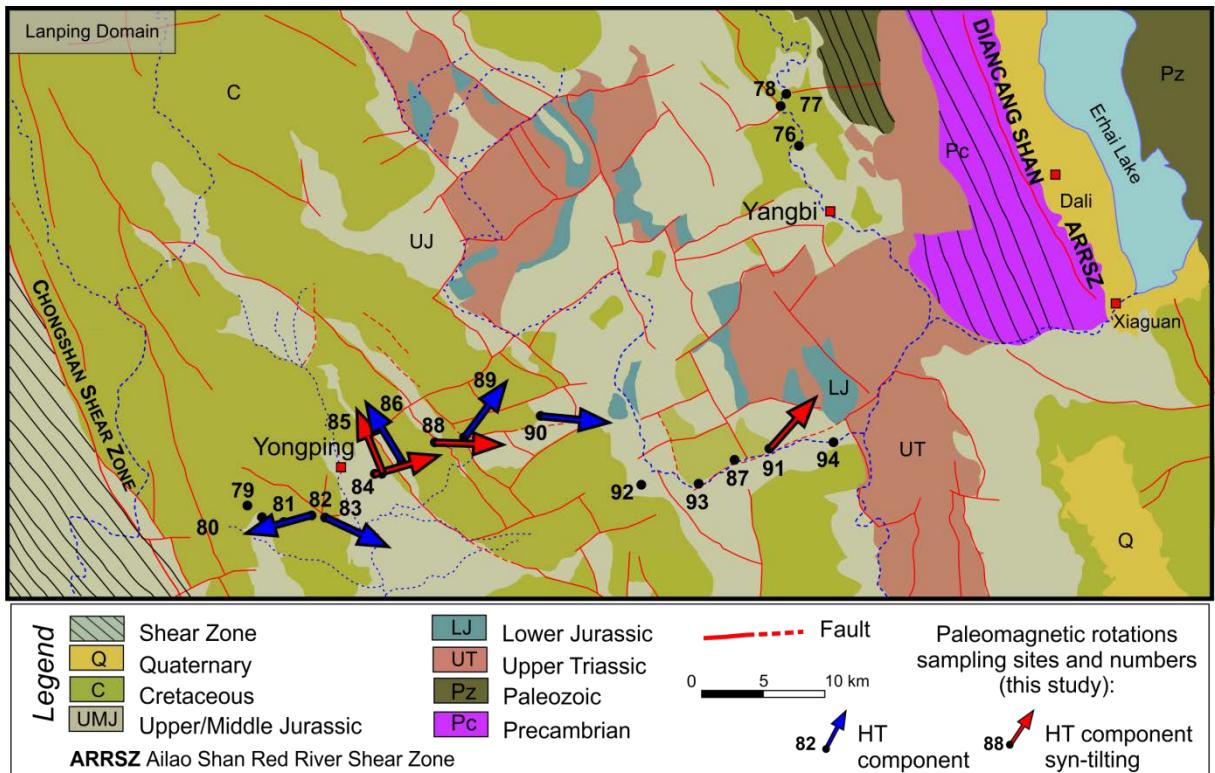


**Figure 71.** Paleomagnetic rotations with respect to Asia from the Lanping domain (geological map modified after *Bureau of Geology and Mineral Resources of Yunnan Province, 1990*).

White and green arrows show the rotations evaluated after high temperature (HT) and characteristic (ChRM) magnetization components, respectively. The medium temperature (MT) components shown to have undergone an overprint at 28% unfolding (see text) are shown in Figure 72a,b.

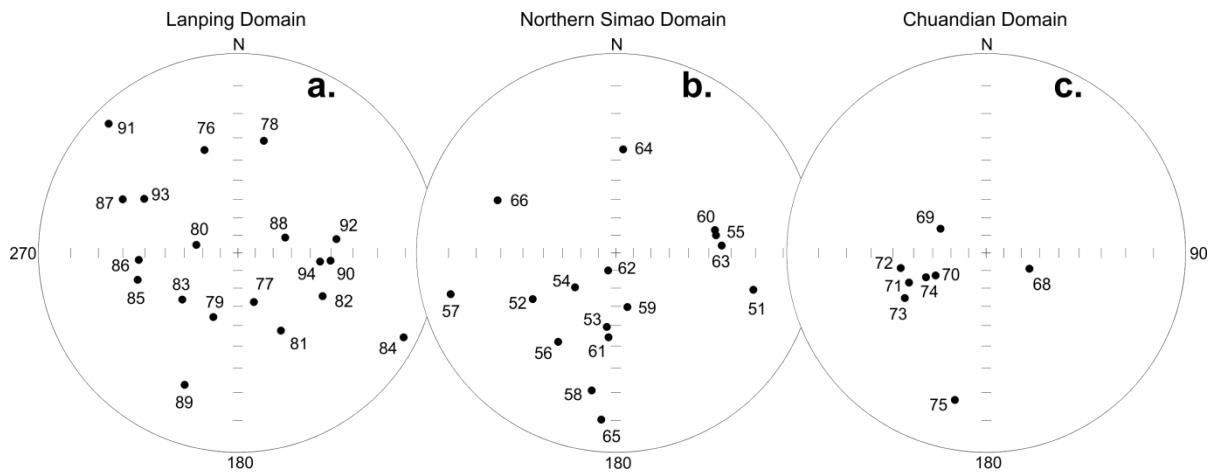


b)



**Figure 72.** Paleomagnetic sampling areas in the north-western Lanping domain of the Ailaoshan Red River Shear Zone modified after *Bureau of Geology and Mineral Resources of Yunnan Province*, 1990. Rotation values are reported in Table 3. a) White, green and pink arrows indicate ChRM/MT components pre-, syn- and post-tilting; b) Blue and red HT components pre- and syn-tilting.

Certainly, our data show that the northern Simao and the Lanping Domain do not rotate rigidly, but are composed of small (few km of size) blocks/domains rotating CW, separated by non-rotating domains, as previously reported by other authors [e.g. *Chen et al.*, 1995; *Tanaka et al.*, 2008; *Tong et al.*, 2013].



**Figure 73.** Poles of bedding at sampling sites from a) Lanping, b) Northern Simao, and c) Chuandian domains.

As clearly shown in Figure 73, the dip of strata in the studied sites is differently oriented being scattered in the Lanping domain, confirming the strong and non-coaxial deformation of the block itself as shown on the geological map; in northern Simao, except for some sites, a SE dip orientation is observed; finally, in the Chuandian domain the poles concentration in SW quadrant is much more clustered.

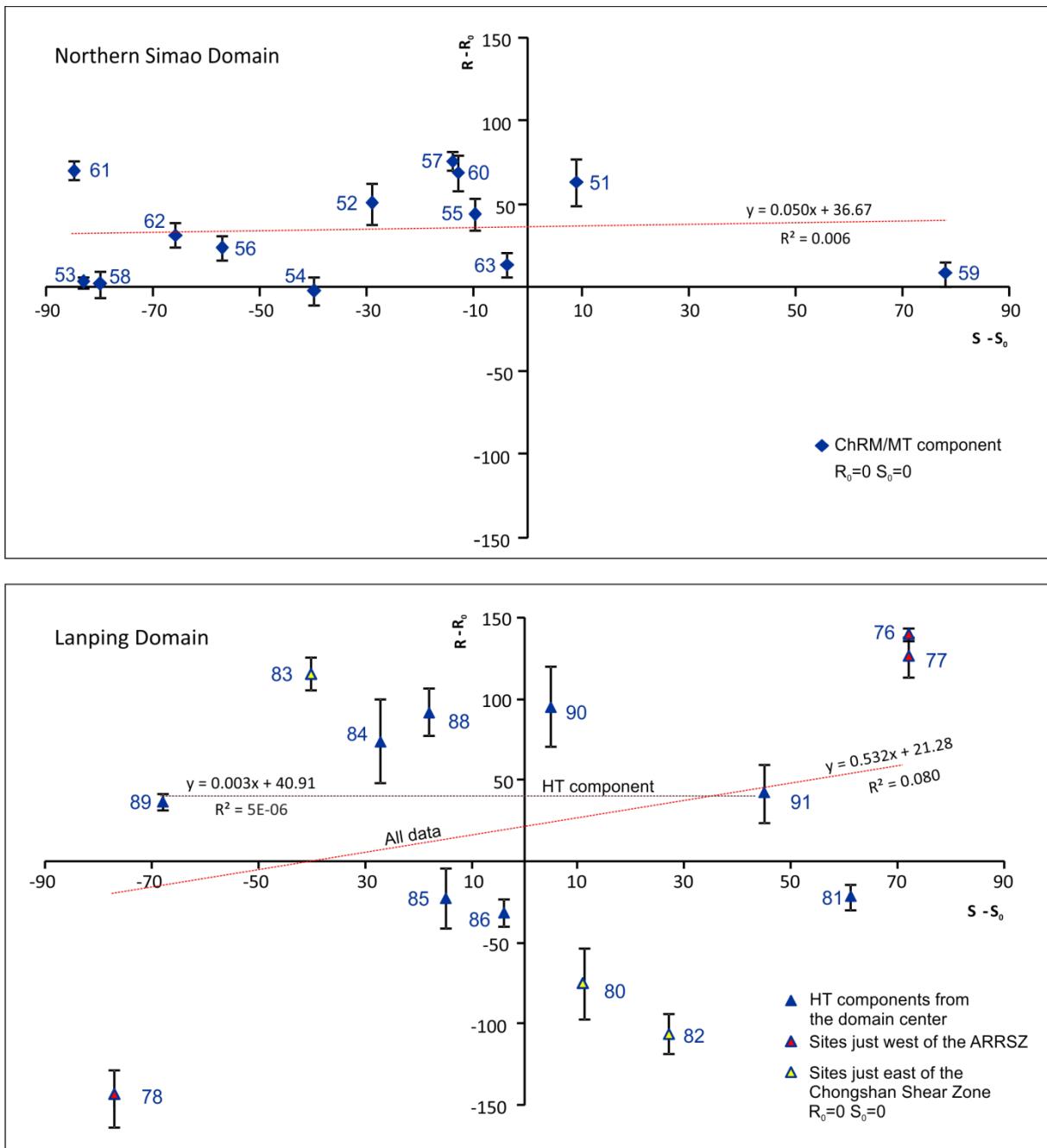
The relation between paleomagnetic rotations and orogenic trends is routinely investigated through the so-called “orocinal plot”, evaluating the correlation between rotations and structural (fold axis) directions (*Schwartz and Van der Voo, 1983; Cifelli and Mattei, 2010; Weil et al., 2010*). A positive correlation indicates orogens that are (progressively) bent, normally by rotational thrust sheet emplacement. Conversely, a lack of correlation indicates chains –or primary arcs - that originated with a curved shape without any rotation (typically around a foreland buttress or an hinterland indenter).

Data from both the Simao and the Lanping domains are very scattered, and show a lack of correlation between rotations and bed strikes (that are used as proxies of unavailable fold axes, Figure 74). The data show that highly variable structural trends arise, if local bed strikes are back-rotated according to paleomagnetic data. Such result makes further unlikely a thrust-

driven rotation scenario or rotations related to orogenic arc formation, and strengthens the hypothesis of local small-block rotations driven by strike-slip tectonics.

To quantify the relationship between structural trends and rotations the correlation between the paleomagnetic rotation values ( $R$ ) of sites with ChRM/MT and HT components and the orientation value of the structure ( $S$ , strike) was tested [Mattei *et al.*, 1995; Speranza *et al.*, 1997; Macrì *et al.*, 2014] using the method originally proposed by Schwartz and Van der Voo [1983] and successively applied by Eldredge *et al.* [1985], Lowie and Hirt [1986], and Hirt and Lowie [1988] (Figure 66). By considering a ~N-S orientation of the structures running parallel to the ARRSZ, as determined from geological maps, we considered a reference value  $S_0=0$ . Likewise we chose  $0^\circ$  for the reference rotation  $R_0=0$ , indicating with positive (negative) value the clockwise (counterclockwise) rotation degree.

Figure 66 shows that in northern Simao and Lanping domain sites with ChRM/MT and HT components respectively, a lack of correlation ( $R^2<0.3$ ) between rotations and bed strikes (that are used as proxies of unavailable fold axes, Figure 74). The data show that highly variable structural trends arise, if local bed strikes are back-rotated according to paleomagnetic data. Such result makes further unlikely a thrust-driven rotation scenario or rotations related to orogenic arc formation, and strengthens the hypothesis of local small-block rotations driven by strike-slip tectonics, even if before rotations occurred, there was no constant tectonic trend with sub-parallel strata and/or folds [Mattei *et al.*, 1995].



**Figure 74.** Paleomagnetic rotation deviations relative to structure direction (approximated by bed strike) deviations in the northern Simao and Lanping domains [e.g. Schwartz and Van der Voo, 1983].  $R_0$  and  $S_0$  are respectively reference rotation ( $R_0=0^\circ$ ) and strike value ( $S_0=0^\circ$ , i.e. N-S trend). Error bars for rotation are the  $\alpha_{95}/\cos l$  site values.  $R^2$  is the correlation coefficient. Red and brown lines are the best fit lines calculated by linear regression analysis on the whole data set and on the sole HT component directions, respectively. Triassic sites from the Simao domain and MT component directions from the Lanping domain are omitted.



*Chapter*  
**VIII**

## 8. AMS RESULTS

AMS was measured for the same sites sampled for paleomagnetic analyses.

The AMS parameters were evaluated using Jelinek statistics [see chapter 2.2; *Jelinek*, 1977; 1978] and are reported in Table 5. The results were processed using the software package ANISOFT supplied with the Kappabridge instrument [*Chadima and Jelinek*, 2009].

The available data associated with each specimen are the geographical orientation and the magnitude of the three principal axes of its AMS ellipsoid.

Sites from Gaoligong Shear Zone (Yun03 to Yun50) are characterized by:

- mean values of magnetic susceptibility Km (between  $2.93 \times 10^{-5}$  to  $1.90 \times 10^{-4}$  SI)
- anisotropy degree ( $1.007 < P < 1.149$ )
- shape of ellipsoid ( $-0.339 < T < 0.786$ )

Sites from Ailaoshan Red River Shear Zone (Yun 51 to Yun 92) are characterized by:

- mean values of magnetic susceptibility Km (between  $4.02 \times 10^{-5}$  to  $1.41 \times 10^{-2}$  SI)
- anisotropy degree ( $1.010 < P < 1.066$ )
- shape of ellipsoid ( $-0.395 < T < 0.772$ )

**Table 5.** Anisotropy factors computed at each sites. Age: UT Upper Triassic, LJ Lower Jurassic, MJ Middle Jurassic, UJ Upper Jurassic, J Jurassic, LK Lower Cretaceous, K Cretaceous, J-K Jurassic-Cretaceous. D and I are *in situ* site-mean declination and inclination, respectively, of the maximum susceptibility axis; km = (kmax + kint + kmin)/3 is mean susceptibility ; P and T are corrected anisotropy degree and shape factor, respectively, according to *Jelinek* [1981]. Bedding is expressed in dip azimuth/dip values. Fabric: PS pure sedimentary, TS triaxial sedimentary, PR prolate, TT triaxial tectonic, N normal, I Inverse

**Table 5.** Anisotropy of Magnetic Susceptibility Results of red beds sites from Baoshan, Lanping, Northern Simao and Chuandian blocks

Site	Age	D,deg (K1)	I,deg (K1)	Baoshan block			Km	Km St.	St. deviation	T	St. deviation	Bedding (deg)
				e <sub>12</sub> , deg	D,deg (K3)	I,deg (K3)						
YUN03	M1	124.2	36.1	22.3	4.5	34.3	32.5	4.69E-05	2.74E-05	1.038	0.054	0.476
YUN24	K	68.6	7.3	31.0	331.3	44.9	43.2	7.02E-05	1.92E-05	1.016	0.005	0.241
YUN25	J-K	—	—	—	—	—	—	—	—	—	—	—
YUN26	K	63.4	12.9	15.6	155.6	9.5	7	6.68E-05	2.81E-05	1.097	0.028	0.551
YUN27	K	66.9	16.1	12.5	336.0	3.2	13.6	7.11E-05	3.59E-05	1.100	0.054	0.321
YUN28	K	99.9	11.0	13.2	6.1	18.7	13.8	8.73E-05	2.66E-05	1.119	0.023	0.677
YUN29	K	258.8	45.7	32.7	148.7	18.6	58.6	3.86E-05	1.18E-05	1.033	0.034	0.338
YUN30	K	86.6	74.2	65.4	310.9	11.5	10.6	6.29E-05	1.50E-05	1.149	0.017	0.786
YUN31	M1	109.2	28.8	44.4	322.2	56.8	14.4	9.10E-05	2.43E-05	1.048	0.016	0.728
YUN32	K	236.3	66.5	30.5	141.8	2.0	20.8	8.97E-05	3.73E-05	1.011	0.003	0.378
YUN33	M1	—	—	—	—	—	—	—	—	—	—	—
YUN34*	M1	181.2	55.7	16.9	84.6	4.5	62.9	5.99E-05	1.57E-05	1.014	0.005	-0.332
YUN35	J-K	29.3	36.5	23.8	151.0	35.4	7.2	1.07E-04	3.31E-05	1.018	0.006	0.446
YUN36	M1	353.2	64.4	23.5	214.1	19.9	35.6	1.13E-04	1.86E-05	1.007	0.003	-0.153
YUN37	M1	253.0	46.1	27.8	147.2	14.7	17.0	1.09E-04	3.02E-05	1.017	0.005	0.251
YUN38	M1	44.2	19.7	34.9	158.4	48.8	19.0	9.51E-05	1.87E-05	1.017	0.005	0.489
YUN39	M1	22.6	62.3	27.8	131.8	9.8	30.7	9.76E-05	2.57E-05	1.043	0.006	0.192
YUN40	K	49.5	70.4	28.9	314.6	1.7	18.7	5.57E-05	1.86E-05	1.054	0.018	0.11
YUN41	K	37.3	27.1	13.6	128.3	2.0	73.4	2.93E-05	7.14E-06	1.048	0.078	0.157
YUN43	M1	270.8	2.7	20.6	180.7	1.0	22.6	1.90E-04	6.24E-05	1.008	0.002	-0.047
YUN44	M1	312.9	3.6	28.6	183.0	41.5	7.1	1.34E-04	2.83E-05	1.019	0.005	0.417
YUN45	M1	336.1	6.0	15.5	245.5	5.9	8.2	1.11E-04	3.12E-05	1.026	0.006	0.486
YUN46	M1	21.7	21.1	6.5	120.9	22.4	53.8	8.23E-05	8.83E-06	1.023	0.007	-0.155
YUN47*	M1	261.3	30.5	20	11.6	30.5	34.0	1.06E-04	1.84E-05	1.012	0.004	-0.339
YUN48	M1	149.4	14.9	10.6	246.3	24.3	10.2	6.12E-05	3.02E-05	1.096	0.053	0.443
YUN49	M1	184.3	67.3	37.7	77.9	6.7	17.2	1.21E-04	1.49E-05	1.017	0.003	0.534
YUN50	M1	162.8	72.6	16.6	305.5	14.0	15.1	1.55E-04	2.55E-05	1.009	0.004	0.179
<b>Northern Simao Block</b>												
YUN51	LK	36.4	15	10.0	272.9	64.1	22.9	1.20E-04	1.81E-05	1.010	0.003	0.190
YUN52	UJ	335.3	15.4	72.0	232.8	38.3	8.6	1.12E-04	1.17E-05	1.033	0.007	0.644
YUN53	UJ	299.4	26.7	11.6	97.2	61.5	35.1	1.28E-04	2.74E-05	1.017	0.007	-0.160
YUN54	M1	124.8	16.5	39.7	312.2	73.4	10.6	1.34E-04	2.61E-05	1.013	0.005	0.374
YUN55	UJ	180.6	23.9	22.7	295.5	43.5	12.8	6.71E-05	3.43E-05	1.029	0.029	0.071
YUN56	LK	—	—	—	—	—	—	—	—	—	—	—
YUN57	LK	—	—	—	—	—	—	—	—	—	—	—
YUN58	LK	295.6	49.9	26.2	189.5	13.2	12.3	9.03E-05	1.57E-05	1.035	0.015	0.508
YUN59	LK	344.9	15.1	7.6	243.1	37.0	43.5	1.13E-04	2.88E-05	1.031	0.013	0.049
YUN60*	UJ	84.2	35.1	42.2	309.8	44.9	32.3	4.02E-05	1.35E-05	1.024	0.011	0.334
YUN61	UJ	7.0	27.0	14.5	151.9	58.1	18.6	1.05E-04	1.17E-05	1.015	0.003	-0.001

Table 5. Anisotropy of Magnetic Susceptibility Results of red beds sites from Baoshan, Lanping, Northern Simao and Chuandian blocks

Site	Age	D,deg	I <sub>1</sub> ,deg (K1)	e <sub>12</sub> , deg	D,deg (K3)	I <sub>1</sub> ,deg (K3)	e <sub>31</sub> , deg	Km	Km St.	St.	T	St.	Bedding (deg)
YUN62	UJ	354.0	7.1	13.1	88.0	29.3	12.5	1.28E-04	1.81E-05	1.019	0.003	-0.13	0.236
YUN63	MJ	20.5	15.5	7.7	235.5	71.3	18.7	1.22E-04	1.19E-05	1.039	0.013	-0.247	0.262
YUN64	UT	201.9	67.2	22.3	94.2	7.3	20.8	1.29E-04	1.02E-05	1.04	0.013	0.040	0.323
YUN65	UT	139.6	60.4	13.7	289.0	26.1	12.0	1.42E-04	2.07E-05	1.034	0.008	0.105	0.236
YUN66	UT	67.4	59.1	17.7	257.2	30.5	37.4	1.14E-04	1.24E-05	1.027	0.008	0.031	0.434
<b>Chuandian block</b>													
YUN67	UT	299.8	35.0	21.2	97.2	52.9	8.3	1.36E-04	2.08E-05	1.195	0.016	0.772	0.127
YUN68	UT	349.7	24.1	40.2	100.2	38.1	14.8	6.16E-05	1.78E-05	1.064	0.035	0.250	0.361
YUN69	UT	244.0	61.4	21.4	81.9	27.4	20.0	1.47E-04	6.63E-05	1.062	0.030	0.343	0.273
YUN70	UJ	—	—	—	—	—	—	—	—	—	—	—	—
YUN71	MJ	79.2	60.2	41.5	243.2	28.9	42.4	1.21E-04	2.13E-05	1.037	0.033	0.169	0.404
YUN72	MJ	164.5	7.0	33.0	261.2	43.4	18.9	1.14E-04	1.77E-05	1.023	0.009	0.394	0.322
YUN73	MJ	24.5	47.6	27.8	246.2	34.3	26.5	9.05E-05	1.44E-05	1.042	0.024	0.656	0.228
YUN74	UJ	330.8	57.9	5.9	233.2	4.7	6.3	1.57E-04	1.04E-05	1.026	0.001	0.292	0.130
YUN75	UJ	296.2	27.0	16.9	192.1	25.6	9.1	1.40E-04	5.97E-06	1.029	0.002	0.557	0.069
<b>Lanping block</b>													
YUN76	LK	203.6	48.7	49.4	342.1	33.3	23.7	7.19E-05	7.60E-06	1.014	0.005	-0.025	0.307
YUN77	LK	317.8	13.4	25.6	199.6	63.2	21.3	9.93E-05	3.74E-05	1.021	0.007	0.311	0.346
YUN78	LK	272.4	17.4	28.4	12.7	29.8	22.7	8.11E-05	1.58E-05	1.020	0.006	0.146	0.403
YUN79	LK	318.1	4.6	17.3	223.5	44.8	13.7	6.63E-05	8.46E-06	1.018	0.005	0.371	0.341
YUN80*	LK	264.7	20.7	34.5	80.7	69.3	8.1	9.55E-04	1.87E-05	1.066	0.009	0.565	0.227
YUN81	LK	57.0	4.5	37.1	151.0	41.8	12.9	9.86E-05	1.06E-05	1.013	0.003	0.374	0.289
YUN82	LK	227.7	22.9	38.1	117.4	39.2	9.3	8.74E-05	3.72E-06	1.046	0.078	0.154	0.291
YUN83	LK	25.7	24.7	13.5	213.8	65.1	15.8	1.17E-04	1.26E-05	1.053	0.009	0.503	0.192
YUN84	UJ	213.0	56.8	24.8	117.3	3.7	14.0	1.17E-04	1.08E-05	1.025	0.004	0.468	0.159
YUN85	UJ	147.6	33.5	18.5	262.1	31.9	28.8	7.45E-05	1.23E-05	1.025	0.012	0.315	0.374
YUN86	UJ	4.9	11.4	33.0	266.1	37.1	5.0	9.37E-05	9.56E-06	1.032	0.003	0.688	0.168
YUN87	LK	178.8	28.8	7.2	273.3	8.2	14.5	8.78E-05	7.61E-06	1.037	0.006	-0.395	0.251
YUN88*	LK	27.1	64.6	16.0	257.0	17.2	9.3	1.49E-04	2.24E-05	1.038	0.007	-0.002	0.256
YUN89	LK	353.3	68.2	15.6	202.4	19.3	13.7	1.52E-04	2.98E-05	1.053	0.008	0.178	0.518
YUN90	UJ	306.2	50.3	22.8	105.2	37.8	7.4	1.59E-04	1.92E-05	1.058	0.016	0.490	0.187
YUN91	LK	222.8	37.6	8.4	324.4	14.6	13.1	8.59E-05	1.90E-05	1.040	0.010	-0.070	0.352
YUN92	UJ	185.6	17.6	16.0	81.8	36.9	9.5	1.41E-02	1.52E-05	1.014	0.005	0.030	0.272
YUN93	LK	—	—	—	—	—	—	—	—	—	—	—	—
YUN94	MJ	—	—	—	—	—	—	—	—	—	—	—	—

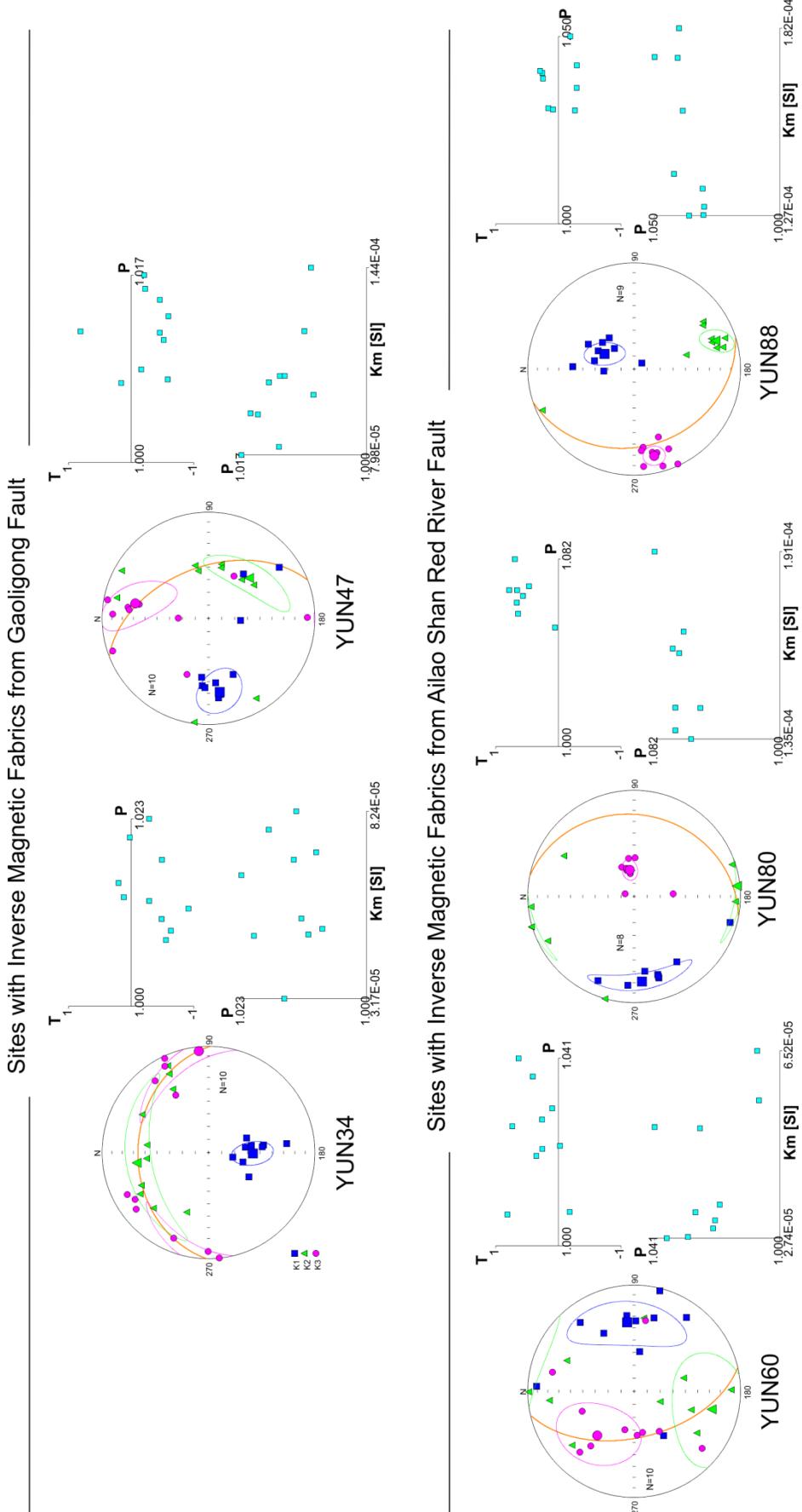
Age: UT Upper Triassic, L1 Lower Jurassic, M1 Middle Jurassic, UJ Upper Jurassic, J1 Lower Cretaceous, K1 Cretaceous, D1 Kimmeridgian-Cretaceous, D and I are in situ site-mean declination and inclination, respectively, of the maximum susceptibility axis; km = (kmax + kint + kmin)/3 is mean susceptibility; P and T are corrected anisotropy degree and shape factor, respectively, according to Jelinek [1981]. Bedding is expressed in dip azimuth/dip values. Fabric: PS pure sedimentary/dip values. Fabric: PR prolate, TT triaxial tectonic, N normal, I inverse

The magnetic lineation relative to GLG data is well defined only in 18 sites, where the  $e_{12}$  confidence angle of the  $K_{\max}$  in the  $K_{\max}$ - $K_{\text{int}}$  plane is  $< 30^\circ$ . In 5 sites (Yun24, 29, 32, 38, 49) the magnetic lineation is still defined but has a larger confidence angle ( $30^\circ < e_{12} < 40^\circ$ ), while the remaining 2 sites (Yun30, 31) have a purely scattered results with no virtually lineation ( $e_{12} >$  or  $>> 40^\circ$ ).

Relatively to ARRSZ, magnetic lineation data is well defined in 28 sites, where the  $e_{12}$  confidence angle of the  $K_{\max}$  in the  $K_{\max}$ - $K_{\text{int}}$  plane is  $< 30^\circ$ . In 6 other sites (Yun54, 72, 80, 81, 82, 86) the magnetic lineation is still defined but has a larger confidence angle ( $30^\circ < e_{12} < 40^\circ$ ), while the remaining 5 sites (Yun52, 60, 68 71, 76) have a purely scattered results with virtually no lineation ( $e_{12} > 40^\circ$ ). Sites with poorly defined lineation have been discarded from further analysis but are shown in Figure 77.

Unusual relationships between structural and magnetic axes occurred in two/three sampled sites along the Gaoligong Shear Zone (Yun34 and Yun47) and ARRSZ (Yun60, Yun80, Yun88) transect respectively, showing an inverse, or tendentially inverse magnetic fabric, with well defined lineation ( $e_{12} < 30^\circ$ ) and prolate ellipsoid shape orthogonal to bedding (\* in Table 5 and plots in Figure 75). As explained in chapter 2.2, this unusual relationships between structural and magnetic axes, can occur because of the presence of certain magnetic minerals, either SD magnetite, iron-bearing carbonates or various paramagnetic minerals like tourmaline, cordierite, goethite or siderite [Rochette *et al.*, 1992]. In particular, rocks with fine-grained magnetite, are prone to anomalous AMS fabric.

**Figure 75.** Sites with inverse (or tendentially inverse) magnetic fabric respectively from Gaoligong and Ailaoshan Red River fault areas (equal area Schmidt projection, lower hemisphere). The squares, triangle, and dots represent  $K_{\max}$ ,  $K_{\text{int}}$  and  $k_{\min}$  respectively. The ellipses indicate the 95% region around the principal susceptibility axes. The orange line indicates the bedding planes (expressed in dip azimuth/dip values).

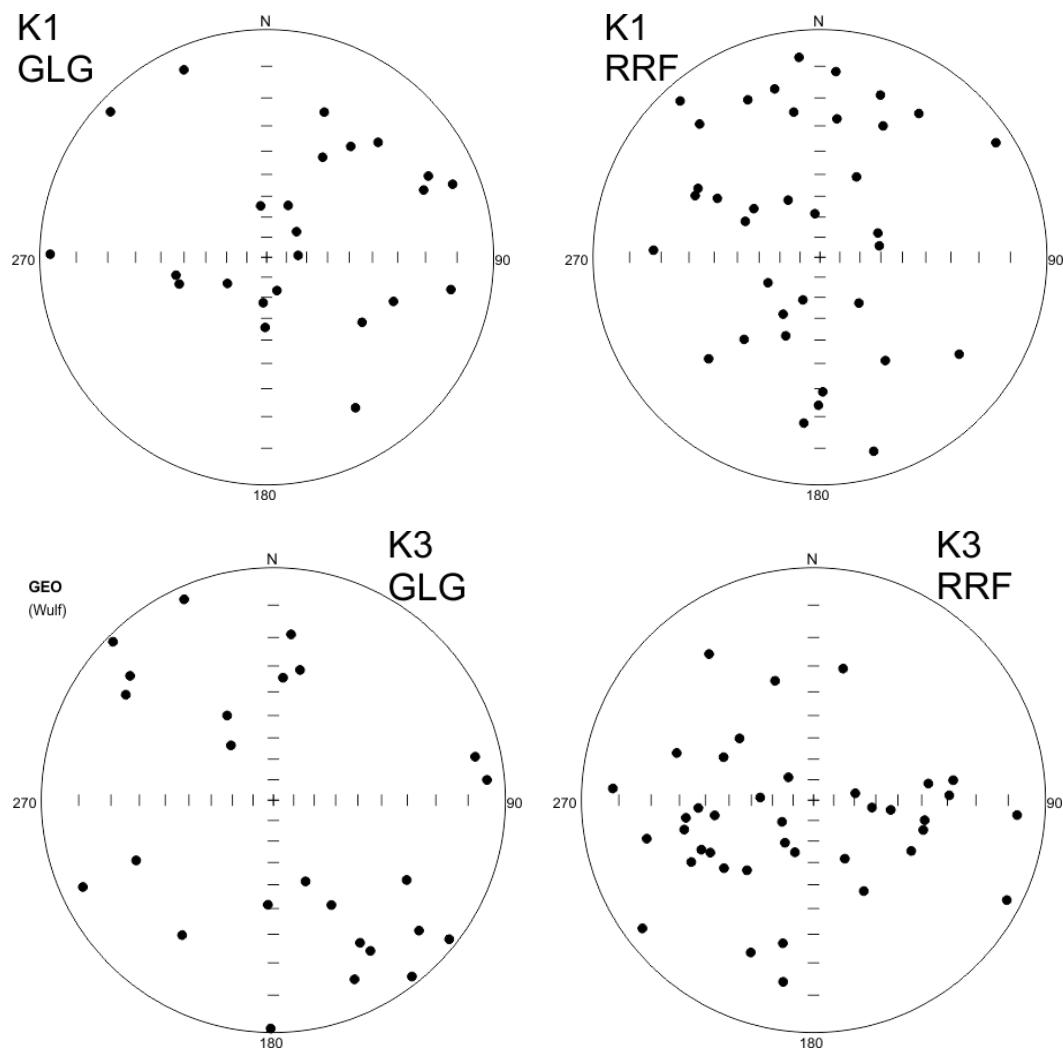


As already described in previous chapters, previous field observations along the Gaoligong Mountains showed that the Gaoligong Shear Zone is characterized by steep foliation and nearly N-S trending sub-horizontal stretching lineations [Wang *et al.*, 2008]. The banded mylonitic rocks are tightly folded with the hinges parallel to the lineations [Zhang *et al.*, 2012a]. To the west of the Gaoligong strike-slip Shear Zone, Xu *et al.* [2015] identified widespread mylonitic ortho- and paragneisses with gently dipping foliation and nearly horizontal stretching lineation defined by long axes of amphibole and sillimanite grains and boudinage of felsic veins. Despite the varying attitude of the foliation, the stretching lineation has a consistent NE-azimuth, parallel to the strike of the Gaoligong Shear Zone.

In the Lanping-Simao Basin, the red beds are generally affected by the NNW-SSE trending folds and thrusts, but an arcuate trend is maintained in the Simao Basin [Leloup, 1995]. Timing of fold and thrust formation remains controversial. In the central sector of the Red River Fault, the foliation is usually steep, and the lineation is usually nearly horizontal, both being almost parallel to the trend of the mylonitization in the gneiss massifs [e.g., Leloup *et al.*, 1993, 2001; Jolivet *et al.*, 2001; Cao *et al.*, 2010; Zhang *et al.*, 2017]. In the southern transect of the ARRSZ in the Zhengyuan area, the Nanxin Formation is separate by several anticlines N-S trending and, to the west others monoclinal stuctural. Also in Mengla locality, a monoclinal structure with N-S trending axis and a dip of 2°-63° always in Nanxin Formation (from northeastward to eastward) has been observed by Tanaka *et al.* [2008].

Observing the distribution of the K1 axes on a stereoplot obtained by our data (Figure 76), it is noteworthy that in both cases ~ N-S orientation is predominant, similar to the orientation of primary tectonic structures analysed by us (GLG and ARRSZ) but also with minor secondary structures bordering the shear zones (eg. Nujiang River Fault; see structural and geological maps in previous chapter). The most striking evidence is the frequently sub-vertical trend of the magnetic lineations that is unusual in fold and thrust tectonics, and might reflect the main influence of strike-slip tectonics on the sediments located within the blocks.

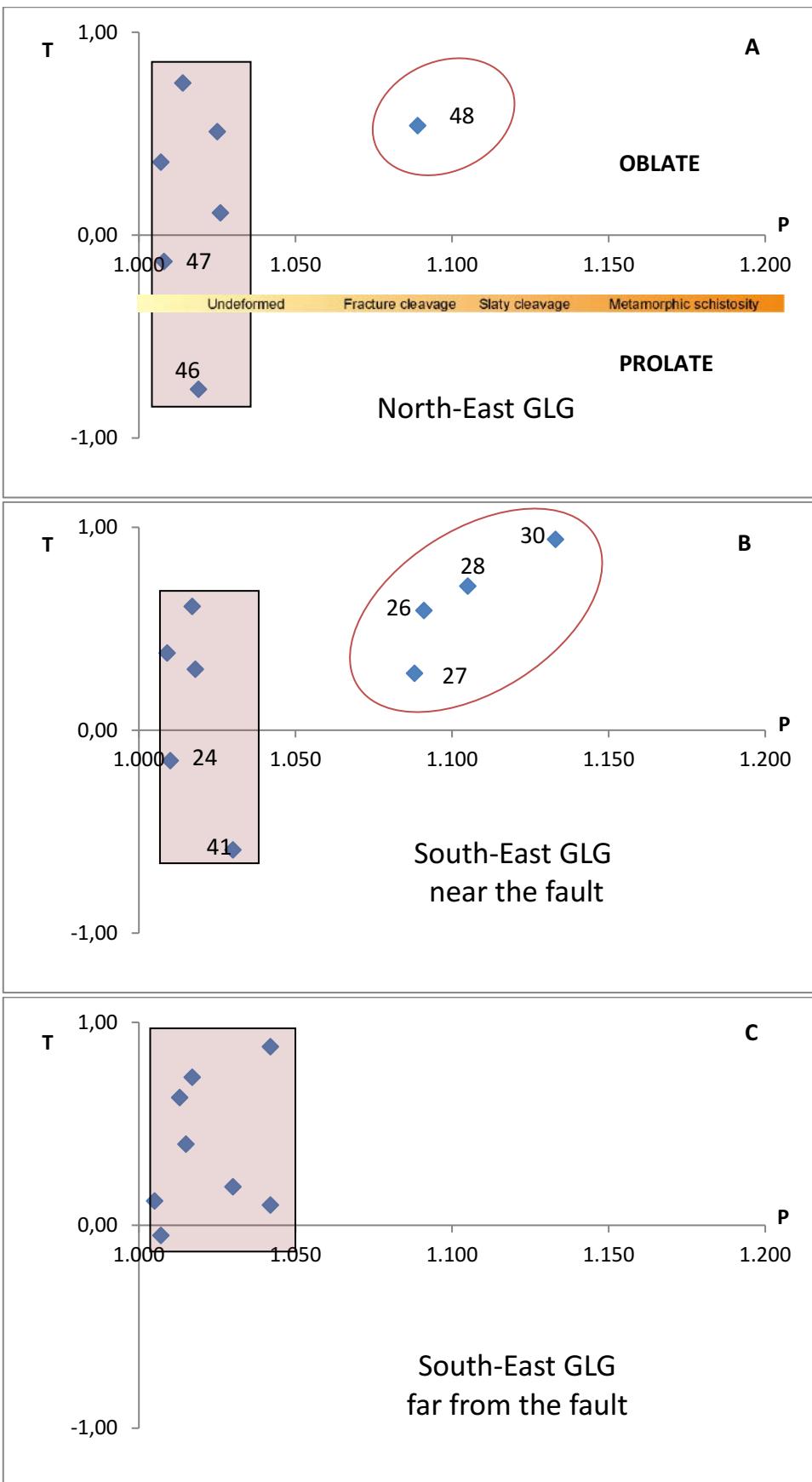
In particular, for samples collected along the GLG, the mean distribution of the K1 axis seem to have a roughly subvertical ~NE-SO orientation and therefore a perpendicular K3 orientation ~ NO-SE. Instead, for the ARRSZ, the mean distribution of the K1 axes seem to have a rough orientation ~N-S and therefore a perpendicular K3 axes distribution, ~E-O oriented (Figure 76). In both cases, data suggest a medium structural compressive deformation phase (from oblate to triaxial fabric according to relations between shape factor (T) of AMS ellipsoid and the anisotropy degree parameter (P), discussed below; Figure 77 A-F).

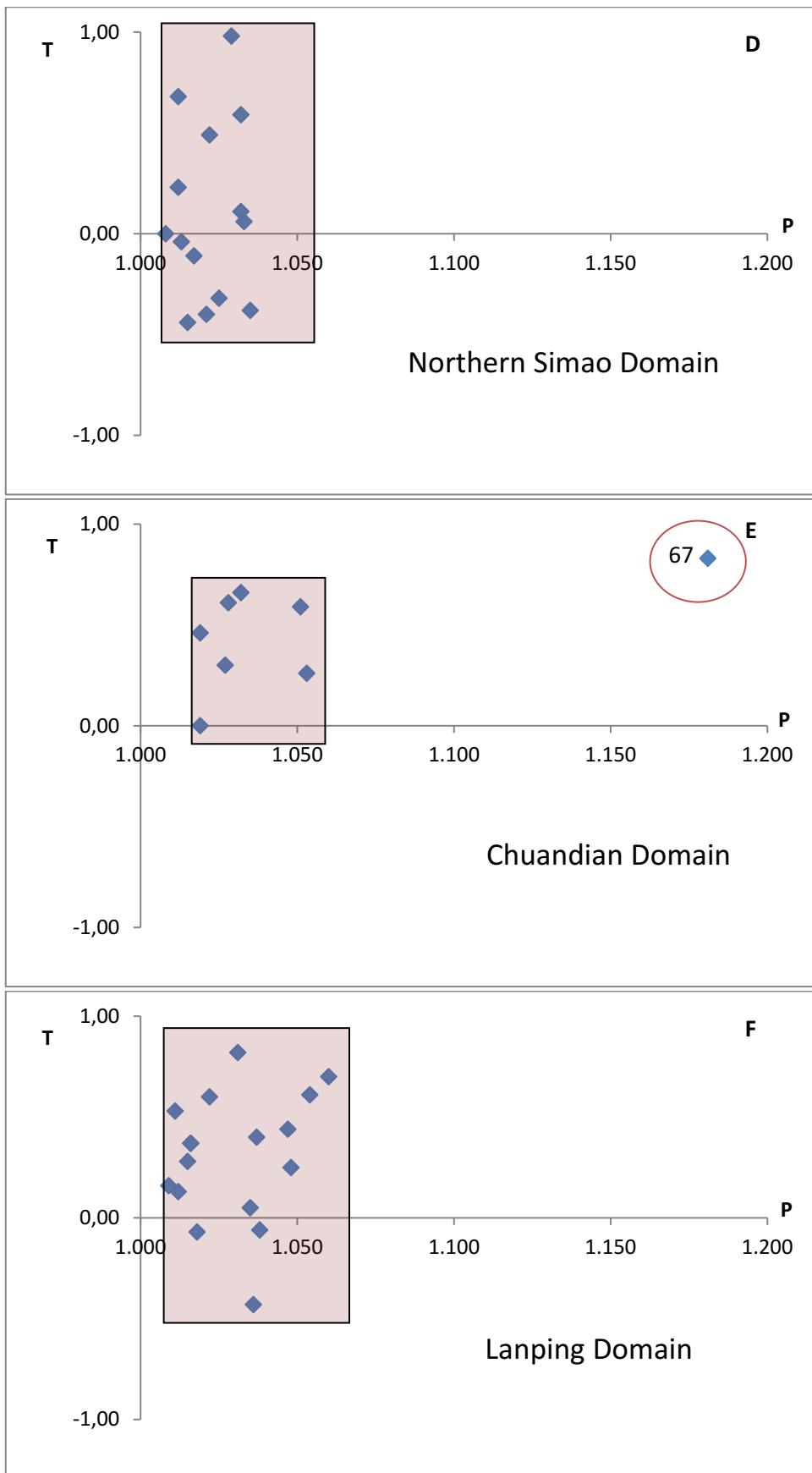


**Figure 76.** Lower hemisphere equal area projection of magnetic lineations (K1) (A) and foliations poles (K3) (B) respectively, in the sites carried out from Gaoligong fault and from Red River fault. Open circles represent samples from red beds.

Moreover, the shape factor (T) of AMS ellipsoid and the anisotropy degree parameter (P) seem to be broadly constant also moving away from the faults. For all sites (Figure 77 A-F) the shape of the AMS ellipsoid is predominantly oblate, with a shape factor  $0 < T < 1$ , and the anisotropy degree parameter is  $1 < P < 1.080$  indicating a prevalent undeformed state. Only in 5 sites (Yun48; Yun26, 27, 28, 30; Figure 77 A-B) located at a distance of less than 2 km from the Gaoligong Shear Zone, respectively to the north-east and south-east of the structure, show an higher anisotropy degree parameter with respect to the others ( $P>1.080$ ) suggesting a degree of deformation between fracture and slaty cleavage. Indeed, site Yun48 showed an unusual strong macroscopic foliation *in situ* (Figure 77 A, 78), not found in other sites (see supplementary information for each detailed plots). This confirms, once again, which the deformation in areas very close to the major structure, is influenced by tectonic structure and that the structural behaviors also change according to the distance from the faults. The same is true for the Yun67 site, located in the Chuandian domain, which present a P value higher than 1.150 and well defined foliation is visible *in situ* (Figure 77E, 80).

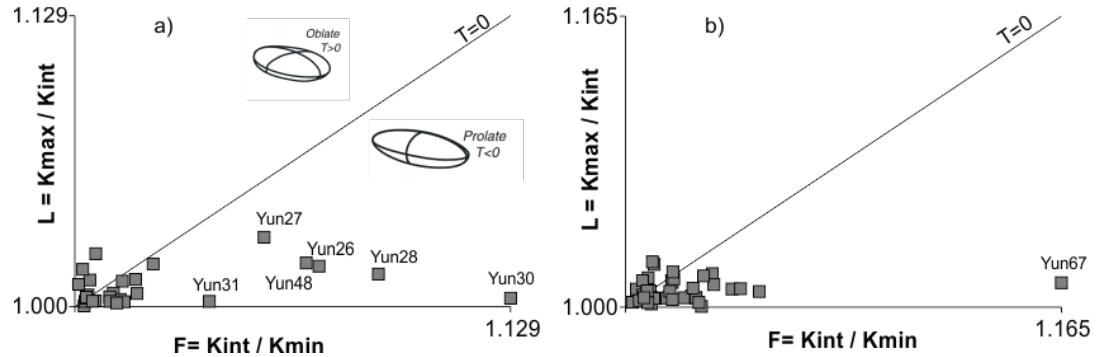
Sites Yun24, 41, 46, 47 have a shape factor  $T < 0$  and a fabric tendentially inverse, except for Yun24 in which we have no evidence of inverse fabric (Table 5 and Figure 77, 79).



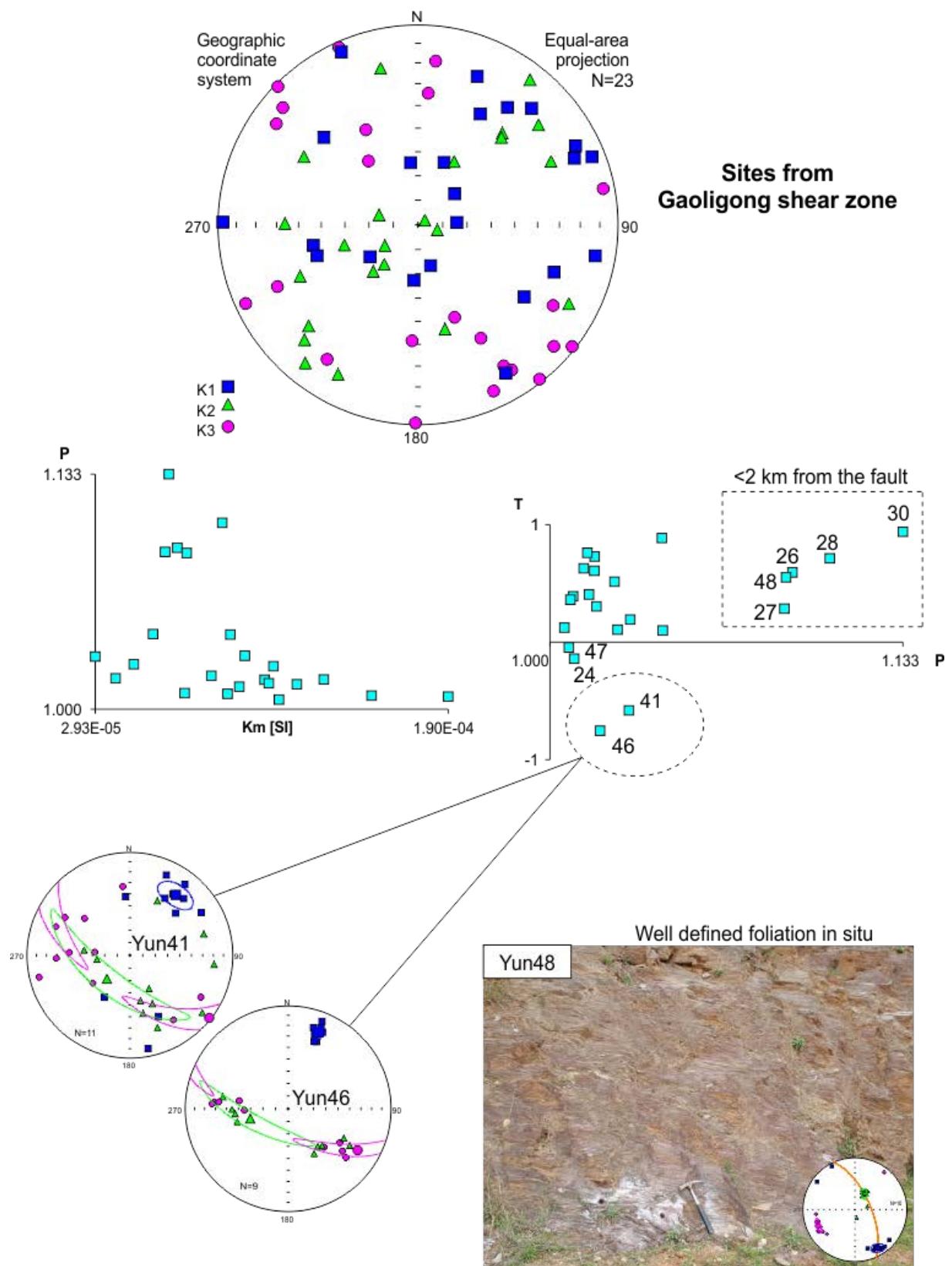


**Figure 77.** A-F Jelinek plot from different sampling areas. T, shape factor of AMS ellipsoid; P, anisotropy degree parameter. See Table 5 for values. the red rectangle indicates the average distribution area of the analyzed sites. The red circle highlights sites with a value of  $P > 1.050$ .

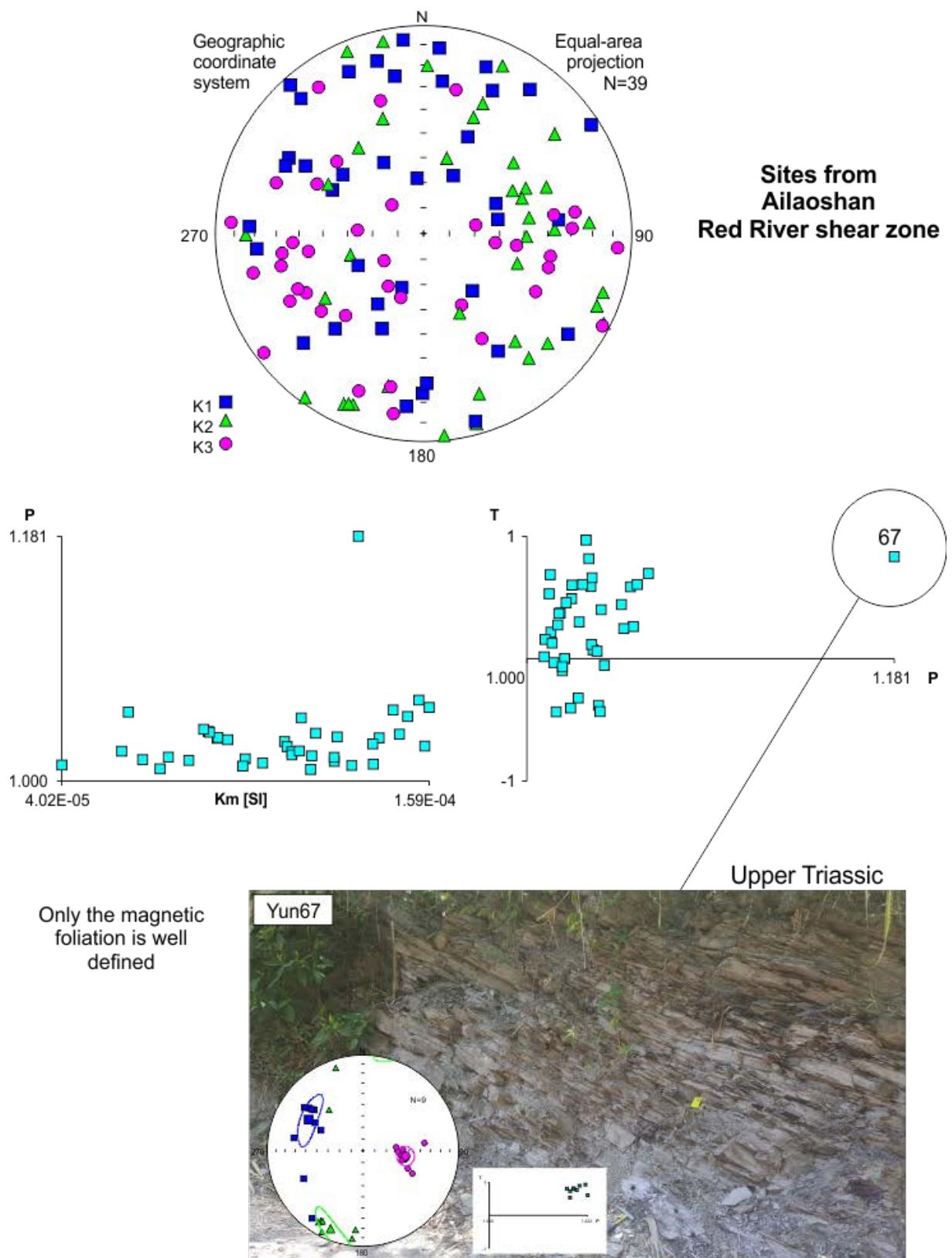
The eccentricity of the susceptibility ellipsoids have been evaluated (Figure 78) by plotting L (lineation) versus F (foliation). The distribution of both the foliation and lineations values, is indicative of predominant triaxial magnetic fabric, except for sites Yun26, 27, 28, 30, 31, 48, cropping out less than 5 km from the Gaoligong Shear Zone (Figure 79), and Yun67 sampled in Chuandian domain (Figure 80).



**Figure 78.** Flinn diagrams [Flinn, 1962] of shape parameters (L lineation, F foliation) relative to AMS results from a) Gaoligong Shear Zone and b) Ailaoshan Red River Shear Zone. Lineation/Foliation values are the results of  $K_{\max}/K_{\text{int}}$  and  $K_{\text{int}}/K_{\min}$ , respectively. T, is the shape factor of AMS ellipsoid (oblate,  $T>0$ ; prolate  $T<0$ ).



**Figure 79.** Summary of AMS results obtained from the Gaoligong Shear Zone. The squares, triangle, and dots represent  $K_{\max}$ ,  $K_{\text{int}}$  and  $k_{\min}$  respectively. The ellipses indicate the 95% region around the principal susceptibility axes. The orange line indicates the bedding planes (expressed in dip azimuth/dip values).



**Figure 80.** Summary of AMS results obtained from the Ailaoshan Red River Shear Zone. The squares, triangle, and dots represent  $K_{\max}$ ,  $K_{\text{int}}$  and  $k_{\min}$  respectively. The ellipses indicate the 95% region around the principal susceptibility axes. The orange line indicates the bedding planes (expressed in dip azimuth/dip values).

Observing plots with projection of the bedding, lineation and the foliation planes, it is possible to summarize that:

(1) along the Gaoligong Shear Zone, in the north-eastern area (Figure 81a), the magnetic foliation (in pink) is predominantly sub-parallel, and sometimes coincident, to the bedding plane (in yellow), exception for sites Yun03 and Yun46 (Middle Jurassic), as is commonly observed in sediments. Furthermore, the magnetic lineation clusters are mostly oriented NW-SE exception for site Yun46 NNE-SSW oriented. Also in the south-eastern transect (Figure 81b), the magnetic foliation is predominantly sub-parallel, and sometimes coincident, to the bedding plane, exception for sites Yun24 and 29 (Cretaceous age), and Yun36, 39, 43 (Middle Jurassic age). But, we note that the axis of maximum elongation, parallel to K1, is sometimes sub-vertical, according to predominantly strike-slip regime, completely different from cases of thrust tectonics, where it appears mainly sub-horizontal [eg. *Speranza et al.*, 1999; *Maffione et al.*, 2015]. The distribution of the magnetic lineations cluster is the southern part, is variable. The sampled sites are located in an area delimited by two normal faults (Figure 81b). *Huang and Opdyke* [1993] describe this area as a monocline with moderate dip. In a group of sites to the north (Yun24, 26, 27, 39, 32, 35) and to the south (Yun37, 38, 39, 40, 41), located near the two main structure, a NE-SW direction seems to prevail, consistent with the orientation of the Gaoligong west detachment fault. An opposite trend respect to two group mentioned above, is shown by sites Yun28, 30, 31, 36, 43, 44;

(2) along the Red River Shear Zone, the situation is more complicated because the results are scattered. In northern Simao block (Figure 81c), from site Yun59 to Yun66, the foliation is perpendicular to the bedding plane, in Yun58 are almost coincident, Yun53, 54, 55 again perpendicular and in Yun51 and Yun52 are respectively perpendicular and parallel. This seems to highlight a different behaviour in different areas but very close to each other. Furthermore, the dispersion of the fabric indicates that it was probably acquired after rotation. Parallelism between K1 and the fold axis direction within the Lanping–Simao Basin, was also

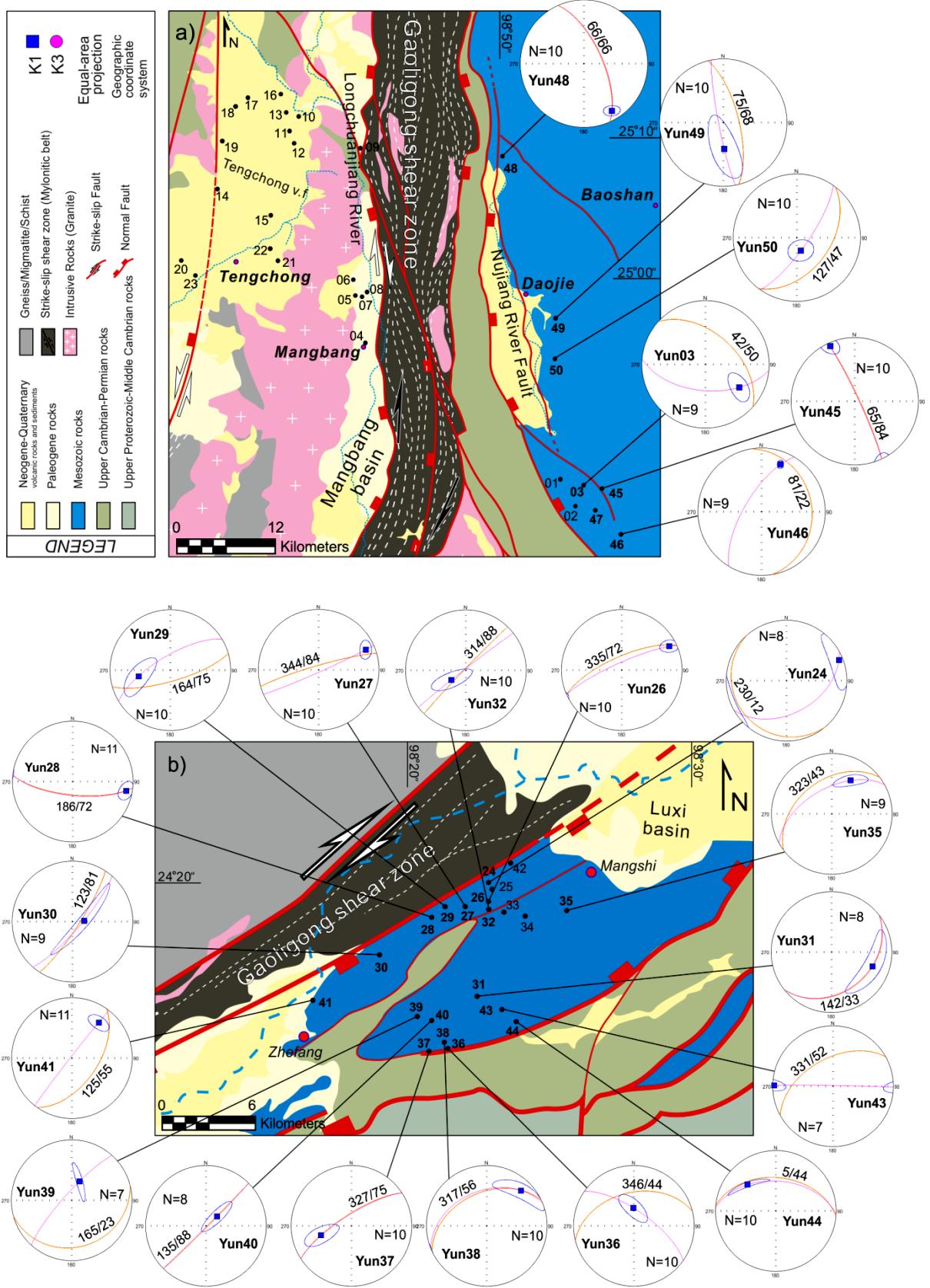
evidenced by *Sato et al.* [2007], confirming the tectonic origin of the observed magnetic lineation, implying that K1 is a passive marker that was rotated during shortening [*Gao et al.*, 2015];

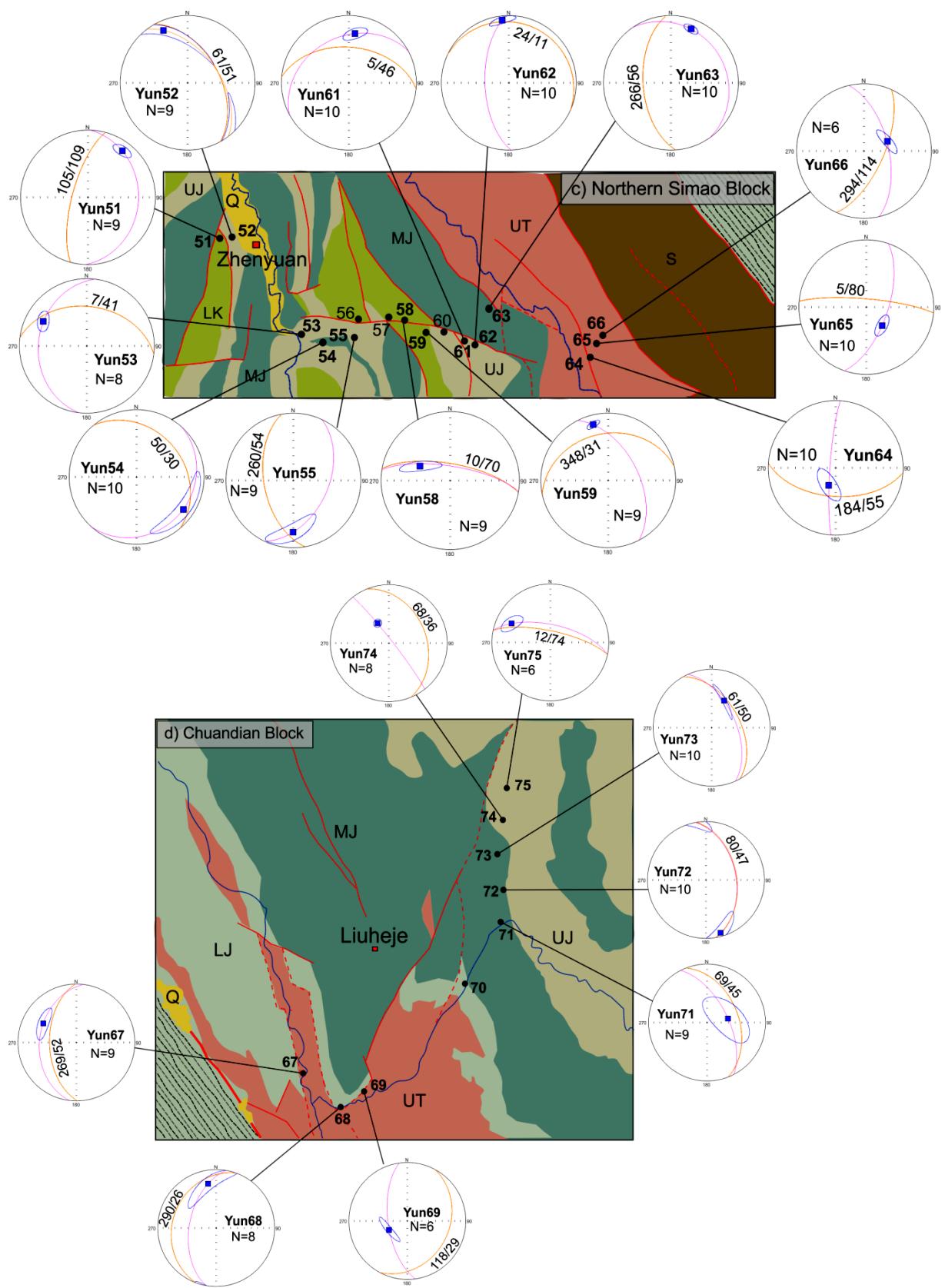
(3) different are the results in Chuandian domain (Figure 81d), where bedding and foliation planes are predominantly parallel, exception for Yun69 and Yun74. Also here we have no other evidence to correlate these two results;

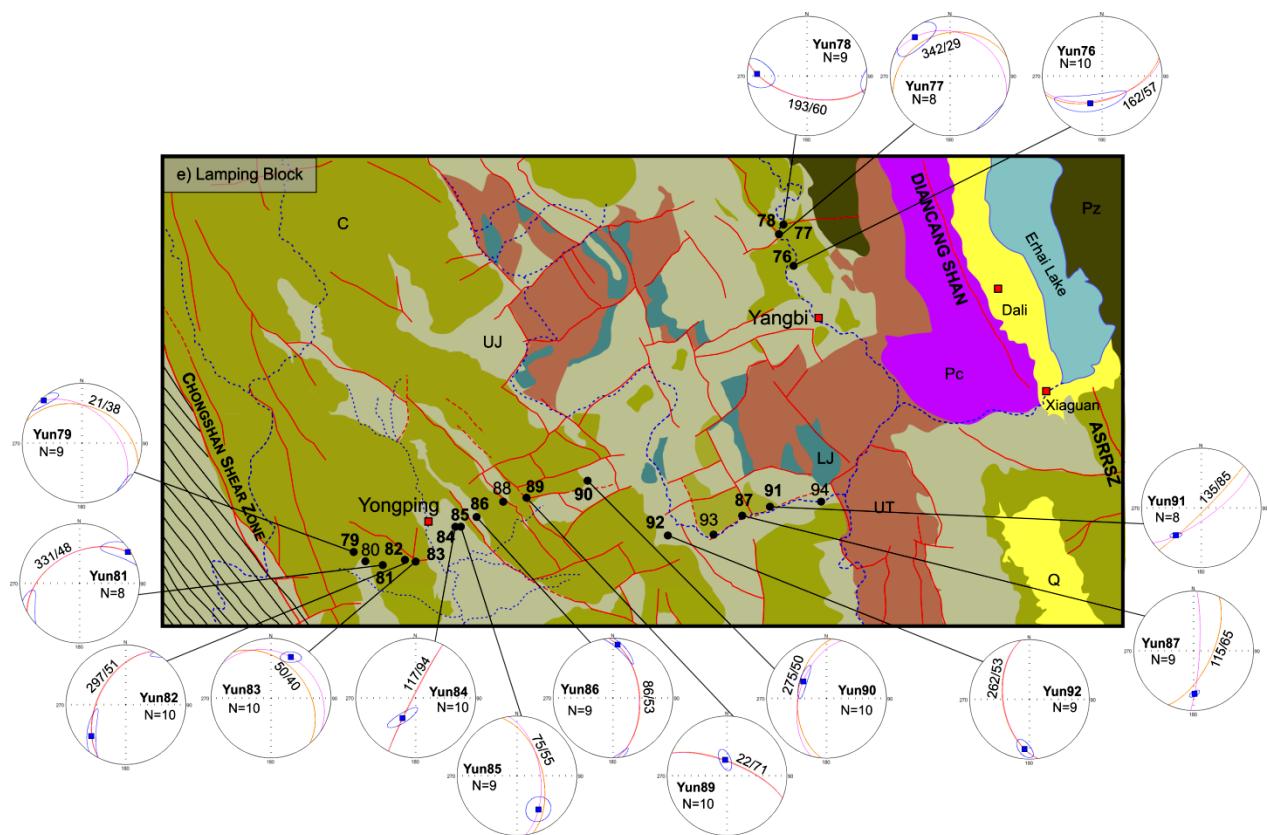
(4) finally, in the Lanping domain, the bedding and foliation planes are almost all parallel (Figure 81e). In both cases, the irregular foliations orthogonal to the bedding plane, indicate that there is a strong tectonic activity inside the blocks and these are deformed inside them.

This confirms, once again, which there are no evidence of rigid rotations of the crustal blocks, but the deformation is widely distributed both along the main structures and within the blocks delimited by them. Indeed, according to our paleomagnetic results and previous data, especially for the Gaoligong strike-slip fault area, it seems clear that tectonic deformation was mainly driven by strike-slip fault activity.

**Figure 81.** Structural sketch maps and plotted AMS results. Pink line represents the foliation planes ( $k_{\min}$ ) and blue square indicate the lineation planes ( $K_{\max}$ ). The ellipses indicate the 95% region around the principal susceptibility axes. The orange line indicates the bedding planes (expressed in dip azimuth/dip values). **a)** and **b)** are relative to the Gaoligong Shear Zone. **c), d)** and **e)** are relative to the Ailaoshan Red River Shear Zone.









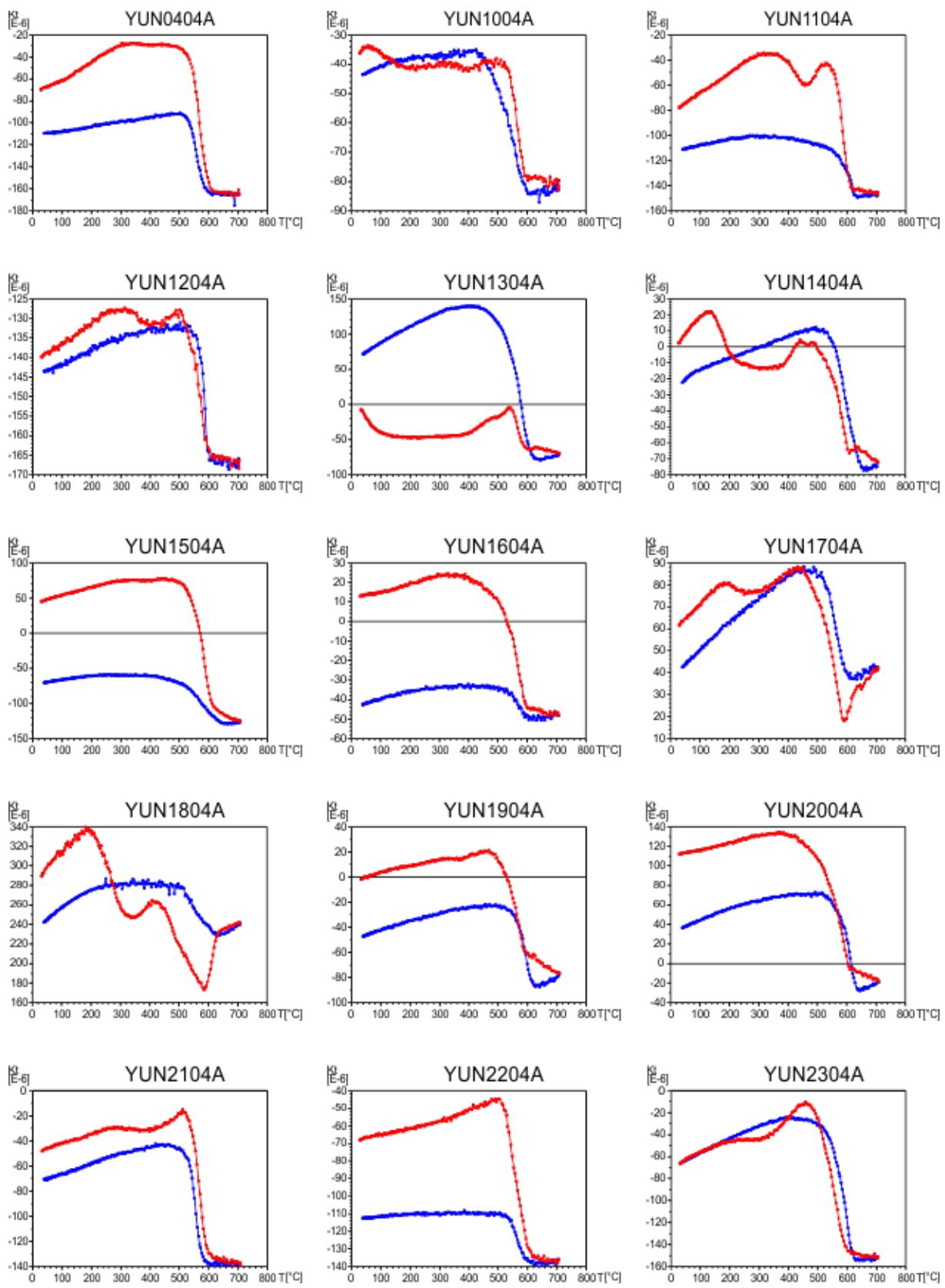
*Chapter*  
***IX***

## 9. MAGNETIC MINERALOGY RESULTS

Additional magnetic analyses were carried to characterize the magnetic mineralogy. The measured hysteresis parameters include saturation magnetization ( $M_s$ ), saturation remanent magnetization ( $M_{rs}$ ), and coercive force ( $B_c$ ) (Table 6 and 7). Acquisition of an isothermal remanent magnetization (IRM) and subsequent back-field DC remagnetization (both in a succession of fields up to 1 T), were also carried out on the same specimens (Figure 86, 87, 88). Data were also used to compute the coercivity of remanence ( $B_{cr}$ ). Furthermore, for lavas specimens, we also measured the variation of the low-field magnetic susceptibility during a heating and cooling cycle performed in air, from room temperature up to 700°C. The Curie point of the magnetic minerals present in the samples was determined from the thermomagnetic curves as the temperature, or range of temperatures, at which paramagnetic behavior starts to dominate.

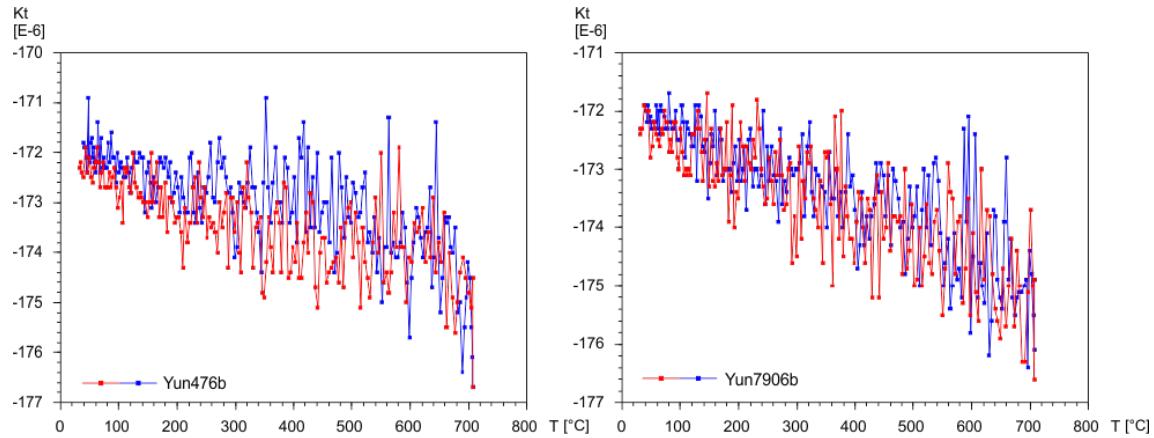
The magnetic mineralogy of the volcanic samples is heterogeneous (Figure 82), ranging from almost paramagnetic (whitish siltstones sites Yun05 to 09 and red beds site Yun32), to prevailing low coercivity and saturated at low field properties (Pliocene-Holocene basalt sites Yun04, 11 to 16, 18 to 23, and red beds sites 33 to 34), up to high coercivity and/or markedly unsaturated features at 1 T (Jurassic basalt sites Yun01 and 02, red beds sites Yun03, 24 to 31, 35 to 50) [Pellegrino *et al.*, 2018].

The main ubiquitous magnetic minerals are magnetite/titanomagnetite in the volcanics and hematite in the red beds; the magnetic susceptibility vs. temperature curves show that hematite is also present in the low coercivity class of samples, considering that k does not completely drop before 680°C (Figure 82). In this case, the presence of hematite in the thermomagnetic curves can also be ascribed to the inversion of ferromagnetic maghemite occurring from  $T>350^\circ$ . For most of the samples, cooling curves are below the heating curves, thus excluding the relevant formation of ferrimagnetic minerals while heating in air [Pellegrino *et al.*, 2018].



**Figure 82.** Thermomagnetic curves for basalt specimens; data are raw and not corrected for the empty furnace.

Different situation occurred in red beds samples in which the heating and cooling curves are not well defined but note a trend of decrease of susceptibility with the increasing temperature.



**Figure 83.** Thermomagnetic curves for representative red beds specimens (Yun47 and Yun79).

The “Day plot” has been drawn only for the low coercivity/saturated samples (Table 6), in order to define the domain state and the consequent magnetic grainsize according to the theoretical hysteresis ratios expected for magnetite and titanomagnetite (Figure 84b).

**Table 6.** Hysteresis Parameters of samples relative to "Day plot" shown in Figure 84b.

Samples	Mass (g)	Mass (kg) (=[Mass]/1000)	Ms (Am <sup>2</sup> )	Mass_Ms (Am <sup>2</sup> /kg) (=[Ms]/[Mass])	Mrs (Am <sup>2</sup> )	Mass_Mrs (Am <sup>2</sup> /kg) (=[Mrs]/[Mass])	Bc (T)	Bcr (T)	Bcr/Bc	Mrs/Ms
04-04a	0.16	0.00016	8.03473E-05	0.502170375	1.15606E-05	0.072253813	0.01230367	0.03086621	2.51	0.14
11-04a	0.26	0.00026	0.000108264	0.416401538	1.81067E-05	0.069641038	0.009290813	0.01977991	2.13	0.17
12-04a	0.22	0.00022	5.98083E-05	0.271855818	1.22364E-05	0.055620045	0.01653307	0.03403819	2.06	0.20
13-04a	0.25	0.00025	0.0000762	0.3048	1.43589E-05	0.05743556	0.006049184	0.01413916	2.34	0.19
15-04a	0.25	0.00025	0.000167472	0.6698868	2.73662E-05	0.1094648	0.01158414	0.02547041	2.20	0.16
16-04a	0.28	0.00028	7.89221E-05	0.281864786	1.74328E-05	0.062260036	0.0192496	0.04005314	2.08	0.22
18-04a	0.19	0.00019	0.000112137	0.590195263	1.21231E-05	0.063806	0.004727123	0.01557542	3.29	0.11
19-04a	0.25	0.00025	5.79663E-05	0.2318652	1.49559E-05	0.05982364	0.01253398	0.01840643	1.47	0.26
20-04a	0.25	0.00025	8.31621E-05	0.33264832	1.81376E-05	0.07255028	0.01471912	0.01878878	1.28	0.22
21-04a	0.21	0.00021	0.000143387	0.682796667	2.10463E-05	0.100220524	0.01352794	0.03143721	2.32	0.15
22-04a	0.24	0.00024	0.000163223	0.680096667	2.03443E-05	0.084767875	0.0116234	0.02649358	2.28	0.12
23-04a	0.26	0.00026	0.000143302	0.551159615	6.38942E-05	0.245746923	0.03123689	0.04238027	1.36	0.45
33-04c	0.21	0.00021	4.44939E-07	0.002118755	1.06661E-07	0.000507911	0.01753115	0.05317518	3.03	0.24
34-03b	0.23	0.00023	3.34486E-07	0.001454289	6.64024E-08	0.000288706	0.02130191	0.06908135	3.24	0.20

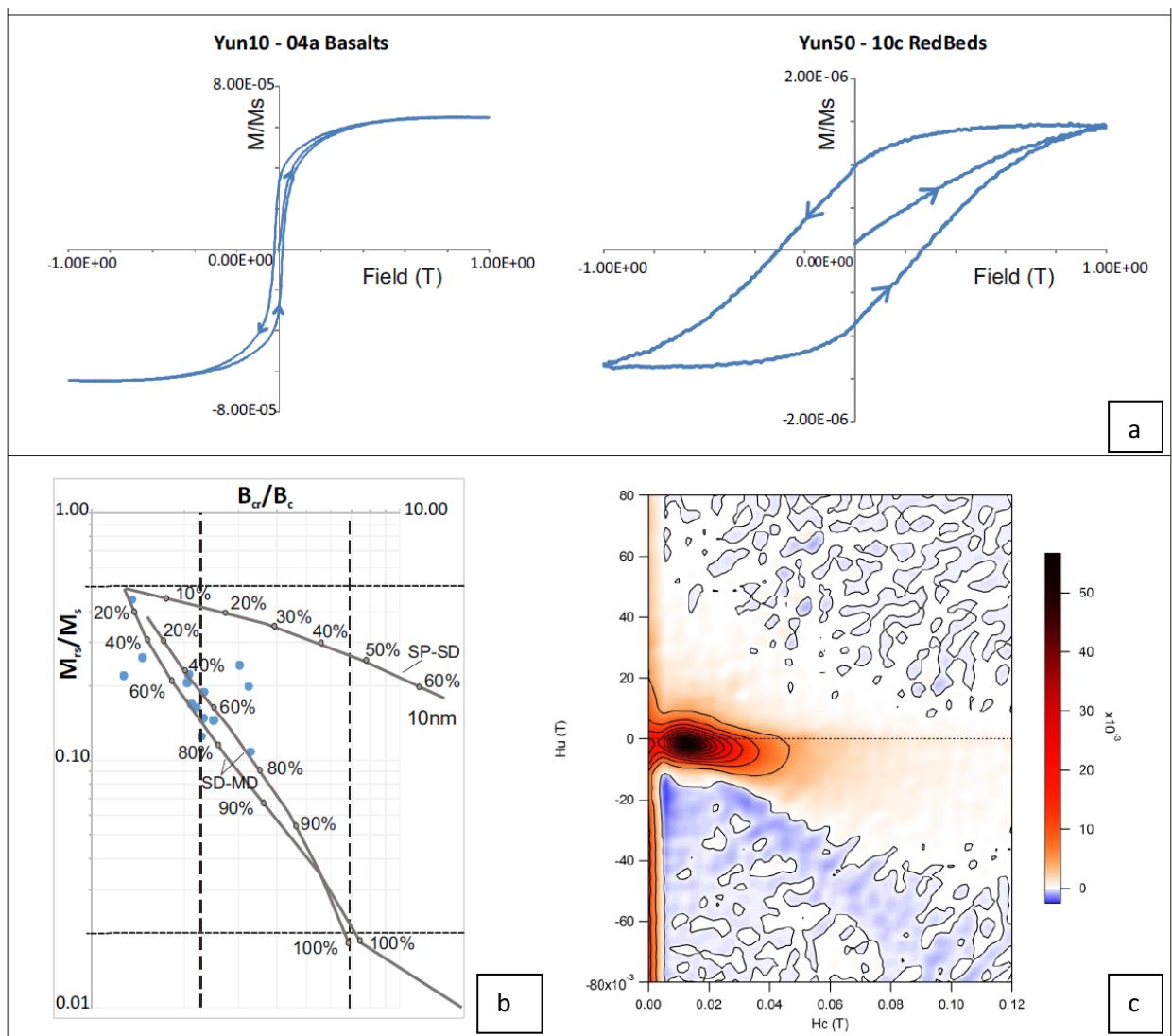
Most of these samples fall in the upper-central region of the plot, not far from the upper segment of the theoretical curves calculated for mixtures of single domain (SD) and multidomain (MD) magnetite grains, thus suggesting a significant contribution of stable SD particles to the overall magnetic mineralogy, according to *Zhao et al.* [2015; *Day et al.*, 1977; *Dunlop*, 2002].

Noteworthy, selected samples from volcanic sites Yun10, 14 and 17 show wasp-waisted hysteresis loops, which can be ascribed both to the coexistence of superparamagnetic (SP) and single-domain (SD) particles or to the mixing of hard and soft coercivity magnetic minerals, as magnetite and hematite [Table 7 and Figure 84a; *Roberts et al.*, 1995].

**Table 7.** Hysteresis Parameters of samples from Gaoligong Shear Zone.

Samples	Mass (g)	Mass (kg) (=[Mass]/1000)	Ms (Am <sup>2</sup> )	Mass_Ms (Am <sup>2</sup> /kg) (=[Ms]/[Mass])	Mrs (Am <sup>2</sup> )	Mass_Mrs (Am <sup>2</sup> /kg) (=[Mrs]/[Mass])	Bc (T)	Bcr (T)	Bcr/Bc	Mrs/Ms
01-03b	0.27	0.00027	1.39E-05	5.13E-02	3.29E-06	1.22E-02	1.85E-02	5.07E-02	2.74	0.24
02-08b	0.29	0.00029	4.31E-06	1.49E-02	1.69E-06	5.81E-03	3.07E-02	1.04E-01	3.38	0.39
05-01a	0.21	0.00021	8.50E-08	4.05E-04	1.03E-08	4.93E-05	8.12E-03	4.52E-02	5.57	0.12
10-04a	0.25	0.00025	6.47E-05	2.59E-01	3.10E-05	1.24E-01	1.92E-02	2.60E-02	1.36	0.48
14-04a	0.21	0.00021	6.81E-05	3.24E-01	2.96E-05	1.41E-01	2.15E-02	3.47E-02	1.61	0.43
17-04a	0.25	0.00025	8.02E-05	3.21E-01	4.23E-05	1.69E-01	3.71E-02	4.04E-02	1.09	0.53
24-08b	0.24	0.00024	1.50E-06	6.23E-03	1.06E-06	4.43E-03	3.71E-01	4.98E-01	1.34	0.71
25-04c	0.26	0.00026	5.43E-07	2.09E-03	3.56E-07	1.37E-03	2.62E-01	4.28E-01	1.63	0.66
26-04c	0.25	0.00025	8.47E-07	3.39E-03	4.96E-07	1.98E-03	1.83E-01	3.73E-01	2.04	0.59
27-08b	0.25	0.00025	6.61E-07	2.64E-03	3.38E-07	1.35E-03	1.40E-01	4.33E-01	3.08	0.51
28-05a	0.2	0.0002	7.10E-07	3.55E-03	3.05E-07	1.52E-03	5.66E-02	3.98E-01	7.04	0.43
29-02c	0.27	0.00027	6.48E-07	2.40E-03	3.80E-07	1.41E-03	2.13E-01	3.30E-01	1.55	0.59
30-10b	0.25	0.00025	8.64E-07	3.46E-03	4.44E-07	1.77E-03	1.41E-01	4.06E-01	2.88	0.51
31-06c	0.19	0.00019	1.17E-06	6.17E-03	8.72E-07	4.59E-03	4.28E-01	5.46E-01	1.28	0.74
32-01c	0.22	0.00022	2.11E-07	9.58E-04	1.45E-08	6.60E-05	7.36E-03	3.65E-02	4.95	0.07
35-02b	0.2	0.0002	5.67E-07	2.83E-03	2.91E-07	1.46E-03	1.73E-01	3.28E-01	1.89	0.51
36-07c	0.27	0.00027	1.04E-06	3.86E-03	7.06E-07	2.61E-03	3.39E-01	5.03E-01	1.48	0.68
37-03c	0.29	0.00029	1.01E-06	3.49E-03	6.17E-07	2.13E-03	2.64E-01	5.10E-01	1.93	0.61
38-04c	0.26	0.00026	9.77E-07	3.76E-03	5.13E-07	1.97E-03	2.01E-01	5.12E-01	2.54	0.53
39-04b	0.24	0.00024	1.40E-06	5.82E-03	8.81E-07	3.67E-03	3.02E-01	4.85E-01	1.61	0.63
40-05c	0.22	0.00022	7.10E-07	3.23E-03	4.65E-07	2.11E-03	2.90E-01	4.88E-01	1.68	0.65
41-04c	0.26	0.00026	4.75E-07	1.83E-03	3.17E-07	1.22E-03	2.97E-01	4.53E-01	1.53	0.67
42-03b	0.27	0.00027	3.07E-06	1.14E-02	1.97E-06	7.31E-03	2.42E-01	3.97E-01	1.64	0.64
43-05c	0.24	0.00024	1.16E-06	4.84E-03	6.87E-07	2.86E-03	1.71E-01	2.75E-01	1.61	0.59
44-10c	0.24	0.00024	1.15E-06	4.78E-03	6.28E-07	2.62E-03	2.02E-01	4.46E-01	2.21	0.55
45-03c	0.22	0.00022	8.27E-07	3.76E-03	4.07E-07	1.85E-03	1.94E-01	4.47E-01	2.30	0.49
46-09c	0.25	0.00025	5.59E-07	2.24E-03	2.90E-07	1.16E-03	1.73E-01	4.82E-01	2.78	0.52
47-08b	0.23	0.00023	8.25E-07	3.58E-03	5.39E-07	2.34E-03	3.32E-01	5.03E-01	1.52	0.65
48-05c	0.25	0.00025	5.74E-07	2.29E-03	3.69E-07	1.48E-03	3.30E-01	5.37E-01	1.63	0.64
49-04c	0.28	0.00028	1.02E-06	3.64E-03	6.87E-07	2.46E-03	4.13E-01	5.75E-01	1.39	0.67
50-10c	0.22	0.00022	1.40E-06	6.38E-03	9.16E-07	4.16E-03	2.85E-01	4.35E-01	1.53	0.65

The FORC diagram (Figure 84c) of a selected sample from site Yun10 shows closed contours peaked at about 15-20 mT, with a concurring secondary peak at the origin and asymmetric nearly vertical contours extending into the lower quadrant [*Pike et al.*, 2001], thus indicating that the wasp-waisted shape of the hysteresis loop, in this case, is due to the coexistence of ultrafine SP and fine SD magnetite particles.



**Figure 84.** a) Representative wasp-waisted and high-coercivity/unsaturated hysteresis loops; b) “Day plot” for the low coercivity/saturated (up to 1 T applied magnetic field) set of samples; the SD, PSD and MD fields and the theoretical mixing lines for SD-MD and SP-SD grains are from *Dunlop* [2002] and refer to magnetite; c) FORC diagram of a specimen with wasp-waisted hysteresis features; the smoothing factor SF = 3. Modified after *Pellegrino et al.* [2018].

Red beds sediments are widely distributed in tropical and subtropical areas [Robb, 1949; Walker, 1967; Van Houten, 1968, 1973; Collinson, 1974] and have been extensively used for paleomagnetic studies [Kent *et al.*, 1986; Tan *et al.*, 2003; Iosifidi *et al.*, 2010]. Often, these sediments, are re-magnetized and when it occurs, the interpretation of paleomagnetic, and magnetic mineralogy data becomes complicated [Collinson, 1965; Kent *et al.*, 1987; McCabe and Elmore, 1989; Suk *et al.*, 1993; Wang and Vander Voo, 1993; Lu *et al.*, 1994; Piper *et al.*, 1999; Kodama and Dekkers, 2004; Roberts and Weaver, 2005; Elmore *et al.*, 2006; Rowan and Roberts, 2006; Deng *et al.*, 2007; Jin and Liu, 2011; Liu *et al.*, 2011; Vander Voo and Torsvik, 2012; Roberts, 2015]. Remagnetization may happen with different mechanisms: 1) mineral transformations associated with redox process, i.e. magnetite, greigite, or pyrrhotite formation [Katz *et al.*, 2000; Weaver *et al.*, 2002; Roberts and Weaver, 2005], 2) deformation-associated to fluid migration and/or pressure solution [McCabe and Elmore, 1989; Elmore and McCabe, 1991; Housen *et al.*, 1993], 3) chemical weathering in moist, tropical climates, and soil-forming environments [Creer, 1961, 1968; Jin and Liu, 2011; Liu *et al.*, 2011], or 4) thermoviscous remanent magnetization acquisition [Scotesse *et al.*, 1982; Kent, 1985; Hashimoto *et al.*, 2008; Jiang *et al.*, 2017]. The natural remanent magnetization (NRM), in fact, in most red beds is a composite of detrital and authigenic hematite [Collinson, 1966, 1974; Tauxe *et al.*, 1980].

In South China, several paleomagnetic investigations were conducted on the red beds, especially for regional paleogeographic reconstructions [e.g. Achache *et al.*, 1984; Zaman and Torii, 1999; Steiner and Lucas, 2000; Sun *et al.*, 2006]. Recently, data reported by Jiang *et al.*, [2017], show that two remanence components carried by secondary and primary hematite are isolated with thermal demagnetization. CRM carried by secondary hematite was interpreted, generated by magnetite oxidation during tectonic activity. Therefore, the secondary magnetization with steep inclination, commonly observed in Triassic samples of the South China Block, was interpreted to have been influenced by a combination of the

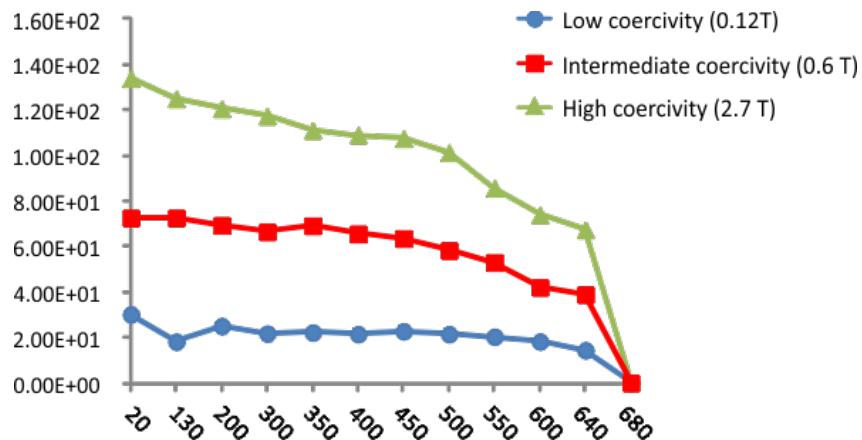
remanence carried by parent magnetite, the orogenic stress field, and the contemporaneous geomagnetic field direction during deformation [Jiang *et al.*, 2017].

Liu *et al.*, [2011] investigated the remagnetization mechanism in red beds from Lower Triassic sandstones in Yunnan Province. They put forward a remagnetization process summarized as follow: initially, the relatively coarse-grained haematite and magnetite of detrital origin were deposited during the early Triassic. These two minerals carried the primary detrital remanent magnetization (DRM). During the Quaternary–Palaeogene period, the region went through a fluid event that accelerated the neoformation of fine-grained haematite and maghemite with variable unblocking temperature spectra. These newly formed magnetic minerals, particularly the haematite, carry the secondary CRM, which fully or partially contaminated the primary DRM. This region was uplifted at the same time as the Tibetan plateau to the northwest as the result of the India–Eurasia collision during the early Palaeogene [e.g. Wang *et al.*, 1998; Najman *et al.*, 2010]. This uplift of the Tibetan plateau could have been responsible for the fluid migration that instigated the formation of the authigenic maghemite and haematite [Liu *et al.*, 2011].

A progressive acquisition of isothermal remanent magnetization (IRM) up to a maximum field of 2.7 T was performed on my red beds, samples along the Red River faults, using a 2G Enterprises 2G 660 Pulse Magnetizer.

The acquisition of **IRM** curves, through the laboratory application of stepwise-increasing uniaxial fields to a rock-magnetic sample, provides an important non-destructive tool for the investigation of coercivity spectra [Dunlop and O’zdemir, 1997]. A ferromagnetic substance subject to the action of an external magnetic field retains a remanent magnetization even when the field is removed. The coexistence of different phases can be highlighted with the procedure proposed by Lowrie [1990]. The samples are magnetized at three decreasing field values in three orthogonal directions  $H_z > H_y > H_x$ . The three components of the IRM are thus

carried by grains with high ( $J_z$ ), intermediate ( $J_y$ ) and low-coercivity ( $J_x$ ). For a more complete interpretation, IRM acquisition must be combined with subsequent thermal demagnetization of the IRM gradually removes the three components and provides indications on their blocking temperature spectrum and Curie point (Figure 85).



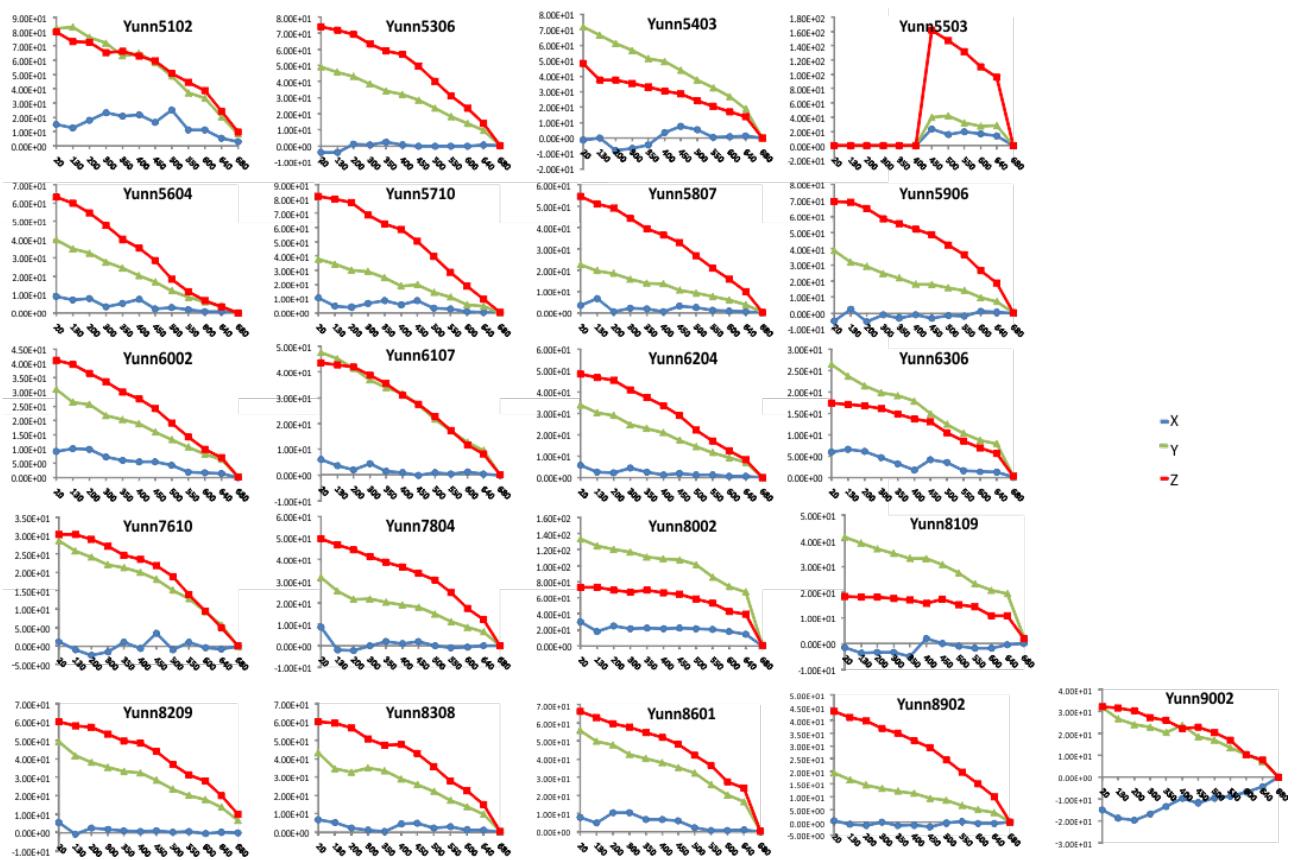
**Figure 85.** Example of thermal demagnetization of three-component isothermal remanent magnetization (IRM) curves, according to the *Lowrie* [1990] method. T (°C) in x axis from 20° to 680°; Intensity of magnetization in y axis. Green, red and blue lines represents respectively Y, Z and X axis of coercivity.

A thermal demagnetization of the composite IRMs (imparted at 2.7 T, 0.6 T and 0.12 T along z, y and x axes of the specimens, respectively) was conducted for samples outcropping along the Ailaoshan Red River Fault Zone, to identify the unblocking temperature spectra of the representative samples [*Lowrie*, 1990].

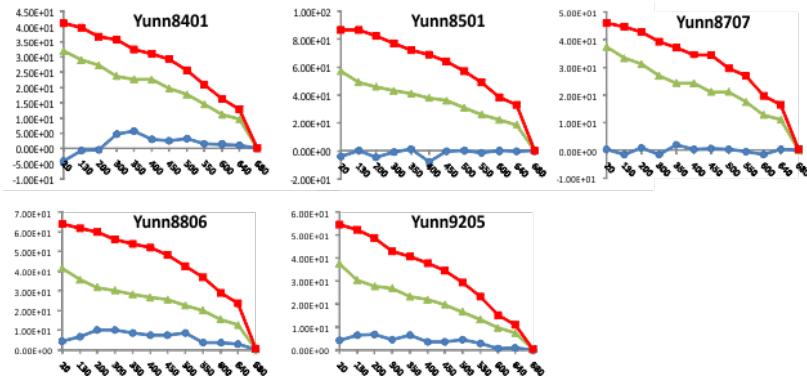
Thermal demagnetization curves of the three-component IRM show that the intensity of all the three components decrease quickly above 650°C and unblock up to 680°C, suggesting the dominance of high-coercivity hematite (the Néel temperature for hematite is approximately 685 °C). This is confirmed also by the results of hysteresis loops. All of these results indicate that hematite is the main magnetic remanence carrier in the sediments.

As shown in Figures 86,87,88 in which the data are clustered respectively in pre-, syn- and post-tilting (indicating post, syn and pre-tilting overprint, respectively – see table 3 for more details), according to our paleomagnetic results (see Table 3), a clear mineralogical fingerprint is evident in almost all analyzed samples. The low coercivity component is mostly absent and the medium and high magnetization components essentially prevail. The shape of prevailing convex decay of high-coercivity component (carried by hematite) is similar for all sediments. Thus such magnetic mineralogy experiments are unfortunately not useful to identify sites undergoing (or not undergoing) magnetic overprint, and sites remagnetized before, during or after strata tilting.

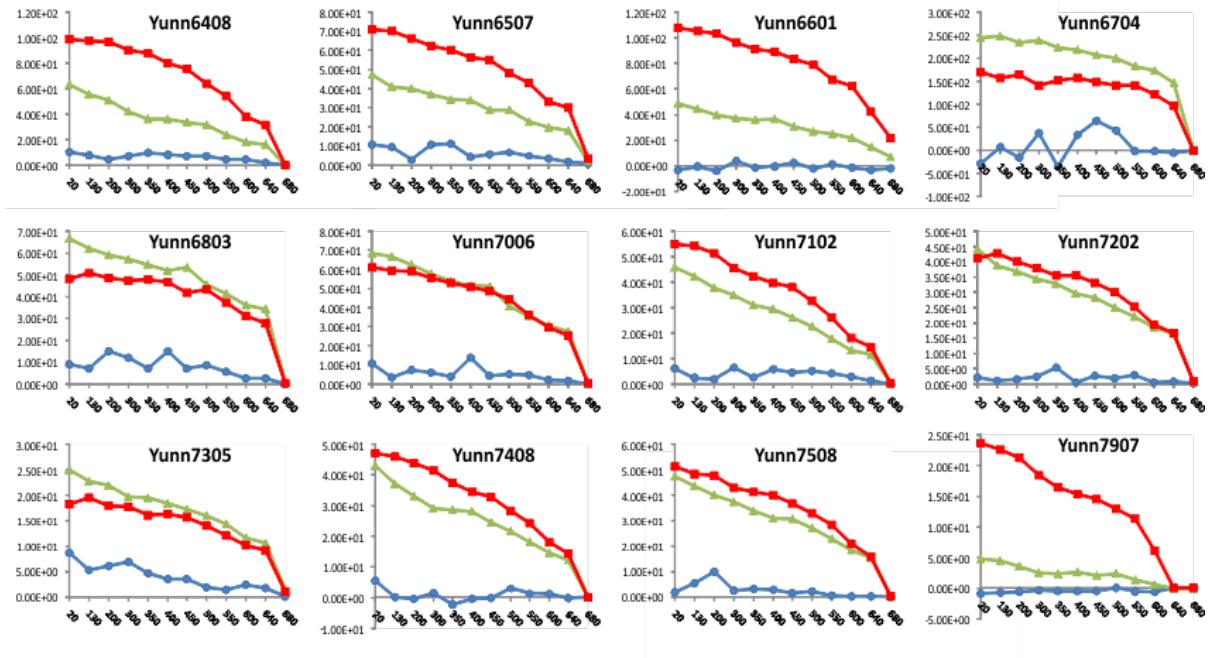
In agreement to *Jiang et al.* [2015], my data suggests that two components of intermediate and high coercivity could indicate a double magnetization case: ChRM/MT and HT components discriminate, respectively, a first chemical (gradual decay between ~200 and 640°C) and then detritic remagnetization (most decay in the 600–680°C range). *Jiang et al.*, [2017] interpreted the CRM to have been generated by magnetite oxidation during tectonic activity. As said, is valid for each step of magnetization respect to tilting events (Figure 86,87,88), and its clearly shown from different sites (Yun80, 85, 88, 65, 70, 75).



**Figure 86.** IRM curves relative to sites with pre-tilting magnetization.



**Figure 87.** IRM curves relative to sites with syn-tilting magnetization.



**Figure 88.** IRM curves relative to sites with post-tilting magnetization.

*Chapter*  
**X**

## 10. CONCLUSIONS

The conclusions of this thesis can be summarized as follows:

- According to *Pellegrino et al.* [2018], the paleomagnetic study of Pliocene-Holocene volcanics and Mesozoic red beds exposed east and west (respectively) of the Gaoligong Shear Zone show that low Ti-magnetite and hematite dominate the respective magnetic mineralogy. Three bits of evidence suggest that the magnetization in the red beds is here of primary origin (detrital hematite), and/or acquired during early diagenesis by chemical hematite growth: 1) the data support a positive fold test at the 99% significance level, indicating that magnetization acquisition is pre-tilting (i.e. pre-Eo-Oligocene) in age; 2) although the majority of the sites yield a normal polarity, three sites show a reverse polarity, and alternating normal and reverse polarity beds were documented in an additional site; 3) a statistically significant positive inclination flattening of the whole red bed data set proves that magnetization acquisition predate diagenesis. Thus we speculate that – assuming a primary magnetization origin- the majority of our sites were deposited during the long normal Cretaceous superchron. Similar paleomagnetic evidence was previously documented at several other red beds sites from Yunnan (e.g. *Li et al.*, 2017b and references therein), and the final resolution of this problem would lie in the accurate site age determination, indeed a problematic issue given the continental nature and the coarse grain size of the studied sediments.
- The clockwise rotations pattern observed in Mesozoic red beds exposed east of the Gaoligong Shear Zone (Baoshan block) points to a quasi-continuous crust kinematics model, characterized by rigid blocks of smaller size ( $\leq 1$  km close to the fault) than the fault deformation zone itself and by rotation values  $>90^\circ$  close to strike-slip faults. CW rotations close to the fault reach  $176^\circ$ , and decrease moving

east up to be virtually annulled at ca. 20 km east from the contact with mylonites exposed along the shear zone. As discussed in chapter 1, block rotations adjacent to shear zones occur in response to the angular velocity of the ductile deformation, taking place at great depth in the lower crust [Beck, 1976; McKenzie and Jackson, 1983; Lamb, 1987; Nelson and Jones, 1987; Salyards et al., 1992; Sonder et al., 1994; Piper et al., 1997; Randall et al., 2011]. We cannot exclude that the same behavior is reflected even at greater depths beyond the crust, but more data (eg. deep seismic data, ect..) would be necessary to prove it. Paleomagnetism is restricted to the examination of the upper crust and does not provide a deep crust/lithospheric depth evaluation.

- By using the equation of Lamb [1987], which relates rotations to fault displacement in a quasi-continuous crust deformation model, we find a total Gaoligong fault dextral offset ranging between 230 and 290 km, which is consistent with the 600 km total fault length, if offset is related to mega-block extrusion due to India-Asia collision [Wang and Burchfiel, 1997; Zhang et al., 2012a]. Three sites sampled adjacent to the fault yielded scattered rotations (from 74° CCW to 116° CW) that may be interpreted as huge CW rotations exceeding 180°. This suggests that a sort of “rotation channel” exists adjacent to the fault, where small block rotate freely by large magnitude, conforming to the “ball bearing” model of Beck [1976].
- Four sites sampled at greatest distance (15-20 km) from the contact with Gaoligong Shear Zone mylonites yield a mean CW rotation of  $20.5^\circ \pm 12.2^\circ$ . Although the trend of rotation values vs. fault distance is suggestive of a complete rotation cessation at ca. 25 km from mylonite contact, a 20° rigid CW rotation of the Baoshan block cannot be excluded by our data set. A 20° Baoshan block rotation indeed does not compare with the 70°-80° CW rotations gathered by Tong et al.

[2016] from two additional localities from the centre of the block itself. A 20° rigid Baoshan block rotation would also reduce to 145° the maximum CW rotation due to Gaoligong fault shear, and this in turn would change to 200-250 km the paleomagnetically-evaluated dextral fault offset. However, the discrepancy between our data and those by *Tong et al.* [2016] strongly suggests that the Baoshan block is broken in additional sub-blocks undergoing independent rotations.

- Pliocene-Holocene volcanic sites (and one Pliocene sedimentary site) located west of the fault (Tengchong block) do not rotate, implying that the Gaoligong dextral Shear Zone yielding CW rotations has been active during the Oligo-Miocene time span (early Oligocene Gaoligong Shear Zone activity onset according to *Wang et al.*, 2006), but not over the last 5 Ma. This is in good agreement with the focal mechanism of the 1976 M=7.4 earthquake, showing a sinistral shear sense along the southern end of the Gaoligong Shear Zone, and prove that the present-day CW rotation of SE Tibet and northern Yunnan shown GPS data [*Liang et al.*, 2013] has nothing to do with significantly older paleomagnetic rotation. The geomagnetic field PSV is not averaged out in our volcanic paleomagnetic data, thus possibly implying a few degree bias on average null rotation value that is however much smaller than the large (up to 176°) rotations observed in Mesozoic red beds.
- The data collected along the Gaoligong fault, show for the first time, which fault shear may represent a significant tectonics contributing to the CW rotations frequently documented along the Indochina blocks [e.g. *Li et al.*, 2017a; *Tong et al.*, 2016]. However, CW rotations were also documented far (i.e. >20 km) from the strike-slip faults and within the rough centre of the blocks themselves, so that it is possible that some blocks - at least the south Indochina, Simao and Lanping

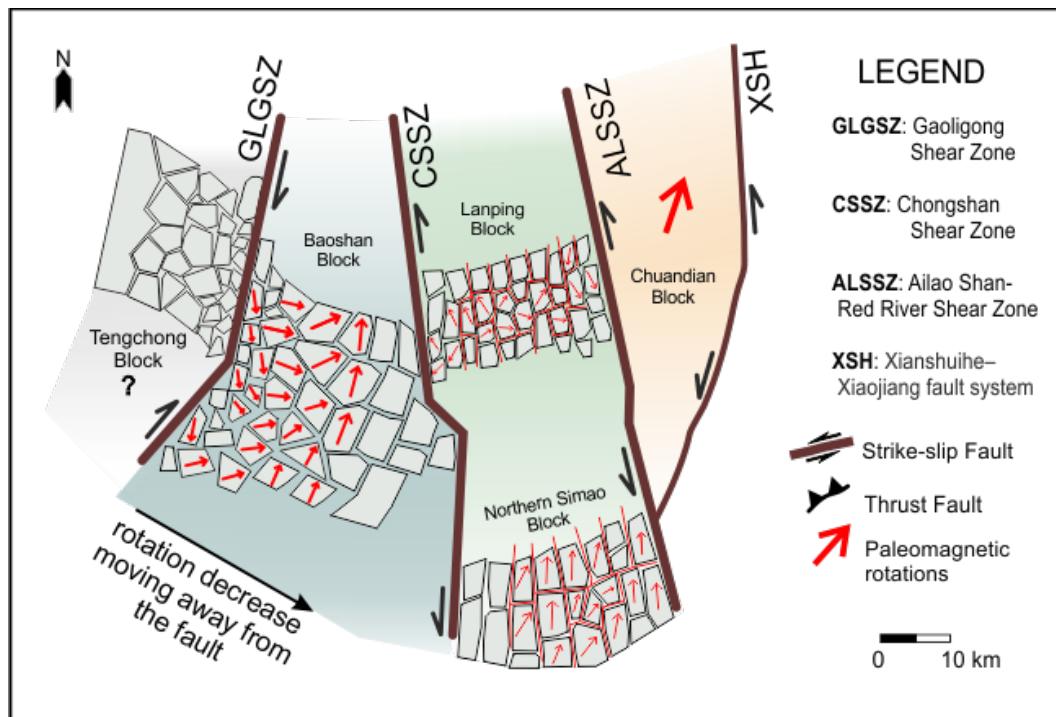
terranes- underwent also semi-rigid rotations. Yet, the tectonics and dynamics of such block rotations (whether a sort of “ball-bearing” or “domino” model) have not been properly elucidated at presently, and it is unclear why predominantly CW rotations occur. Furthermore, several works [*Chen et al.*, 1995; *Sato et al.*, 2001, 2007; *Tanaka et al.*, 2008; *Otofuji et al.*, 2012] seem to support internal deformation (and variable rotations) in the Simao domain, so that no rigid mega-block rotation would exist. *Kondo et al.* [2012], *Tong et al.* [2013, 2017], and *Gao et al.* [2015] even report data suggesting a folding-related oroclinal bending within the Simao domain itself, which again would be at odds with the rigid “microplate” rotation model proposed by *Li et al.* [2017a].

- Trying to solve this question, Triassic-Cretaceous sites were sampled, at both sides of the Ailao-Shan Red River Shear Zone (ARRSZ), within the Chuandian, Lanping and Northern Simao domains. The sites yielded measurable and stable magnetization components, but magnetization acquisition timing was different in the three blocks. Sites from the Chuandian domain show a normal polarity and were remagnetized after folding that probably occurred from Late Eocene to Early Oligocene (about 33 Ma ago) [*Gao et al.*, 2017]. The tectonic rotations calculated by us and previous authors within this block, are mostly CW and do not exceed 20°. Thus, we have no evidence of fragmentation of the Chuandian domain.

In the northern Simao domain the magnetization was acquired before folding event, but the ubiquitous normal polarity in Jurassic-Cretaceous sites suggests a pre-folding magnetic overprint. Unlike to the Chuandian block, here data do not support a rigid block rotation, but suggest that the northern Simao domain is made of small (few km of size) sub-blocks rotating CW, separated by non-rotating

domains of similar size, according to previous authors [Figure 81, *Chen et al.*, 1995; *Sato et al.*, 2007].

Finally, a high-temperature magnetization component (640-680°C) suggests a similar rotational behaviour (CW-rotating and non-rotating sub-blocks) in the centre of the Lanping domain. Conversely, a 300-640°C component was later acquired at 28% unfolding and subsequently underwent no rotation. The folding events took place in Paleogene, resulting in north-south trending fold axes in Mesozoic system [*Funahara et al.*, 1993]. The sites close to the ARRSZ yield great rotations on the order of 180° that confirm past occurrence of significant left strike-slip shear along the ARRSZ itself. Conversely, sites located at 10-15 km distance from the Chongshan Shear Zone show ca. 90° CCW rotations that imply also a left-lateral shear along the fault zone, consistently with recent geological evidence.



**Figure 89.** Cartoon explaining the blocks rotational model in Yunnan (China) in Oligocene-Miocene age, along the major strike-slip fault zones and inside the block itself.

- Magnetic anisotropy data confirm the different behaviour of the crustal blocks, in particular, in the investigated areas, we highlight the presence of foliated planes oriented NNE-SSO, approximately parallel to the orientation of the main structures. Along the Gaoligong fault zone the direction of the lineation is often strikingly sub-vertical, a setting uncommon in fold and thrust belts. Instead, along the Red River fault zone, the direction of the foliation planes is predominantly N-S. This implies a compression oriented mainly E-W probably responsible of the strong deformation inside the blocks.
- Hematite is the dominant magnetic carrier in the Jurassic-Cretaceous red beds and it also plays a fundamental role in the magnetic anisotropy fabric setting (typically normal fabric) of these sediments.
- The results of this PhD work along with previous paleomagnetic data, indicate that crust deformation of the Yunnan is extremely complex and still to be completely elucidated. The “blocks” certainly underwent a strong internal deformation and were fragmented in smaller independent sub-blocks whose kinematics and tectonics are still a matter of speculation. Additional structural and paleomagnetic data are necessary to fully constrain the kinematics and tectonics of the small size blocks. In any case, this work shows that no “big” (hundred of km wide) rigid crustal blocks exist between Tibet and Indochina.

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## ACKNOWLEDGEMENTS/ RINGRAZIAMENTI

*I'm grateful to Prof. Zhang, Congyuan Yin, Siyu Chen (Peking University) and Mr. Mu for their kindness and for safely driving me in the whole Yunnan during my field campaigns.*

Un ringraziamento particolare va alla *Prof.ssa Rosanna Maniscalco*, mia relatrice di tesi di laurea triennale, laurea magistrale e dottorato di ricerca. Un percorso lunghissimo, che ho scelto di percorrere sempre al suo fianco, e che rifarei altre mille volte! Da sempre ha creduto in me, dimostrandomi disponibilità e attenzioni costanti che mi hanno consentito non solo di raggiungere questo traguardo ma anche di fare innumerevoli esperienze formative e umane;

Al *Prof. Fabio Speranza*, mio correlatore, per la pazienza e l'ospitalità dimostratami durante i lunghi soggiorni presso l'INGV di Roma, ma soprattutto per l'incoraggiamento e i preziosi consigli necessari “per diventare un buon Ricercatore”;

Entrambi hanno sempre cercato di trasmettermi il loro sapere e questo non può che essere la base per definirli dei bravi Docenti. Se sono cresciuta professionalmente è solo grazie a loro .... a loro va la mia infinita gratitudine ed il mio affetto!

A tutti i componenti del gruppo di paleomagnetismo dell'INGV di Roma, in particolare Aldo Winkler e Patrizia Macrì, per il tempo trascorso insieme al “fresco” in laboratorio e per il supporto tecnico/morale durante i lunghi cicli di misure.

Ai revisori della mia Tesi di Dottorato, Prof. Massimo Mattei e Prof. Rodolfo Carosi, per gli attenti commenti e per i preziosi consigli che hanno contribuito certamente a valorizzare e migliorare il lavoro finale.

Al coordinatore Prof.ssa Agata Di Stefano, per il costante supporto e per essere sempre stata attiva nella formazione e nella crescita di tutti i dottorandi.

Ai miei amici e colleghi (vecchi e nuovi), ed a Catalina Hernandez-Moreno, per aver condiviso con me i momenti più belli e difficili di questi tre anni.

E poi non finirò mai di ringraziare mio padre, mia madre, i miei fratelli e Mirko, mio più grande amore e compagno di vita, che mi hanno sempre sostenuto in tutte le scelte che ho fatto, mi sono sempre stati vicino e mi hanno “sopportato” dandomi ogni giorno un motivo in più per andare avanti.

*Università degli Studi di Catania, Catania, Italy*

Supporting Information for

**UNDERSTANDING BLOCK ROTATION ALONG  
STRIKE-SLIP FAULT ZONES IN YUNNAN (CHINA):  
PALEOMAGNETIC AND STRUCTURAL APPROACH**

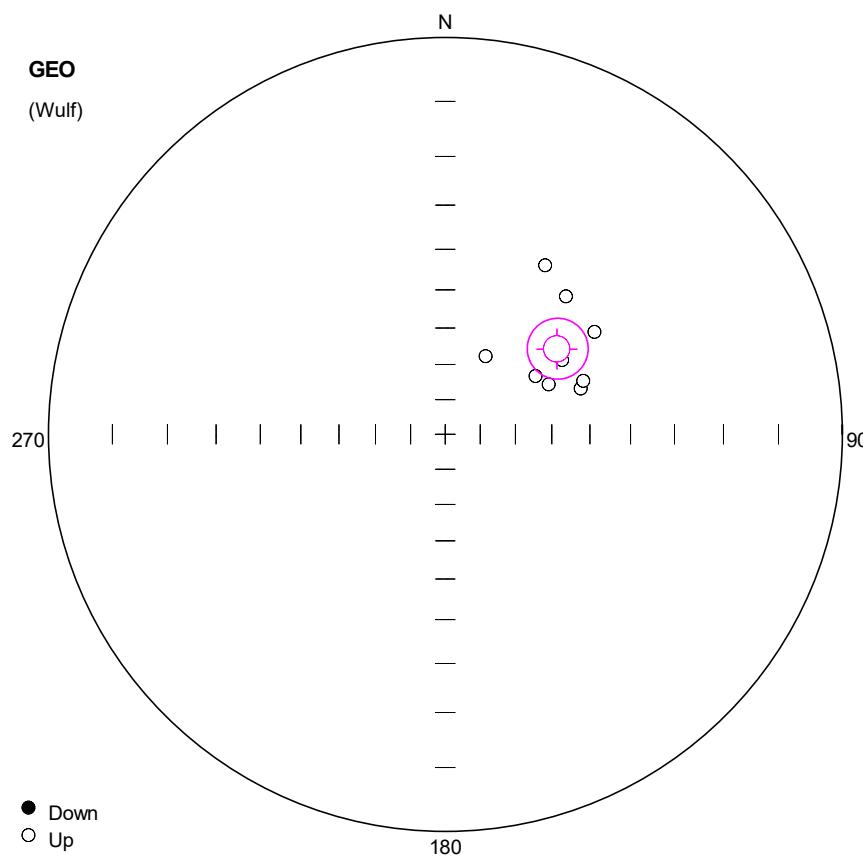
Ph.D. Thesis

***Dott.ssa Alessandra Giovanna Pellegrino***

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**Contents of this file : Caption for Data Set**

- Equal angle projections (Wulf), both sample mean (ChRM/ High-Temperature Component) and site mean paleomagnetic directions, gathered by us along **a)** Gaoligong shear zone and **b)** Red River shear zone. Geographic or In-situ (GEO) and Tilt correction coordinate system (TILT CORR). The magnetization components were identified by principal component analysis [Kirschvink, 1980], and the site mean paleomagnetic directions were computed using Fisher [1953] statistics. Software: Remasoft 3.0 [Chadima and Hrouda, 2007].
- Equal area projections on lower hemisphere of the Anisotropy of magnetic susceptibility (AMS) directions for sites gathered by us along **c)** Gaoligong shear zone and **d)** Red River shear zone. The squares, triangle, and dots represent  $K_{\max}$ ,  $K_{\text{int}}$  and  $k_{\min}$  respectively. The ellipses indicate the 95% region around the principal susceptibility axes. The orange line (open circle) indicates the bedding planes (poles) (expressed in dip azimuth/dip values). Software: Anisoft 4.2 [Chadima and Jelinek, 2009].
- **e)** Wasp-waisted hysteresis loops from basalts and red beds sites sampled along Gaoligong shear zone. The hysteresis parameters for these samples are shown in Table 6 and 7.



Number of data points: 10

#### Fisher statistics

Mean vector: 52.59 / -51.02

Resultant vector: 9.77

(X= 3.73, Y= 4.88, Z= -7.6)

Estimated precision, k: 39.4

95% Confidence limit: 7.79

95% Confidence limit, approximation: 7.05

#### Orientation matrix

1st eigenvalue: 0.96

1st eigenvector: 232.82 / 51.02

2nd eigenvalue: 0.03

2nd eigenvector: 353.26 / 22.29

3rd eigenvalue: 0.02

3rd eigenvector: 97.01 / 30.12

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -14.95 (14.95)

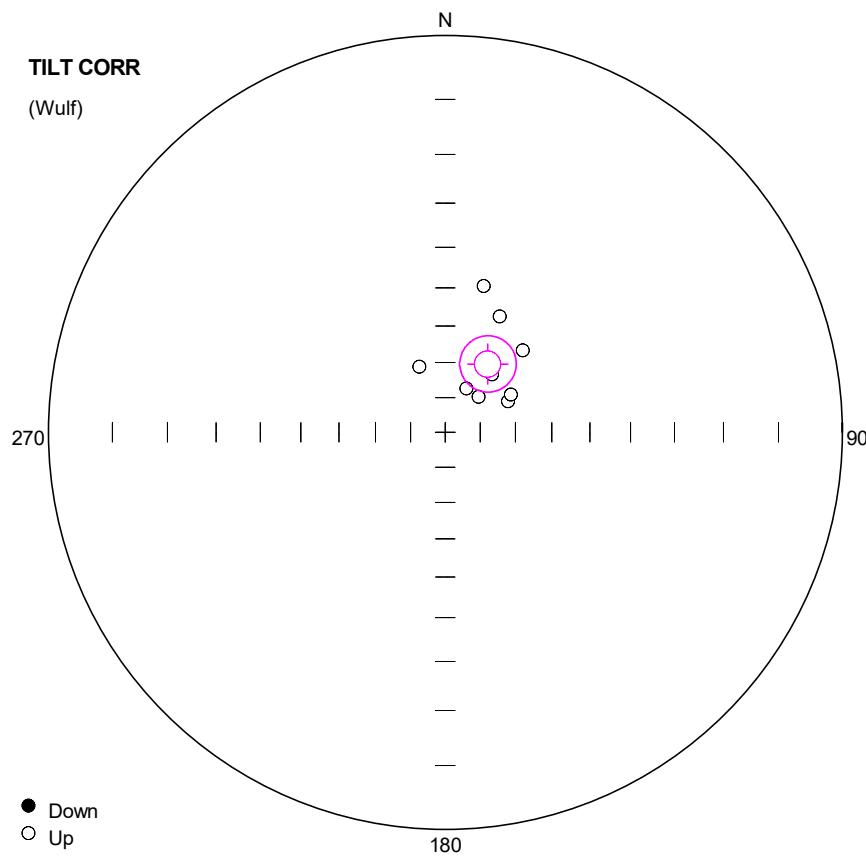
Pole longitude: 149.66 (-30.34)

Paleolatitude: -31.71 (31.71)

dp: 7.13

dm: 10.54

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0101	CCH	41.2	-40.4	12.1	35	64
2	YUN0102	CCH	30.6	-37.4	12.5	30	64
3	YUN0103	CCH	71.4	-50.4	9.0	35	64
4	YUN0104	CCH	68.9	-49.1	7.5	35	64
5	YUN0105	CCH	55.5	-41.0	10.8	48	64 O
6	YUN0106	CCH	52.4	-49.5	8.6	40	64
7	YUN0107	CCH	57.6	-51.5	12.2	35	64
8	YUN0108	CCH	57.2	-59.7	14.2	35	64 O
9	YUN0109	CCH	64.2	-57.7	17.5	40	64
10	YUN0110	CCH	27.4	-65.0	3.0	35	64



**Number of data points: 10**

## Fisher statistics

Mean vector: 31.81 / -67.12

Resultant vector: 9.77

(X= 3.23, Y= 2, Z= -9)

Estimated precision, k: 39.3

95% Confidence limit: 7.8

95% Confidence limit, approximation: 7.06

## Orientation matrix

1st eigenvalue: 0.96

1st eigenvector: 212.14 / 67.19

2nd eigenvalue: 0.03

2nd eigenvector: 1.09 / 19.82

3rd eigenvalue: 0.02

3rd eigenvector: 95.05 / 10.84

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -27.45 (27.45)

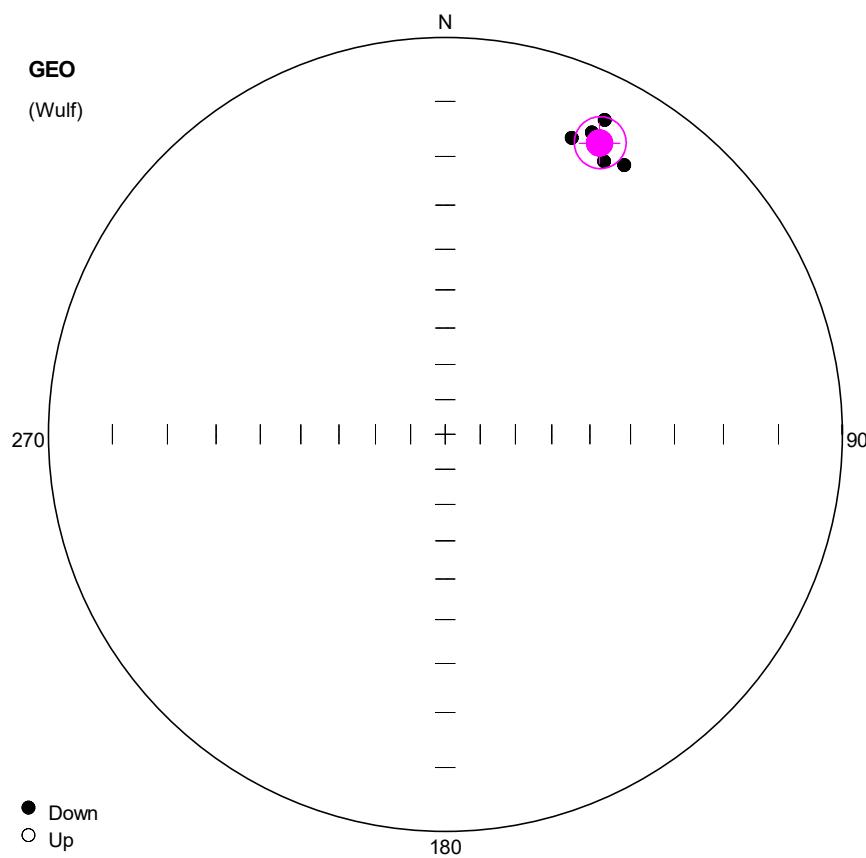
Pole longitude: 171.51 (-8.49)

Paleolatitude: -49.84 (49.84)

dp: 10.74

dm: 12.94

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0101	CCH	25.2	-54.3	12.1	35	64
2	YUN0102	CCH	14.8	-48.3	12.5	30	64
3	YUN0103	CCH	63.8	-70.0	9.0	35	64
4	YUN0104	CCH	60.0	-68.4	7.5	35	64
5	YUN0105	CCH	43.4	-58.3	10.8	48	64 O
6	YUN0106	CCH	33.1	-65.7	8.6	40	64
7	YUN0107	CCH	39.0	-68.8	12.2	35	64
8	YUN0108	CCH	25.9	-76.0	14.2	35	64 O
9	YUN0109	CCH	43.1	-76.0	17.5	40	64
10	YUN0110	CCH	338.5	-69.9	3.0	35	64



**Number of data points: 5**

**Fisher statistics**

Mean vector: 27.93 / 10.62

Resultant vector: 4.99

(X= 4.33, Y= 2.3, Z= 0.92)

Estimated precision, k: 300.49

95% Confidence limit: 4.42

95% Confidence limit, approximation: 3.61

**Orientation matrix**

1st eigenvalue: 0.99

1st eigenvector: 27.93 / 10.62

2nd eigenvalue: 0

2nd eigenvector: 120.78 / 14.85

3rd eigenvalue: 0

3rd eigenvector: 263.62 / 71.6

**VGP**

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 57.17 (-57.17)

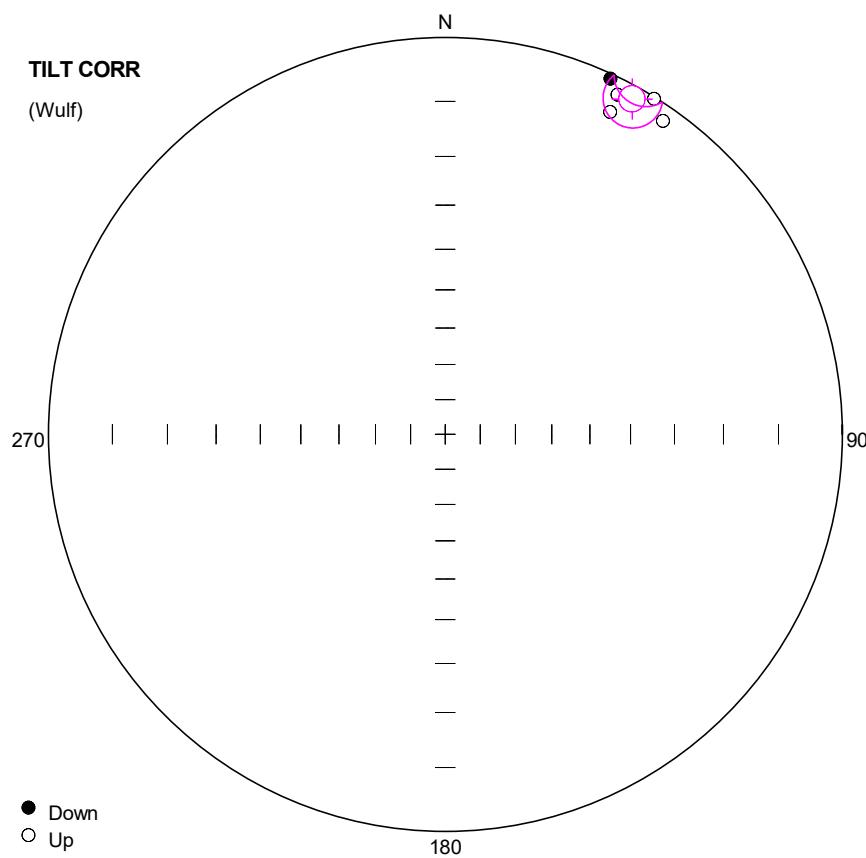
Pole longitude: -140.95 (39.05)

Paleolatitude: 5.36 (-5.36)

dp: 2.27

dm: 4.48

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0201	CCH	26.9	6.8	10.5	40	64
2	YUN0202	CCH	23.1	11.9	8.4	48	64
3	YUN0203	CCH	25.9	9.6	4.6	30	64 O
4	YUN0209	CCH	33.6	11.7	7.6	40	64 O
5	YUN0210	CCH	30.2	13.0	8.3	40	64 O



**Number of data points: 5**

**Fisher statistics**

Mean vector: 29.12 / -1.92  
 Resultant vector: 4.99  
 $(X= 4.35, Y= 2.43, Z= -0.17)$   
 Estimated precision,  $k$ : 298.55  
 95% Confidence limit: 4.44  
 95% Confidence limit, approximation: 3.62

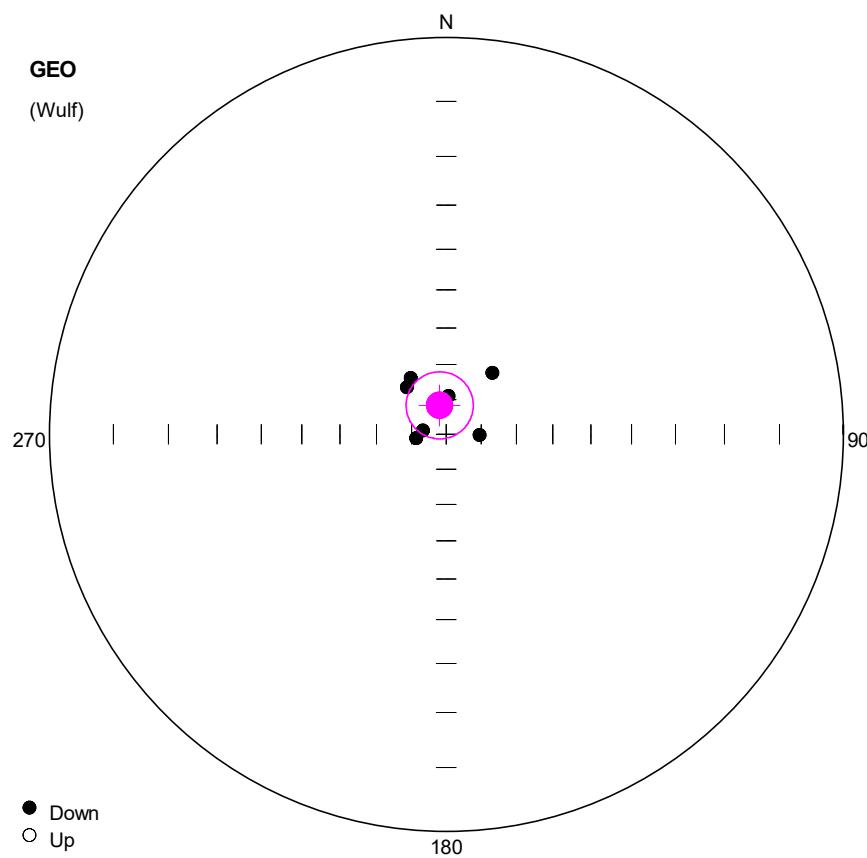
**Orientation matrix**

1st eigenvalue: 0.99  
 1st eigenvector: 209.12 / 1.92  
 2nd eigenvalue: 0  
 2nd eigenvector: 299.15 / 1.1  
 3rd eigenvalue: 0  
 3rd eigenvector: 58.83 / 87.78

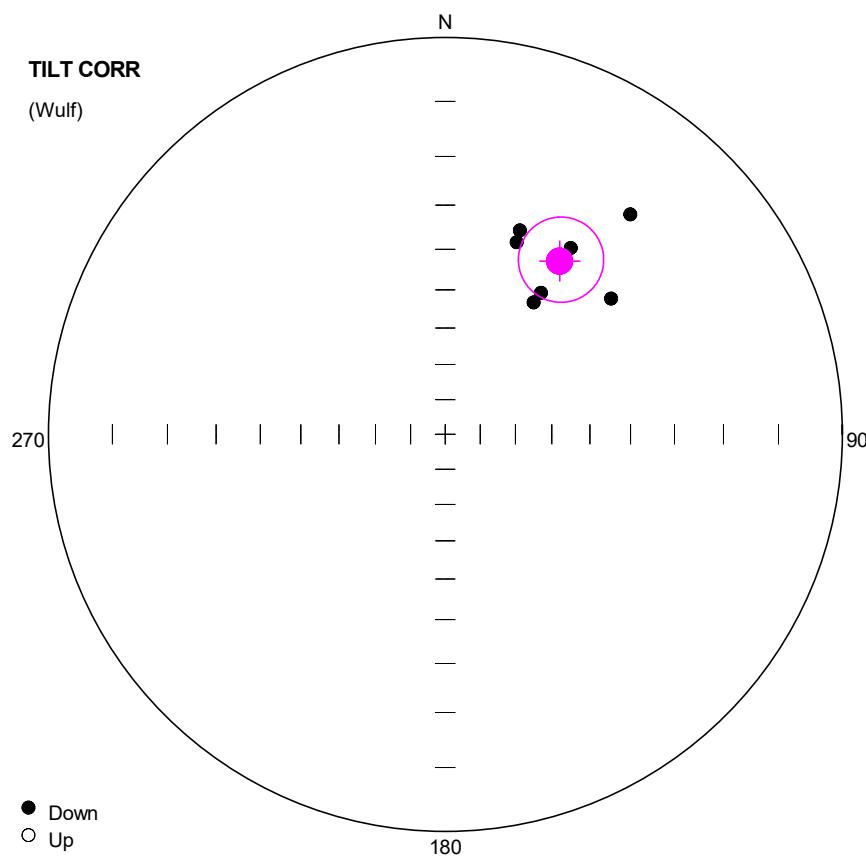
**VGP**

Site latitude: 24.27  
 Site longitude: 98.39  
 Pole latitude: 52.13 (-52.13)  
 Pole longitude: -134.04 (45.96)  
 Paleolatitude: -0.96 (0.96)  
 dp: 2.22  
 dm: 4.44

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0201	CCH	27.1	-5.3	10.5	40	64
2	YUN0202	CCH	24.9	0.7	8.4	48	64
3	YUN0203	CCH	26.9	-2.4	4.6	30	64 O
4	YUN0209	CCH	34.8	-2.3	7.6	40	64 O
5	YUN0210	CCH	31.9	-0.3	8.3	40	64 O



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0301	CCH	91.0	80.4	12.3	35	64
2	YUN0302	CCH		3.5	78.9	9.1	35
3	YUN0303	CCH		37.0	68.1	12.6	35
4	YUN0304	CCH		279.1	83.2	7.2	35
5	YUN0306	CCH		320.1	72.4	11.6	40
6	YUN0307	CCH		262.3	81.3	10.1	35
7	YUN0310	CCH		327.8	71.0	3.2	35



**Number of data points: 7**

**Fisher statistics**

Mean vector: 33.45 / 34.77

Resultant vector: 6.85

(X= 4.7, Y= 3.1, Z= 3.91)

Estimated precision, k: 40.33

95% Confidence limit: 9.62

95% Confidence limit, approximation: 8.33

**Orientation matrix**

1st eigenvalue: 0.96

1st eigenvector: 33.39 / 34.84

2nd eigenvalue: 0.03

2nd eigenvector: 272.64 / 36.3

3rd eigenvalue: 0.02

3rd eigenvector: 152.14 / 34.65

**VGP**

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 58.58 (-58.58)

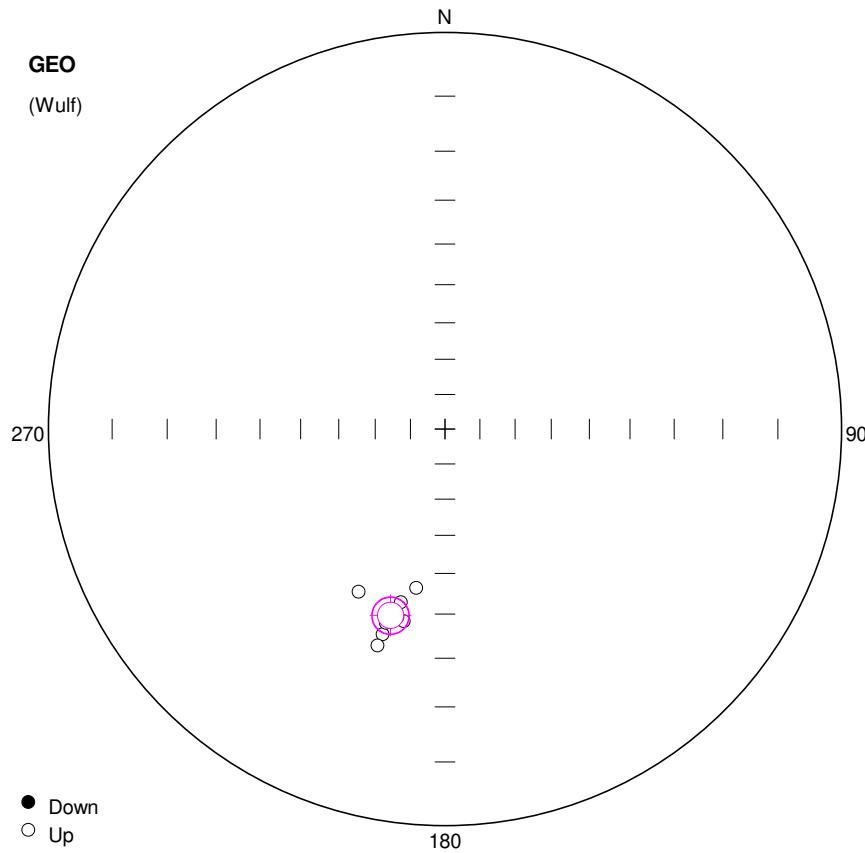
Pole longitude: -168.86 (11.14)

Paleolatitude: 19.14 (-19.14)

dp: 6.36

dm: 11.06

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0301	CCH	50.7	33.3	12.3	35	64
2	YUN0302	CCH	34.0	31.0	9.1	35	64
3	YUN0303	CCH	40.1	18.2	12.6	35	60
4	YUN0304	CCH	34.1	43.4	7.2	35	64
5	YUN0306	CCH	20.4	35.4	11.6	40	64
6	YUN0307	CCH	33.8	46.4	10.1	35	60
7	YUN0310	CCH	20.1	32.7	3.2	35	64



**Number of data points: 8**

## Fisher statistics

Mean vector: 196.3 / -37.86

Resultant vector: 7.96

(X= -6.03, Y= -1.76, Z= -4.88)

Estimated precision, k: 164.83

95% C

95% Confidence limit, approximation: 3.86

## Orientation matrix

1st eigenvalue: 0.99

1st eigenvector: 16.29 / 37.85

2nd eigenvalue: 0.01

2nd eigenvector: 227.28 / 47.8

3rd eigenvalue: 0

3rd eigenvector: 119.03 / 15.85

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -44.95 (44.95)

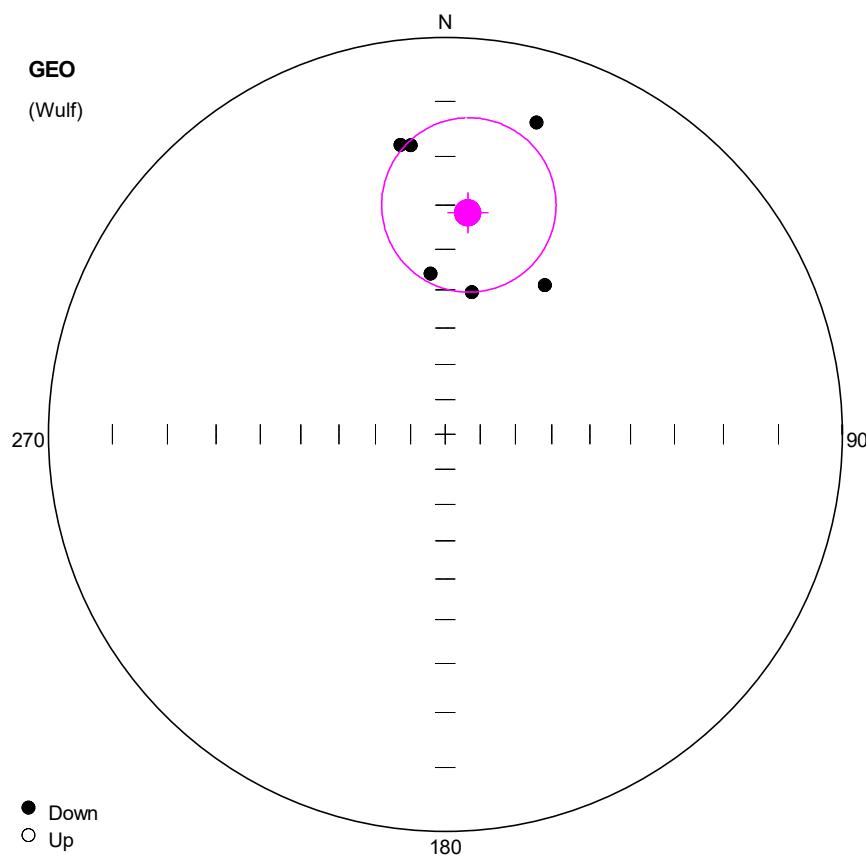
Pole longitude: -7.65 (172.35)

Paleolatitude: -21.24 (21.24)

dp: 3.02

dm: 5.11

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0401	CPCA	196.9	-33.2	0.4	30	100
2	YUN0402	CPCA	207.9	-40.2	1.5	20	120
3	YUN0403	CPCA	197.0	-35.9	1.3	20	120
4	YUN0404	CPCA	192.1	-37.3	0.9	20	120
5	YUN0405	CPCA	197.3	-30.5	2.2	20	120
6	YUN0406	CPCA	194.2	-37.8	1.5	20	120
7	YUN0407	CPCA	194.3	-41.5	1.6	20	120
8	YUN0408	CPCA	190.2	-45.7	2.0	20	120



**Number of data points: 6**

**Fisher statistics**

Mean vector: 5.82 / 31.44

Resultant vector: 5.63

(X= 4.78, Y= 0.49, Z= 2.94)

Estimated precision, k: 13.65

95% Confidence limit: 18.8

95% Confidence limit, approximation: 15.47

**Orientation matrix**

1st eigenvalue: 0.88

1st eigenvector: 5.58 / 31.45

2nd eigenvalue: 0.08

2nd eigenvector: 141.91 / 49.78

3rd eigenvalue: 0.04

3rd eigenvector: 261.01 / 22.35

**VGP**

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 80.92 (-80.92)

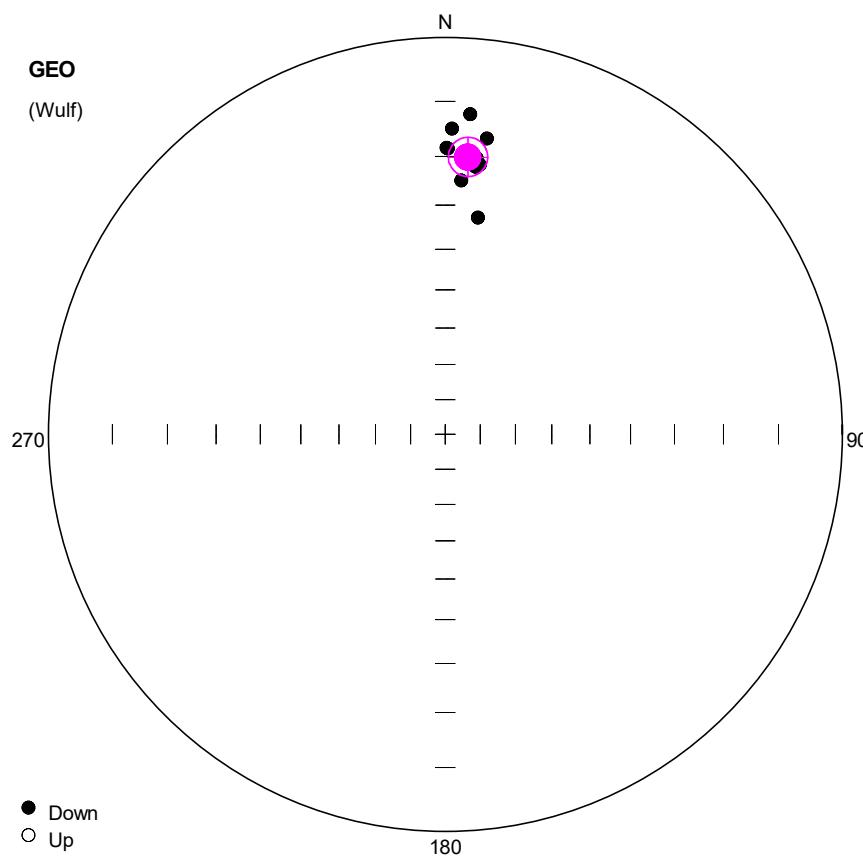
Pole longitude: -119.5 (60.5)

Paleolatitude: 17 (-17)

dp: 11.81

dm: 21.07

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN0701	CPCA	33.7	41.4	23.2	25	70
2	YUN0702	CPCA	351.2	17.2	34.3	25	70
3	YUN0703	CPCA	354.8	45.8	19.0	25	70
4	YUN0706	CPCA	10.6	50.0	27.3	25	70
5	YUN0707	CPCA	16.3	11.4	40.9	25	70
6	YUN0708	CPCA	353.2	17.5	30.4	25	70



Number of data points: 10

#### Fisher statistics

Mean vector: 4.65 / 19.99  
 Resultant vector: 9.94  
 $(X= 9.32, Y= 0.76, Z= 3.4)$   
 Estimated precision, k: 163.34  
 95% Confidence limit: 3.79  
 95% Confidence limit, approximation: 3.46

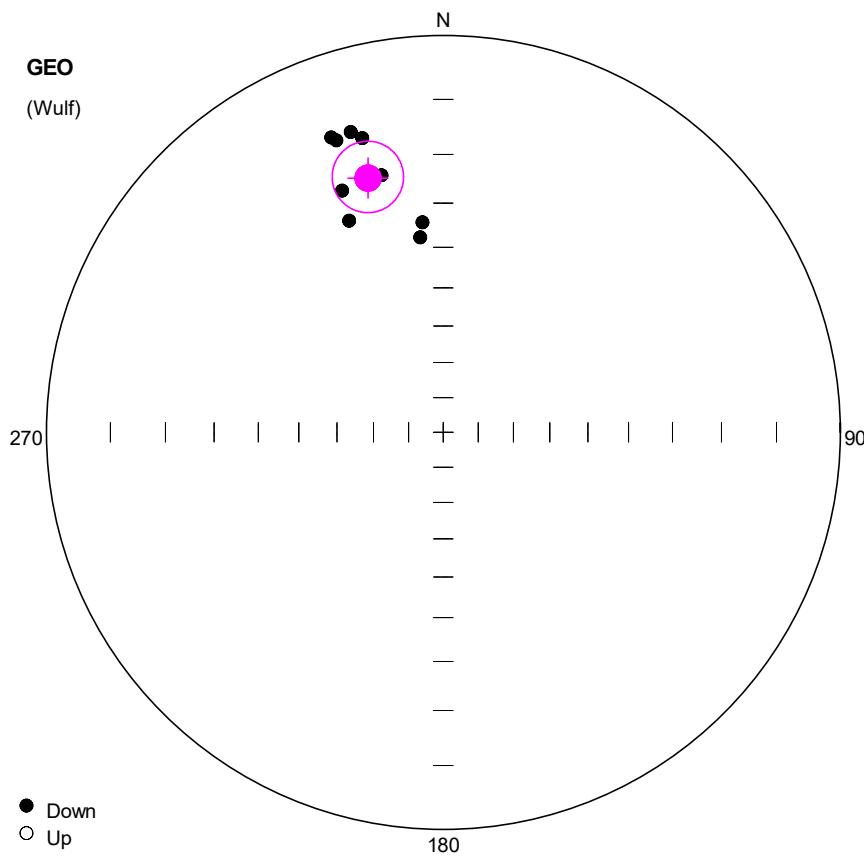
#### Orientation matrix

1st eigenvalue: 0.99  
 1st eigenvector: 4.64 / 19.97  
 2nd eigenvalue: 0.01  
 2nd eigenvector: 146.43 / 65.18  
 3rd eigenvalue: 0  
 3rd eigenvector: 269.4 / 14.12

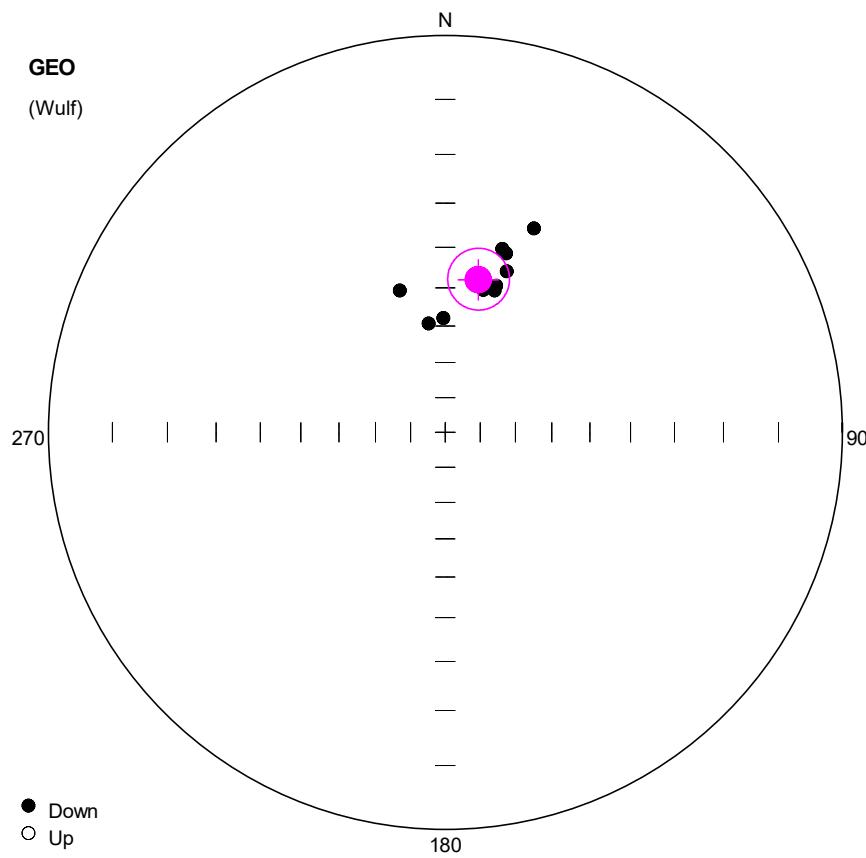
#### VGP

Site latitude: 24.27  
 Site longitude: 98.39  
 Pole latitude: 75.35 (-75.35)  
 Pole longitude: -100 (80)  
 Paleolatitude: 10.31 (-10.31)  
 dp: 2.08  
 dm: 3.97

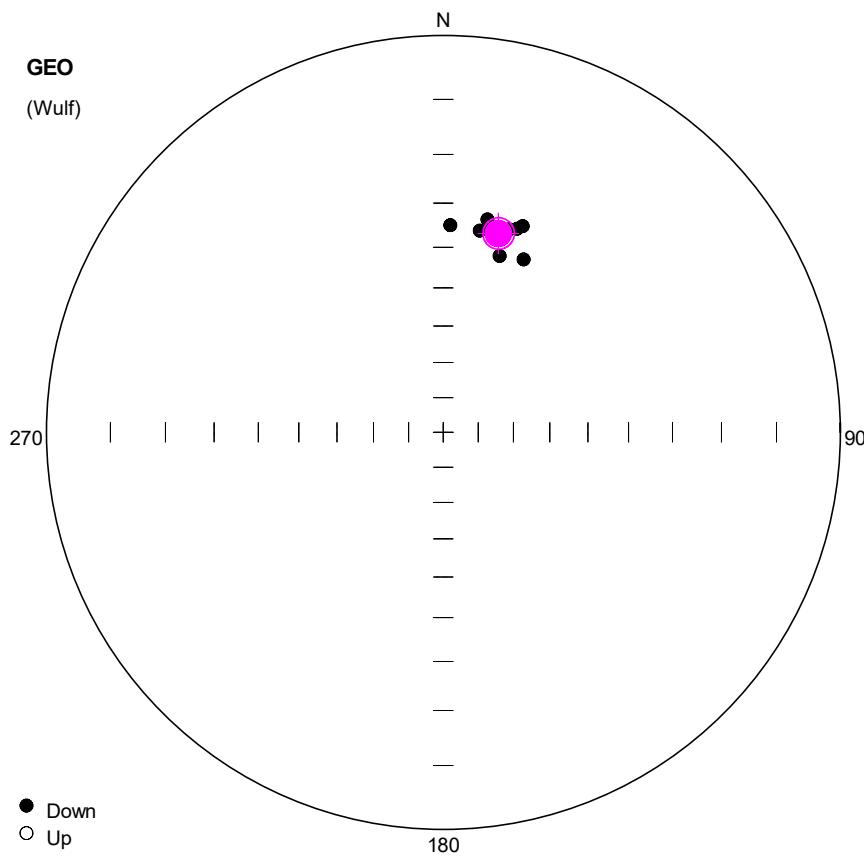
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1001	CPCA	3.6	24.7	0.6	30	120
2	YUN1002	CPCA	8.6	32.2	0.4	30	120
3	YUN1003	CPCA	0.3	18.4	0.5	30	120
4	YUN1004	CPCA	7.3	21.2	0.6	30	120
5	YUN1005	CPCA	6.4	21.7	0.5	30	120
6	YUN1006	CPCA	1.3	14.8	0.9	40	120
7	YUN1007	CPCA	6.5	20.0	1.0	40	120
8	YUN1008	CPCA	8.0	16.1	0.9	30	120
9	YUN1009	CPCA	0.5	18.5	0.5	20	120
10	YUN1010	CPCA	4.5	12.1	0.3	30	120



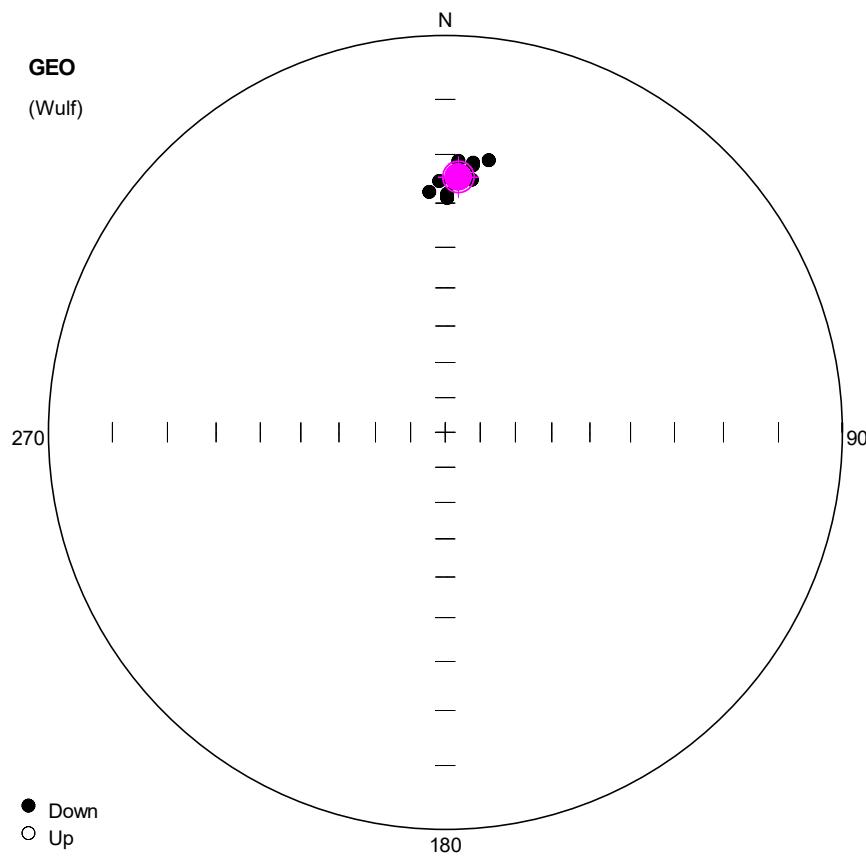
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1101	CPCA	339.9	13.9	3.1	20	120
2	YUN1102	CPCA	339.2	13.1	3.7	30	100
3	YUN1103	CPCA	344.6	14.9	4.3	30	120
4	YUN1104	CPCA	346.5	22.7	3.4	30	120
5	YUN1105	CPCA	342.9	13.3	3.7	30	120
6	YUN1106	CPCA	336.0	29.5	3.3	30	120
7	YUN1107	CPCA	353.3	37.4	2.0	30	120
8	YUN1108	CPCA	337.3	23.2	4.2	30	120
9	YUN1110	CPCA	354.4	34.0	1.7	30	120



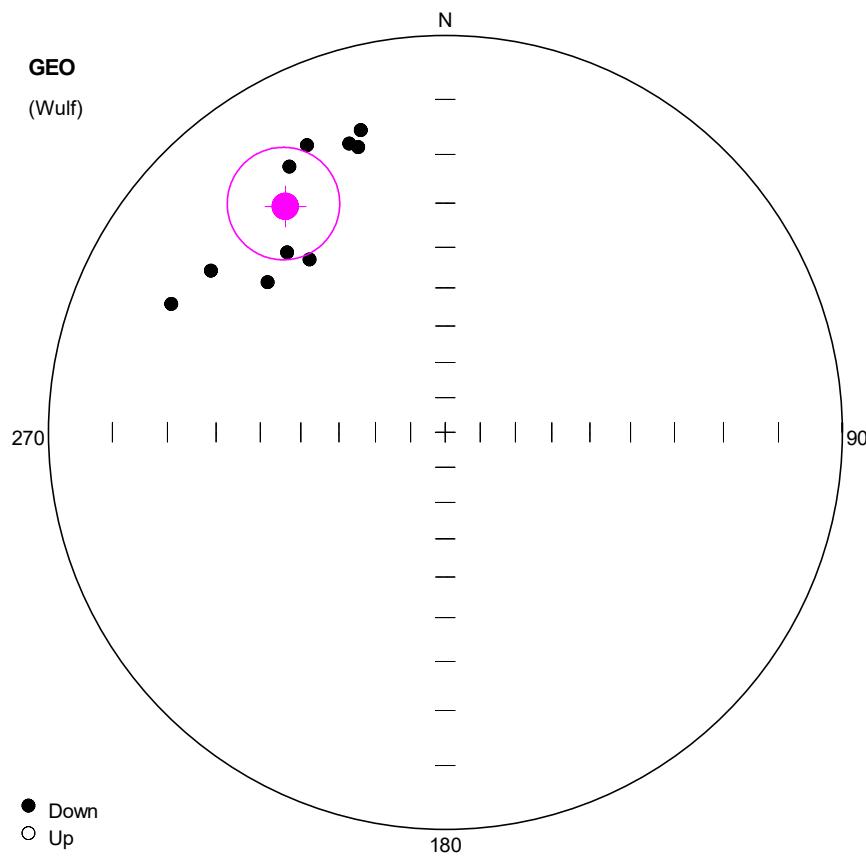
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1201	CPCA	15.0	49.3	1.4	30	120
2	YUN1202	CPCA	19.2	47.3	1.2	10	120
3	YUN1203	CPCA	18.8	39.1	1.3	20	120
4	YUN1204	CPCA	17.3	38.4	1.1	20	120
5	YUN1205	CPCA	342.2	48.9	0.8	30	120
6	YUN1206	CPCA	20.9	43.1	2.7	30	120
7	YUN1207	CPCA	23.5	31.5	0.6	20	120
8	YUN1208	CPCA	351.4	59.0	1.7	30	120
9	YUN1209	CPCA	359.0	57.9	0.6	30	120
10	YUN1210	CPCA	19.2	48.6	0.7	20	120



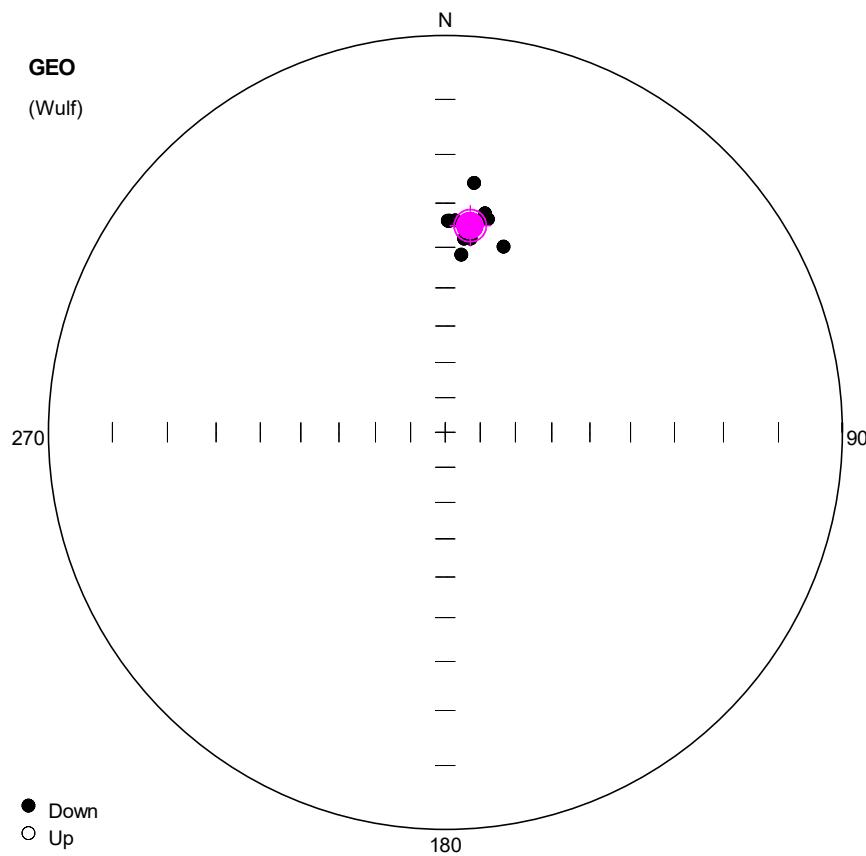
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1301	CPCA	25.0	38.7	0.7	30	120
2	YUN1302	CPCA	10.3	35.4	1.1	30	120
3	YUN1303	CPCA	20.0	32.8	0.6	30	120
4	YUN1304	CPCA	21.1	31.8	1.4	30	120
5	YUN1305	CPCA	18.0	33.5	1.0	20	120
6	YUN1306	CPCA	16.9	35.1	0.7	20	120
7	YUN1307	CPCA	11.7	32.6	0.7	20	120
8	YUN1308	CPCA	2.0	34.9	0.6	10	120
9	YUN1309	CPCA	12.4	34.4	0.8	10	120
10	YUN1310	CPCA	17.7	40.0	0.5	10	120



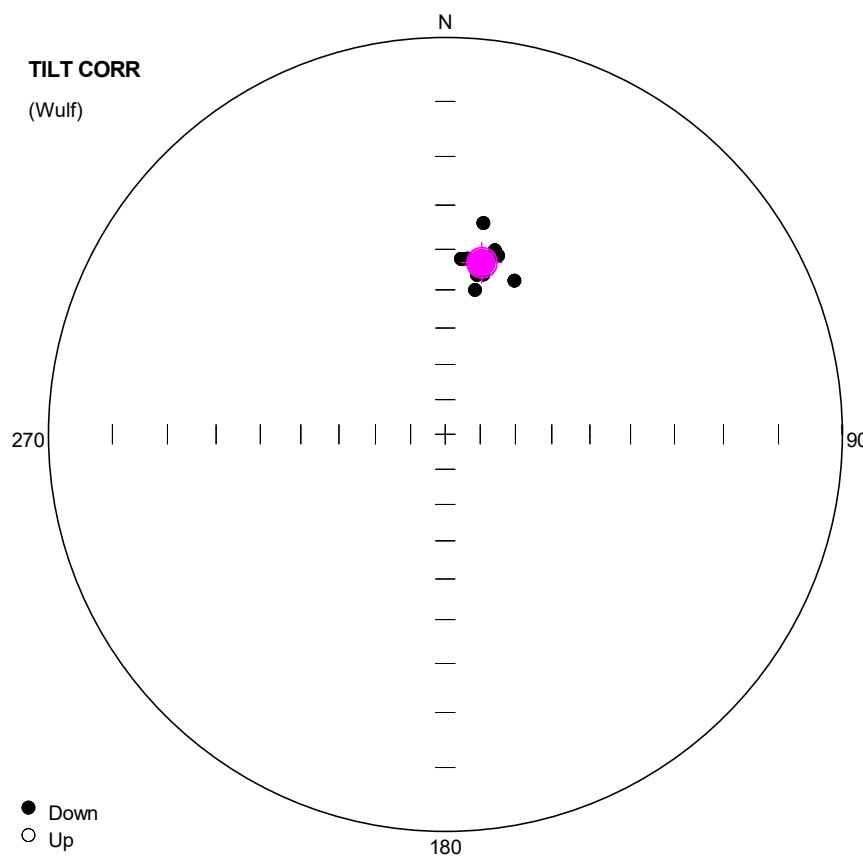
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1401	CPCA	5.9	21.9	0.6	10	120
2	YUN1402	CPCA	9.1	20.5	0.7	10	120
3	YUN1403	CPCA	6.0	24.8	0.5	10	120
4	YUN1404	CPCA	5.9	21.4	0.4	20	120
5	YUN1405	CPCA	358.6	25.3	0.2	10	120
6	YUN1406	CPCA	2.8	21.2	0.2	30	120
7	YUN1407	CPCA	356.2	27.5	0.5	30	120
8	YUN1409	CPCA	0.5	28.9	0.3	30	120
9	YUN1410	CPCA	0.5	28.0	0.2	30	120



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1501	CPCA	321.9	32.1	5.6	50	120
2	YUN1502	CPCA	318.7	27.8	6.2	40	120
3	YUN1503	CPCA	343.0	16.2	2.5	30	80
4	YUN1504	CPCA	341.6	15.1	4.5	30	80
5	YUN1505	CPCA	344.4	13.4	4.8	40	120
6	YUN1506	CPCA	334.3	12.5	2.6	30	120
7	YUN1507	CPCA	329.6	14.4	2.2	20	50
8	YUN1508	CPCA	310.2	29.3	3.6	30	120
9	YUN1509	CPCA	304.6	18.7	5.9	60	120
10	YUN1510	CPCA	295.1	15.4	11.6	60	120



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1601	CPCA	6.6	25.4	1.0	20	120
2	YUN1602	CPCA	5.6	37.8	1.1	20	120
3	YUN1603	CPCA	0.7	33.9	1.1	20	120
4	YUN1604	CPCA	5.2	41.6	1.0	20	120
5	YUN1605	CPCA	2.7	33.8	0.7	30	120
6	YUN1606	CPCA	1.2	33.9	0.6	30	120
7	YUN1607	CPCA	11.3	32.6	0.5	30	120
8	YUN1608	CPCA	10.3	31.5	0.7	30	120
9	YUN1609	CPCA	7.5	37.6	0.5	30	120
10	YUN1610	CPCA	17.4	37.8	0.4	30	120



Number of data points: 10

**Fisher statistics**

Mean vector: 12.02 / 42.31

Resultant vector: 9.95

(X= 7.2, Y= 1.53, Z= 6.7)

Estimated precision, k: 177.11

95% Confidence limit: 3.64

95% Confidence limit, approximation: 3.33

**Orientation matrix**

1st eigenvalue: 0.99

1st eigenvector: 12.01 / 42.31

2nd eigenvalue: 0.01

2nd eigenvector: 147.38 / 38.02

3rd eigenvalue: 0

3rd eigenvector: 257.91 / 24.16

**VGP**

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 79.06 (-79.06)

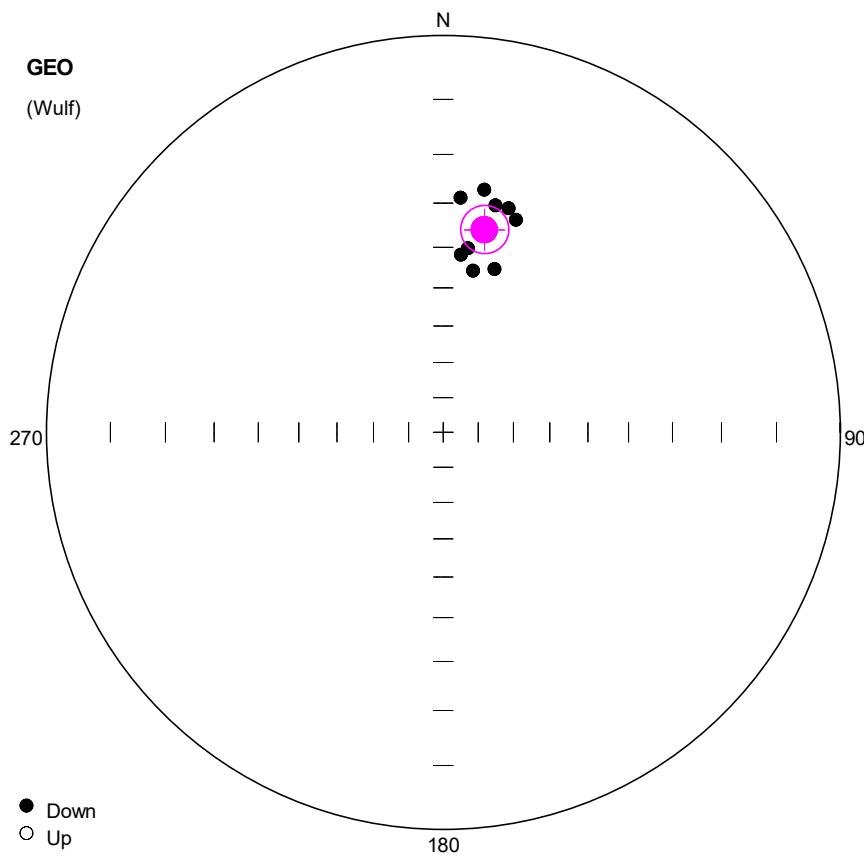
Pole longitude: -175.15 (4.85)

Paleolatitude: 24.47 (-24.47)

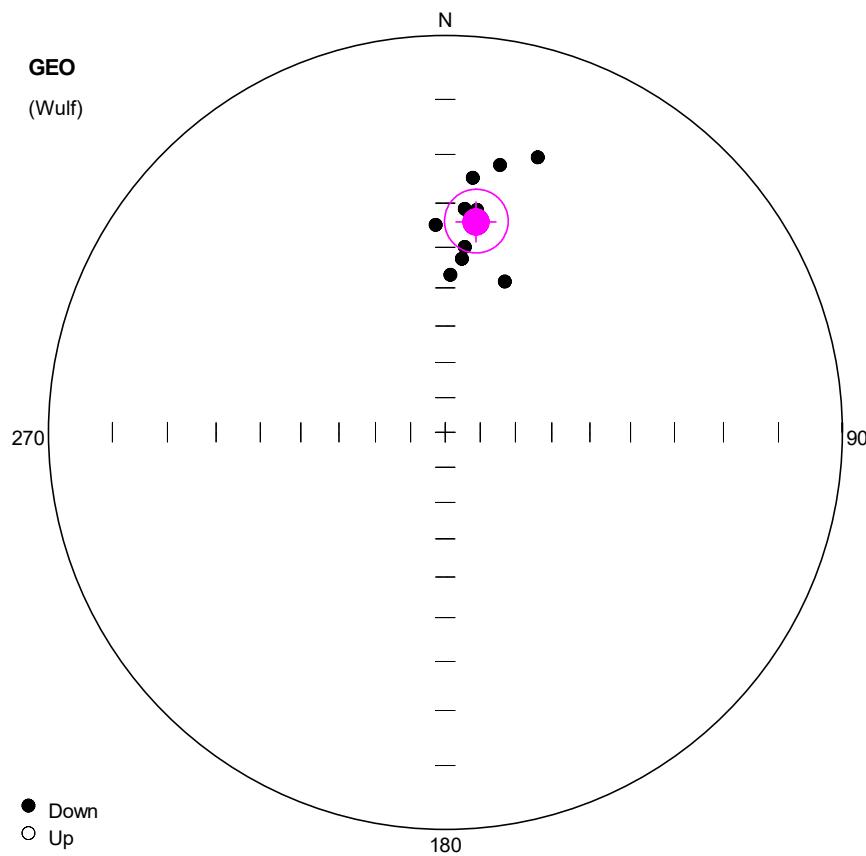
dp: 2.76

dm: 4.48

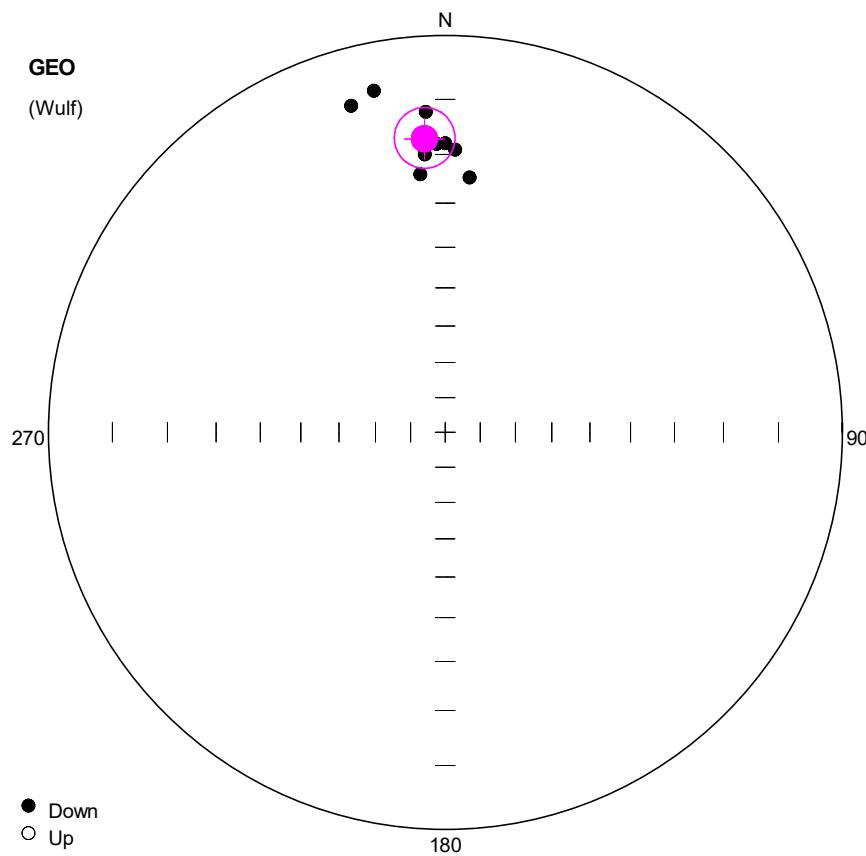
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1601	CPCA	10.2	33.2	1.0	20	120
2	YUN1602	CPCA	11.3	45.5	1.1	20	120
3	YUN1603	CPCA	5.1	42.2	1.1	20	120
4	YUN1604	CPCA	11.7	49.3	1.0	20	120
5	YUN1605	CPCA	7.3	41.9	0.7	30	120
6	YUN1606	CPCA	5.6	42.2	0.6	30	120
7	YUN1607	CPCA	16.5	39.7	0.5	30	120
8	YUN1608	CPCA	15.1	38.7	0.7	30	120
9	YUN1609	CPCA	13.4	45.1	0.5	30	120
10	YUN1610	CPCA	24.3	44.0	0.4	30	120



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1701	CPCA	18.9	31.0	0.8	30	120
2	YUN1702	CPCA	17.4	43.4	0.6	20	120
3	YUN1703	CPCA	5.7	41.6	0.3	30	120
4	YUN1704	CPCA	10.4	45.0	0.6	20	120
5	YUN1705	CPCA	4.2	28.7	1.0	20	120
6	YUN1706	CPCA	7.7	39.9	1.0	20	120
7	YUN1708	CPCA	9.6	26.4	0.5	20	120
8	YUN1709	CPCA	13.0	29.2	0.6	20	120
9	YUN1710	CPCA	16.3	29.1	0.4	20	120



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1801	CPCA	5.0	31.1	0.4	20	120
2	YUN1802	CPCA	8.1	31.0	0.4	30	120
3	YUN1803	CPCA	6.0	39.7	0.3	30	120
4	YUN1804	CPCA	11.6	21.0	2.3	30	120
5	YUN1805	CPCA	18.6	17.7	1.0	20	120
6	YUN1806	CPCA	6.2	24.4	1.4	30	120
7	YUN1807	CPCA	1.9	46.7	1.0	30	120
8	YUN1808	CPCA	357.4	34.8	2.2	40	120
9	YUN1810	CPCA	5.5	42.6	0.2	30	120
10	YUN1811	CPCA	21.6	45.6	0.1	30	120



**Number of data points: 9**

## Fisher statistics

Mean vector: 355.95 / 16.92

Resultant vector: 8.91

(X= 8.5, Y= -0.6, Z= 2.59)

Estimated precision, k: 84.25

95% Confidence limit: 5.64

95% Confidence limit, approximation: 5.08

## Orientation matrix

1st eigenvalue: 0.98

1st eigenvector: 355.99 / 16.95

2nd eigenvalue: 0.02

2nd eigenvector: 101.03 / 40.4

3rd eigenvalue: 0

3rd eigenvector: 248.43 / 44.71

VGP

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 73.91 (-73.91)

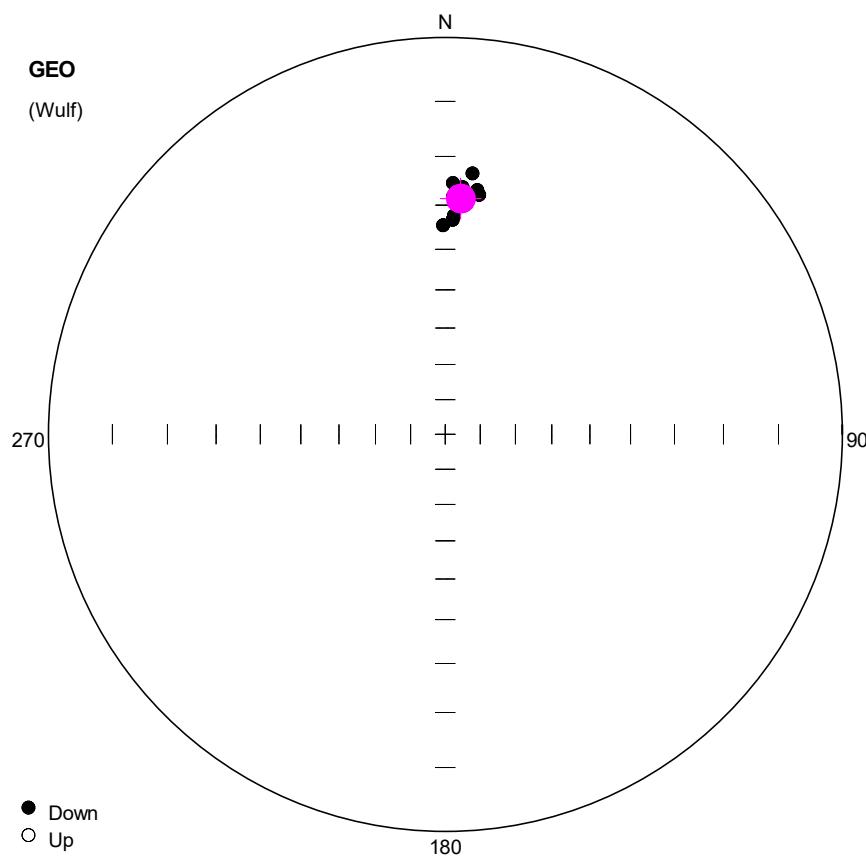
Pole longitude: -67.03 (112.97)

Paleolatitude: 8.65 (-8.65)

dp: 3.01

dm: 5.83

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN1901	CPCA	343.9	8.9	1.0	30	120
2	YUN1902	CPCA	348.2	7.4	1.2	30	120
3	YUN1904	CPCA	358.3	18.0	0.6	30	120
4	YUN1905	CPCA	360.0	17.9	0.4	20	120
5	YUN1906	CPCA	5.5	24.4	0.7	30	120
6	YUN1907	CPCA	2.0	19.1	0.8	30	120
7	YUN1908	CPCA	354.5	23.7	0.4	20	120
8	YUN1909	CPCA	356.5	12.1	0.5	20	120
9	YUN1910	CPCA	355.8	19.9	0.4	20	120



**Number of data points:** 9

**Fisher statistics**

Mean vector: 3.73 / 28.49  
 Resultant vector: 8.97  
 $(X= 7.87, Y= 0.51, Z= 4.28)$   
 Estimated precision, k: 296.96  
 95% Confidence limit: 2.99  
 95% Confidence limit, approximation: 2.71

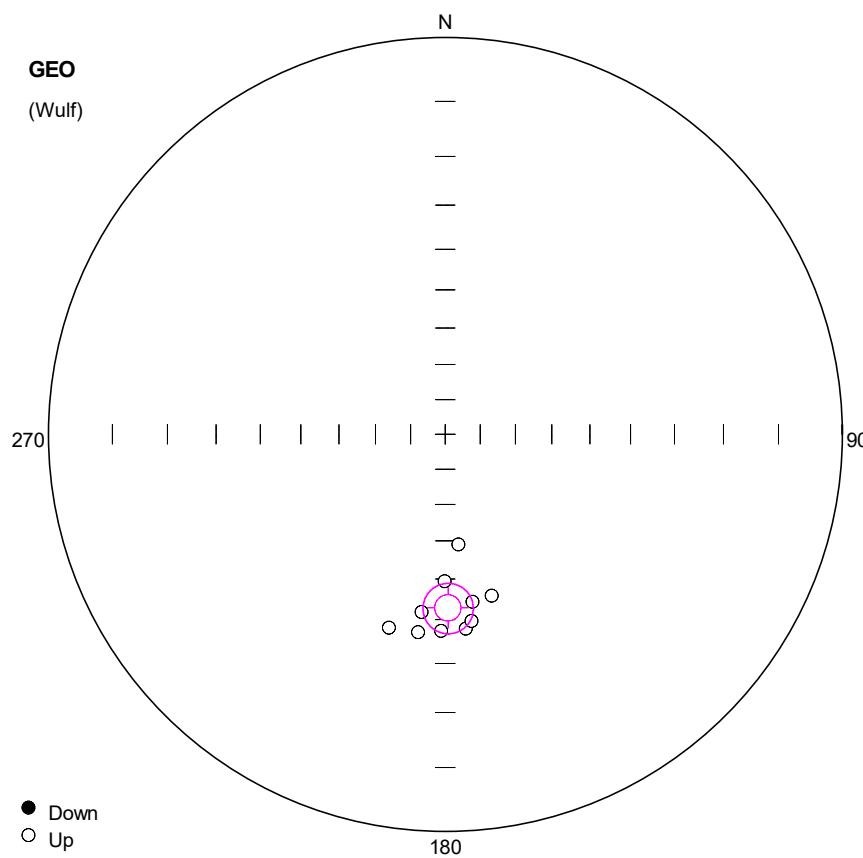
**Orientation matrix**

1st eigenvalue: 0.99  
 1st eigenvector: 3.73 / 28.49  
 2nd eigenvalue: 0.01  
 2nd eigenvector: 230.93 / 51.39  
 3rd eigenvalue: 0  
 3rd eigenvector: 107.53 / 23.73

**VGP**

Site latitude: 24.27  
 Site longitude: 98.39  
 Pole latitude: 80.26 (-80.26)  
 Pole longitude: -103.4 (76.6)  
 Paleolatitude: 15.18 (-15.18)  
 dp: 1.8  
 dm: 3.29

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2002	CPCA	8.1	27.3	0.2	30	120
2	YUN2003	CPCA	359.4	34.5	0.3	30	120
3	YUN2004	CPCA	6.0	23.1	0.4	30	120
4	YUN2005	CPCA	7.5	26.4	0.3	30	120
5	YUN2006	CPCA	1.9	28.0	0.5	30	120
6	YUN2007	CPCA	2.0	33.2	0.4	30	120
7	YUN2008	CPCA	4.0	26.1	0.6	30	120
8	YUN2009	CPCA	1.8	25.3	0.7	30	120
9	YUN2010	CPCA	2.3	32.3	0.2	30	120



**Number of data points: 10**

**Fisher statistics**

Mean vector: 179.09 / -42.79

Resultant vector: 9.86

(X= -7.23, Y= 0.11, Z= -6.7)

Estimated precision, k: 63.67

95% Confidence limit: 6.1

95% Confidence limit, approximation: 5.55

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 359.07 / 42.75

2nd eigenvalue: 0.02

2nd eigenvector: 239.05 / 28.42

3rd eigenvalue: 0.01

3rd eigenvector: 127.65 / 34

**VGP**

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: -89 (89)

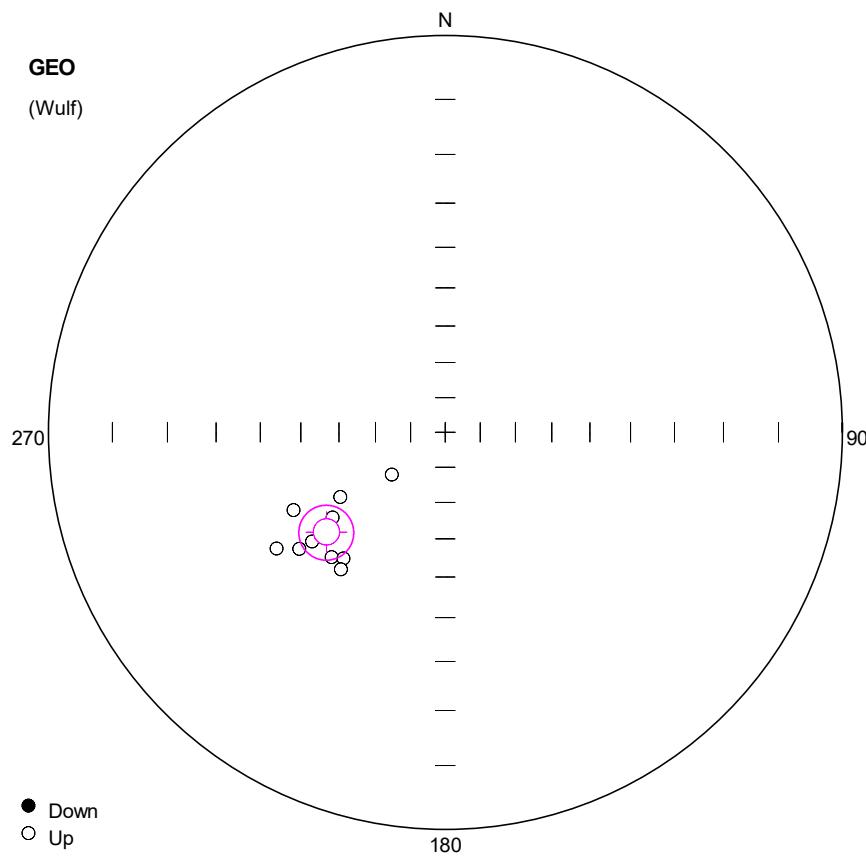
Pole longitude: -136.75 (43.25)

Paleolatitude: -24.84 (24.84)

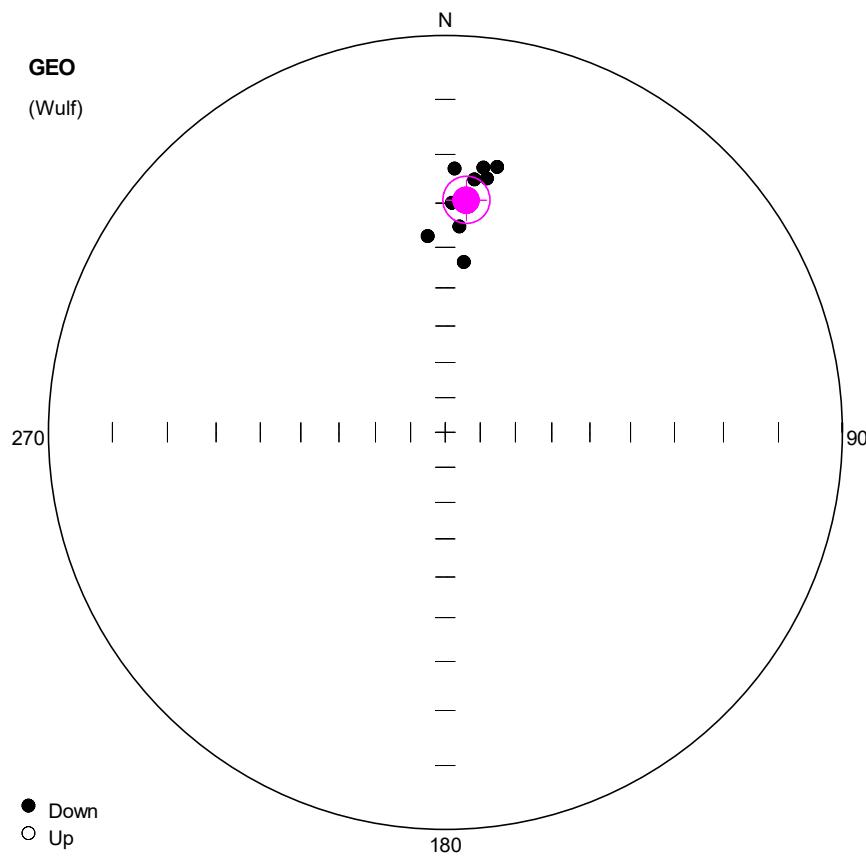
dp: 4.67

dm: 7.54

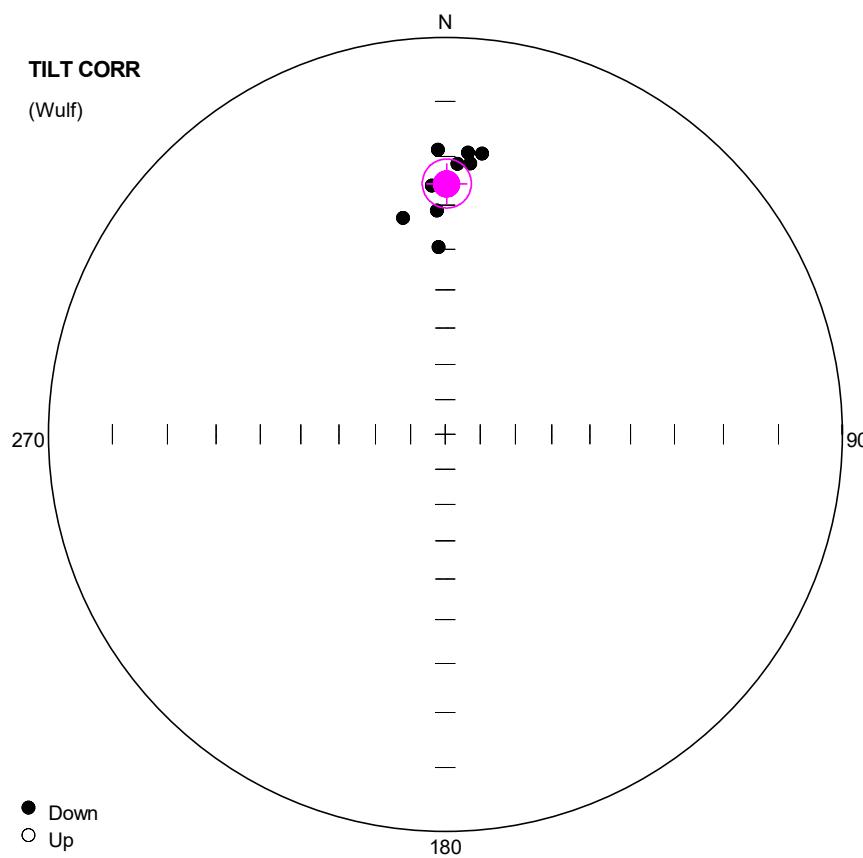
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2101	CPCA	181.2	-37.3	1.5	30	120
2	YUN2102	CPCA	187.6	-41.4	0.9	30	120
3	YUN2103	CPCA	174.0	-37.6	1.1	40	120
4	YUN2104	CPCA	170.8	-43.7	1.0	30	120
5	YUN2105	CPCA	172.0	-39.2	0.8	30	120
6	YUN2106	CPCA	180.2	-49.4	0.6	30	120
7	YUN2107	CPCA	187.8	-36.6	1.7	30	120
8	YUN2108	CPCA	196.2	-36.2	0.6	30	120
9	YUN2109	CPCA	163.9	-44.1	0.6	30	120
10	YUN2110	CPCA	173.2	-58.8	1.0	30	120



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2201	CPCA	232.9	-50.8	1.5	30	120
2	YUN2202	CPCA	238.4	-55.5	1.4	30	120
3	YUN2203	CPCA	231.4	-39.6	2.0	30	120
4	YUN2204	CPCA	222.3	-43.9	1.8	30	120
5	YUN2205	CPCA	231.5	-70.5	1.1	30	120
6	YUN2206	CPCA	242.8	-43.5	1.1	30	120
7	YUN2207	CPCA	235.4	-35.4	1.1	30	120
8	YUN2208	CPCA	218.9	-45.6	1.4	30	120
9	YUN2210	CPCA	217.3	-43.1	1.1	30	120
10	YUN2211	CPCA	230.6	-43.1	2.0	30	120



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2301	CPCA	3.9	35.1	0.2	30	120
2	YUN2302	CPCA	2.0	22.8	0.2	30	120
3	YUN2303	CPCA	1.7	29.9	0.4	30	120
4	YUN2304	CPCA	11.1	21.5	0.2	30	120
5	YUN2305	CPCA	8.2	22.1	0.4	30	120
6	YUN2306	CPCA	6.2	43.3	0.2	30	120
7	YUN2307	CPCA	6.6	24.7	0.1	30	120
8	YUN2308	CPCA	9.3	24.1	0.4	30	120
9	YUN2309	CPCA	354.9	37.2	0.6	30	120
10	YUN2310	CPCA	6.1	30.4	0.2	30	120



Number of data points: 10

#### Fisher statistics

Mean vector: 0.32 / 25.53  
Resultant vector: 9.9  
(X= 8.94, Y= 0.05, Z= 4.27)  
Estimated precision, k: 92.56  
95% Confidence limit: 5.05  
95% Confidence limit, approximation: 4.6

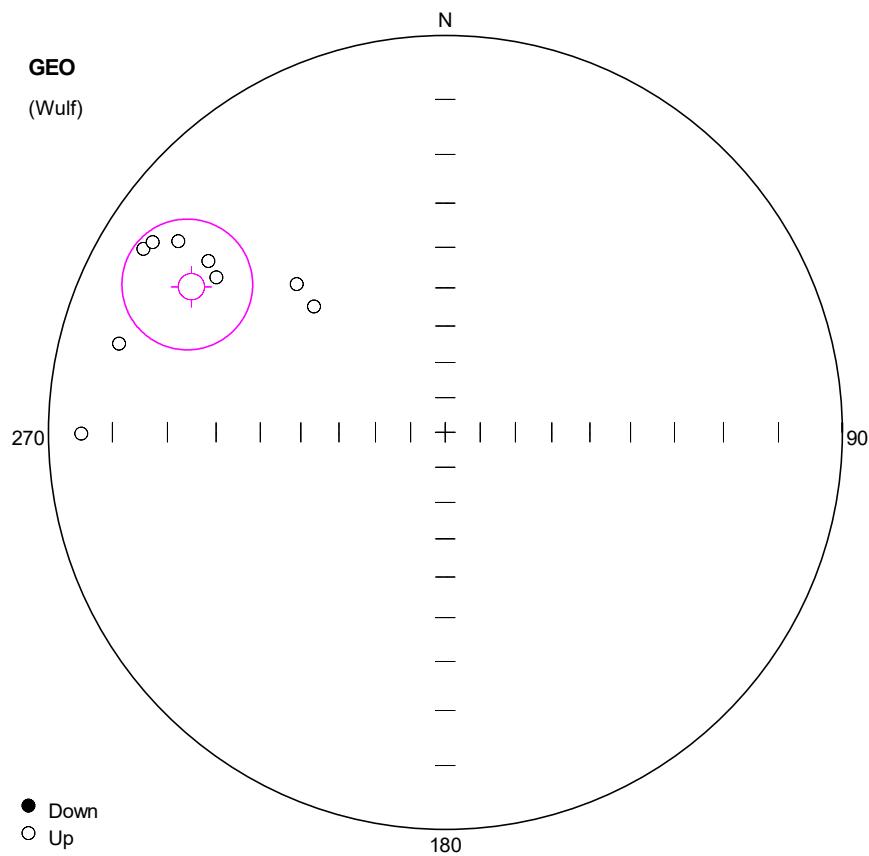
#### Orientation matrix

1st eigenvalue: 0.98  
1st eigenvector: 0.33 / 25.5  
2nd eigenvalue: 0.02  
2nd eigenvector: 232.47 / 52.15  
3rd eigenvalue: 0  
3rd eigenvector: 103.74 / 25.93

#### VGP

Site latitude: 24.27  
Site longitude: 98.39  
Pole latitude: 79.16 (-79.16)  
Pole longitude: -83.25 (96.75)  
Paleolatitude: 13.43 (-13.43)  
dp: 2.93  
dm: 5.44

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2301	CPCA	357.9	31.2	0.2	30	120
2	YUN2302	CPCA	358.5	18.7	0.2	30	120
3	YUN2303	CPCA	356.9	25.8	0.4	30	120
4	YUN2304	CPCA	7.5	19.0	0.2	30	120
5	YUN2305	CPCA	4.6	19.1	0.4	30	120
6	YUN2306	CPCA	358.0	39.5	0.2	30	120
7	YUN2307	CPCA	2.6	21.4	0.1	30	120
8	YUN2308	CPCA	5.3	21.2	0.4	30	120
9	YUN2309	CPCA	349.0	31.9	0.6	30	120
10	YUN2310	CPCA	1.0	26.8	0.2	30	120



**Number of data points: 9**

## Fisher statistics

Mean vector: 299.83 / -17.23

Resultant vector: 8.58

(X= 4.08, Y= -7.11, Z= -2.54)

Estimated precision, k: 19.04

95% Confidence limit: 12.11

95% Confidence limit, approx

## Orientation matrix

1st eigenvalue: 0.91

1st eigenvector: 120.19 / 17.05

2nd eigenvalue: 0.07

2nd eigenvector: 225.27 / 40.33

3rd eigenvalue: 0.02

3rd eigenvector: 12.51 / 44.73

VGP

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 22.65 (-22.65)

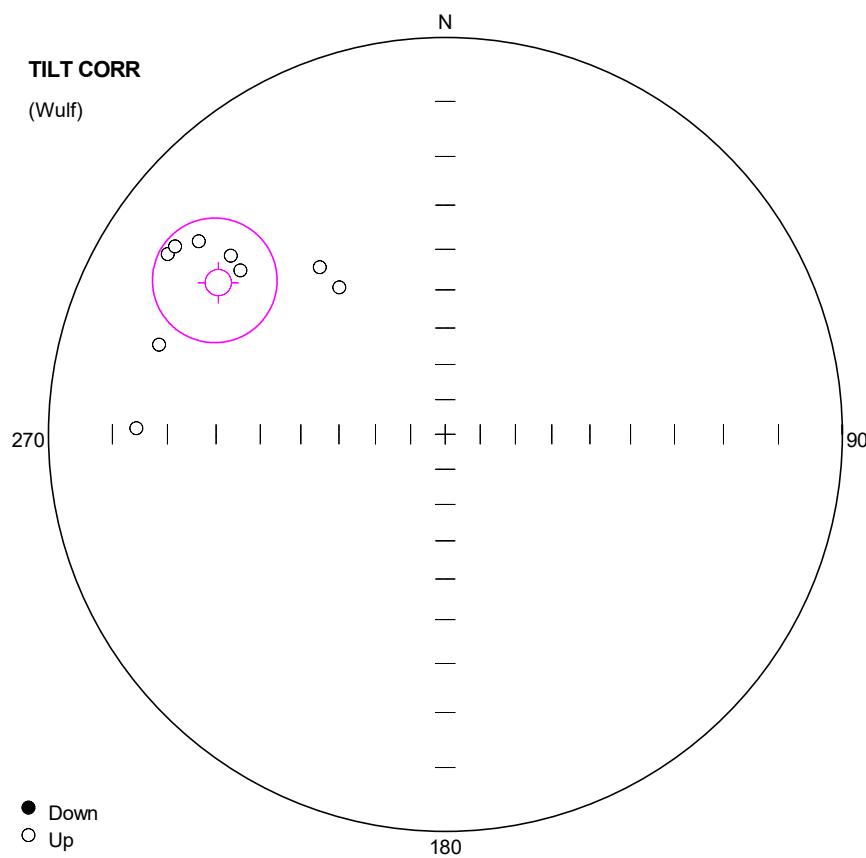
Pole longitude: -13.34 (166.66)

Paleolatitude: -8.81 (8.81)

dp: 6.48

dm: 12.53

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2401	CCH	313.8	-40.8	9.1	44	64
2	YUN2402	CCH	269.8	-5.0	4.7	35	52
3	YUN2403	CCH	304.1	-20.3	7.0	30	52
4	YUN2404	CCH	314.9	-34.3	5.2	44	60
5	YUN2405	CCH	301.3	-6.7	4.8	35	52
6	YUN2408	CCH	305.9	-17.3	16.0	20	48
7	YUN2409	CCH	303.0	-7.4	5.6	20	44
8	YUN2410	CCH	285.2	-9.2	4.0	20	52
9	YUN2412	CCH	305.6	-10.8	14.5	40	56



**Number of data points: 9**

**Fisher statistics**

Mean vector: 303.75 / -20.98  
Resultant vector: 8.58  
(X= 4.45, Y= -6.66, Z= -3.07)  
Estimated precision, k: 19.06  
95% Confidence limit: 12.1  
95% Confidence limit, approximation: 10.69

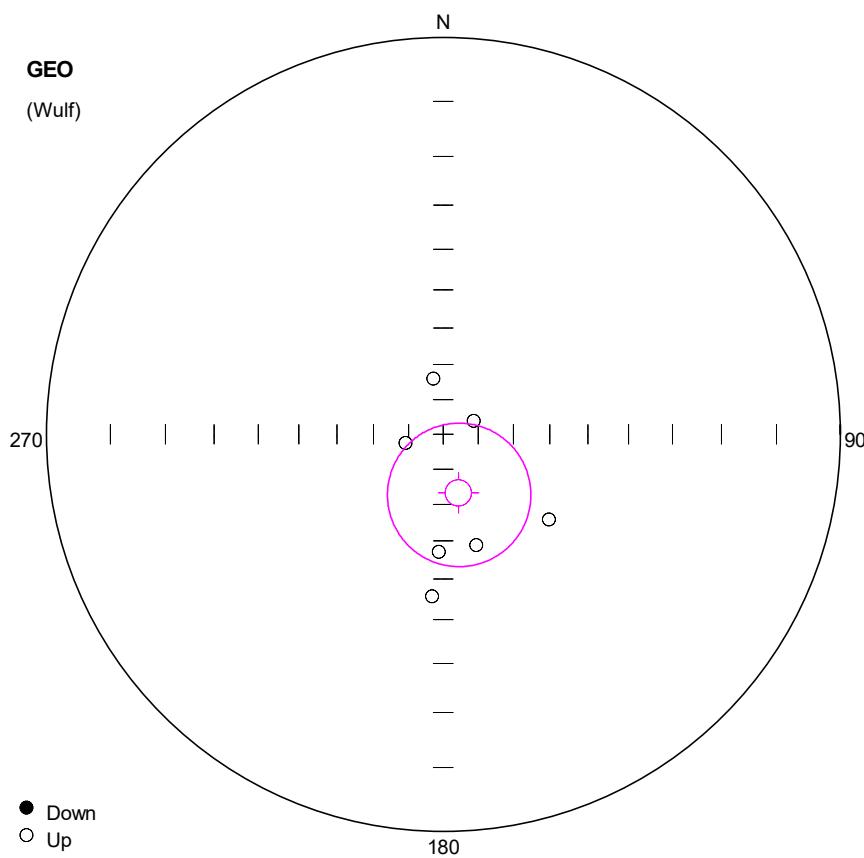
**Orientation matrix**

1st eigenvalue: 0.91  
1st eigenvector: 124.07 / 20.73  
2nd eigenvalue: 0.07  
2nd eigenvector: 225.86 / 28.35  
3rd eigenvalue: 0.02  
3rd eigenvector: 3.08 / 53.68

**VGP**

Site latitude: 24.27  
Site longitude: 98.39  
Pole latitude: 24.84 (-24.84)  
Pole longitude: -17.48 (162.52)  
Paleolatitude: -10.85 (10.85)  
dp: 6.69  
dm: 12.73

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2401	CCH	324.2	-41.0	9.1	44	64
2	YUN2402	CCH	271.1	-14.2	4.7	35	52
3	YUN2403	CCH	308.7	-23.1	7.0	30	52
4	YUN2404	CCH	323.1	-34.5	5.2	44	60
5	YUN2405	CCH	303.0	-10.4	4.8	35	52
6	YUN2408	CCH	309.8	-19.8	16.0	20	48
7	YUN2409	CCH	304.8	-10.7	5.6	20	44
8	YUN2410	CCH	287.4	-15.9	4.0	20	52
9	YUN2412	CCH	308.1	-13.5	14.5	40	56



**Number of data points: 7**

## Fisher statistics

Mean vector: 165.46 / -72.68

Resultant vector: 6.4

(X= -1.84, Y= 0.48, Z= -6.11)

Estimated precision, k: 10.06

95% Confidence limit: 20.01

95% Confidence limit, approx

## Orientation matrix

1st eigenvalue: 0.84

1st eigenvector: 345.64 / 72.25

2nd eigenvalue: 0.12

2nd eigenvector: 170.4 / 17.69

3rd eigenvalue: 0.04

3rd eigenvector: 79.96 / 1.38

VGP

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: -54.66 (54.66)

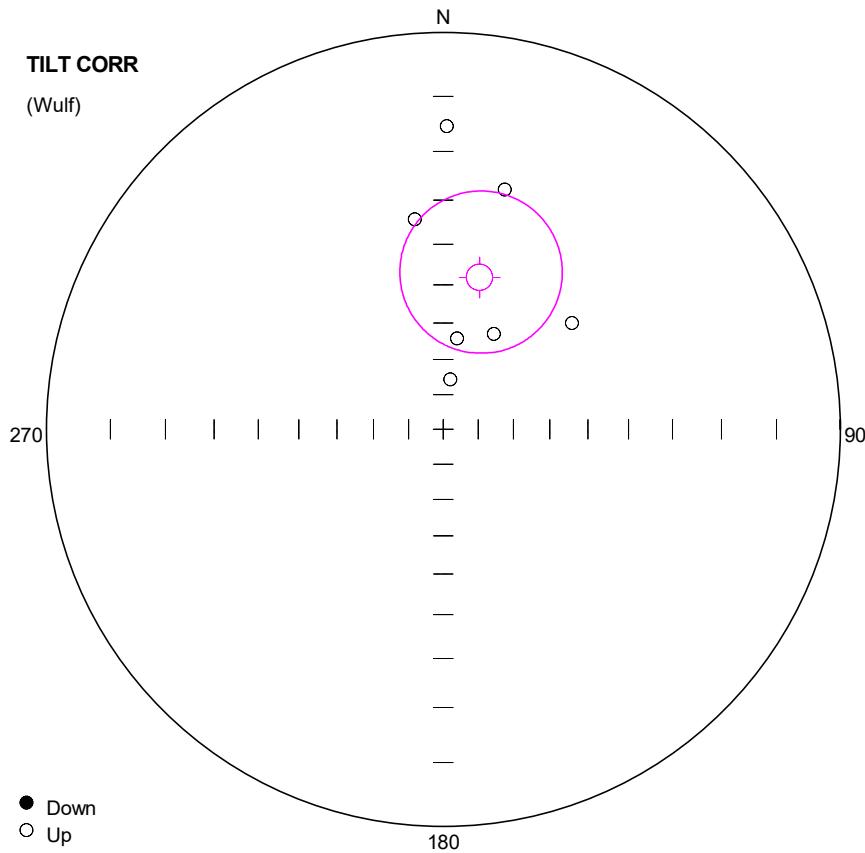
Pole longitude: -94.89 (85.11)

Paleolatitude: -58.05 (58.05)

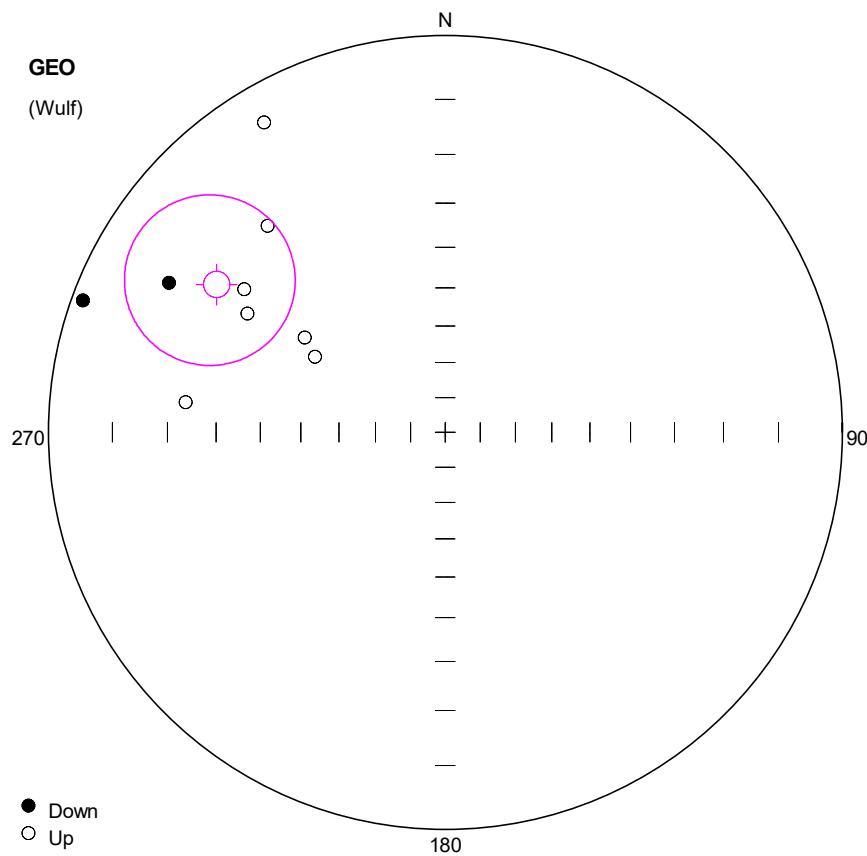
dp: 31.61

dm: 35.56

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2503	CCH	257.1	-78.9	14.8	35	64 O
2	YUN2504	CCH	182.1	-57.0	11.9	30	64 O
3	YUN2505	CCH	66.4	-80.4	6.7	30	64 O
4	YUN2507	CCH	183.9	-45.5	15.3	30	64 O
5	YUN2509	CCH	128.9	-52.2	14.3	30	60 O
6	YUN2510	CCH	163.4	-57.5	15.7	30	60 O
7	YUN2511	CCH	349.8	-73.8	7.2	35	60 O



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2503	CCH	352.3	-33.8	14.8	35	64 C
2	YUN2504	CCH	8.7	-63.9	11.9	30	64 C
3	YUN2505	CCH	14.4	-26.1	6.7	30	64 C
4	YUN2507	CCH	8.0	-75.5	15.3	30	64 C
5	YUN2509	CCH	50.4	-44.4	14.3	30	60 C
6	YUN2510	CCH	28.0	-59.6	15.7	30	60 C
7	YUN2511	CCH	0.7	-15.3	7.2	35	60 C



**Number of data points: 9**

## Fisher statistics

Mean vector: 302.88 / -21.14

Resultant vector: 8.25

(X= 4.18, Y= -6.46, Z= -2.98)

Estimated precision, k: 10.72

95% Confidence limit: 16.48

95% Confidence limit, approximation: 14.25

## Orientation matrix

1st eigenvalue: 0.84

1st eigenvector: 122.92 / 21.82

2nd eigenvalue: 0.1

2nd eigenvector: 306.59 / 68.14

3rd eigenvalue: 0.06

3rd eigenvector: 213.43 / 1.27

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 3.05 (-3.05)

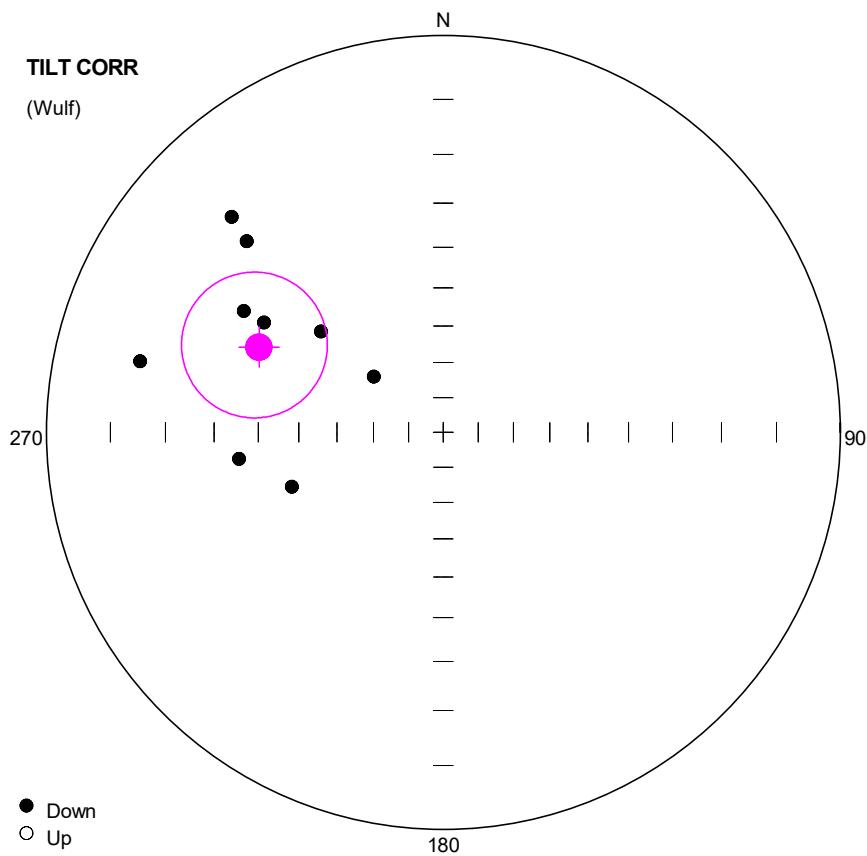
Pole longitude: -110.3 (69.7)

Paleolatitude: -10.94 (10.94)

dp: 9.13

dm: 17.35

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2901	CCH	298.4	13.3	14.2	30	56
2	YUN2902	CCH	301.0	-29.7	12.0	30	60
3	YUN2903	CCH	319.3	-21.1	20.2	35	60
4	YUN2904	CCH	304.0	-43.8	8.4	35	60
5	YUN2905	CCH	329.7	-5.8	9.5	44	64
6	YUN2906	CCH	290.0	1.7	12.8	30	56
7	YUN2908	CCH	300.1	-48.5	18.7	30	64
8	YUN2909	CCH	276.6	-23.3	18.1	30	52
9	YUN2911	CCH	305.4	-26.3	8.1	40	52



**Number of data points: 9**

## Fisher statistics

Mean vector: 294.84 / 35.85

Resultant vector: 8.25

(X= 2.81, Y= -6.07, Z= 4.83)

Estimated precision, k: 10.72

95% CI

95% Confidence limit, approximation: 14.25

## Orientation matrix

1st eigenvalue: 0.84

1st eigenvector: 295.52 / 35.46

2nd eigenvalue: 0.1

2nd eigenvector: 179.32 / 31.79

3rd eigenvalue: 0.05

3rd eigenvector: 59.89 / 38.4

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 28.35 (-28.35)

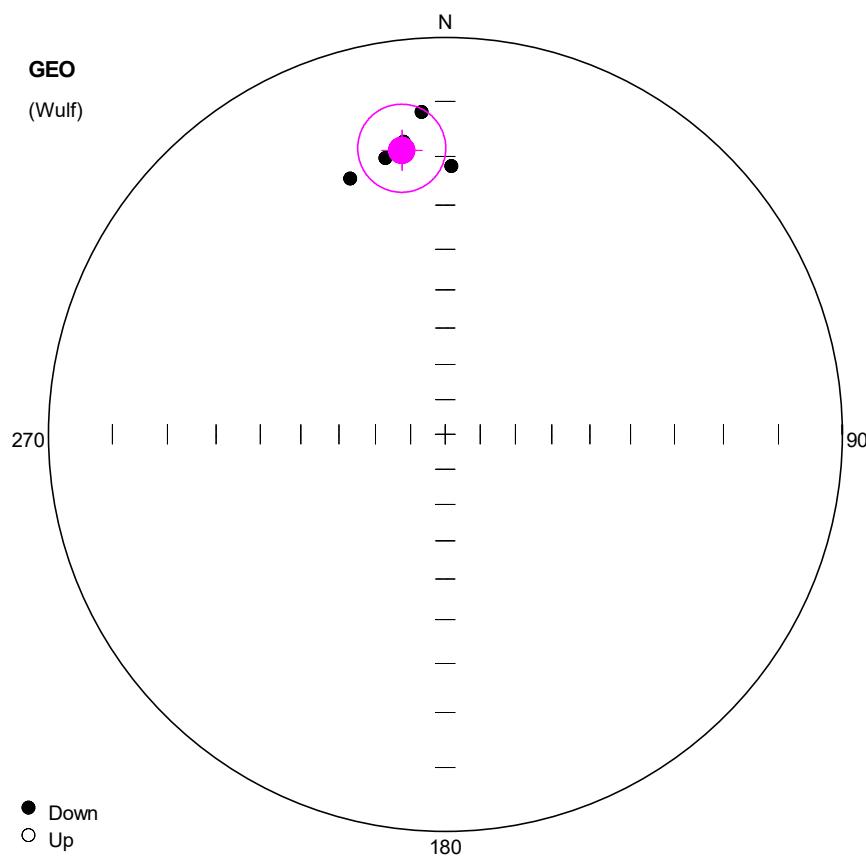
Pole longitude: -90.06 (89.94)

Paleolatitude: 19.86 (-19.86)

dp: 11.09

dm: 19.12

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN2901	CCH	250.2	45.9	14.2	30	56
2	YUN2902	CCH	301.3	29.1	12.0	30	60
3	YUN2903	CCH	309.5	46.5	20.2	35	60
4	YUN2904	CCH	314.2	20.8	8.4	35	60
5	YUN2905	CCH	308.7	64.8	9.5	44	64
6	YUN2906	CCH	262.6	35.2	12.8	30	56
7	YUN2908	CCH	315.5	15.5	18.7	30	64
8	YUN2909	CCH	283.2	13.8	18.1	30	52
9	YUN2911	CCH	301.5	34.2	8.1	40	52



**Number of data points: 5**

**Fisher statistics**

Mean vector: 351.31 / 18.22

Resultant vector: 4.95

(X= 4.65, Y= -0.71, Z= 1.55)

Estimated precision, k: 85.87

95% Confidence limit: 8.3

95% Confidence limit, approximation: 6.76

**Orientation matrix**

1st eigenvalue: 0.98

1st eigenvector: 351.32 / 18.21

2nd eigenvalue: 0.01

2nd eigenvector: 258.15 / 9.55

3rd eigenvalue: 0

3rd eigenvector: 141.74 / 69.28

**VGP**

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 72.93 (-72.93)

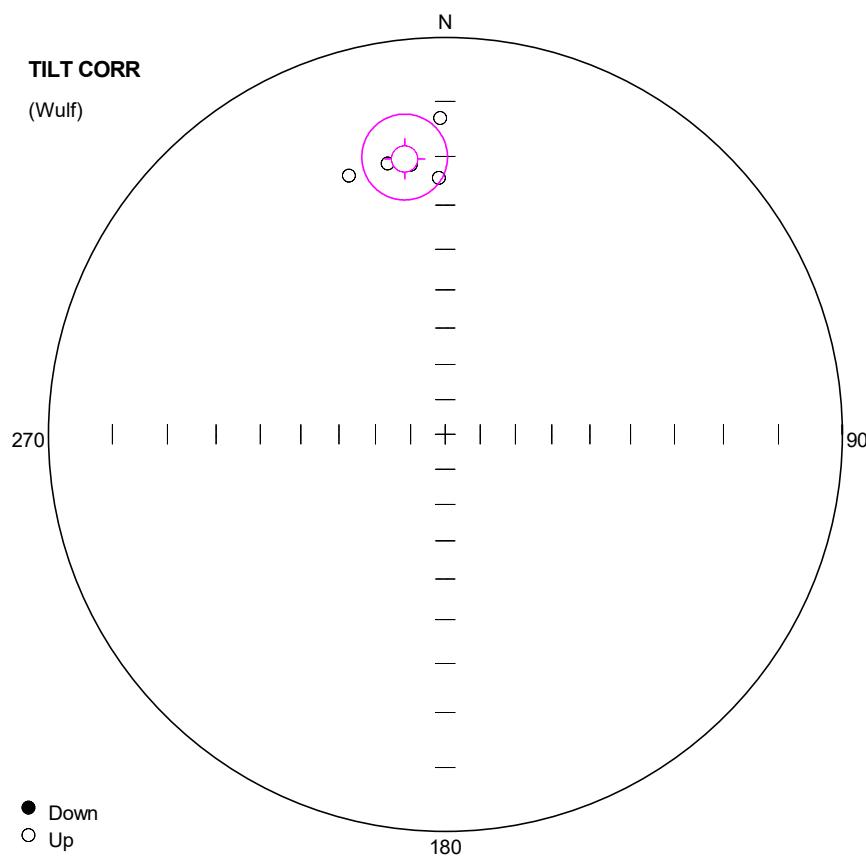
Pole longitude: -51.09 (128.91)

Paleolatitude: 9.35 (-9.35)

dp: 4.48

dm: 8.63

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3501	CCH	351.9	16.7	4.8	30	68
2	YUN3502	CCH	355.8	11.7	4.4	30	68
3	YUN3503	CCH	1.3	21.9	4.7	30	68
4	YUN3504	CCH	339.6	21.0	5.6	40	68
5	YUN3505	CCH	347.8	19.1	5.7	30	68



**Number of data points: 5**

**Fisher statistics**

Mean vector: 351.63 / -19.95

Resultant vector: 4.95

(X= 4.61, Y= -0.68, Z= -1.69)

Estimated precision, k: 86.77

95% Confidence limit: 8.26

95% Confidence limit, approximation: 6.72

**Orientation matrix**

1st eigenvalue: 0.98

1st eigenvector: 171.64 / 19.96

2nd eigenvalue: 0.01

2nd eigenvector: 78.04 / 9.83

3rd eigenvalue: 0

3rd eigenvector: 323.24 / 67.57

**VGP**

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: 54.49 (-54.49)

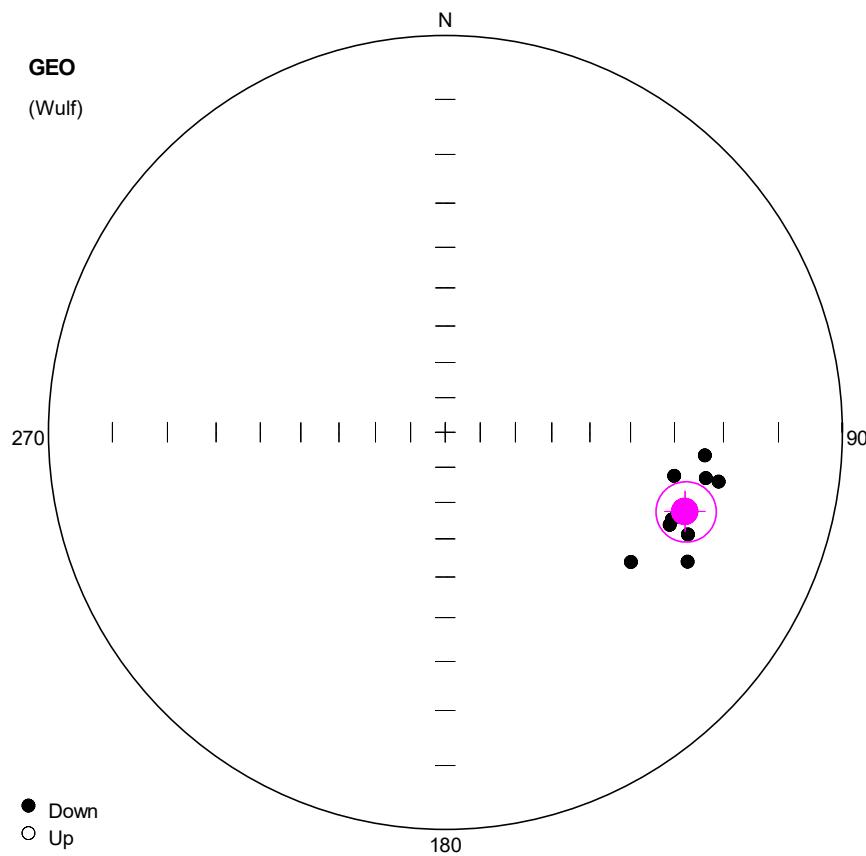
Pole longitude: -67.33 (112.67)

Paleolatitude: -10.29 (10.29)

dp: 4.53

dm: 8.65

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3501	CCH	352.7	-21.2	4.8	30	68
2	YUN3502	CCH	358.6	-24.3	4.4	30	68
3	YUN3503	CCH	359.1	-12.9	4.7	30	68
4	YUN3504	CCH	339.6	-20.4	5.6	40	68
5	YUN3505	CCH	348.0	-20.2	5.7	30	68



**Number of data points: 9**

## Fisher statistics

Mean vector: 108.26 / 25.12

Resultant vector: 8.89

(X= -2.52, Y= 7.64, Z= 3.77)

Estimated precision, k: 70.24

95% C

95% Confidence limit, approximation: 5.57

## Orientation matrix

1st eigenvalue: 0.97

1st eigenvector: 108.23 / 25.11

2nd eigenvalue: 0.02

2nd eigenvector: 202.68 / 9.39

3rd eigenvalue: 0

3rd eigenvector: 311.58 / 62.96

VGP

Site latitude: 24.27

Site longitude: 98.39

Pole latitude: -10.62 (10.62)

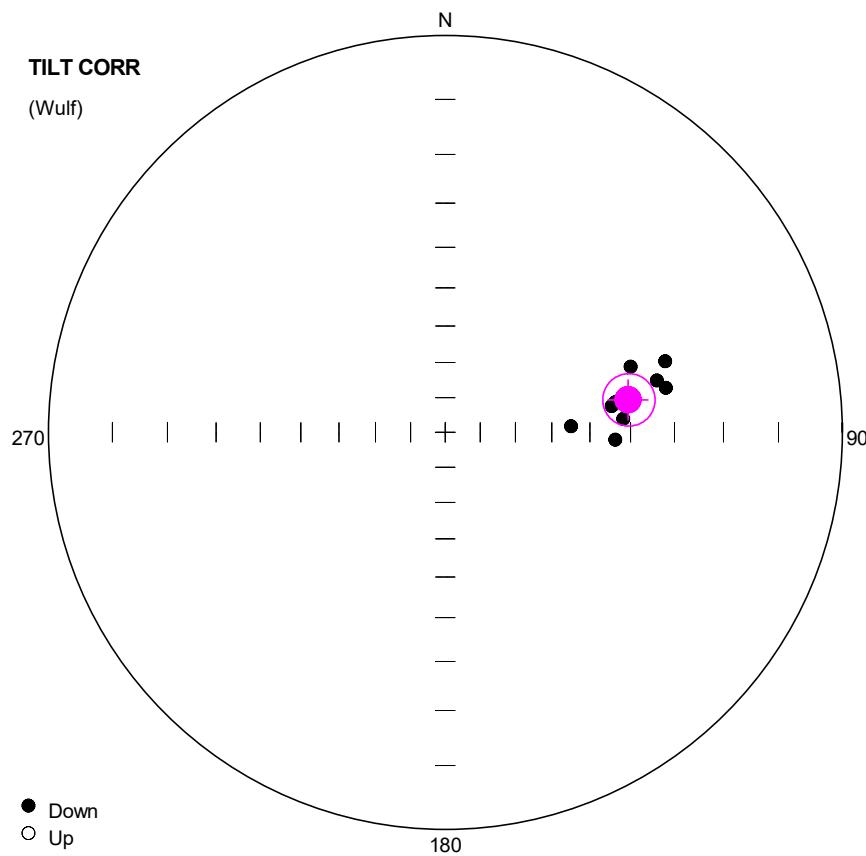
Pole longitude: 168.56 (-11.44)

Paleolatitude: 13.19 (-13.19)

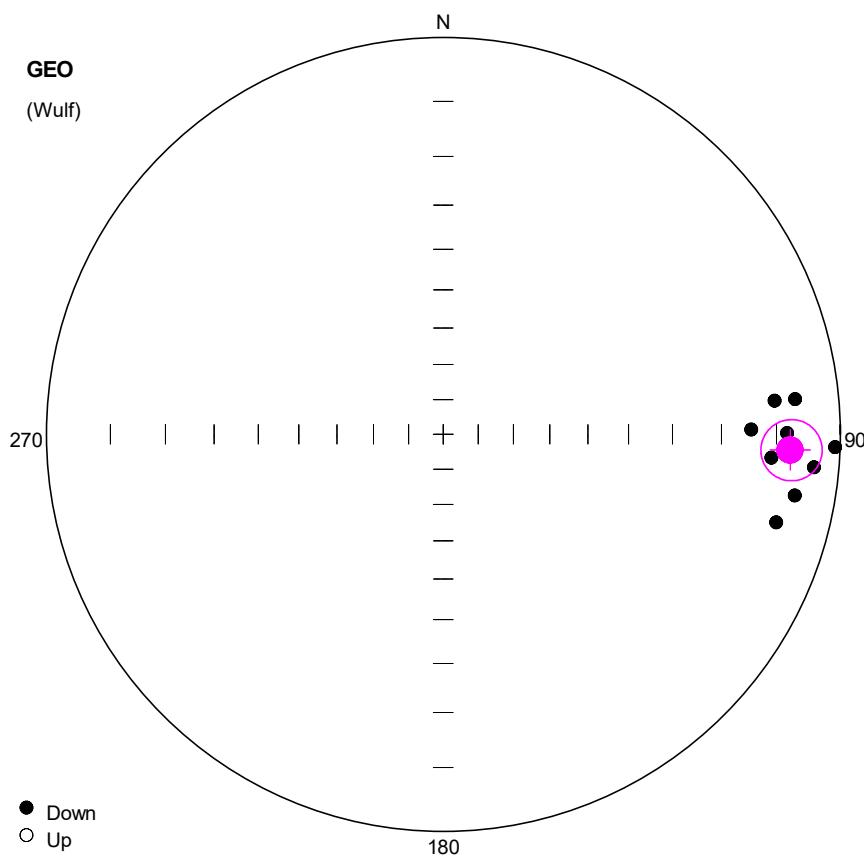
dp: 3.58

dm: 6.65

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3601	CCH	112.4	27.1	6.4	35	64
2	YUN3602	CCH	100.8	29.2	4.5	44	64
3	YUN3603	CCH	112.8	22.9	7.4	35	64
4	YUN3604	CCH	111.0	27.1	8.2	40	64
5	YUN3605	CCH	118.1	20.6	5.3	30	64
6	YUN3606	CCH	95.1	23.4	5.3	35	64
7	YUN3608	CHT	124.9	30.6	6.8	60	68 OA
8	YUN3609	CCH	100.0	22.6	5.0	40	64
9	YUN3610	CCH	100.2	20.0	8.2	35	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3601	CCH	81.1	44.0	6.4	35	64
2	YUN3602	CCH	70.5	37.3	4.5	44	64
3	YUN3603	CCH	85.7	41.6	7.4	35	64
4	YUN3604	CCH	80.0	43.0	8.2	40	64
5	YUN3605	CCH	92.5	43.6	5.3	30	64
6	YUN3606	CCH	72.1	29.6	5.3	35	64
7	YUN3608	CHT	87.3	54.8	6.8	60	68 OA
8	YUN3609	CCH	76.2	32.5	5.0	40	64
9	YUN3610	CCH	78.6	30.9	8.2	35	64



**Number of data points: 9**

## Fisher statistics

Mean vector: 92.59 / 7.66

Resultant vector: 8.93

(X= -0.4, Y= 8.84, Z= 1.19)

Estimated precision, k: 107

95% Confidence limit: 5

95% Confidence limit, approximation: 4.51

## Orientation matrix

1st eigenvalue: 0.98

1st eigenvector: 92.57 / 7.66

2nd eigenvalue: 0.01

2nd eigenvector: 1.28 / 9.51

3rd eigenvalue: 0

3rd eigenvector: 220.82 / 77.75

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 2.39 (-2.39)

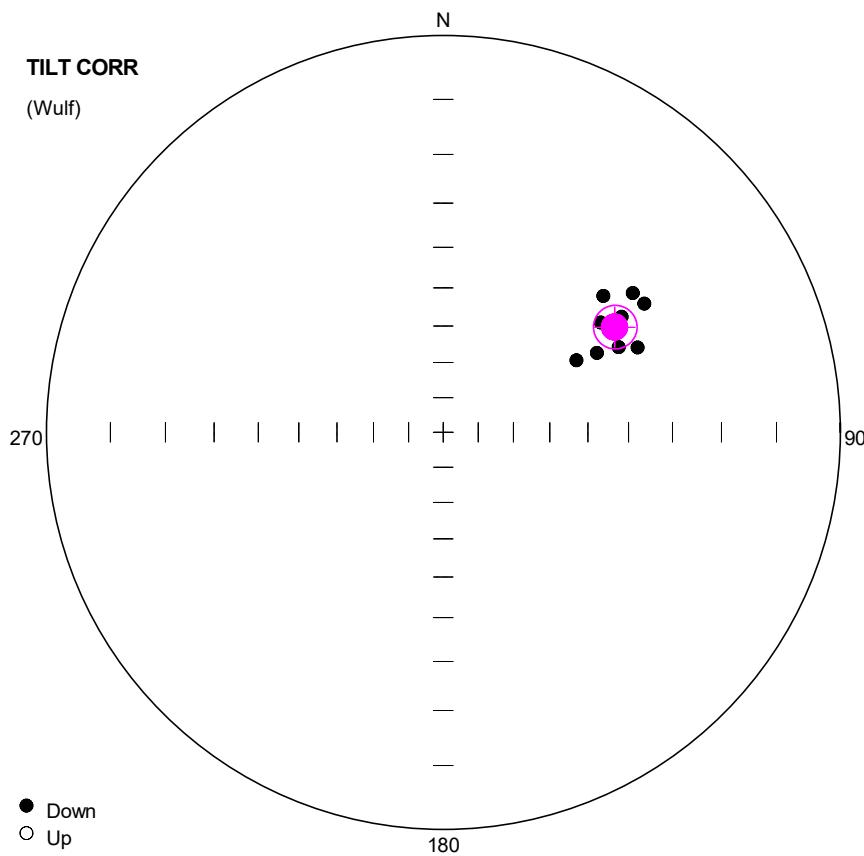
Pole longitude: 100.07 (-79.93)

Paleolatitude: 3.85 (-3.85)

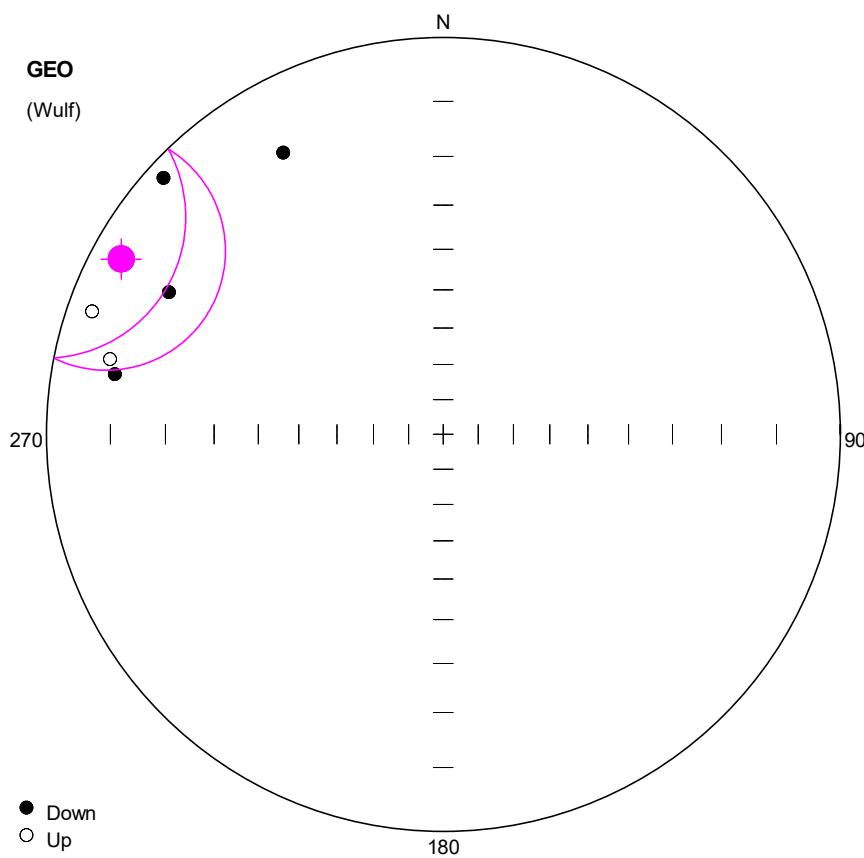
dp: 2.53

dm: 5.03

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3701	CCH	104.8	8.1	4.8	40	64
2	YUN3702	CCH	84.2	10.0	7.3	40	64
3	YUN3704	CCH	89.1	14.4	5.9	44	64
4	YUN3705	CCH	99.9	6.1	8.9	40	64
5	YUN3706	CCH	94.1	10.7	9.3	40	64
6	YUN3707	CCH	89.8	8.2	6.4	35	60
7	YUN3708	CCH	84.3	6.6	6.1	40	64
8	YUN3709	CCH	95.1	3.7	7.2	35	64
9	YUN3710	CCH	91.9	0.7	5.4	35	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3701	CCH	61.6	48.2	4.8	40	64
2	YUN3702	CCH	53.7	28.7	7.3	40	64
3	YUN3704	CCH	49.6	34.2	5.9	44	64
4	YUN3705	CCH	62.7	42.9	8.9	40	64
5	YUN3706	CCH	55.1	38.4	9.3	40	64
6	YUN3707	CCH	57.1	33.7	6.4	35	60
7	YUN3708	CCH	57.4	28.0	6.1	40	64
8	YUN3709	CCH	64.1	37.7	7.2	35	64
9	YUN3710	CCH	66.4	33.8	5.4	35	64



**Number of data points: 6**

## Fisher statistics

Mean vector: 298.56 / 4.56

Resultant vector: 5.66

(X= 2.7, Y= -4.96, Z= 0.45)

Estimated precision, k: 14.75

95% C

95% Confidence limit, approximation: 14.88

## Orientation matrix

1st eigenvalue: 0.89

1st eigenvector: 297.97 / 4.44

2nd eigenvalue: 0.09

2nd eigenvector: 29.16 / 14.94

3rd eigenvalue: 0.02

3rd eigenvector: 191.85 / 74.39

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 13.77 (-13.77)

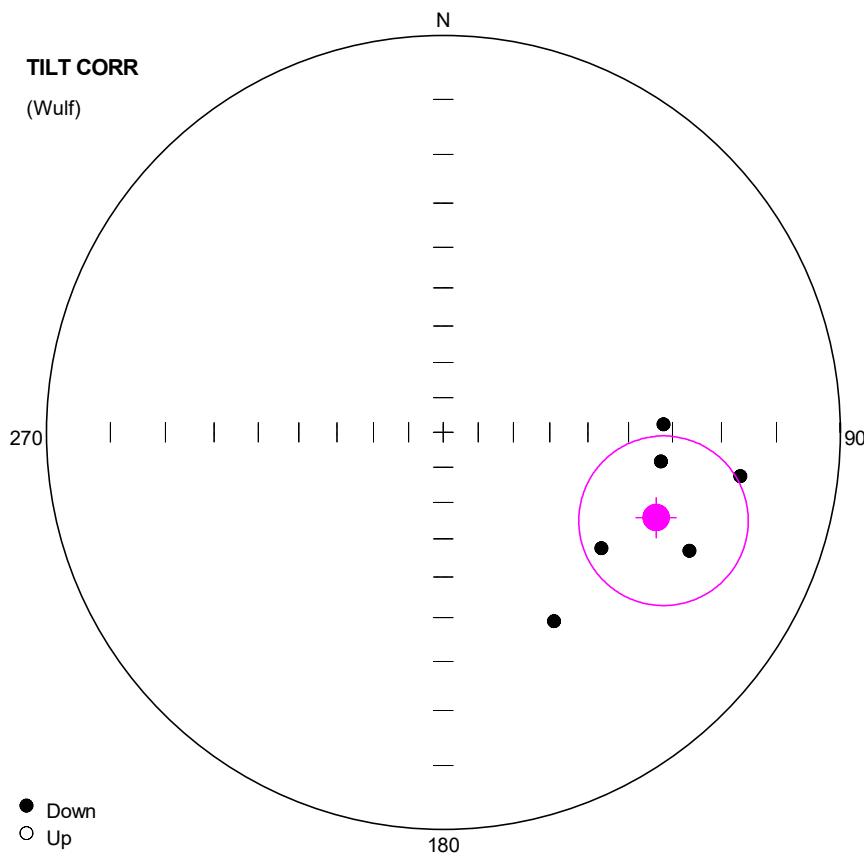
Pole longitude: -101.33 (78.67)

Paleolatitude: 2.28 (-2.28)

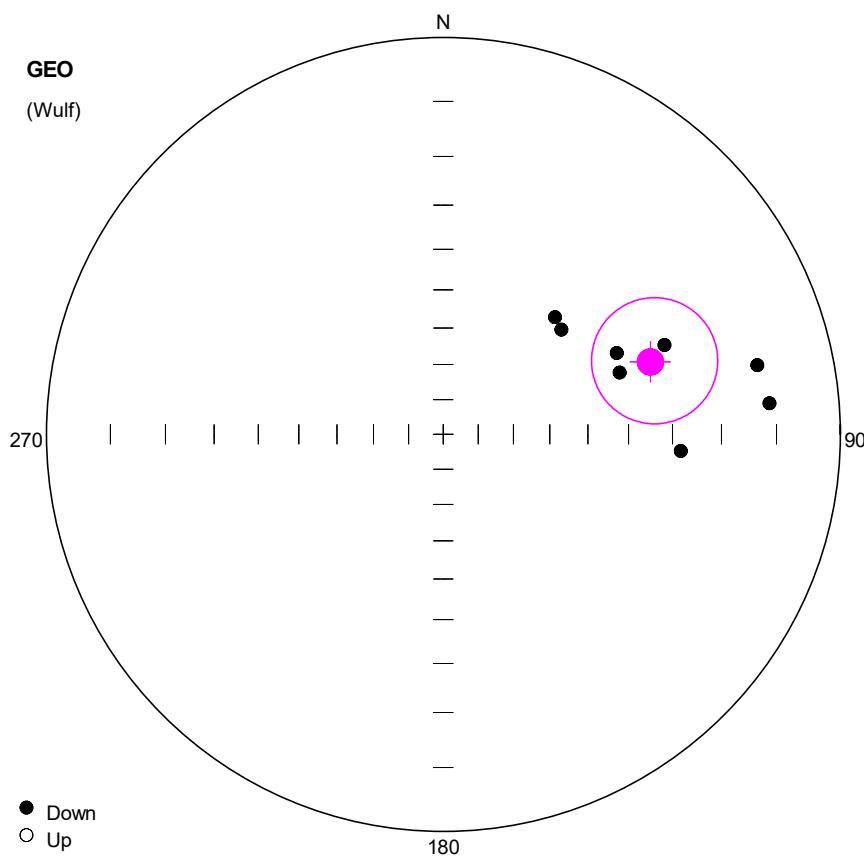
dp: 9.06

dm: 18.08

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3803	CCH	297.4	14.2	5.0	30	64
2	YUN3804	CCH	312.5	2.6	11.2	30	52
3	YUN3805	CCH	289.3	-3.7	19.1	40	56
4	YUN3807	CCH	330.4	11.6	6.9	40	64
5	YUN3808	CCH	282.7	-8.6	2.6	30	56
6	YUN3809	CCH	280.4	9.9	2.3	35	56



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3803	CCH	115.7	20.9	5.0	30	64
2	YUN3804	CCH	126.2	37.4	11.2	30	52
3	YUN3805	CCH	97.6	32.1	19.1	40	56
4	YUN3807	CCH	149.6	32.2	6.9	40	64
5	YUN3808	CCH	87.9	31.9	2.6	30	56
6	YUN3809	CCH	98.4	15.8	2.3	35	56



**Number of data points: 8**

## Fisher statistics

Mean vector: 70.74 / 32.15

Resultant vector: 7.59

(X= 2.12, Y= 6.07, Z= 4.04)

Estimated precision, k: 17.03

95% Confidence limit: 13.82

95% Confidence limit, approximation: 12

## Orientation matrix

1st eigenvalue: 0.9

1st eigenvector: 70.72 / 32.23

2nd eigenvalue: 0.09

2nd eigenvector: 310.35 / 38.73

3rd eigenvalue: 0.01

3rd eigenvector: 186.61 / 34.71

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 23.87 (-23.87)

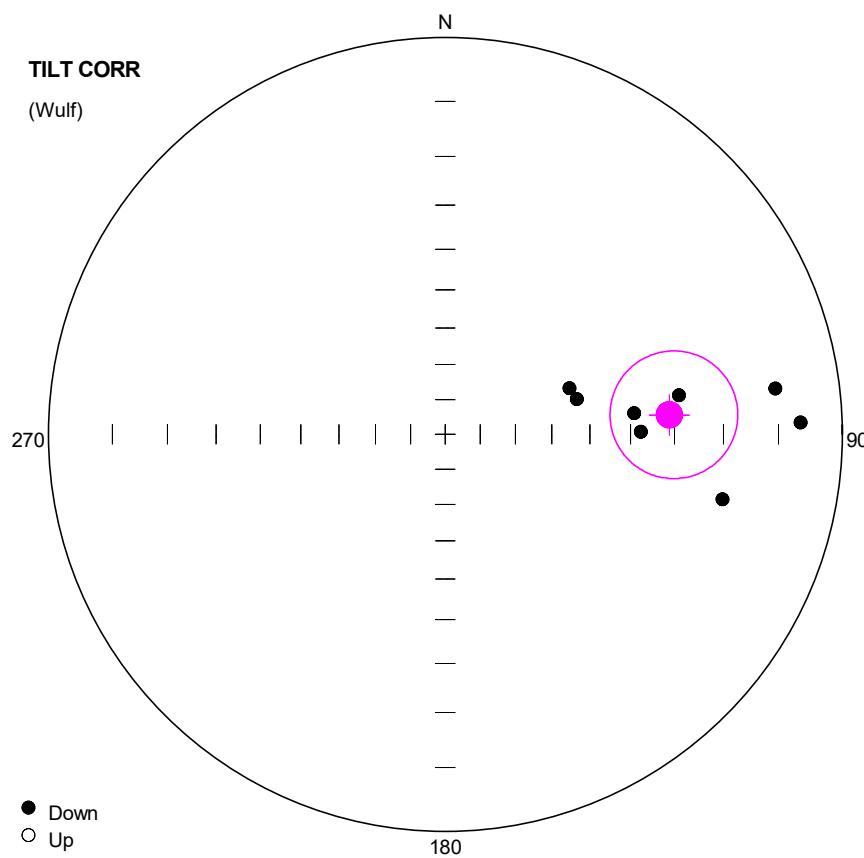
Pole longitude: 114.02 (-65.98)

Paleolatitude: 17.45 (-17.45)

dp: 8.77

dm: 15.57

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3901	CCH	43.7	45.6	3.4	48	64
2	YUN3902	CCH	48.5	46.6	3.6	44	64
3	YUN3903	CCH	68.0	28.0	4.4	40	64
4	YUN3904	CCH	84.6	10.9	12.0	30	52
5	YUN3905	CCH	64.9	38.5	13.3	40	64
6	YUN3906	CCH	77.6	12.0	4.4	35	60
7	YUN3907	CCH	70.7	39.6	11.7	48	60
8	YUN3908	CCH	94.0	28.1	7.7	40	56



Number of data points: 8

#### Fisher statistics

Mean vector: 85.06 / 30.95

Resultant vector: 7.59

(X= 0.56, Y= 6.48, Z= 3.9)

Estimated precision, k: 17.01

95% Confidence limit: 13.83

95% Confidence limit, approximation: 12

#### Orientation matrix

1st eigenvalue: 0.9

1st eigenvector: 85.08 / 31.02

2nd eigenvalue: 0.09

2nd eigenvector: 293.01 / 55.76

3rd eigenvalue: 0.01

3rd eigenvector: 183.1 / 13.05

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 17.16 (-17.16)

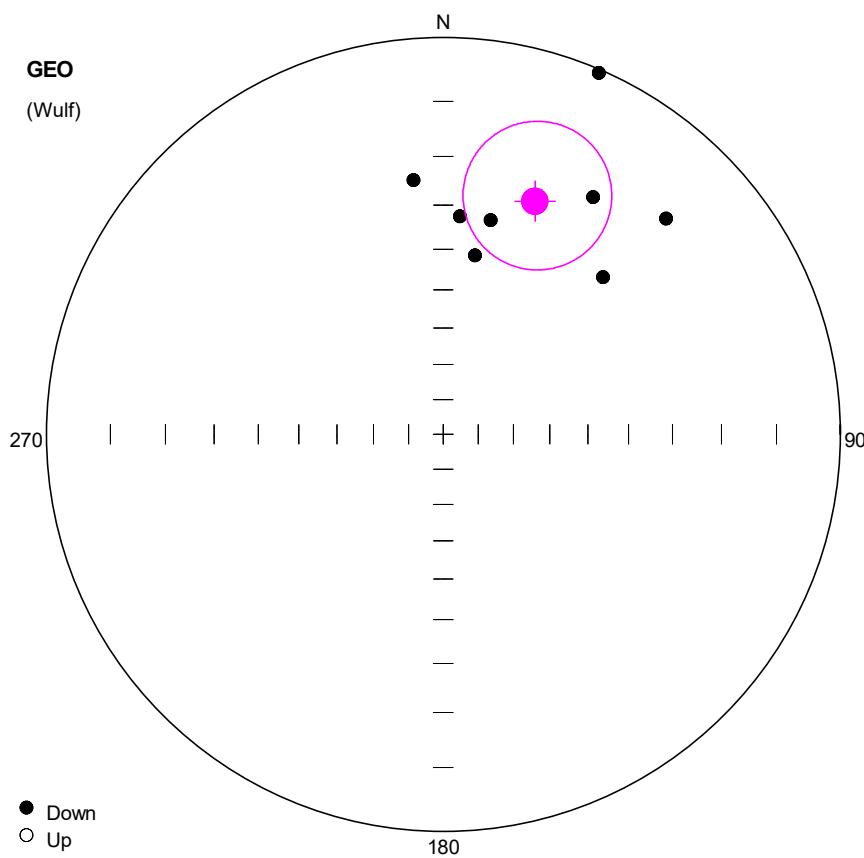
Pole longitude: 101.24 (-78.76)

Paleolatitude: 16.69 (-16.69)

dp: 8.62

dm: 15.44

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN3901	CCH	69.7	53.1	3.4	48	64
2	YUN3902	CCH	75.1	52.1	3.6	44	64
3	YUN3903	CCH	80.5	28.3	4.4	40	64
4	YUN3904	CCH	88.1	6.3	12.0	30	52
5	YUN3905	CCH	83.6	38.8	13.3	40	64
6	YUN3906	CCH	82.1	10.0	4.4	35	60
7	YUN3907	CCH	89.3	37.5	11.7	48	60
8	YUN3908	CCH	103.2	18.7	7.7	40	56



**Number of data points: 8**

## Fisher statistics

Mean vector: 21.47 / 25.51

Resultant vector: 7.51

(X= 6.31, Y= 2.48, Z= 3.23)

Estimated precision, k: 14.34

95% Confidence limit: 15.14

95% Confidence limit, approximation: 13.07

## Orientation matrix

1st eigenvalue: 0.88

1st eigenvector: 21.37 / 25.72

2nd eigenvalue: 0.08

2nd eigenvector: 280.19 / 21.91

3rd eigenvalue: 0.03

3rd eigenvector: 155.01 / 55.08

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 36.36 (-36.36)

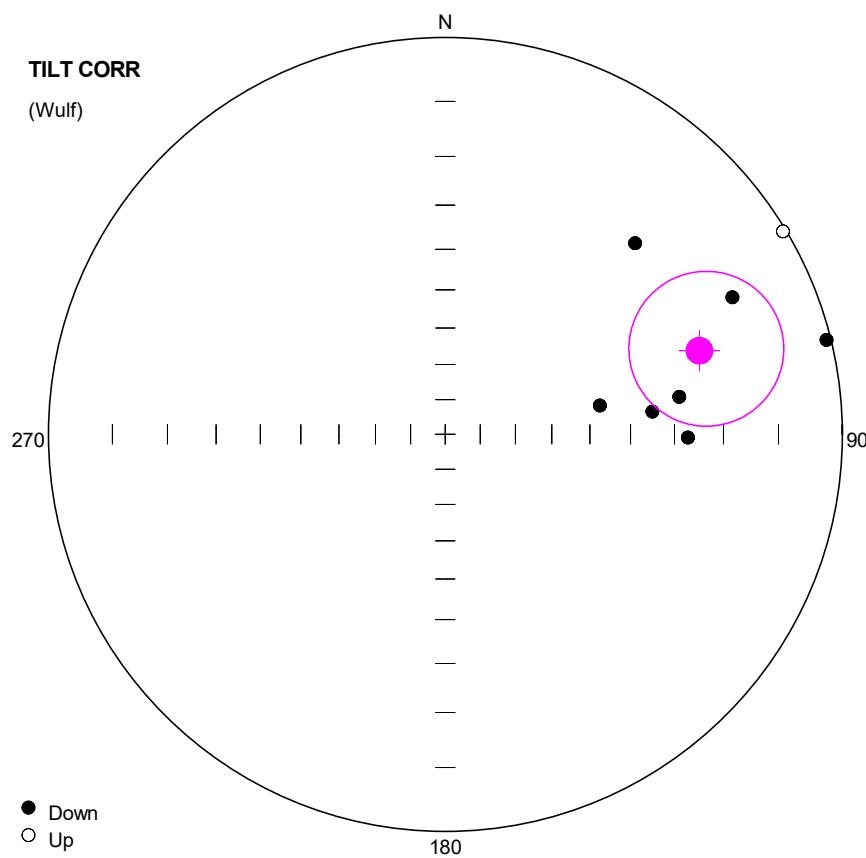
Pole longitude: 167.8 (-12.2)

Paleolatitude: 13.42 (-13.42)

dp: 8.79

dm: 16.32

#	Name	State	Dec	Inc	MAD	Limit1	Limit2	
1	YUN4001	CCH	23.3	0.5	12.9	30	64	
2	YUN4003	CCH		4.4	32.3	9.8	40	60
3	YUN4004	CCH		10.1	40.8	15.0	40	60
4	YUN4005	CCH		32.3	19.5	16.4	30	60
5	YUN4006	CCH	353.3	24.4	8.6	30	64	
6	YUN4007	CCH	12.5	32.2	8.6	30	64	
7	YUN4009	CCH	45.5	31.1	7.7	30	60	
8	YUN4010	CCH	45.9	14.0	15.0	35	64	



Number of data points: 8

#### Fisher statistics

Mean vector: 71.8 / 22.03

Resultant vector: 7.51

(X= 2.18, Y= 6.61, Z= 2.82)

Estimated precision, k: 14.32

95% Confidence limit: 15.15

95% Confidence limit, approximation: 13.08

#### Orientation matrix

1st eigenvalue: 0.88

1st eigenvector: 72.04 / 22.09

2nd eigenvalue: 0.08

2nd eigenvector: 191.81 / 50.73

3rd eigenvalue: 0.03

3rd eigenvector: 328.15 / 30.6

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 18.01 (-18.01)

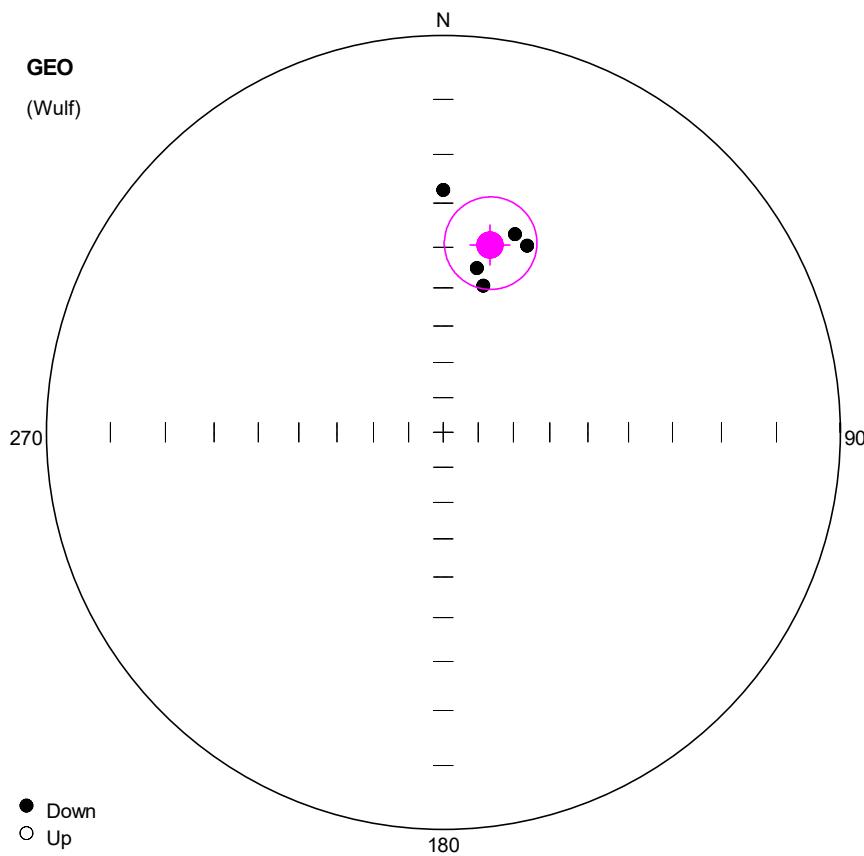
Pole longitude: 115.79 (-64.21)

Paleolatitude: 11.44 (-11.44)

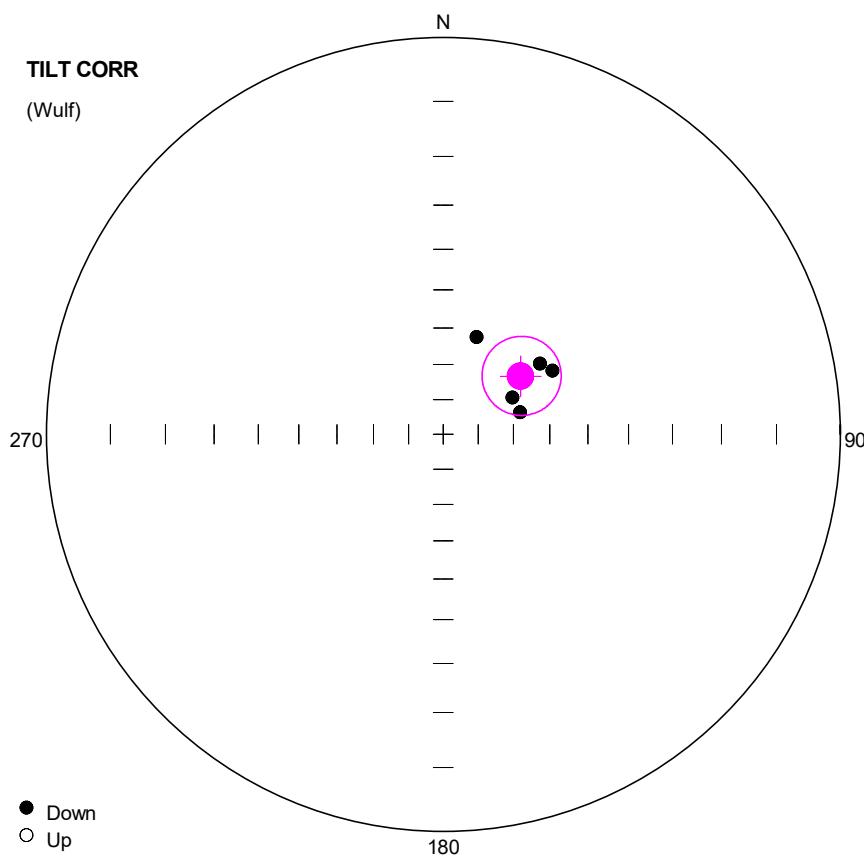
dp: 8.47

dm: 16.02

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4001	CCH	44.8	21.7	12.9	30	64
2	YUN4003	CCH	83.8	34.6	9.8	40	60
3	YUN4004	CCH	90.8	27.1	15.0	40	60
4	YUN4005	CCH	64.5	12.6	16.4	30	60
5	YUN4006	CCH	79.5	46.8	8.6	30	64
6	YUN4007	CCH	80.9	28.3	8.6	30	64
7	YUN4009	CCH	76.1	0.6	7.7	30	60
8	YUN4010	CCH	59.0	-0.4	15.0	35	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4201	Cch	15.3	48.1	9.0	40	68 OA
2	YUN4202	Cch	0.0	27.2	8.3	35	64 OA
3	YUN4203	Cch	11.6	44.2	4.8	40	64 OA
4	YUN4204	Cch	24.2	35.5	13.5	40	60
5	YUN4205	Cch	19.9	34.1	8.8	35	68



**Number of data points: 5**

## Fisher statistics

Mean vector: 53.1 / 62.61

Resultant vector: 4.92

(X= 1.36, Y= 1.81, Z= 4.37)

Estimated precision, k: 51.73

95% C

95% Confidence limit, approximation: 8.7

## Orientation matrix

1st eigenvalue: 0.97

1st eigenvector: 53.29 / 62.59

2nd eigenvalue: 0.02

2nd eigenvector: 320.54 / 1.43

3rd eigenvalue: 0.01

3rd eigenvector: 229.8 / 27.37

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 54.28 (-54.28)

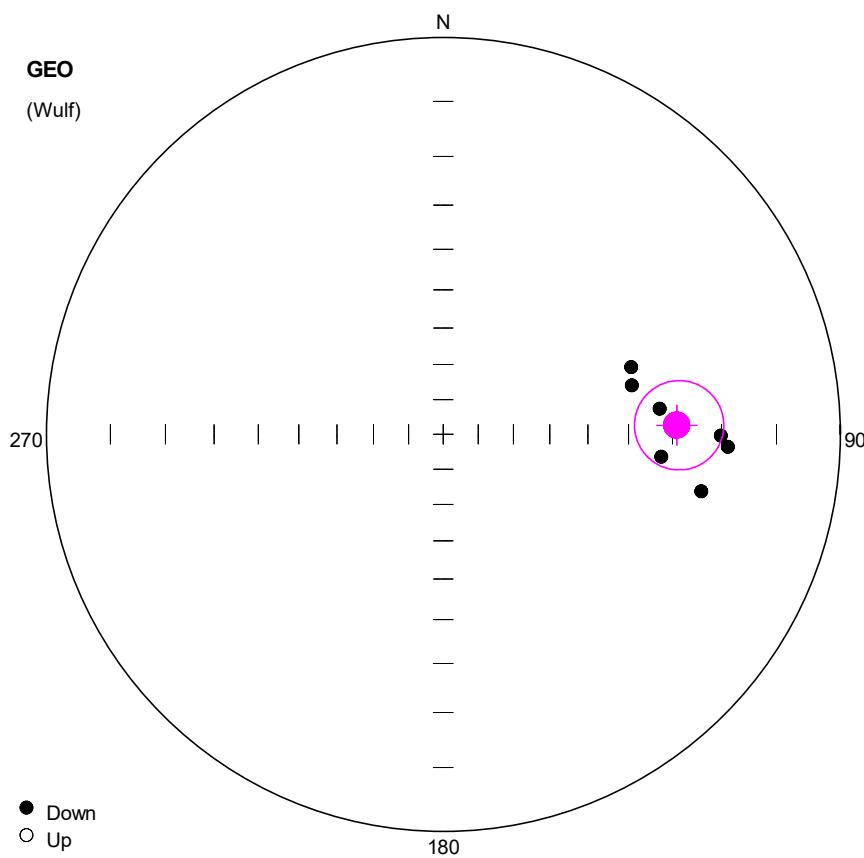
Pole longitude: 113.72 (-66.28)

Paleolatitude: 43.98 (-43.98)

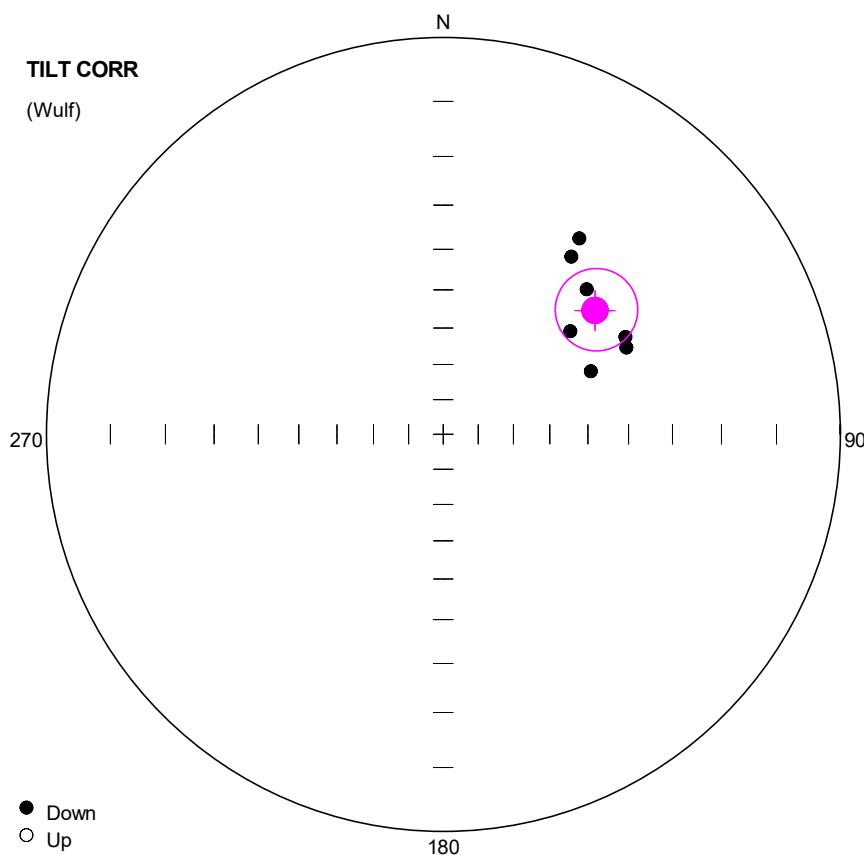
dp: 13.13

dm: 16.79

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4201	Cch	74.1	67.2	9.0	40	68 OA
2	YUN4202	Cch	18.9	61.0	8.3	35	64 OA
3	YUN4203	Cch	62.1	67.7	4.8	40	64 OA
4	YUN4204	Cch	59.8	54.7	13.5	40	60
5	YUN4205	Cch	53.8	56.4	8.8	35	68



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4301	CCH	70.3	36.6	2.3	40	68
2	YUN4302	CCH	83.3	32.5	2.9	44	68
3	YUN4303	CCH	90.3	20.1	4.2	40	68
4	YUN4304	CCH	75.5	37.7	6.0	40	68
5	YUN4305	CCH	95.9	32.2	3.1	44	68
6	YUN4306	CCH	92.5	18.7	3.1	30	68
7	YUN4307	CCH	102.5	22.7	7.1	30	60



**Number of data points: 7**

## Fisher statistics

Mean vector: 50.79 / 37.51

Resultant vector: 6.85

(X= 3.44, Y= 4.21, Z= 4.17)

Estimated precision, k: 41.06

95% Confidence limit: 9.53

95% Confidence limit, approximation: 8.26

## Orientation matrix

1st eigenvalue: 0.96

1st eigenvector: 50.83 / 37.53

2nd eigenvalue: 0.04

2nd eigenvector: 155.97 / 18.78

3rd eigenvalue: 0.01

3rd eigenvector: 266.93 / 46.45

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 35.04 (-35.04)

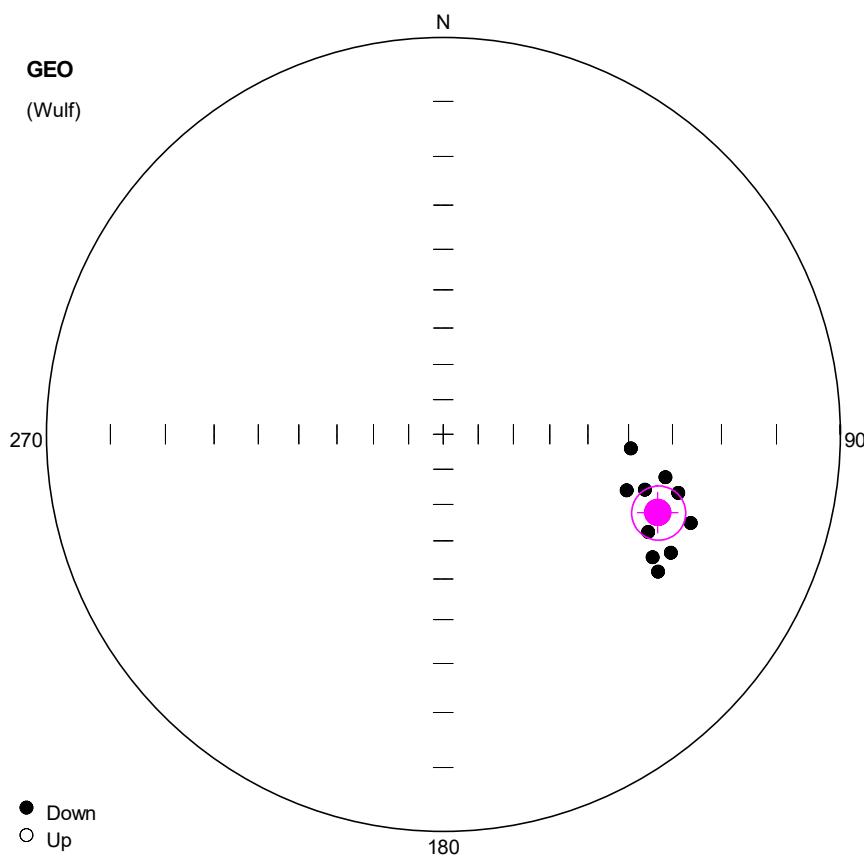
Pole longitude: 131.97 (-48.0)

Paleolatitude: 21 (-21)

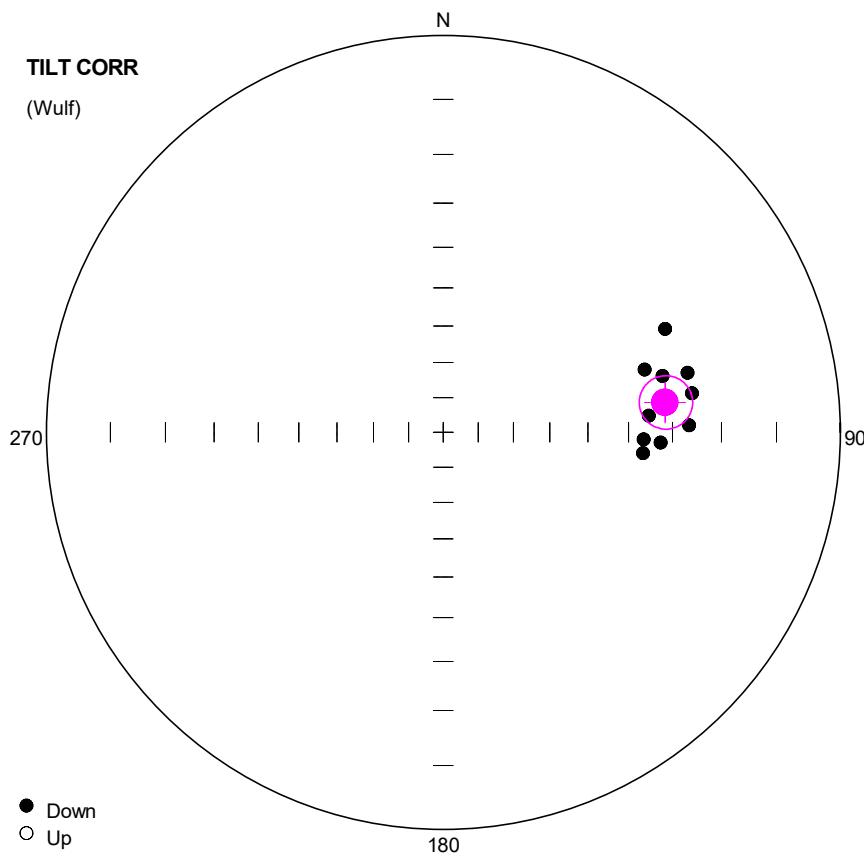
dp: 6.6

dm: 11.22

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4301	CCH	34.8	28.0	2.3	40	68
2	YUN4302	CCH	44.7	35.6	2.9	44	68
3	YUN4303	CCH	61.9	35.0	4.2	40	68
4	YUN4304	CCH	35.8	32.2	6.0	40	68
5	YUN4305	CCH	51.0	45.2	3.1	44	68
6	YUN4306	CCH	64.7	35.9	3.1	30	68
7	YUN4307	CCH	66.9	46.0	7.1	30	60



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4401	CCH	117.5	24.2	5.9	30	52
2	YUN4402	CCH	105.4	34.4	5.4	30	60
3	YUN4403	CCH	100.9	30.6	5.9	35	60
4	YUN4404	CCH	120.4	27.1	5.8	30	56
5	YUN4405	CCH	122.6	24.6	5.7	35	60
6	YUN4406	CCH	115.5	30.5	2.9	30	60
7	YUN4407	CCH	94.3	39.3	5.5	35	64
8	YUN4408	CCH	104.0	27.2	4.9	35	60
9	YUN4409	CCH	107.0	38.4	8.2	40	64
10	YUN4410	CCH	109.7	23.0	5.6	35	60



**Number of data points: 10**

## Fisher statistics

Mean vector: 82.3 / 31.2

Resultant vector: 9.87

(X= 1.13, Y= 8.37, Z= 5.11)

Estimated precision, k: 69.3

95% Confidence limit: 5.84

95% Confidence limit, approximation: 5.32

## Orientation matrix

1st eigenvalue: 0.97

1st eigenvector: 82.33 / 31.2

2nd eigenvalue: 0.02

2nd eigenvector: 182.11 / 15.67

3rd eigenvalue: 0

3rd eigenvector: 295.04 / 54.26

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 18.47 (-18.47)

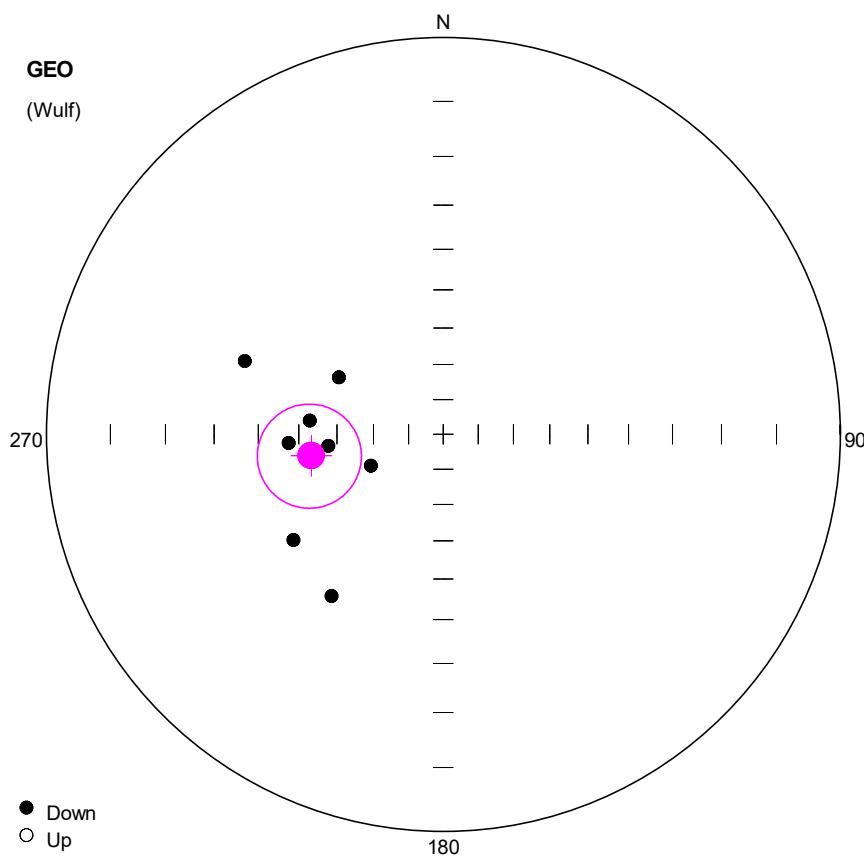
Pole longitude: 103.66 (-76.34)

Paleolatitude: 16.85 (-16.85)

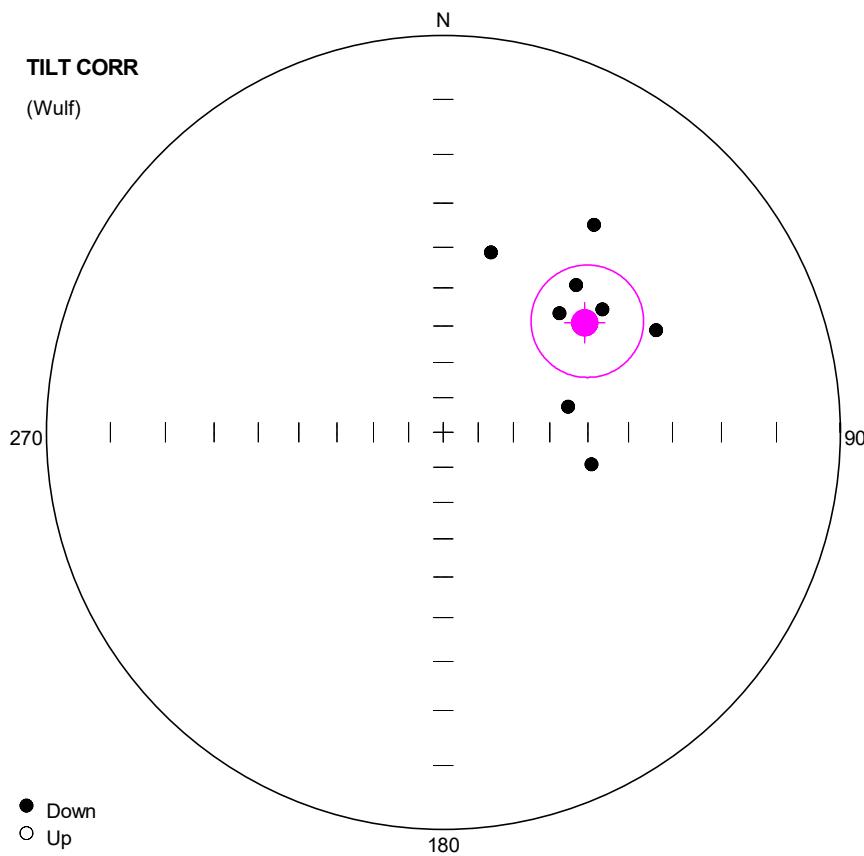
dp: 3.66

dm: 6.54

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4401	CCH	92.7	32.5	5.9	30	52
2	YUN4402	CCH	75.6	30.6	5.4	30	60
3	YUN4403	CCH	76.3	25.3	5.9	35	60
4	YUN4404	CCH	92.0	36.4	5.8	30	56
5	YUN4405	CCH	96.0	36.3	5.7	35	60
6	YUN4406	CCH	85.4	35.1	2.9	30	60
7	YUN4407	CCH	65.0	26.7	5.5	35	64
8	YUN4408	CCH	81.1	25.2	4.9	35	60
9	YUN4409	CCH	72.7	34.0	8.2	40	64
10	YUN4410	CCH	88.4	26.4	5.6	35	60



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4501	CCH	264.3	57.6	6.3	30	64
2	YUN4502	CCH	246.5	67.5	5.7	35	64
3	YUN4503	CCH	298.6	56.7	12.7	35	60
4	YUN4505	CCH	290.3	33.9	9.3	44	64
5	YUN4506	CCH	214.6	37.3	19.5	35	60
6	YUN4507	CCH	275.9	52.7	10.1	30	64
7	YUN4508	CCH	234.8	40.4	8.6	35	64
8	YUN4509	CCH	260.6	54.2	5.4	30	64
9	YUN4510	CCH	266.8	47.4	5.3	40	64



**Number of data points: 9**

## Fisher statistics

Mean vector: 52.28 / 41.5

Resultant vector: 8.49

(X= 3.89, Y= 5.03, Z= 5.63)

Estimated precision, k: 15.72

95% C

95% Confidence limit, approximation: 11.77

## Orientation matrix

1st eigenvalue: 0.89

1st eigenvector: 51.82 / 41.13

2nd eigenvalue: 0.08

2nd eigenvector: 157.93 / 17.64

3rd eigenvalue: 0.02

3rd eigenvector: 265.56 / 43.6

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 37.09 (-37.09)

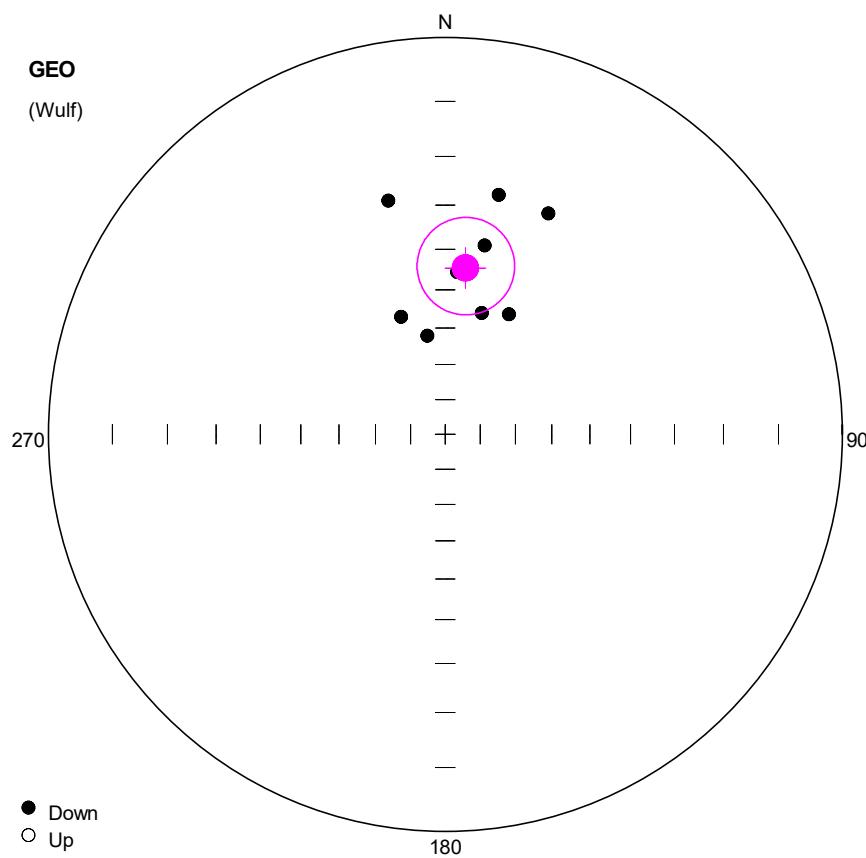
Pole longitude: 128.96 (-51.04)

Paleolatitude: 23.86 (-23.86)

dp: 9.99

dm: 16.36

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4501	CCH	52.3	36.3	6.3	30	64
2	YUN4502	CCH	64.4	28.5	5.7	35	64
3	YUN4503	CCH	36.0	24.3	12.7	35	60
4	YUN4505	CCH	14.9	39.8	9.3	44	64
5	YUN4506	CCH	102.2	48.2	19.5	35	60 C
6	YUN4507	CCH	42.1	36.9	10.1	30	64
7	YUN4508	CCH	78.4	54.4	8.6	35	64
8	YUN4509	CCH	53.1	40.2	5.4	30	64
9	YUN4510	CCH	44.3	44.5	5.3	40	64



**Number of data points: 9**

**Fisher statistics**

Mean vector: 6.92 / 44.25

Resultant vector: 8.6

(X= 6.11, Y= 0.74, Z= 6)

Estimated precision, k: 19.78

95% Confidence limit: 11.87

95% Confidence limit, approximation: 10.49

**Orientation matrix**

1st eigenvalue: 0.91

1st eigenvector: 7.06 / 44.34

2nd eigenvalue: 0.05

2nd eigenvector: 223.88 / 39.33

3rd eigenvalue: 0.03

3rd eigenvector: 117.14 / 19.36

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 50.72 (-50.72)

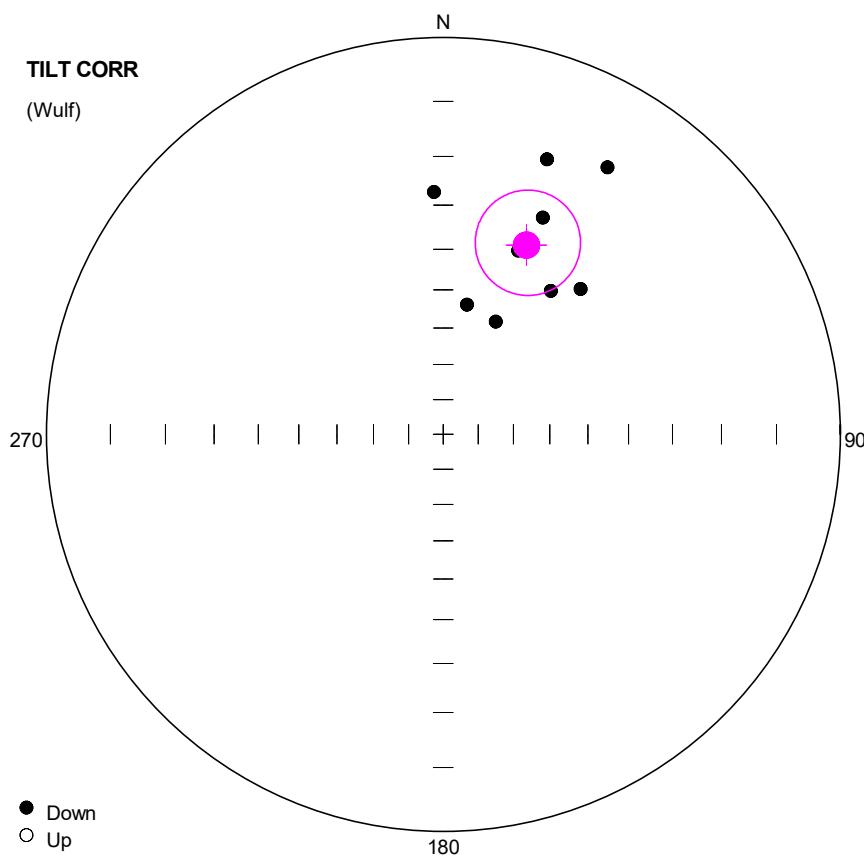
Pole longitude: -175.8 (4.2)

Paleolatitude: 25.97 (-25.97)

dp: 9.35

dm: 14.89

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4601	CCH	25.0	26.9	8.6	48	64
2	YUN4602	CCH	346.3	27.6	10.3	35	60
3	YUN4603	CCH	11.8	38.2	5.3	35	60
4	YUN4604	CCH	12.6	26.6	9.2	40	52
5	YUN4605	CCH	4.2	45.4	12.2	35	60
6	YUN4606	CCH	27.9	52.2	11.3	35	56
7	YUN4607	CCH	349.7	61.7	10.1	35	56 O
8	YUN4608	CCH	339.4	54.9	8.4	44	60
9	YUN4609	CCH	16.8	54.6	7.3	40	68



**Number of data points: 9**

## Fisher statistics

Mean vector: 23.76 / 35

Resultant vector: 8.59

(X= 6.44, Y= 2.84, Z= 4.93)

Estimated precision, k: 19.74

95% Confidence limit: 1

95% Confidence limit, approximation: 10.5

## Orientation matrix

1st eigenvalue: 0.91

1st eigenvector: 23.92 / 35.04

2nd eigenvalue: 0.05

2nd eigenvector: 207.41 / 54.91

3rd eigenvalue: 0.03

3rd eigenvector: 115.07 / 1.64

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 41.65 (-41.65)

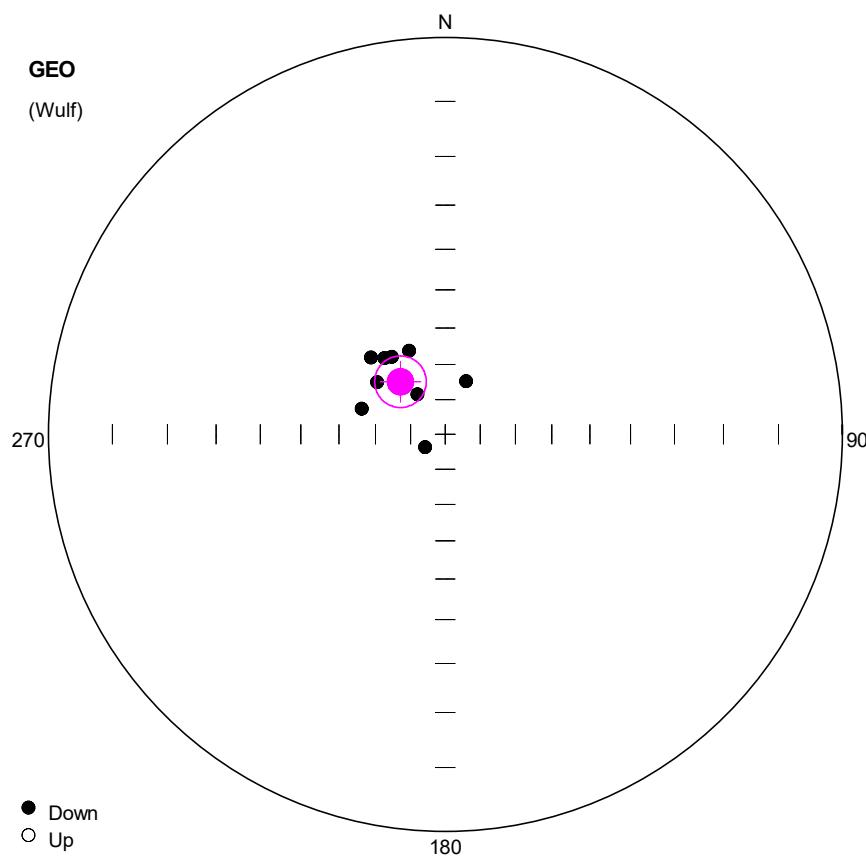
Pole longitude: 163.45 (-16.5)

Paleolatitude: 19.3 (-19.3)

dp: 7.89

dm: 13.69

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4601	CCH	31.6	13.4	8.6	48	64
2	YUN4602	CCH	357.8	27.2	10.3	35	60
3	YUN4603	CCH	24.7	28.0	5.3	35	60
4	YUN4604	CCH	20.7	17.0	9.2	40	52
5	YUN4605	CCH	22.2	36.9	12.2	35	60
6	YUN4606	CCH	43.4	36.5	11.3	35	56
7	YUN4607	CCH	25.0	55.2	10.1	35	56 O
8	YUN4608	CCH	10.4	53.3	8.4	44	60
9	YUN4609	CCH	36.9	41.4	7.3	40	68



Number of data points: 10

#### Fisher statistics

Mean vector: 319.54 / 70.27

Resultant vector: 9.8

(X= 2.52, Y= -2.15, Z= 9.23)

Estimated precision, k: 45.76

95% Confidence limit: 7.22

95% Confidence limit, approximation: 6.54

#### Orientation matrix

1st eigenvalue: 0.96

1st eigenvector: 319.52 / 70.13

2nd eigenvalue: 0.02

2nd eigenvector: 129.56 / 19.59

3rd eigenvalue: 0.02

3rd eigenvector: 220.69 / 3.17

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 67.5 (-67.5)

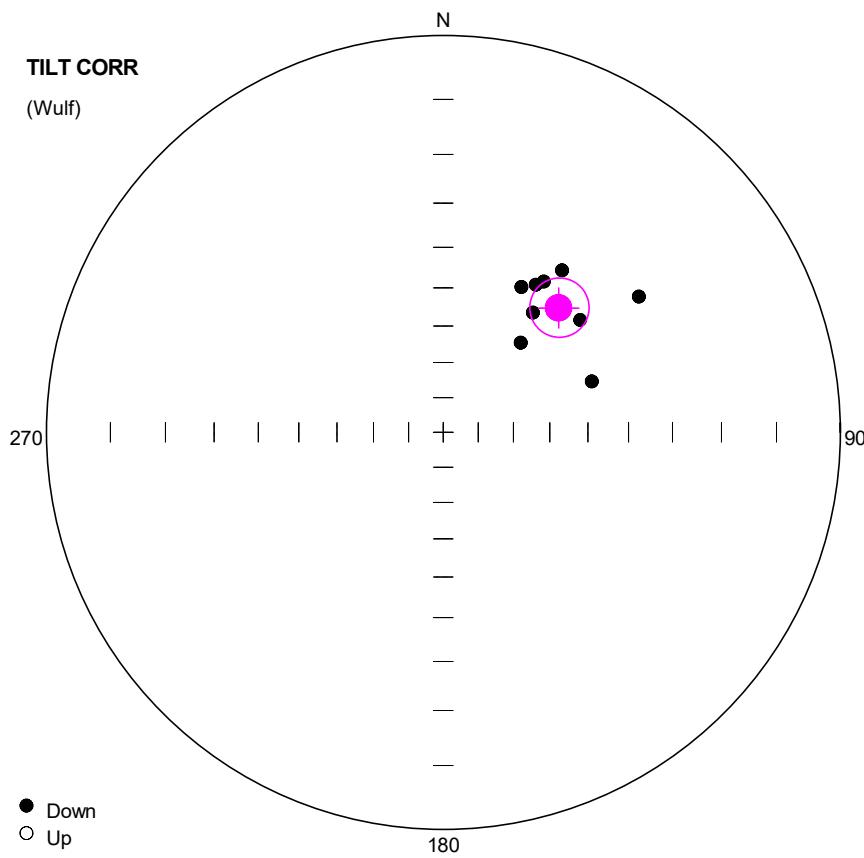
Pole longitude: -84.75 (95.25)

Paleolatitude: 54.35 (-54.35)

dp: 10.76

dm: 12.46

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4701	CCH	321.4	62.4	6.9	35	64
2	YUN4702	CCH	325.4	63.3	8.6	35	60
3	YUN4703	CCH	325.3	76.0	7.8	40	64
4	YUN4704	CCH	319.3	70.4	4.5	40	64
5	YUN4705	CCH	307.5	65.6	3.5	40	64
6	YUN4706	CHT	287.1	65.2	1.7	56	68
7	YUN4707	CHT	237.1	83.1	6.7	44	68
8	YUN4708	CCH	316.0	59.9	5.7	40	60 O
9	YUN4709	CCH	336.7	64.2	5.3	40	60
10	YUN4710	CCH	21.5	73.6	5.4	40	60



**Number of data points: 10**

## Fisher statistics

Mean vector: 42.87 / 43.74

Resultant vector: 9.8

(X= 5.19, Y= 4.82, Z= 6.78)

Estimated precision, k: 45.75

95% Confidence limit: 7.22

95% Confidence limit, approximation: 6.55

## Orientation matrix

1st eigenvalue: 0.96

1st eigenvector: 42.68 / 43.75

2nd eigenvalue: 0.02

2nd eigenvector: 305.42 / 7.52

3rd eigenvalue: 0.02

3rd eigenvector: 207.76 / 45.27

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 42.11 (-42.11)

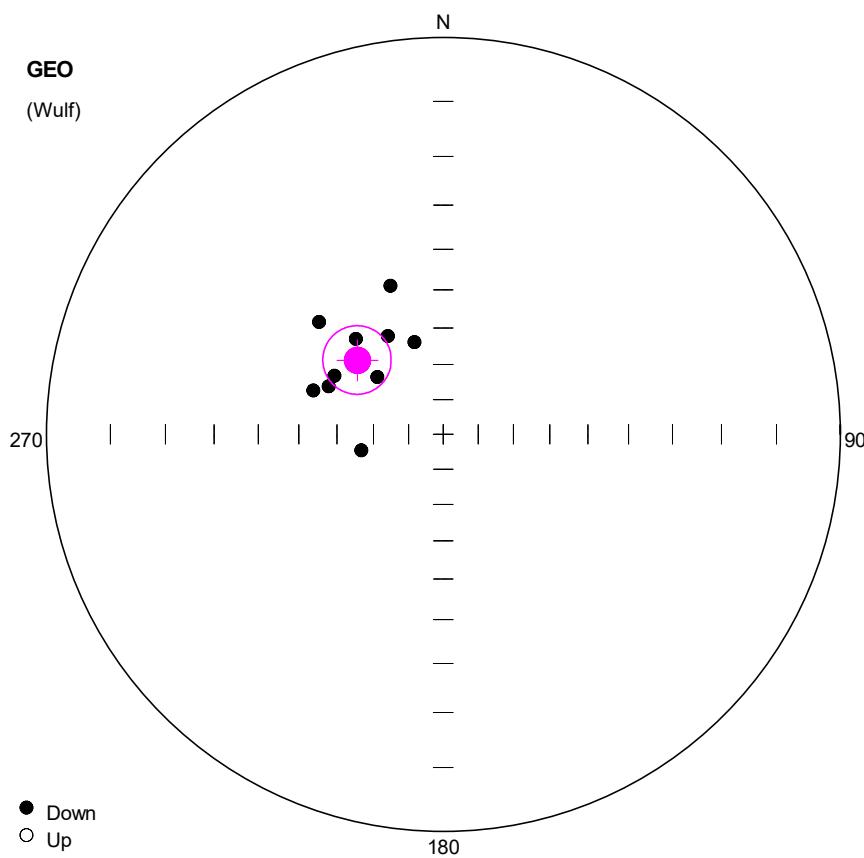
Pole longitude: 138.23 (-41.77)

Paleolatitude: 25.57 (-25.57)

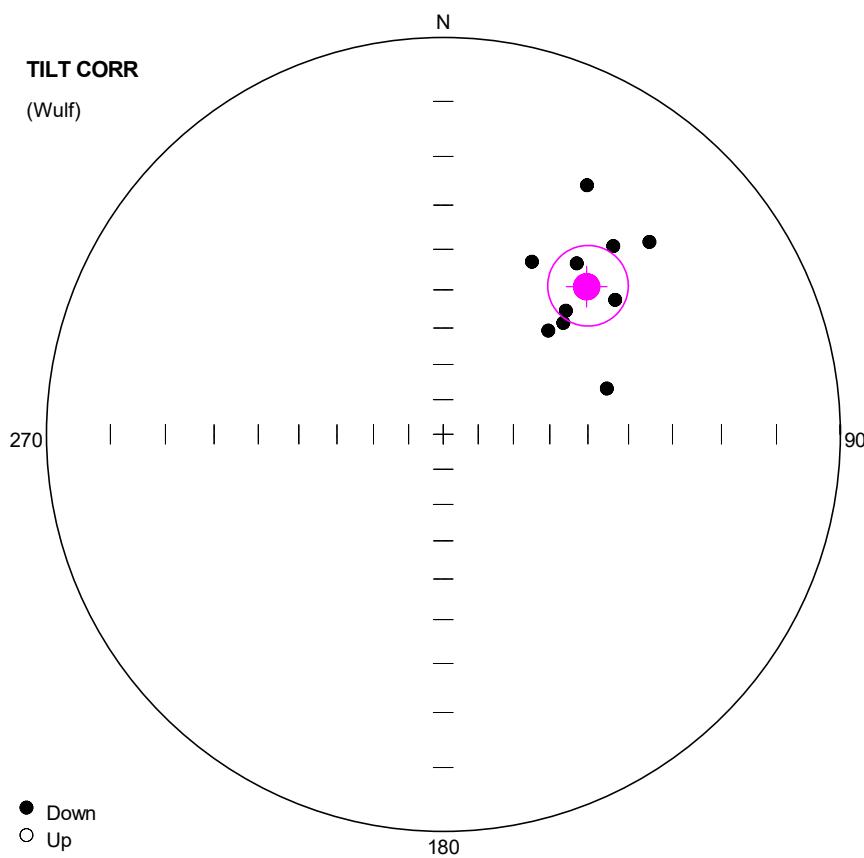
dp: 5.63

dm: 9.01

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4701	CCH	32.1	42.6	6.9	35	64
2	YUN4702	CCH	33.7	40.9	8.6	35	60
3	YUN4703	CCH	50.6	41.9	7.8	40	64
4	YUN4704	CCH	43.0	43.8	4.5	40	64
5	YUN4705	CCH	36.9	48.7	3.5	40	64
6	YUN4706	CHT	40.9	56.8	1.7	56	68
7	YUN4707	CHT	71.1	46.8	6.7	44	68
8	YUN4708	CCH	28.3	44.9	5.7	40	60 C
9	YUN4709	CCH	36.3	36.3	5.3	40	60
10	YUN4710	CCH	55.3	28.1	5.4	40	60



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4901	CCH	258.9	66.3	10.4	35	64
2	YUN4902	CCH	292.7	55.3	14.3	30	64
3	YUN4903	CCH	288.7	51.9	14.5	40	64
4	YUN4904	CCH	298.3	55.4	8.1	35	64
5	YUN4905	CCH	312.1	44.3	5.3	40	64
6	YUN4906	CCH	340.4	46.7	6.6	35	60
7	YUN4907	CCH	342.7	62.7	4.8	40	64
8	YUN4908	CCH	330.7	58.3	4.0	35	64
9	YUN4909	CCH	317.5	53.9	3.0	44	64
10	YUN4910	CCH	311.0	65.2	12.9	40	64



**Number of data points: 10**

## Fisher statistics

Mean vector: 44.16 / 35.21

Resultant vector: 9.69

(X= 5.68, Y= 5.52, Z= 5.59)

Estimated precision, k: 28.95

95% Confidence limit: 9.13

95% Confidence limit, approximation: 8.23

## Orientation matrix

1st eigenvalue: 0.94

1st eigenvector: 44.06 / 35.22

2nd eigenvalue: 0.04

2nd eigenvector: 171.65 / 40.83

3rd eigenvalue: 0.02

3rd eigenvector: 290.69 / 29.33

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 35.98 (-35.98)

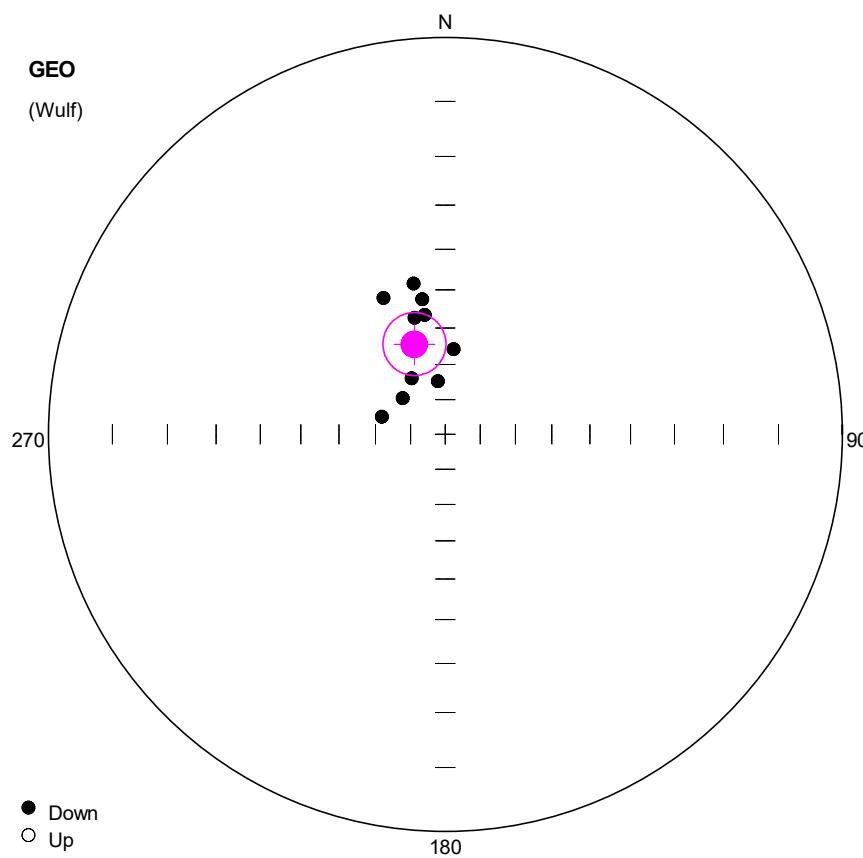
Pole longitude: 139.76 (-40.24)

Paleolatitude: 19.44 (-19.44)

dp: 6.08

dm: 10.54

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN4901	CCH	74.4	43.7	10.4	35	64
2	YUN4902	CCH	47.2	45.2	14.3	30	64
3	YUN4903	CCH	45.4	49.2	14.5	40	64
4	YUN4904	CCH	44.8	42.6	8.1	35	64
5	YUN4905	CCH	27.2	37.9	5.3	40	64
6	YUN4906	CCH	30.0	18.2	6.6	35	60
7	YUN4907	CCH	47.0	19.2	4.8	40	64
8	YUN4908	CCH	42.1	24.9	4.0	35	64
9	YUN4909	CCH	38.0	32.7	3.0	44	64
10	YUN4910	CCH	52.0	32.4	12.9	40	64



Number of data points: 10

#### Fisher statistics

Mean vector: 341.08 / 63.06

Resultant vector: 9.72

(X= 4.17, Y= -1.43, Z= 8.67)

Estimated precision, k: 32.72

95% Confidence limit: 8.57

95% Confidence limit, approximation: 7.74

#### Orientation matrix

1st eigenvalue: 0.95

1st eigenvector: 341.31 / 62.98

2nd eigenvalue: 0.04

2nd eigenvector: 188.13 / 24.47

3rd eigenvalue: 0.01

3rd eigenvector: 93.17 / 10.75

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 67.01 (-67.01)

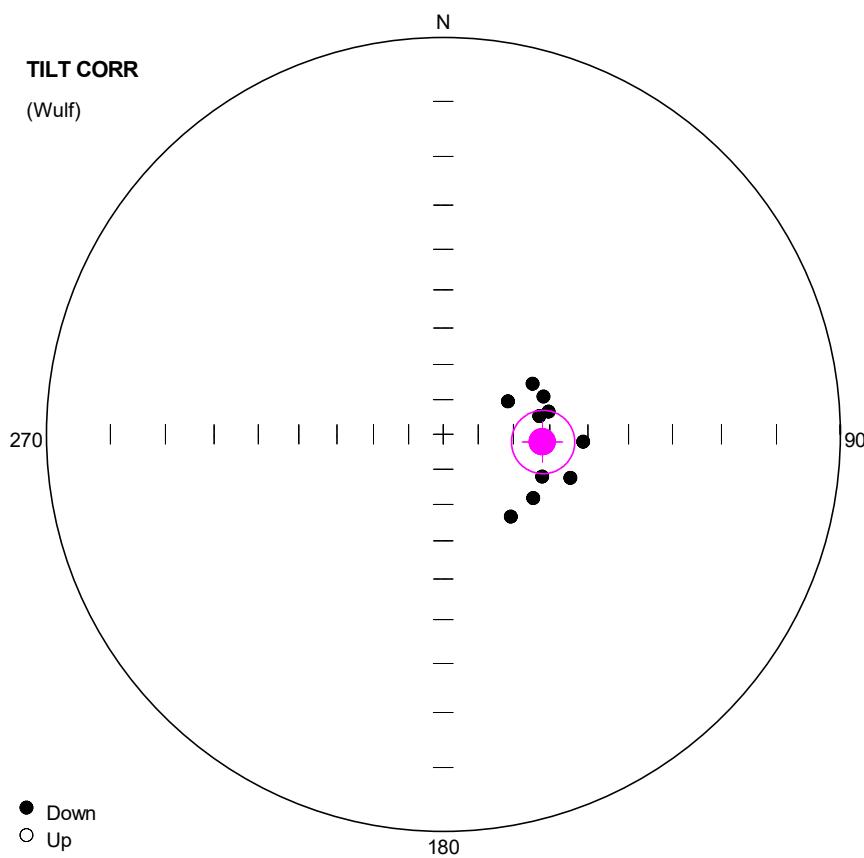
Pole longitude: -129.67 (50.33)

Paleolatitude: 44.53 (-44.53)

dp: 10.61

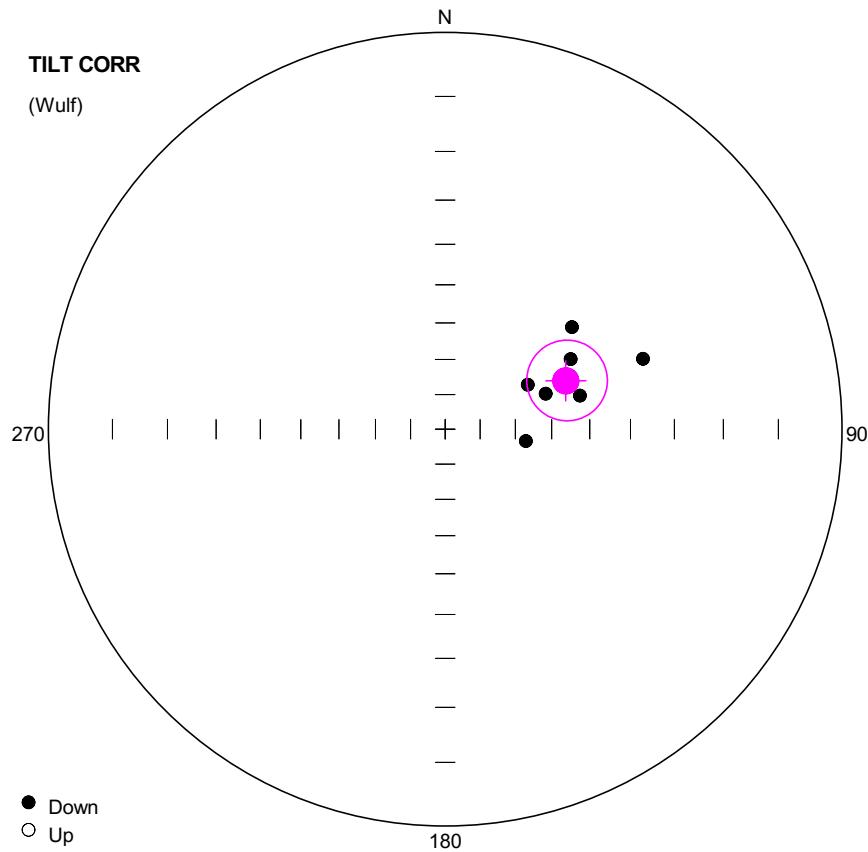
dm: 13.49

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN5001	CCH	310.3	74.0	8.7	48	60
2	YUN5002	CCH	285.5	71.2	4.2	44	60 O
3	YUN5003	CCH	348.2	47.6	5.3	40	64 O
4	YUN5004	CCH	329.1	71.3	12.9	48	64 O
5	YUN5005	CCH	352.1	74.6	8.7	35	60
6	YUN5006	CCH	335.7	48.7	4.4	30	60
7	YUN5007	CCH	350.4	51.9	3.7	35	60
8	YUN5008	CCH	350.3	56.1	5.8	35	64
9	YUN5009	CCH	345.3	56.3	5.8	35	64
10	YUN5010	CCH	5.6	65.7	8.2	35	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUN5001	CCH	125.3	58.9	8.7	48	60
2	YUN5002	CCH	140.6	59.9	4.2	44	60
3	YUN5003	CCH	60.5	61.0	5.3	40	64
4	YUN5004	CCH	113.1	59.7	12.9	48	64
5	YUN5005	CCH	109.0	52.6	8.7	35	60
6	YUN5006	CCH	63.0	69.3	4.4	30	60
7	YUN5007	CCH	69.4	59.8	3.7	35	60
8	YUN5008	CCH	77.9	59.6	5.8	35	64
9	YUN5009	CCH	79.2	62.3	5.8	35	64
10	YUN5010	CCH	93.0	51.1	8.2	35	64





**Number of data points: 7**

## Fisher statistics

Mean vector: 68.16 / 53.73

Resultant vector: 6.82

(X= 1.5, Y= 3.75, Z= 5.5)

Estimated precision, k: 34.21

95% C

95% Confidence limit, approximation: 9.05

## Orientation matrix

1st eigenvalue: 0.95

1st eigenvector: 68.06 / 53.77

2nd eigenvalue: 0.04

2nd eigenvector: 223.31 / 33.64

3rd eigenvalue: 0.01

3rd eigenvector: 321.36 / 11.89

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 39.81 (-39.81)

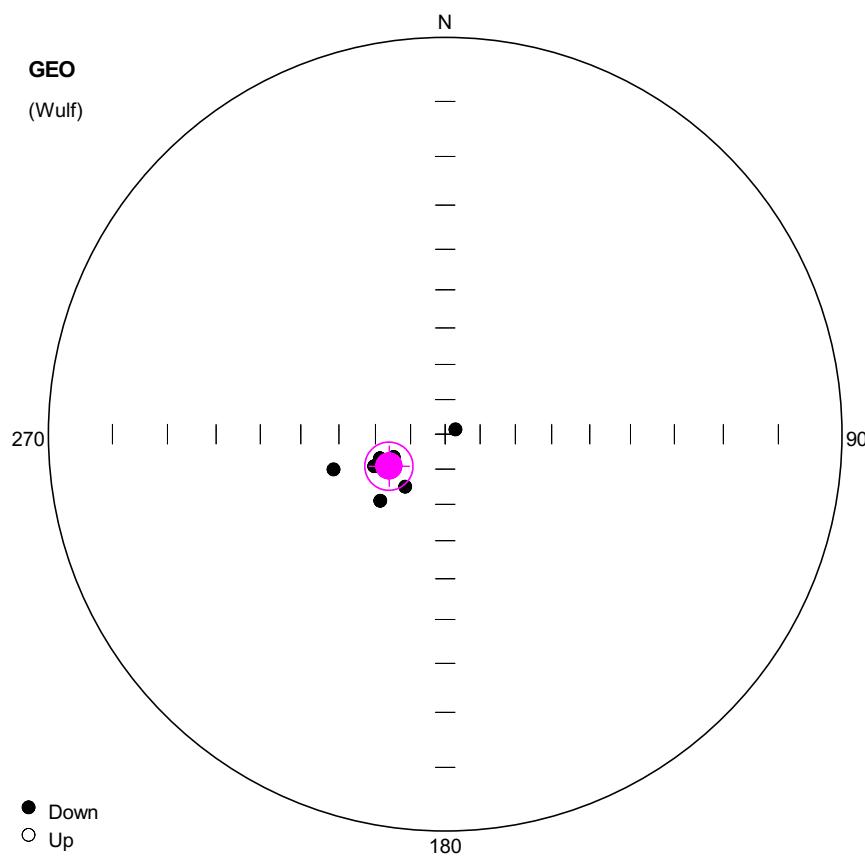
Pole longitude: 107.07 (-72.93)

Paleolatitude: 34.27 (-34.27)

dp: 10.21

dm: 14.62

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5101	Ch	60.8	50.2	13.8	30	60
2	YUNN5102	Cht	98.4	66.8	5.9	44	60
3	YUNN5103	Cht	76.0	51.4	3.6	48	60
4	YUNN5104	Ch	61.7	63.4	12.8	30	60
5	YUNN5105	Cht	70.4	34.2	8.2	40	56
6	YUNN5106	Ch	51.2	45.4	17.4	30	56
7	YUNN5108	Ch	70.6	59.9	8.9	44	56



**Number of data points:** 9

**Fisher statistics**

Mean vector: 240.51 / 71.51

Resultant vector: 8.86

(X= -1.38, Y= -2.45, Z= 8.41)

Estimated precision, k: 58.83

95% Confidence limit: 6.77

95% Confidence limit, approximation: 6.08

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 240.49 / 71.4

2nd eigenvalue: 0.02

2nd eigenvector: 66.96 / 18.49

3rd eigenvalue: 0.01

3rd eigenvector: 336.31 / 1.95

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 39.63 (-39.63)

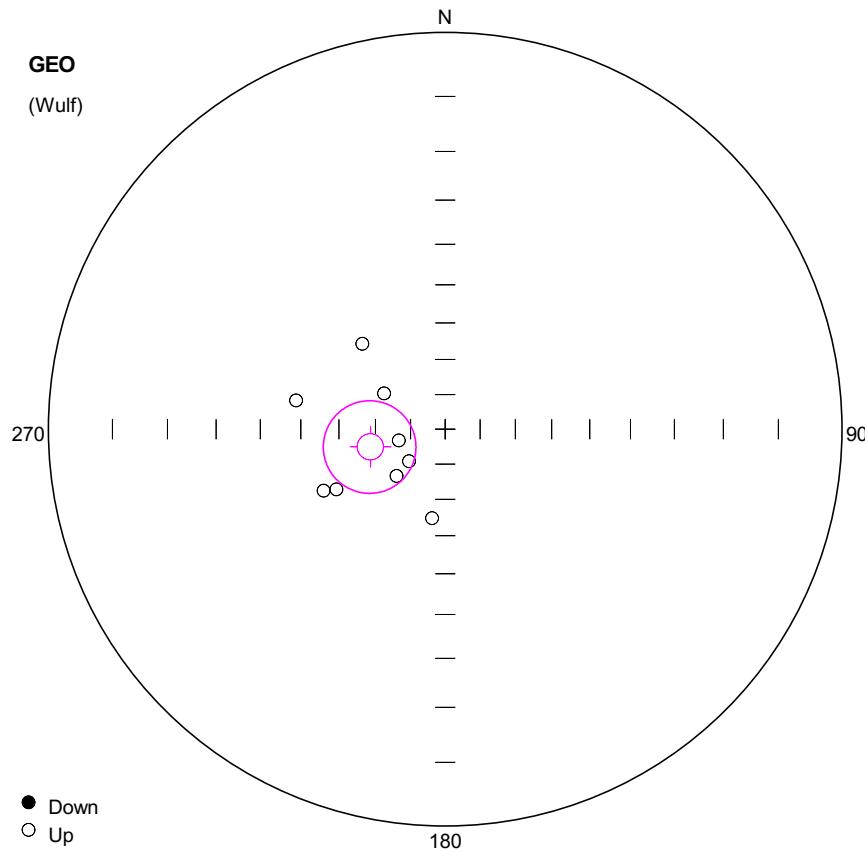
Pole longitude: -24.87 (155.13)

Paleolatitude: 56.23 (-56.23)

dp: 10.4

dm: 11.86

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5201	Clt	249.7	70.1	7.9	35	48
2	YUNN5202	Ch	244.4	72.0	3.8	30	60
3	YUNN5204	Ch	65.7	86.7	4.5	35	60
4	YUNN5205	Cht	224.3	63.6	3.6	44	60
5	YUNN5206	Cht	217.2	71.1	4.8	40	60
6	YUNN5207	Cht	245.7	67.8	3.8	35	60
7	YUNN5208	Ch	252.5	57.1	7.7	30	60
8	YUNN5209	Ch	241.9	71.6	3.9	30	60
9	YUNN5210	Ch	246.2	73.9	6.4	35	60



**Number of data points: 9**

## Fisher statistics

Mean vector: 256.79 / -68.12

Resultant vector: 8.53

(X= -0.73, Y= -3.1, Z= -7.92)

Estimated precision, k: 17.06

95% Confidence limit: 12.83

### 95% Confidence limit, app

## Orientation matrix

1st eigenvalue: 0.9

1st eigenvector: 76.19 / 68.16

2nd eigenvalue: 0.07

2nd eigenvector: 335.3 / 4.33

3rd eigenvalue: 0.03

3rd eigenvector: 243.6 / 21.36

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -50.09 (50.09)

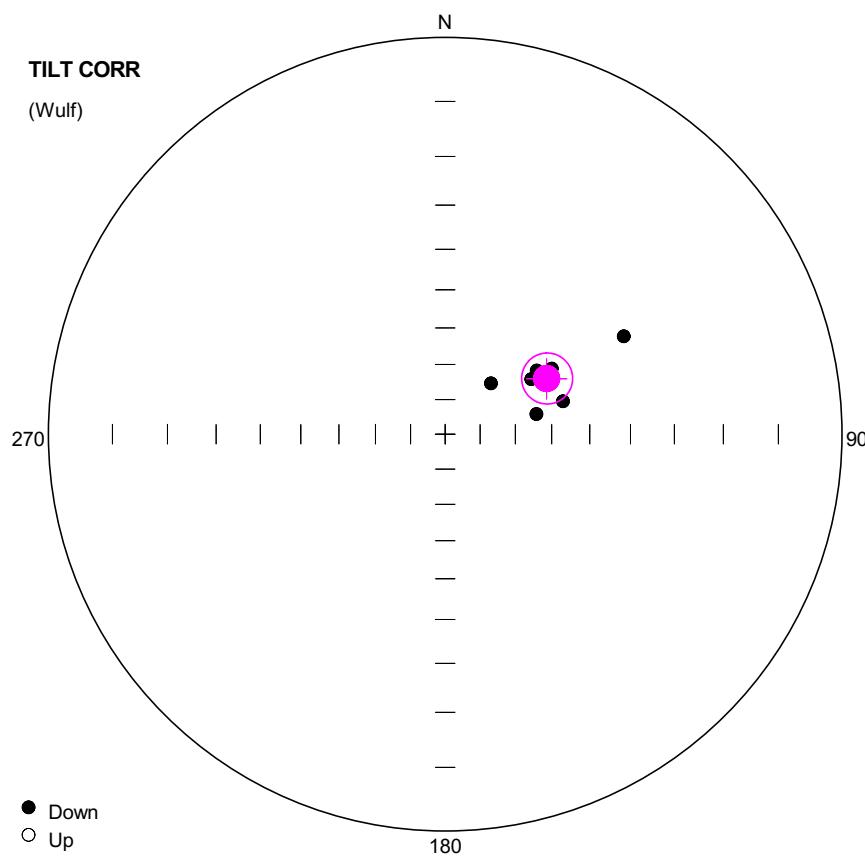
Pole longitude: -94.1 (85.9)

Paleolatitude: -51.23 (51.23)

dp: 18.11

dm: 21.56

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5201	CHT	228.3	-76.1	1.0	64	64 OA
2	YUNN5202	CHT	226.2	-70.7	1.0	64	64 OA
3	YUNN5204	CHT	243.2	-52.1	1.0	64	64 OA
4	YUNN5205	CHT	300.3	-69.8	1.0	64	64 OA
5	YUNN5206	CHT	241.1	-55.2	1.0	64	64 OA
6	YUNN5207	CHT	280.9	-48.1	1.0	64	64 OA
7	YUNN5208	CHT	188.3	-64.5	1.0	64	64 OA
8	YUNN5209	CHT	256.5	-76.3	1.0	64	64 OA
9	YUNN5210	CHT	315.9	-56.6	1.0	64	64 OA



**Number of data points:** 9

**Fisher statistics**

Mean vector: 61.29 / 57.49

Resultant vector: 8.86

(X= 2.29, Y= 4.18, Z= 7.47)

Estimated precision, k: 58.75

95% Confidence limit: 6.77

95% Confidence limit, approximation: 6.09

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 61.3 / 57.6

2nd eigenvalue: 0.02

2nd eigenvector: 247.59 / 32.24

3rd eigenvalue: 0.01

3rd eigenvector: 155.79 / 2.85

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 45.99 (-45.99)

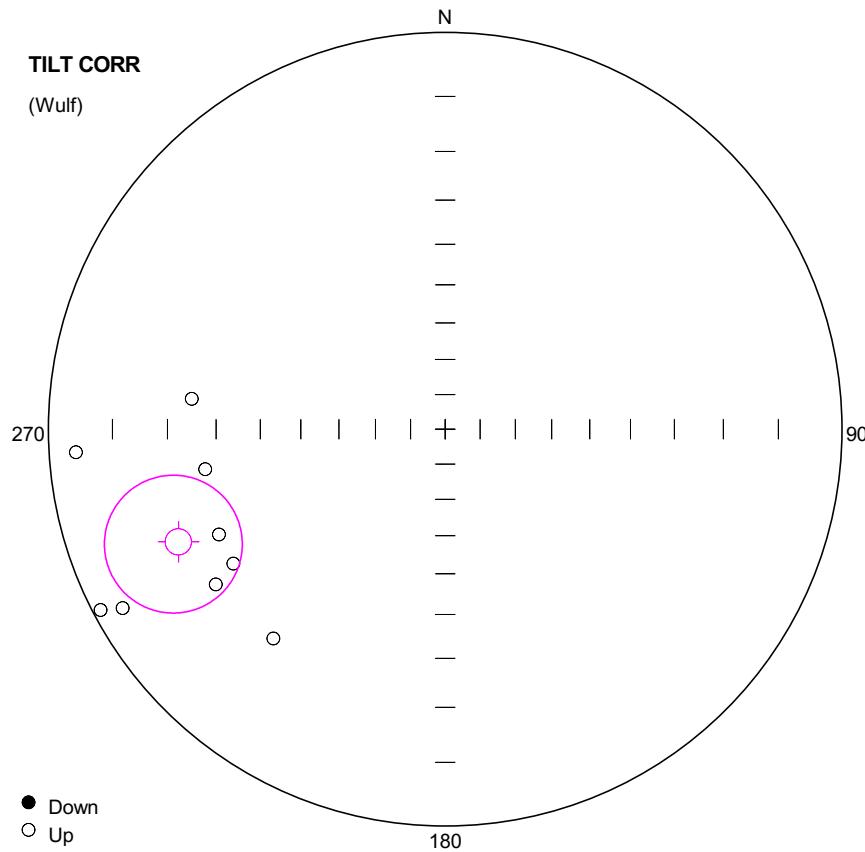
Pole longitude: 110.79 (-69.21)

Paleolatitude: 38.12 (-38.12)

dp: 7.26

dm: 9.91

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5201	Clt	55.3	58.6	7.9	35	48
2	YUNN5202	Ch	59.1	56.9	3.8	30	60
3	YUNN5204	Ch	61.3	35.7	4.5	35	60
4	YUNN5205	Cht	77.6	63.5	3.6	44	60
5	YUNN5206	Cht	74.4	55.7	4.8	40	60
6	YUNN5207	Cht	57.4	61.1	3.8	35	60
7	YUNN5208	Ch	42.1	70.4	7.7	30	60
8	YUNN5209	Ch	60.5	57.4	3.9	30	60
9	YUNN5210	Ch	58.5	55.0	6.4	35	60



**Number of data points: 9**

## Fisher statistics

Mean vector: 247.1 / -17.77

Resultant vector: 8.53

(X= -3.16, Y= -7.48, Z= -2.6)

Estimated precision, k: 17.03

95% Confidence limit: 12.84

95% Confidence limit, approximation: 11.31

## Orientation matrix

1st eigenvalue: 0.9

1st eigenvector: 66.86 / 17.77

2nd eigenvalue: 0.07

2nd eigenvector: 157.12 / 0.82

3rd eigenvalue: 0.03

3rd eigenvector: 249.68 / 72.21

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -17.81 (17.81)

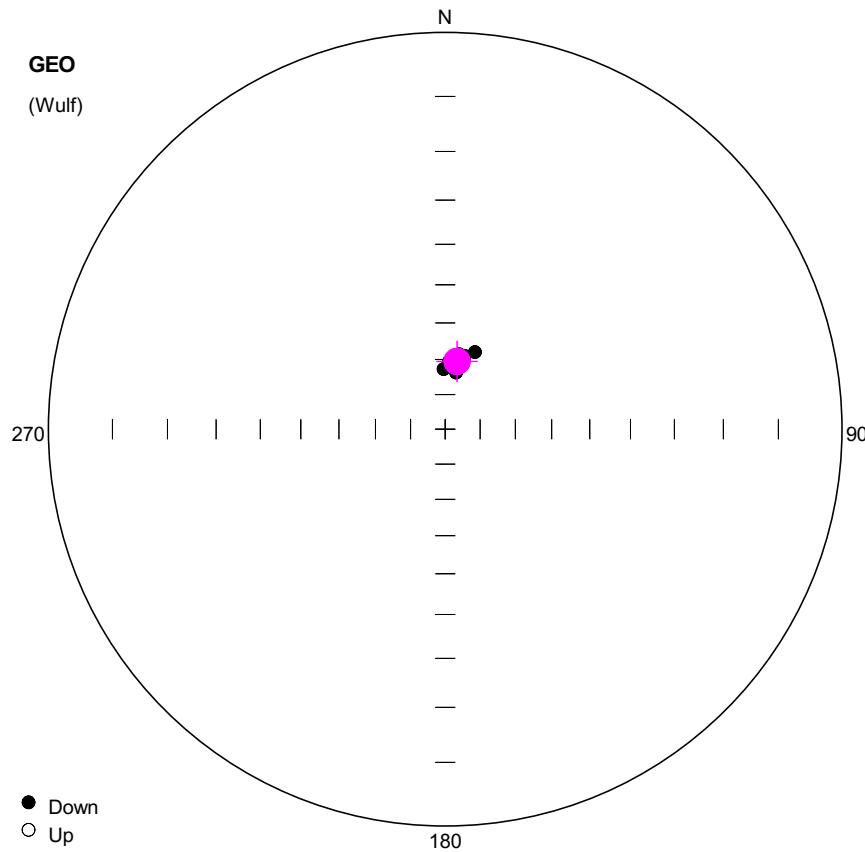
Pole longitude: -58.77 (121.23)

Paleolatitude: -9.1 (9.1)

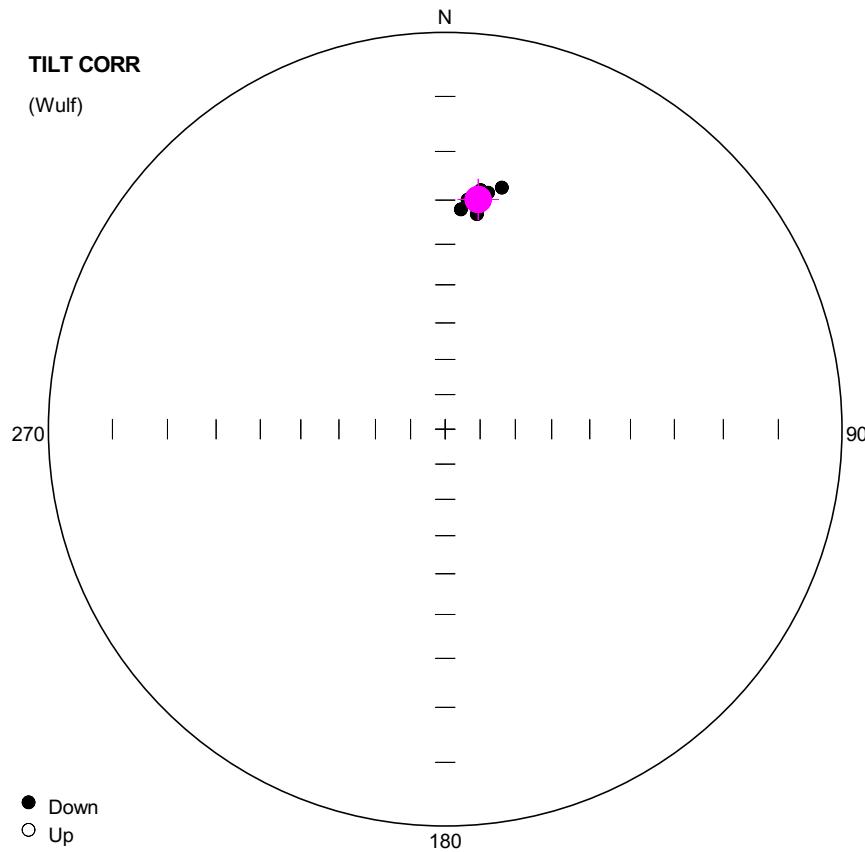
dp: 6.9

dm: 13.32

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5201	CHT	237.6	-25.4	1.0	64	64 OA
2	YUNN5202	CHT	235.9	-20.2	1.0	64	64 OA
3	YUNN5204	CHT	242.3	-1.1	1.0	64	64 OA
4	YUNN5205	CHT	260.5	-27.0	1.0	64	64 OA
5	YUNN5206	CHT	241.0	-4.2	1.0	64	64 OA
6	YUNN5207	CHT	266.4	-4.0	1.0	64	64 OA
7	YUNN5208	CHT	219.4	-21.4	1.0	64	64 OA
8	YUNN5209	CHT	245.0	-25.7	1.0	64	64 OA
9	YUNN5210	CHT	276.8	-24.5	1.0	64	64 OA



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5301	Ch	10.6	68.3	5.1	35	60
2	YUNN5302	Ch	21.1	66.5	5.6	20	68
3	YUNN5303	Ch	11.1	73.5	5.0	20	68
4	YUNN5304	Ch	5.9	70.2	3.1	20	68
5	YUNN5305	Ch	5.3	71.1	2.6	20	68
6	YUNN5306	Ch	14.9	68.4	2.5	35	68
7	YUNN5307	Ch	13.3	70.1	5.0	40	68
8	YUNN5308	Ch	3.4	70.7	4.4	35	68
9	YUNN5309	Ch	11.2	70.6	3.7	30	68
10	YUNN5310	Ch	358.6	72.8	4.2	20	68



**Number of data points: 10**

## Fisher statistics

Mean vector: 8.15 / 29.34

Resultant vector: 9.99

(X= 8.62, Y= 1.23, Z= 4.89)

Estimated precision, k: 713.21

95% Confidence limit: 1.8

95% Confidence limit, approximation: 1.66

## Orientation matrix

1st eigenvalue: 1

1st eigenvector: 8.15 / 29.34

2nd eigenvalue: 0

2nd eigenvector: 255.96 / 33.9

3rd eigenvalue: 0

3rd eigenvector: 128.63 / 42.06

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 40.39 (-40.39)

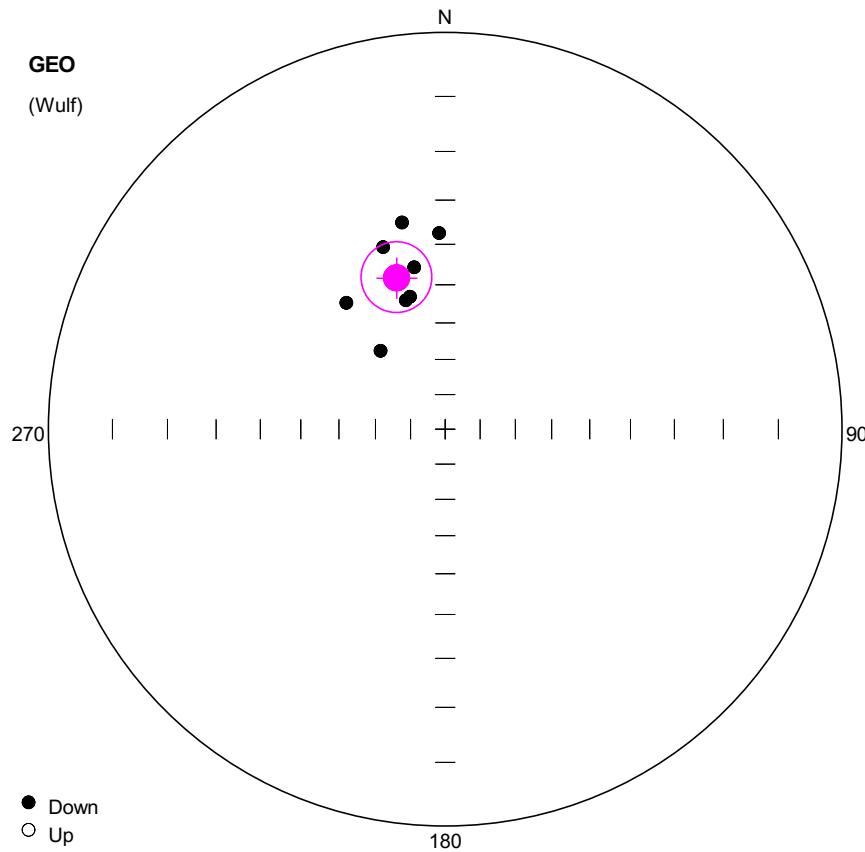
Pole longitude: -176.28 (3.72)

Paleolatitude: 15.7 (-15.7)

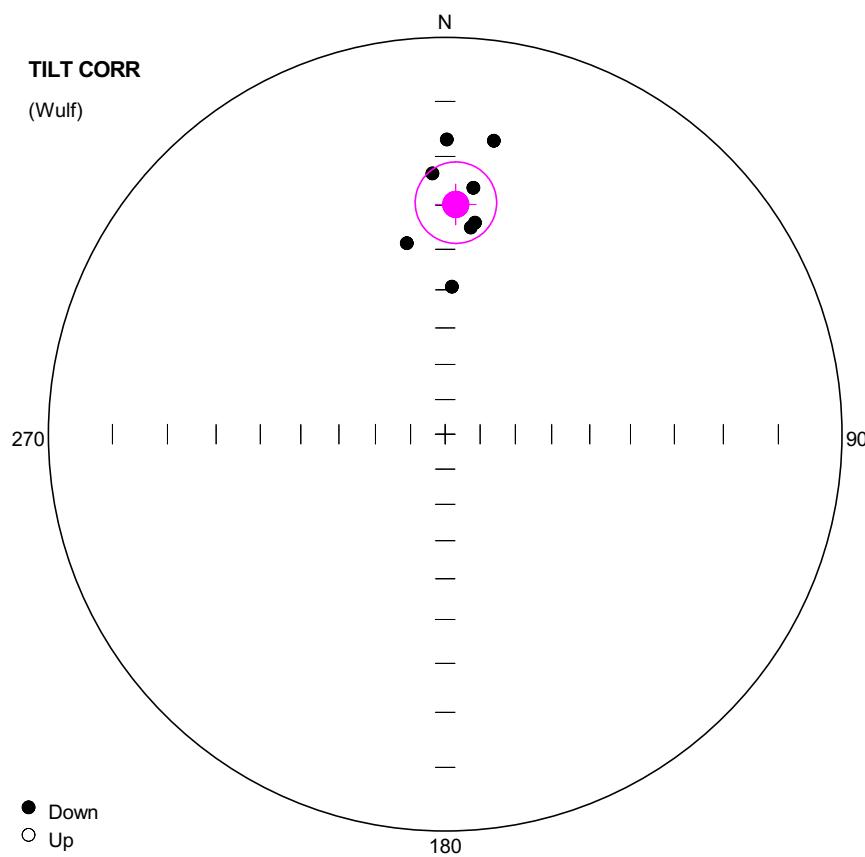
dp: 1.1

dm: 2

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5301	Ch	8.5	27.3	5.1	35	60
2	YUNN5302	Ch	13.2	26.0	5.6	20	68
3	YUNN5303	Ch	8.4	32.6	5.0	20	68
4	YUNN5304	Ch	6.6	29.2	3.1	20	68
5	YUNN5305	Ch	6.4	30.1	2.6	20	68
6	YUNN5306	Ch	10.3	27.6	2.5	35	68
7	YUNN5307	Ch	9.5	29.1	5.0	40	68
8	YUNN5308	Ch	5.6	29.7	4.4	35	68
9	YUNN5309	Ch	8.6	29.7	3.7	30	68
10	YUNN5310	Ch	4.1	31.9	4.2	20	68



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5401	Ch	341.2	38.3	12.1	20	52
2	YUNN5402	Ch	345.1	51.9	7.9	20	60
3	YUNN5403	Ch	349.2	44.9	3.7	20	60
4	YUNN5404	Ch	358.2	37.4	8.0	35	68
5	YUNN5405	Ch	348.2	34.0	5.4	30	68
6	YUNN5406	Ch	343.0	52.5	5.5	30	64
7	YUNN5409	Clt	321.9	46.0	20.0	20	48
8	YUNN5410	Ch	320.6	61.3	3.8	40	64



**Number of data points: 8**

**Fisher statistics**

Mean vector: 2.63 / 29.8  
Resultant vector: 7.83  
(X= 6.79, Y= 0.31, Z= 3.89)  
Estimated precision, k: 40.93  
95% Confidence limit: 8.76  
95% Confidence limit, approximation: 7.74

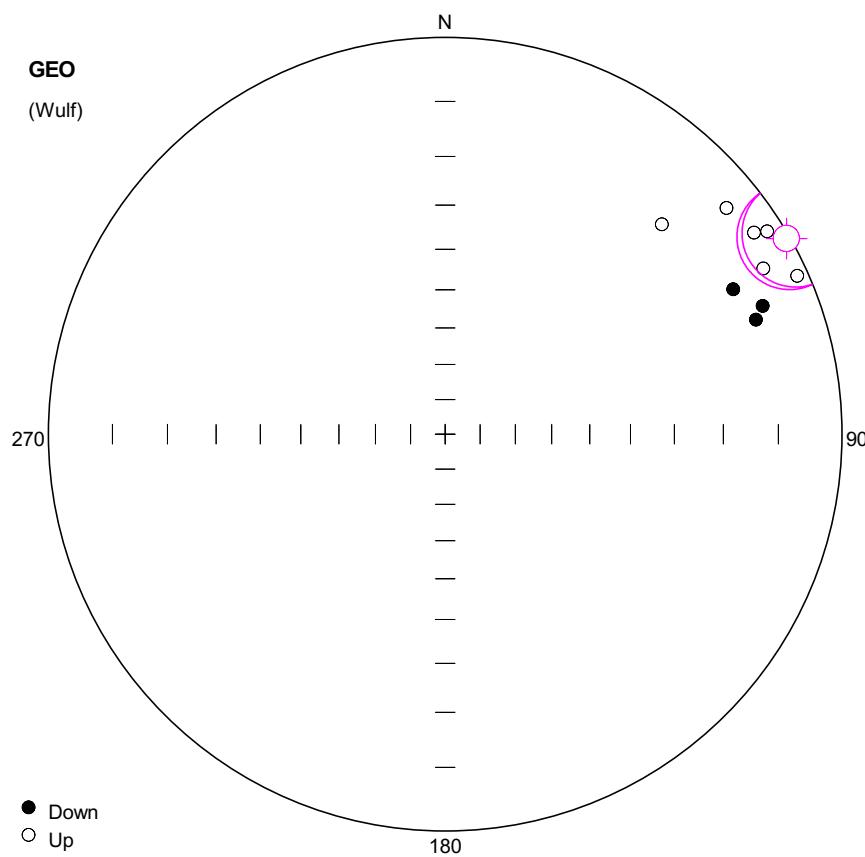
**Orientation matrix**

1st eigenvalue: 0.96  
1st eigenvector: 2.65 / 29.73  
2nd eigenvalue: 0.03  
2nd eigenvector: 199.54 / 59.17  
3rd eigenvalue: 0.01  
3rd eigenvector: 96.92 / 7.43

**VGP**

Site latitude: 65  
Site longitude: 14.04  
Pole latitude: 40.95 (-40.95)  
Pole longitude: -169.3 (10.7)  
Paleolatitude: 15.98 (-15.98)  
dp: 5.38  
dm: 9.71

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5401	Ch	357.2	23.3	12.1	20	52
2	YUNN5402	Ch	8.0	33.4	7.9	20	60
3	YUNN5403	Ch	6.5	26.0	3.7	20	60
4	YUNN5404	Ch	9.4	16.3	8.0	35	68
5	YUNN5405	Ch	0.3	16.8	5.4	30	68
6	YUNN5406	Ch	7.1	34.6	5.5	30	64
7	YUNN5409	Clt	348.7	37.7	20.0	20	48
8	YUNN5410	Ch	2.6	49.2	3.8	40	64



**Number of data points: 9**

**Fisher statistics**

Mean vector: 60.17 / -0.49

Resultant vector: 8.83

(X= 4.39, Y= 7.66, Z= -0.08)

Estimated precision, k: 46.48

95% Confidence limit: 7.63

95% Confidence limit, approximation: 6.84

**Orientation matrix**

1st eigenvalue: 0.96

1st eigenvector: 240.23 / 0.46

2nd eigenvalue: 0.03

2nd eigenvector: 149.67 / 49.97

3rd eigenvalue: 0

3rd eigenvector: 330.61 / 40.02

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 11.91 (-11.91)

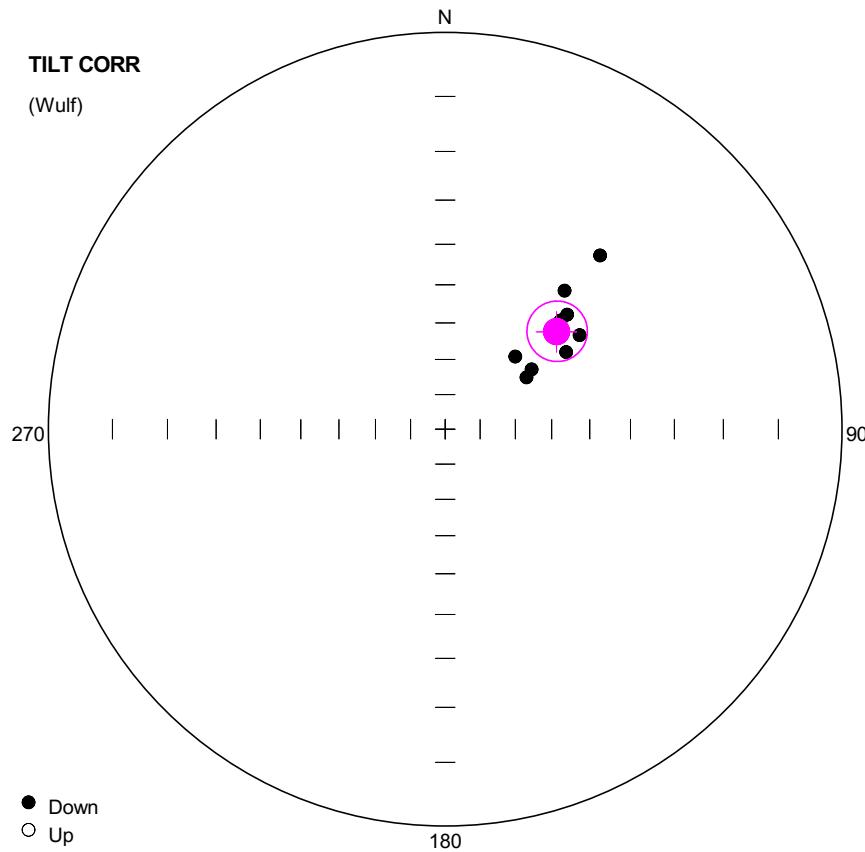
Pole longitude: 131.59 (-48.41)

Paleolatitude: -0.25 (0.25)

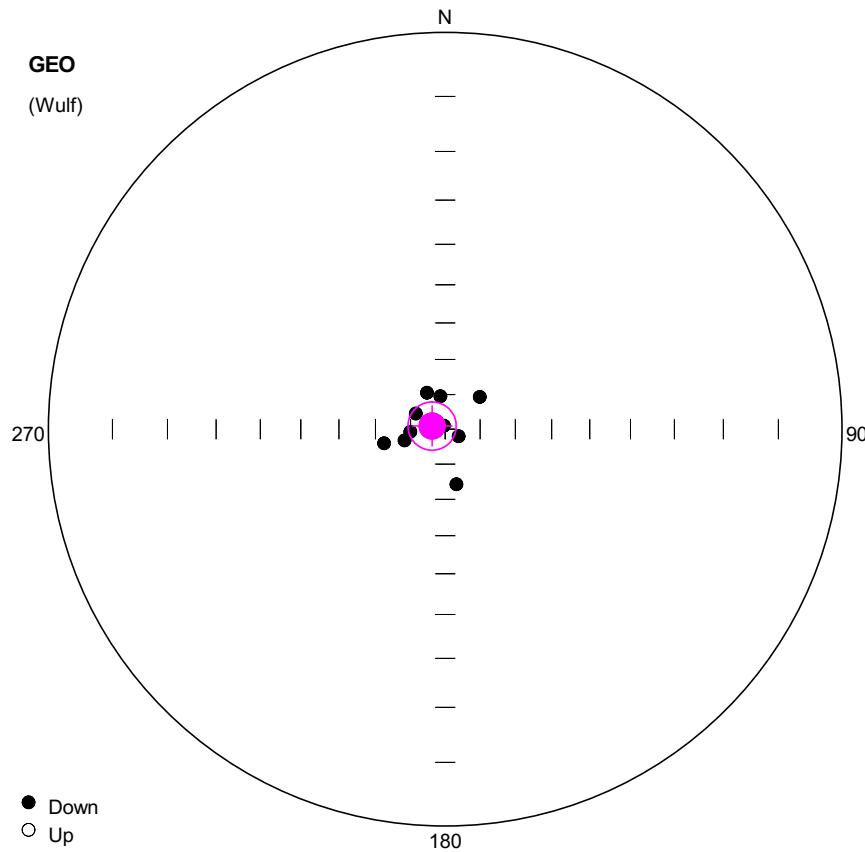
dp: 3.82

dm: 7.63

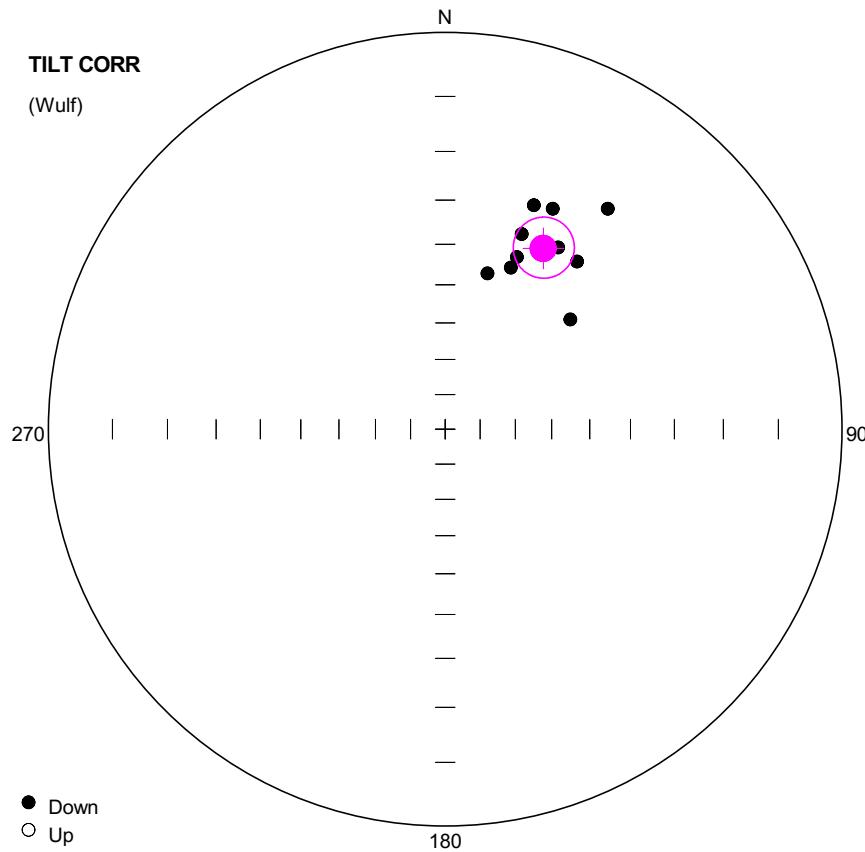
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5502	Ch	45.9	-15.5	7.2	30	68
2	YUNN5503	Ch	69.8	10.3	4.5	35	68
3	YUNN5504	Ch	63.3	11.8	14.0	20	68
4	YUNN5505	Ch	57.8	-2.4	10.2	30	68
5	YUNN5506	Ch	62.5	-5.8	6.6	30	68
6	YUNN5507	Cht	65.8	-1.6	5.7	48	68
7	YUNN5508	Ch	56.9	-4.2	8.5	30	68
8	YUNN5509	Ch	68.0	8.4	10.0	30	68
9	YUNN5510	Ch	51.2	-5.4	16.8	30	64



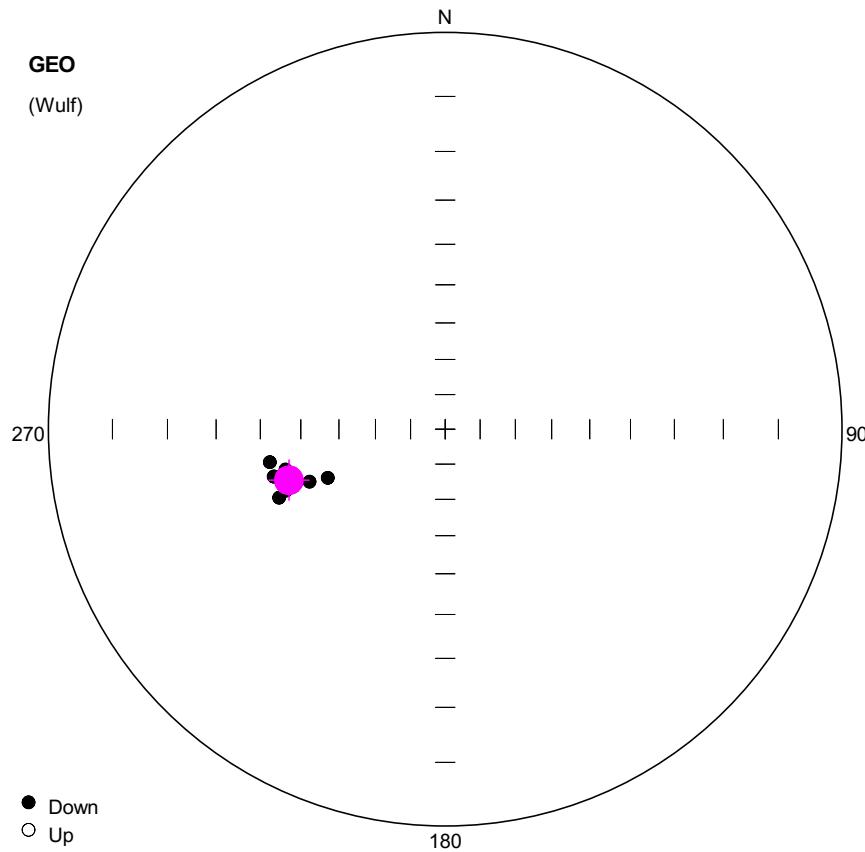
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5502	Ch	41.7	29.2	7.2	30	68
2	YUNN5503	Ch	57.6	62.7	4.5	35	68
3	YUNN5504	Ch	44.0	61.5	14.0	20	68
4	YUNN5505	Ch	46.8	46.4	10.2	30	68
5	YUNN5506	Ch	55.0	45.1	6.6	30	68
6	YUNN5507	Cht	57.5	50.2	5.7	48	68
7	YUNN5508	Ch	46.8	44.3	8.5	30	68
8	YUNN5509	Ch	55.4	60.3	10.0	30	68
9	YUNN5510	Ch	40.8	40.5	16.8	30	64



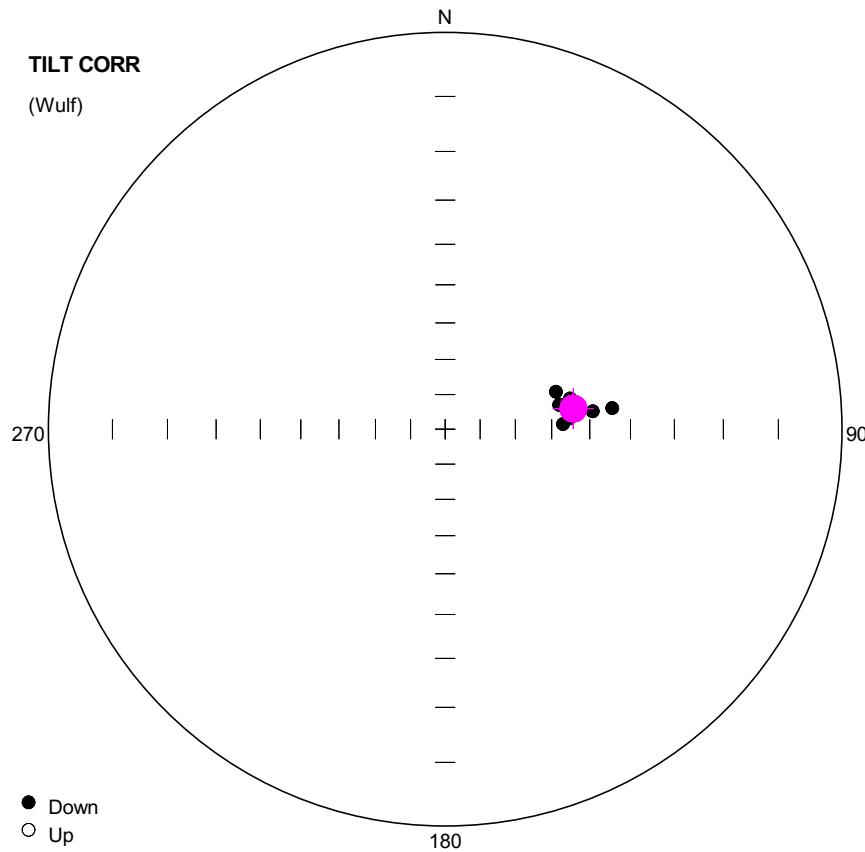
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5601	Ch	254.5	77.9	3.3	30	68
2	YUNN5602	Ch	265.7	79.9	1.0	44	64
3	YUNN5603	Ch	168.5	73.9	5.2	30	64
4	YUNN5604	Ch	117.3	85.6	3.4	20	64
5	YUNN5605	Ch	333.5	78.3	2.6	40	68
6	YUNN5606	Ch	352.1	80.4	3.0	30	64
7	YUNN5607	Ch	46.8	76.4	2.4	35	60
8	YUNN5608	Ch	339.8	88.9	4.4	35	68
9	YUNN5609	Ch	257.0	72.1	5.6	20	68
10	YUNN5610	Ch	297.8	80.4	4.6	20	64



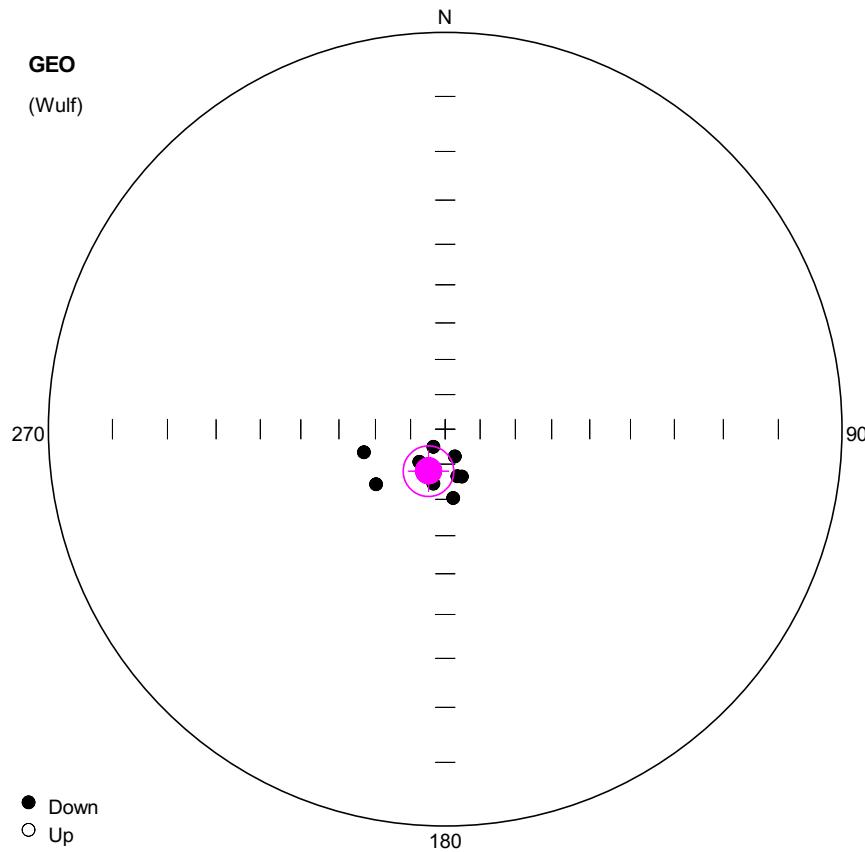
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5601	Ch	22.1	42.6	3.3	30	68
2	YUNN5602	Ch	22.6	39.7	1.0	44	64
3	YUNN5603	Ch	48.8	44.5	5.2	30	64
4	YUNN5604	Ch	38.2	33.5	3.4	20	64
5	YUNN5605	Ch	21.6	27.5	2.6	40	68
6	YUNN5606	Ch	26.0	26.6	3.0	30	64
7	YUNN5607	Ch	36.4	20.8	2.4	35	60
8	YUNN5608	Ch	31.9	33.3	4.4	35	68
9	YUNN5609	Ch	15.2	45.7	5.6	20	68
10	YUNN5610	Ch	21.4	34.3	4.6	20	64



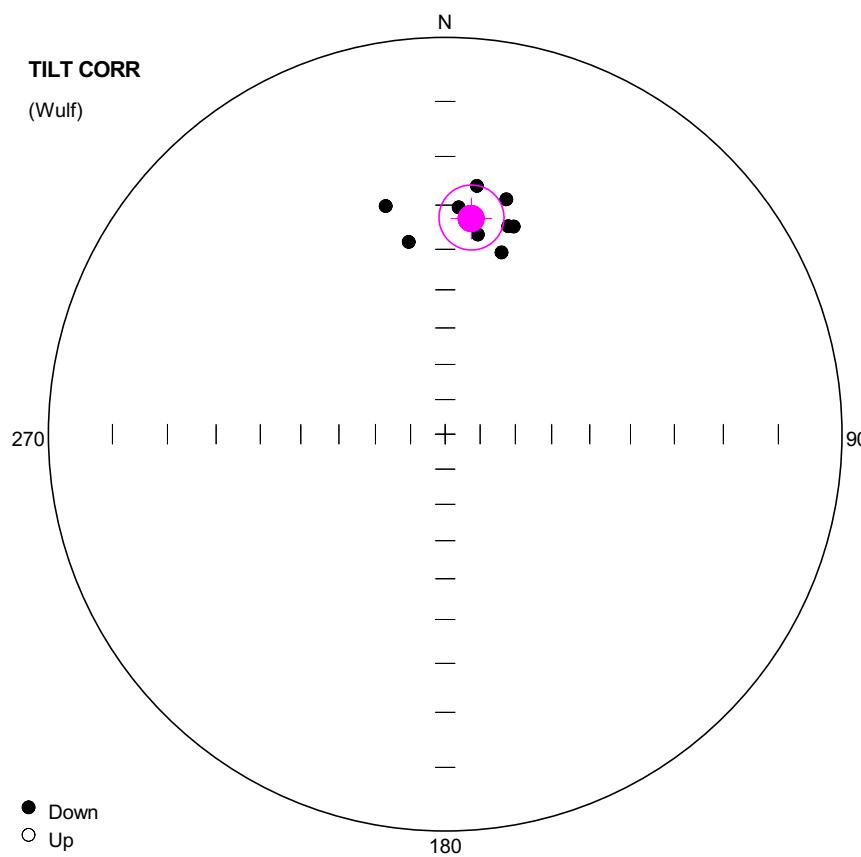
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5701	Ch	248.8	49.7	14.6	20	60
2	YUNN5702	Ch	259.3	41.6	2.2	35	64 OA
3	YUNN5703	Ch	254.5	41.9	7.8	30	60
4	YUNN5704	Ch	254.5	41.7	5.6	20	64
5	YUNN5705	Ch	247.6	41.3	2.9	30	68
6	YUNN5706	Ch	247.5	54.5	6.1	35	68
7	YUNN5708	Ch	249.3	45.5	5.0	44	64
8	YUNN5709	Ch	255.8	44.9	6.3	30	60
9	YUNN5710	Ch	248.9	43.4	4.6	35	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5701	Ch	83.1	48.9	14.6	20	60
2	YUNN5702	Ch	71.4	57.2	2.2	35	64 OA
3	YUNN5703	Ch	78.0	57.0	7.8	30	60
4	YUNN5704	Ch	78.0	57.3	5.6	20	64
5	YUNN5705	Ch	87.6	56.9	2.9	30	68
6	YUNN5706	Ch	82.8	44.0	6.1	35	68
7	YUNN5708	Ch	83.8	53.1	5.0	44	64
8	YUNN5709	Ch	76.2	54.1	6.3	30	60
9	YUNN5710	Ch	85.0	55.0	4.6	35	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5801	Ch	173.3	70.2	4.6	20	68
2	YUNN5802	Ch	160.3	81.7	7.1	35	68
3	YUNN5803	Ch	213.9	83.9	10.7	30	68
4	YUNN5804	Ch	165.3	76.0	6.0	35	64
5	YUNN5805	Ch	231.4	64.9	10.5	30	64
6	YUNN5806	Cht	254.2	66.0	5.8	40	68
7	YUNN5807	Ch	218.2	78.0	7.9	30	68
8	YUNN5808	Ch	160.6	75.6	7.6	35	64
9	YUNN5809	Ch	192.2	74.0	9.5	20	68



**Number of data points: 9**

**Fisher statistics**

Mean vector: 6.88 / 32.61

Resultant vector: 8.85

(X= 7.4, Y= 0.89, Z= 4.77)

Estimated precision, k: 52.08

95% Confidence limit: 7.2

95% Confidence limit, approximation: 6.47

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 7.03 / 32.6

2nd eigenvalue: 0.03

2nd eigenvector: 276.43 / 0.93

3rd eigenvalue: 0.01

3rd eigenvector: 184.98 / 57.39

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 42.51 (-42.51)

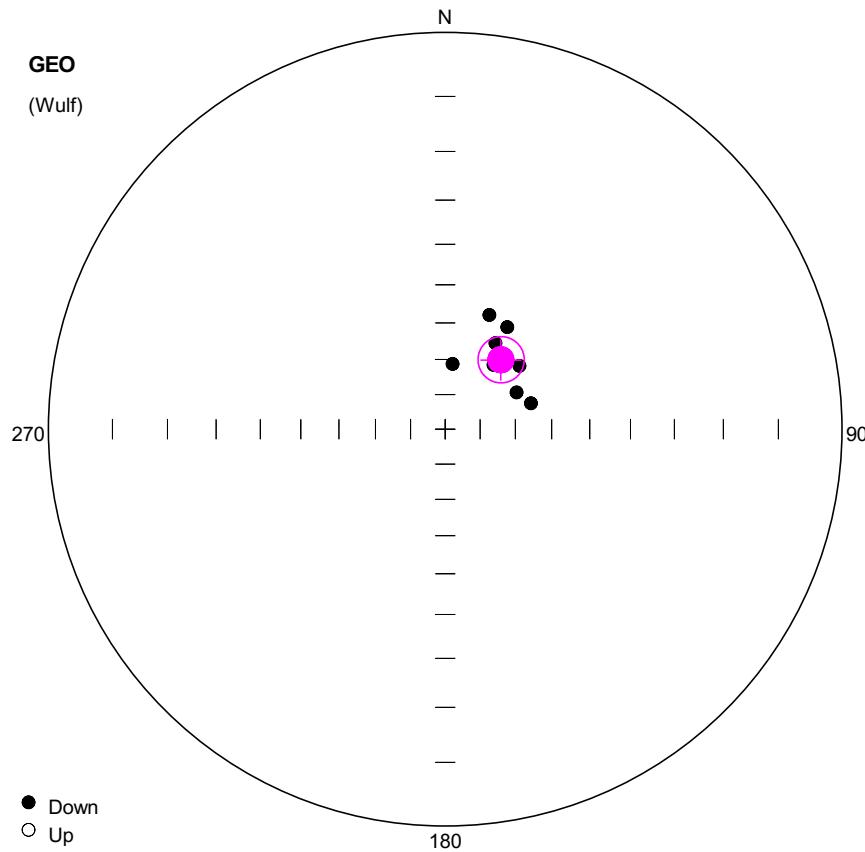
Pole longitude: -174.87 (5.13)

Paleolatitude: 17.74 (-17.74)

dp: 4.6

dm: 8.14

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5801	Ch	17.2	38.8	4.6	20	68
2	YUNN5802	Ch	14.6	27.1	7.1	35	68
3	YUNN5803	Ch	7.3	25.5	10.7	30	68
4	YUNN5804	Ch	16.9	32.6	6.0	35	64
5	YUNN5805	Ch	349.3	37.5	10.5	30	64
6	YUNN5806	Cht	345.4	28.6	5.8	40	68
7	YUNN5807	Ch	3.4	30.4	7.9	30	68
8	YUNN5808	Ch	18.3	32.3	7.6	35	64
9	YUNN5809	Ch	9.3	36.0	9.5	20	68



**Number of data points: 9**

## Fisher statistics

Mean vector: 38.75 / 64.74

Resultant vector: 8.88

(X= 2.95, Y= 2.37, Z= 8.03)

Estimated precision, k: 66.7

95% Confidence limit: 6.35

95% Confidence limit, app

## Orientation matrix

1st eigenvalue: 0.97

1st eigenvector: 38.74 / 64.73

2nd eigenvalue: 0.02

2nd eigenvector: 146.06 / 8

3rd eigenvalue: 0.01

3rd eigenvector: 239.62 / 23.8

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 62.3 (-62.3)

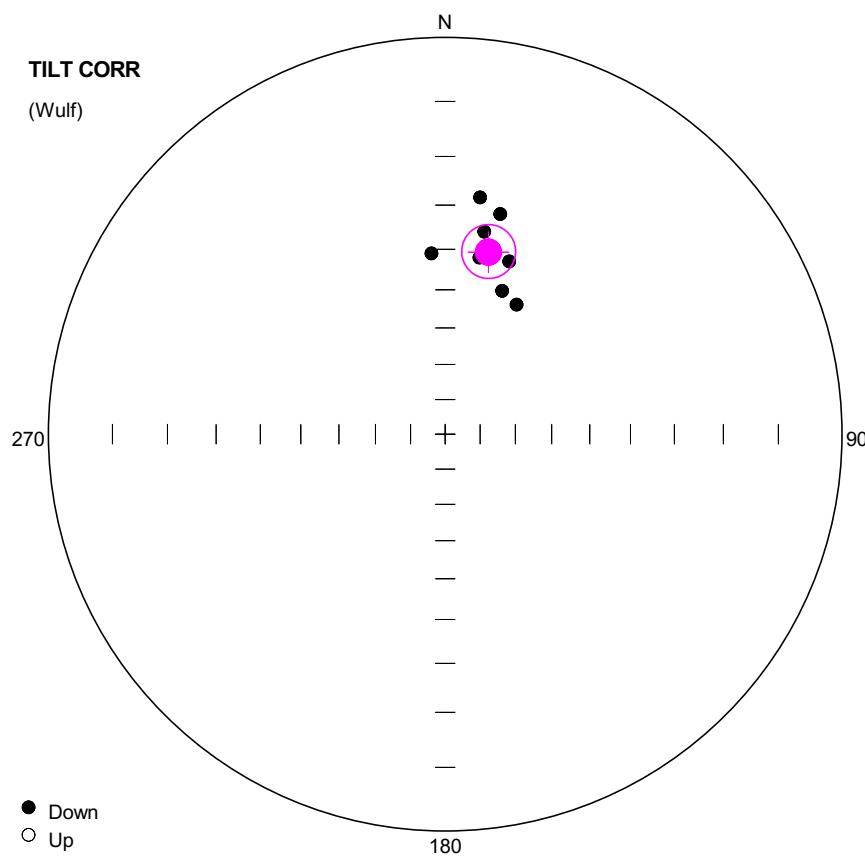
Pole longitude: 126.51 (-53.49)

Paleolatitude: 46.67 (-46.67)

dp: 8.21

dm: 10.21

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5901	Cht	49.6	62.4	4.1	52	68
2	YUNN5902	Cht	73.3	64.6	7.7	52	68
3	YUNN5903	Cht	31.3	56.5	4.1	60	68
4	YUNN5904	Cht	39.5	65.7	2.6	60	68
5	YUNN5905	Cht	37.1	67.1	2.8	60	68
6	YUNN5906	Cht	6.4	71.2	5.2	60	68
7	YUNN5907	Cht	30.4	61.8	2.4	60	68
8	YUNN5908	Cht	21.2	55.7	7.1	60	68
9	YUNN5910	Cht	62.7	67.1	13.0	48	68



**Number of data points: 9**

**Fisher statistics**

Mean vector: 13.35 / 39.52

Resultant vector: 8.88

(X= 6.67, Y= 1.58, Z= 5.65)

Estimated precision, k: 66.68

95% Confidence limit: 6.35

95% Confidence limit, approximation: 5.72

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 13.36 / 39.5

2nd eigenvalue: 0.02

2nd eigenvector: 140.63 / 36.3

3rd eigenvalue: 0.01

3rd eigenvector: 255.36 / 29.66

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 46.53 (-46.53)

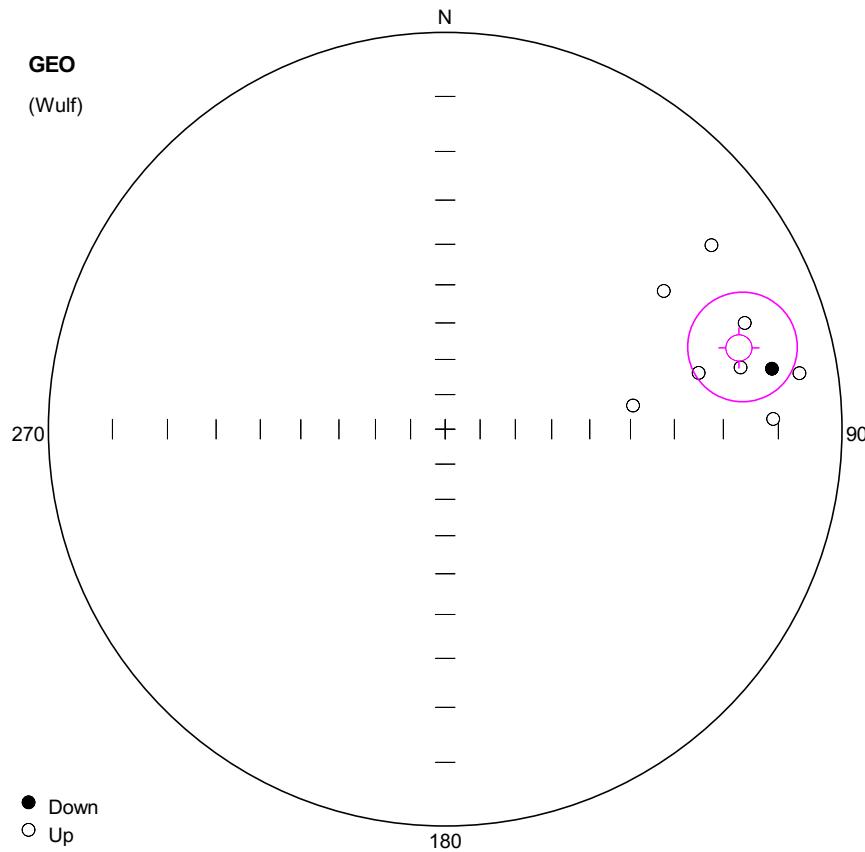
Pole longitude: 175.96 (-4.04)

Paleolatitude: 22.41 (-22.41)

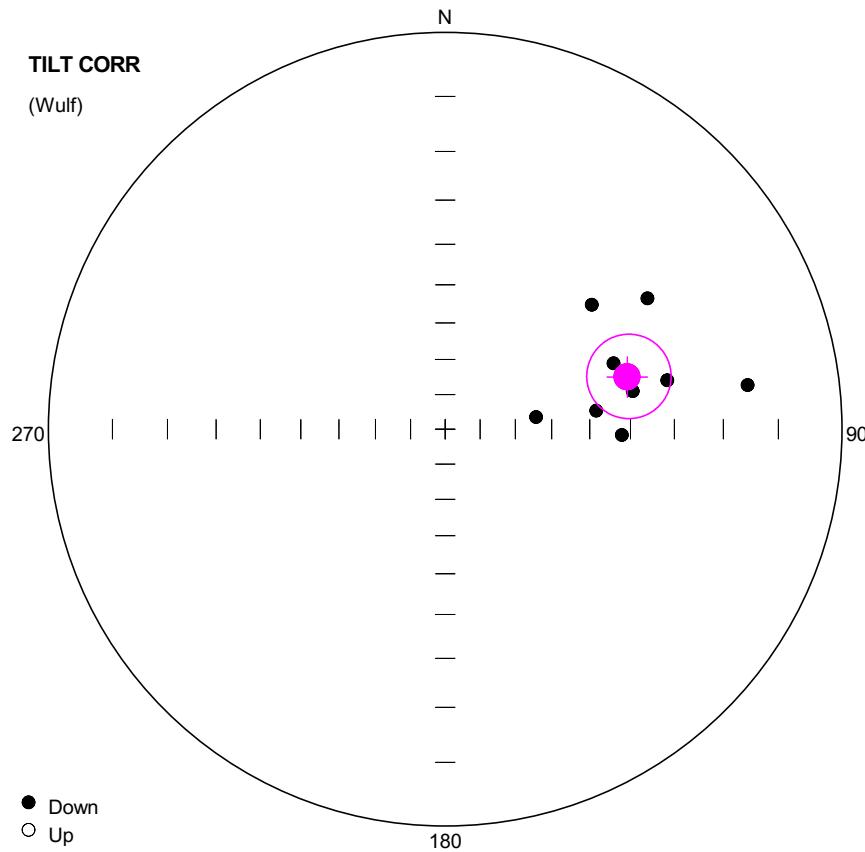
dp: 4.56

dm: 7.61

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN5901	Cht	20.3	40.2	4.1	52	68
2	YUNN5902	Cht	28.8	49.1	7.7	52	68
3	YUNN5903	Cht	14.1	30.5	4.1	60	68
4	YUNN5904	Cht	13.0	40.5	2.6	60	68
5	YUNN5905	Cht	11.0	41.2	2.8	60	68
6	YUNN5906	Cht	355.7	40.9	5.2	60	68
7	YUNN5907	Cht	10.9	35.1	2.4	60	68
8	YUNN5908	Cht	8.4	27.8	7.1	60	68
9	YUNN5910	Cht	21.7	47.5	13.0	48	68



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6001	Ch	55.4	-11.6	18.5	30	64
2	YUNN6002	Clt	57.7	-23.8	9.3	20	52
3	YUNN6003	Cht	70.5	-12.6	17.7	48	68
4	YUNN6004	Ch	78.2	-15.5	14.2	20	64
5	YUNN6005	Cht	74.6	-14.8	23.0	40	64
6	YUNN6006	Cht	77.5	-23.6	24.6	44	64
7	YUNN6007	Cht	79.5	10.1	9.2	56	64
8	YUNN6008	Cht	82.8	-39.0	29.8	56	68
9	YUNN6009	Ch	81.0	-5.8	9.0	35	60
10	YUNN6010	Cht	88.2	-10.8	8.3	44	60 OA



**Number of data points: 10**

## Fisher statistics

Mean vector: 73.93 / 38.98

Resultant vector: 9.64

(X= 2.07, Y= 7.2, Z= 6.06)

Estimated precision, k: 24.88

95% Confidence limit: 9.88

95% Confidence limit, approximation: 8.88

## Orientation matrix

1st eigenvalue: 0.93

1st eigenvector: 73.95 / 38.98

2nd eigenvalue: 0.04

2nd eigenvector: 236.92 / 49.76

3rd eigenvalue: 0.03

3rd eigenvector: 337.04 / 8.46

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 26.64 (-26.64)

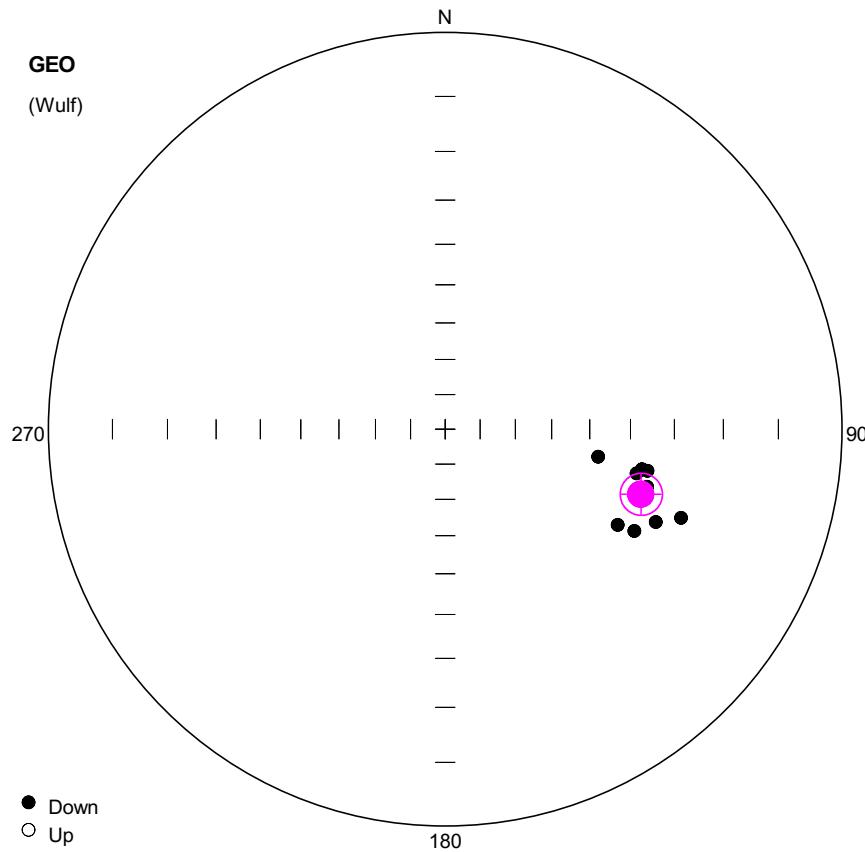
Pole longitude: 108.79 (-71.21)

Paleolatitude: 22.03 (-22.03)

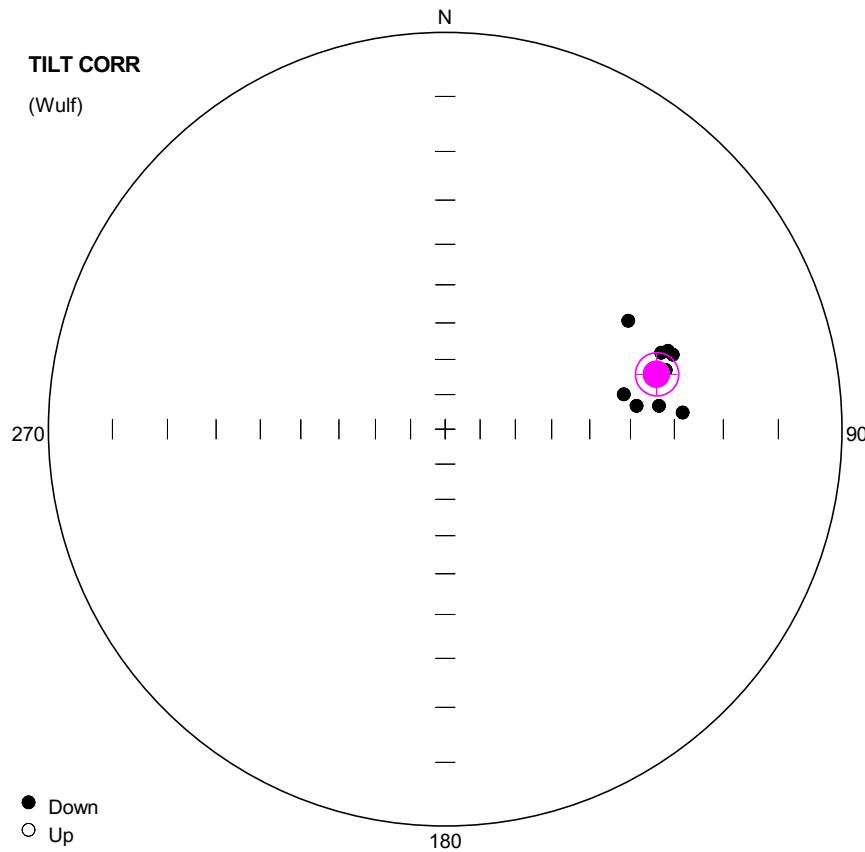
dp: 7.02

dm: 11.78

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6001	Ch	49.7	38.3	18.5	30	64
2	YUNN6002	Clt	57.1	27.5	9.3	20	52
3	YUNN6003	Cht	68.6	41.0	17.7	48	68
4	YUNN6004	Ch	78.5	38.5	14.2	20	64
5	YUNN6005	Cht	74.0	39.1	23.0	40	64
6	YUNN6006	Cht	77.6	30.4	24.6	44	64
7	YUNN6007	Cht	82.5	64.0	9.2	56	64
8	YUNN6008	Cht	81.7	14.8	29.8	56	68
9	YUNN6009	Ch	83.0	48.0	9.0	35	60
10	YUNN6010	Cht	91.9	42.0	8.3	44	60 OA



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6101	Ch	110.6	25.2	4.8	40	52
2	YUNN6102	Ch	105.9	34.2	3.5	35	56
3	YUNN6103	Ch	103.0	37.3	4.6	40	56
4	YUNN6104	Ch	119.0	37.1	8.0	35	56
5	YUNN6105	Ch	101.6	35.0	7.1	35	56
6	YUNN6106	Ch	107.5	34.2	10.5	40	56
7	YUNN6107	Ch	113.8	29.8	6.4	40	56
8	YUNN6108	Ch	101.5	36.3	3.6	35	56
9	YUNN6109	Ch	118.3	33.1	8.5	40	56
10	YUNN6110	Ch	100.2	47.2	2.5	40	56



**Number of data points: 10**

## Fisher statistics

Mean vector: 75.45 / 32.37

Resultant vector: 9.91

(X= 2.1, Y= 8.1, Z= 5.31)

Estimated precision, k: 102.82

95% Confidence limit: 4.7%

95% Confidence limit, approximation: 4.37

## Orientation matrix

1st eigenvalue: 0.98

1st eigenvector: 75.47 / 32.36

2nd eigenvalue: 0.01

2nd eigenvector: 170.88 / 8.47

3rd eigenvalue: 0.01

3rd eigenvector: 273.78 / 56.28

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 22.03 (-22.03)

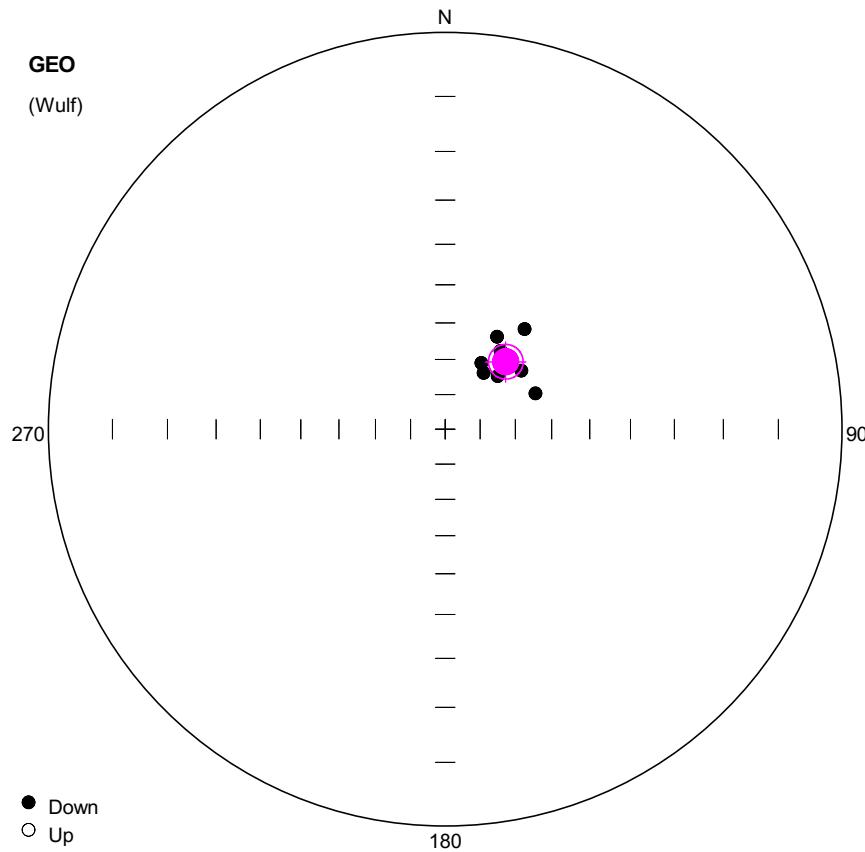
Pole longitude: 109.57 (-70.43)

Paleolatitude: 17.59 (-17.59)

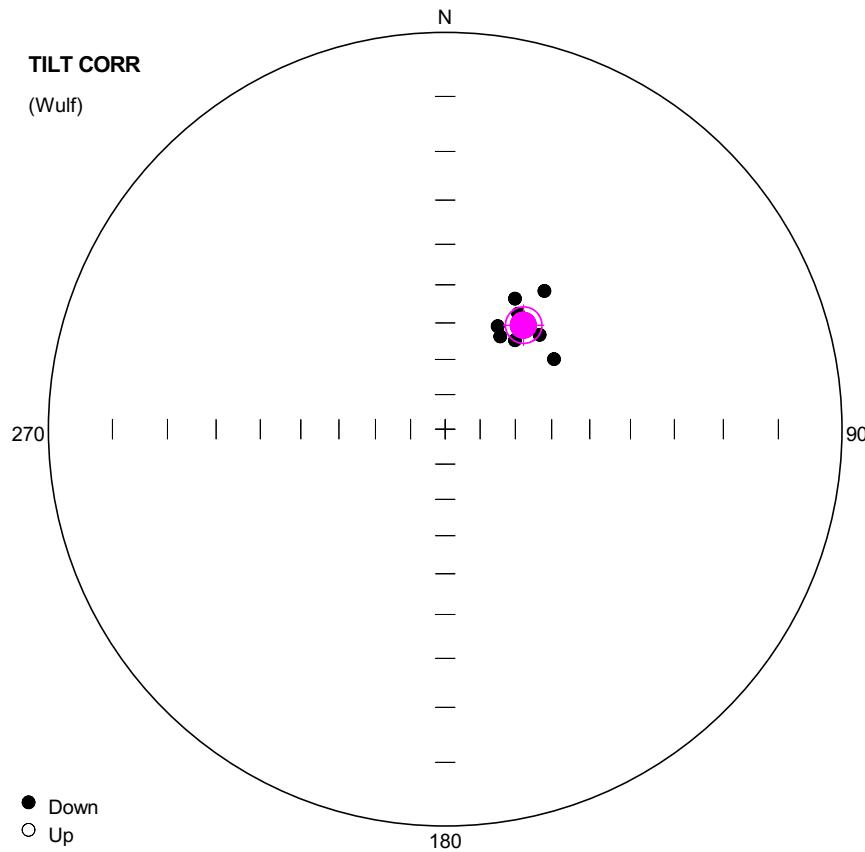
dp: 3.05

dm: 5.4

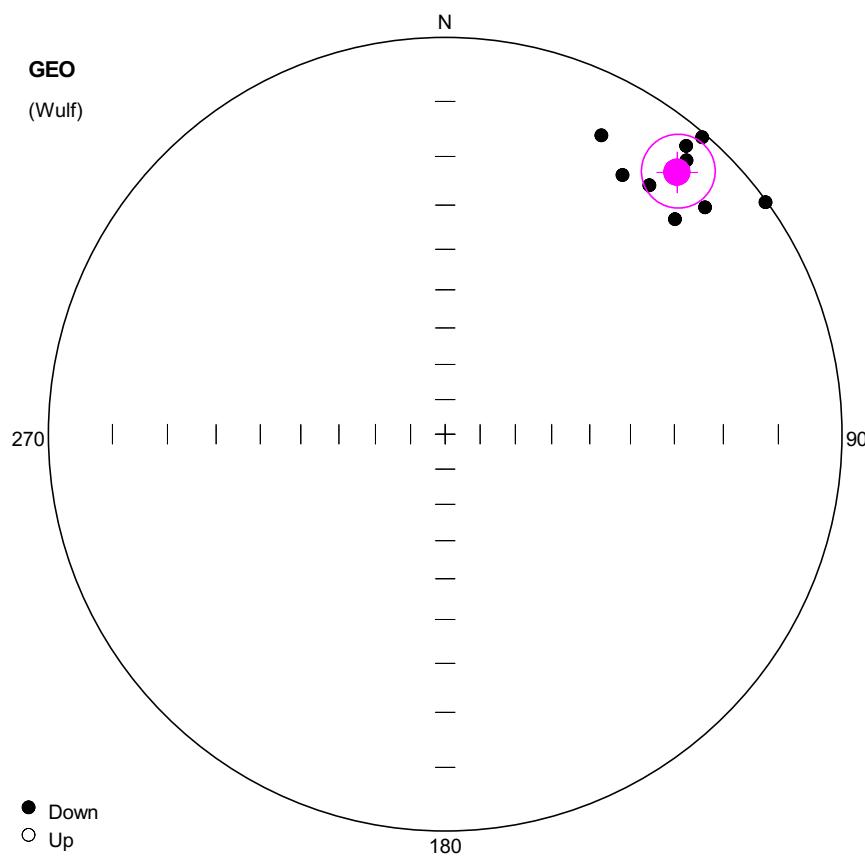
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6101	Ch	86.0	28.1	4.8	40	52
2	YUNN6102	Ch	75.0	30.2	3.5	35	56
3	YUNN6103	Ch	70.5	30.0	4.6	40	56
4	YUNN6104	Ch	79.0	40.7	8.0	35	56
5	YUNN6105	Ch	71.9	27.8	7.1	35	56
6	YUNN6106	Ch	75.9	31.3	10.5	40	56
7	YUNN6107	Ch	83.8	33.1	6.4	40	56
8	YUNN6108	Ch	70.6	28.5	3.6	35	56
9	YUNN6109	Ch	83.1	38.2	8.5	40	56
10	YUNN6110	Ch	59.4	33.6	2.5	40	56



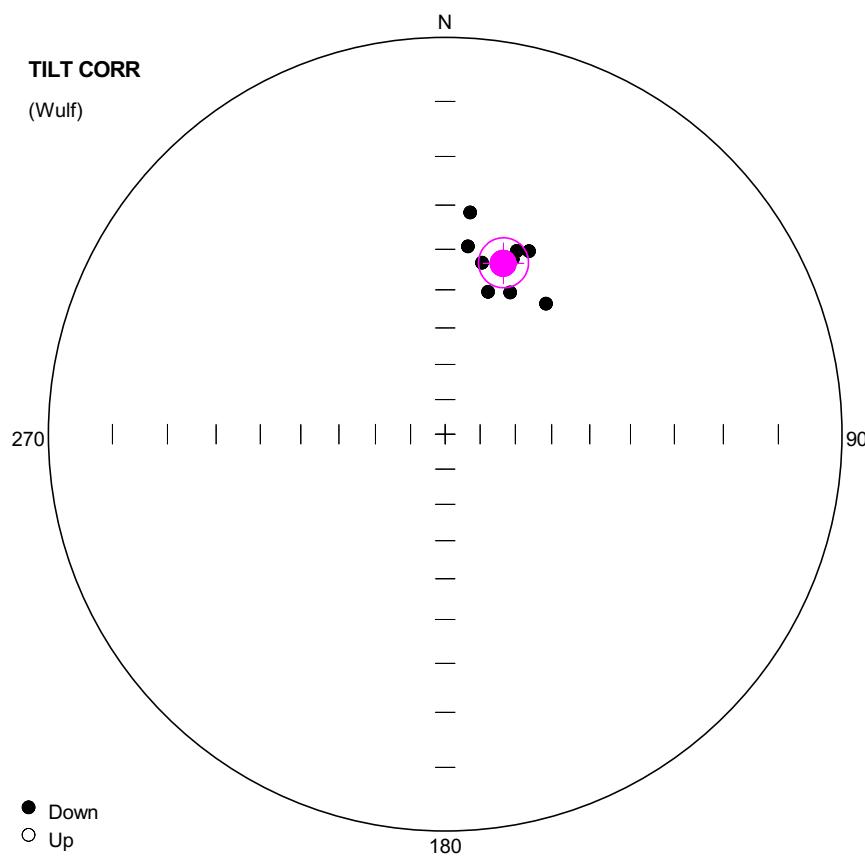
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6201	Ch	43.8	64.1	3.0	40	64
2	YUNN6202	Ch	38.5	54.3	1.9	40	64
3	YUNN6203	Ch	34.2	70.5	6.0	30	60
4	YUNN6204	Ch	52.4	62.8	3.5	40	64
5	YUNN6206	Ch	28.6	68.5	4.2	35	68
6	YUNN6207	Ch	35.2	62.8	3.7	35	68
7	YUNN6208	Ch	68.4	62.5	3.0	30	56
8	YUNN6209	Ch	29.4	60.1	3.4	30	68
9	YUNN6210	Ch	44.7	68.7	5.0	30	56



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6201	Ch	38.4	53.6	3.0	40	64
2	YUNN6202	Ch	35.7	43.6	1.9	40	64
3	YUNN6203	Ch	30.7	59.6	6.0	30	60
4	YUNN6204	Ch	45.1	52.8	3.5	40	64
5	YUNN6206	Ch	27.1	57.5	4.2	35	68
6	YUNN6207	Ch	32.3	52.0	3.7	35	68
7	YUNN6208	Ch	57.2	53.9	3.0	30	56
8	YUNN6209	Ch	28.1	49.1	3.4	30	68
9	YUNN6210	Ch	38.1	58.2	5.0	30	56



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6301	Ch	40.9	0.6	6.3	40	64
2	YUNN6303	Ch	48.9	8.0	5.1	35	56
3	YUNN6304	Ch	39.4	11.9	6.3	40	60
4	YUNN6305	Ch	41.4	4.8	5.7	35	64
5	YUNN6306	Ch	39.9	3.1	5.8	35	56
6	YUNN6307	Ch	34.4	13.3	7.4	40	64
7	YUNN6308	Ch	46.9	13.2	12.8	40	68
8	YUNN6309	Ch	54.1	0.2	7.5	35	68
9	YUNN6310	Ch	27.6	9.3	8.8	35	68



**Number of data points:** 9

**Fisher statistics**

Mean vector: 18.76 / 41.08

Resultant vector: 8.89

(X= 6.35, Y= 2.16, Z= 5.84)

Estimated precision, k: 75.31

95% Confidence limit: 5.97

95% Confidence limit, approximation: 5.38

**Orientation matrix**

1st eigenvalue: 0.98

1st eigenvector: 18.76 / 41.08

2nd eigenvalue: 0.02

2nd eigenvector: 128.84 / 21.5

3rd eigenvalue: 0.01

3rd eigenvector: 239.01 / 41.2

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 46.8 (-46.8)

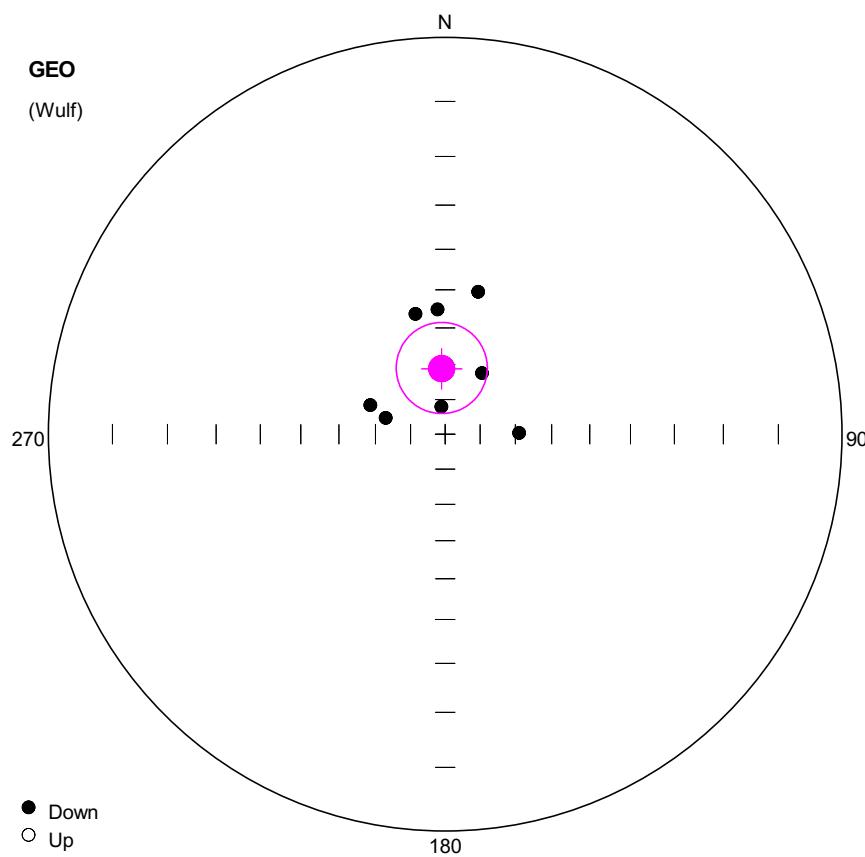
Pole longitude: 168.53 (-11.47)

Paleolatitude: 23.55 (-23.55)

dp: 4.42

dm: 7.26

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6301	Ch	24.6	36.2	6.3	40	64
2	YUNN6303	Ch	24.7	47.1	5.1	35	56
3	YUNN6304	Ch	12.1	42.3	6.3	40	60
4	YUNN6305	Ch	21.1	39.4	5.7	35	64
5	YUNN6306	Ch	21.4	37.2	5.8	35	56
6	YUNN6307	Ch	6.9	39.0	7.4	40	64
7	YUNN6308	Ch	16.8	48.9	12.8	40	68
8	YUNN6309	Ch	37.7	44.9	7.5	35	68
9	YUNN6310	Ch	6.4	31.3	8.8	35	68



**Number of data points: 9**

**Fisher statistics**

Mean vector: 356.97 / 71.24  
 Resultant vector: 8.54  
 $(X= 2.74, Y= -0.15, Z= 8.08)$   
 Estimated precision,  $k$ : 17.3  
 95% Confidence limit: 12.74  
 95% Confidence limit, approximation: 11.22

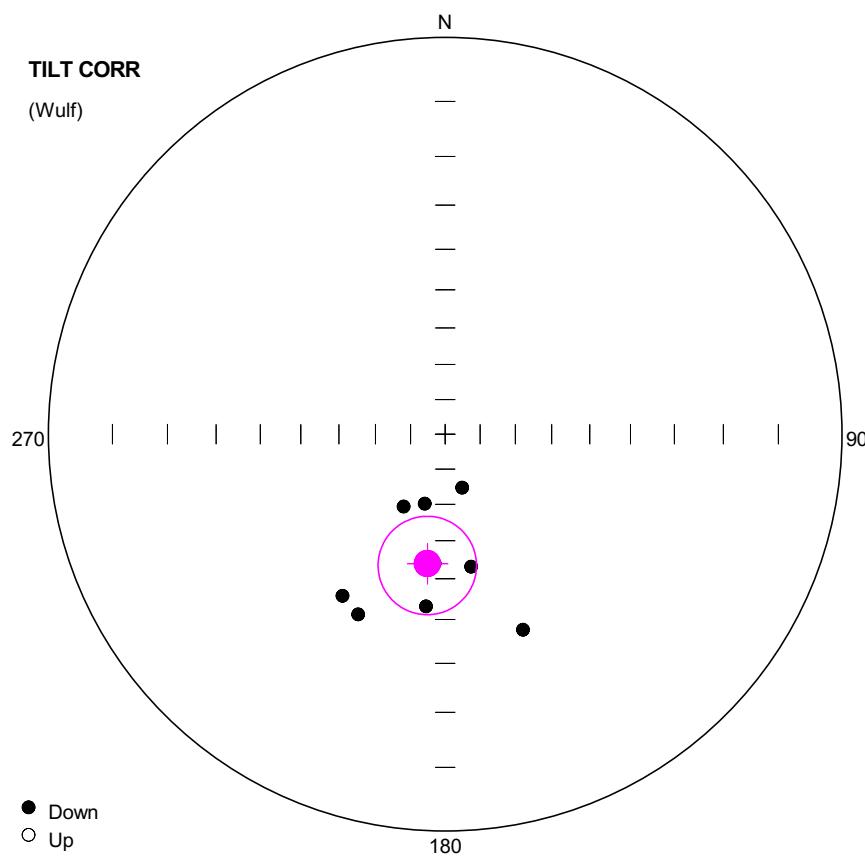
**Orientation matrix**

1st eigenvalue: 0.9  
 1st eigenvector: 356.74 / 71.1  
 2nd eigenvalue: 0.05  
 2nd eigenvector: 204.37 / 16.87  
 3rd eigenvalue: 0.05  
 3rd eigenvector: 111.84 / 8.26

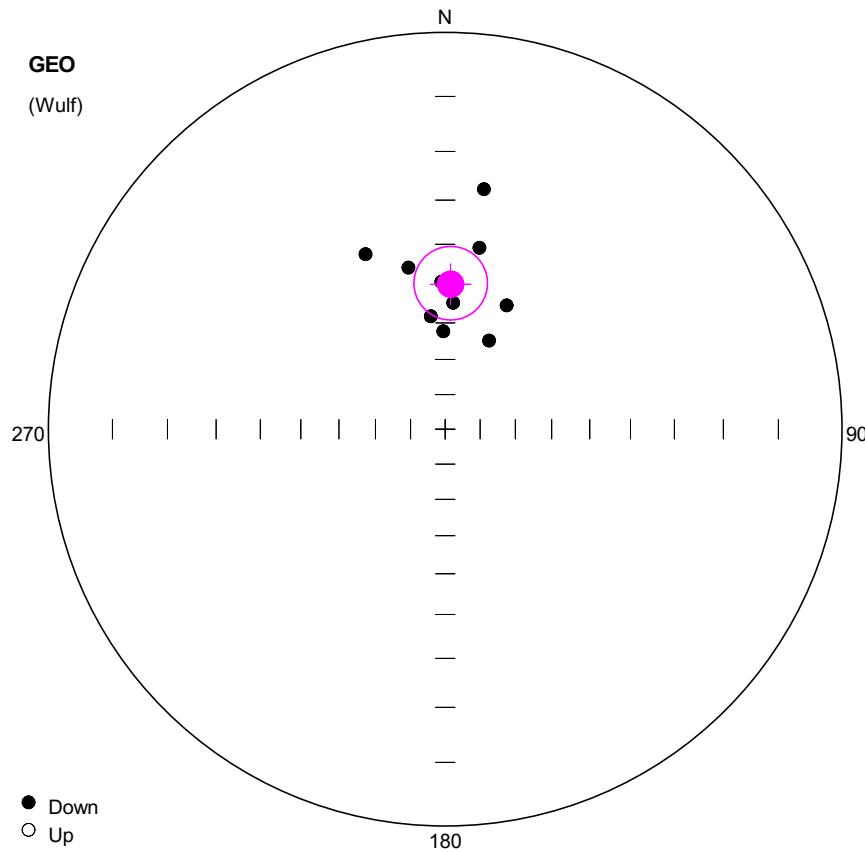
**VGP**

Site latitude: 65  
 Site longitude: 14.04  
 Pole latitude: 80.69 (-80.69)  
 Pole longitude: -155.38 (24.62)  
 Paleolatitude: 55.81 (-55.81)  
 dp: 19.44  
 dm: 22.25

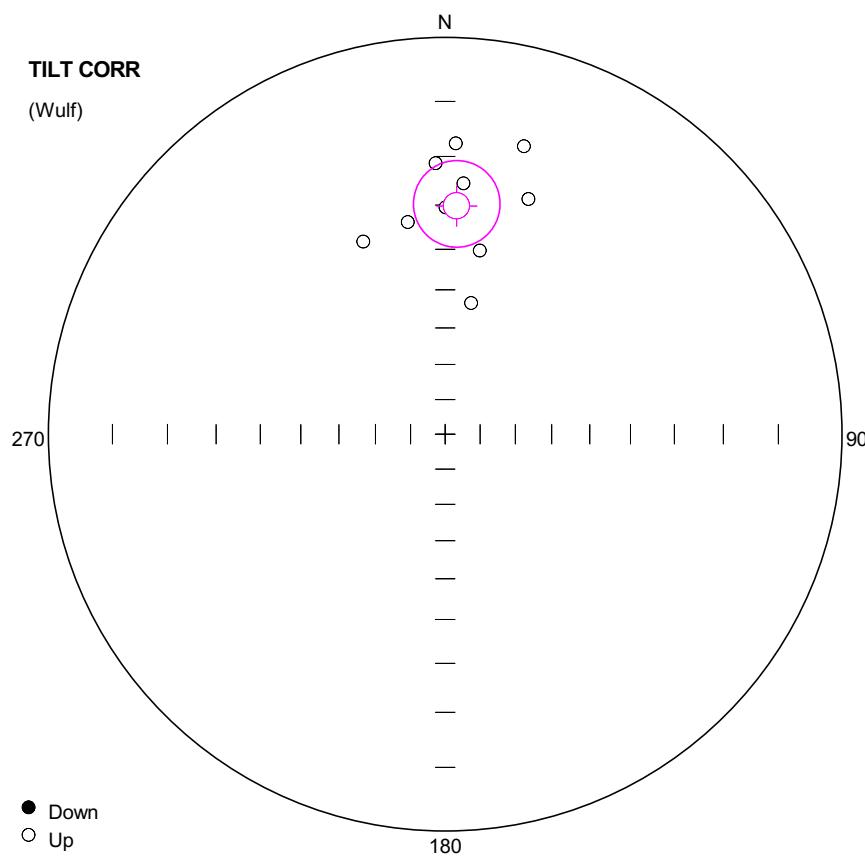
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6401	Cht	2.7	71.2	5.0	56	68
2	YUNN6402	Cht	31.0	69.6	8.0	48	68
3	YUNN6403	Cht	356.5	55.0	0.0	64	68
4	YUNN6404	Cht	346.1	55.3	8.4	56	68
5	YUNN6405	Cht	291.3	67.1	6.5	56	68
6	YUNN6406	Cht	352.1	82.0	7.1	60	68 OA
7	YUNN6407	Cht	13.1	49.6	18.7	60	68
8	YUNN6408	Cht	89.1	68.9	4.0	52	68
9	YUNN6410	Cht	285.3	72.3	13.0	60	68



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6401	Cht	184.7	53.8	5.0	56	68
2	YUNN6402	Cht	168.9	52.4	8.0	48	68
3	YUNN6403	Cht	196.3	69.3	0.0	64	68
4	YUNN6404	Cht	209.8	66.3	8.4	56	68
5	YUNN6405	Cht	212.4	38.5	6.5	56	68
6	YUNN6406	Cht	186.3	42.8	7.1	60	68 OA
7	YUNN6407	Cht	162.2	73.9	18.7	60	68
8	YUNN6408	Cht	158.3	34.1	4.0	52	68
9	YUNN6410	Cht	205.8	36.5	13.0	60	68



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6501	Cht	359.0	62.3	0.0	64	68
2	YUNN6502	Cht	352.7	58.0	0.0	64	68
3	YUNN6503	Cht	347.2	44.7	0.0	64	68
4	YUNN6504	Cht	358.5	49.2	0.0	64	68
5	YUNN6505	Cht	335.5	38.3	0.0	64	68
6	YUNN6506	Cht	3.7	54.6	0.0	64	68
7	YUNN6507	Cht	9.2	27.0	0.0	64	68
8	YUNN6508	Cht	26.4	62.0	0.0	64	68
9	YUNN6509	Cht	10.7	40.1	0.0	64	68
10	YUNN6510	Cht	26.4	51.6	8.8	60	68



**Number of data points: 10**

**Fisher statistics**

Mean vector: 2.86 / -30.11

Resultant vector: 9.68

(X= 8.36, Y= 0.42, Z= -4.85)

Estimated precision, k: 27.7

95% Confidence limit: 9.34

95% Confidence limit, approximation: 8.41

**Orientation matrix**

1st eigenvalue: 0.94

1st eigenvector: 182.88 / 29.96

2nd eigenvalue: 0.04

2nd eigenvector: 59.21 / 43.89

3rd eigenvalue: 0.03

3rd eigenvector: 293.4 / 31.31

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 8.8 (-8.8)

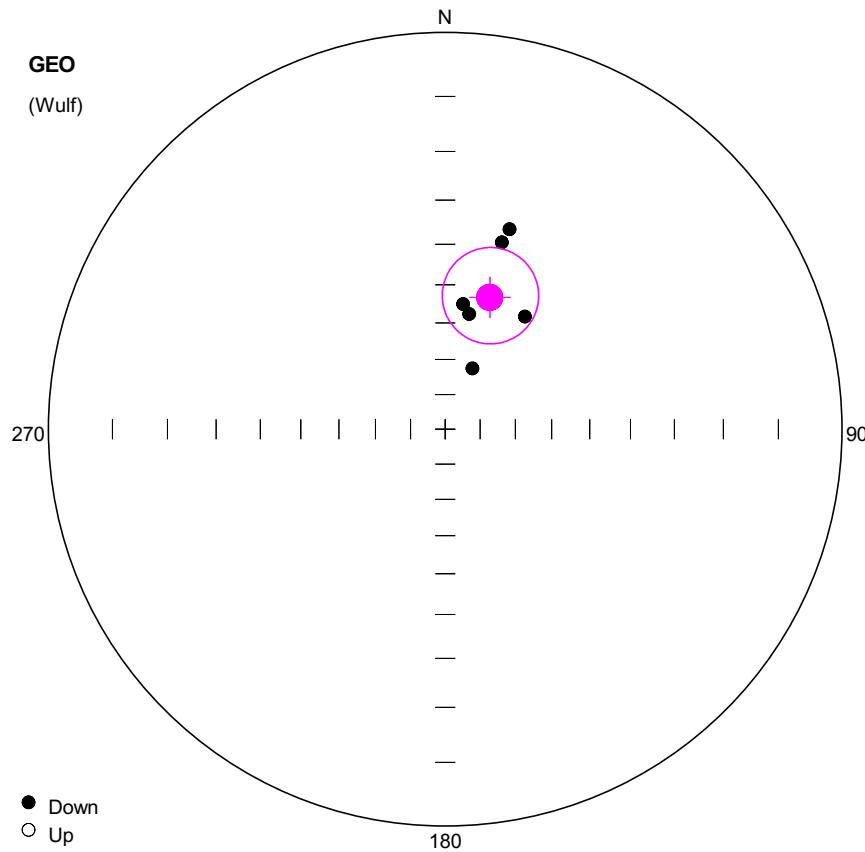
Pole longitude: -168.73 (11.27)

Paleolatitude: -16.17 (16.17)

dp: 5.76

dm: 10.37

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6501	Cht	2.1	-17.5	0.0	64	68
2	YUNN6502	Cht	358.0	-21.3	0.0	64	68
3	YUNN6503	Cht	350.0	-33.0	0.0	64	68
4	YUNN6504	Cht	0.1	-30.5	0.0	64	68
5	YUNN6505	Cht	337.0	-34.4	0.0	64	68
6	YUNN6506	Cht	4.2	-25.3	0.0	64	68
7	YUNN6507	Cht	11.2	-52.8	0.0	64	68
8	YUNN6508	Cht	15.3	-16.1	0.0	64	68
9	YUNN6509	Cht	10.7	-39.6	0.0	64	68
10	YUNN6510	Cht	19.5	-25.7	8.8	60	68



**Number of data points: 6**

## Fisher statistics

Mean vector: 18.69 / 51.34

Resultant vector: 5.84

(X= 3.45, Y= 1.17, Z= 4.56)

Estimated precision, k: 30.53

95% Confidence limit:

95% Confidence limit, approximation: 10.34

## Orientation matrix

1st eigenvalue: 0.95

1st eigenvector: 18.65 / 51.32

2nd eigenvalue: 0.05

2nd eigenvector: 195.07 / 38.62

3rd eigenvalue: 0.01

3rd eigenvector: 286.46 / 1.75

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 55.07 (-55.07)

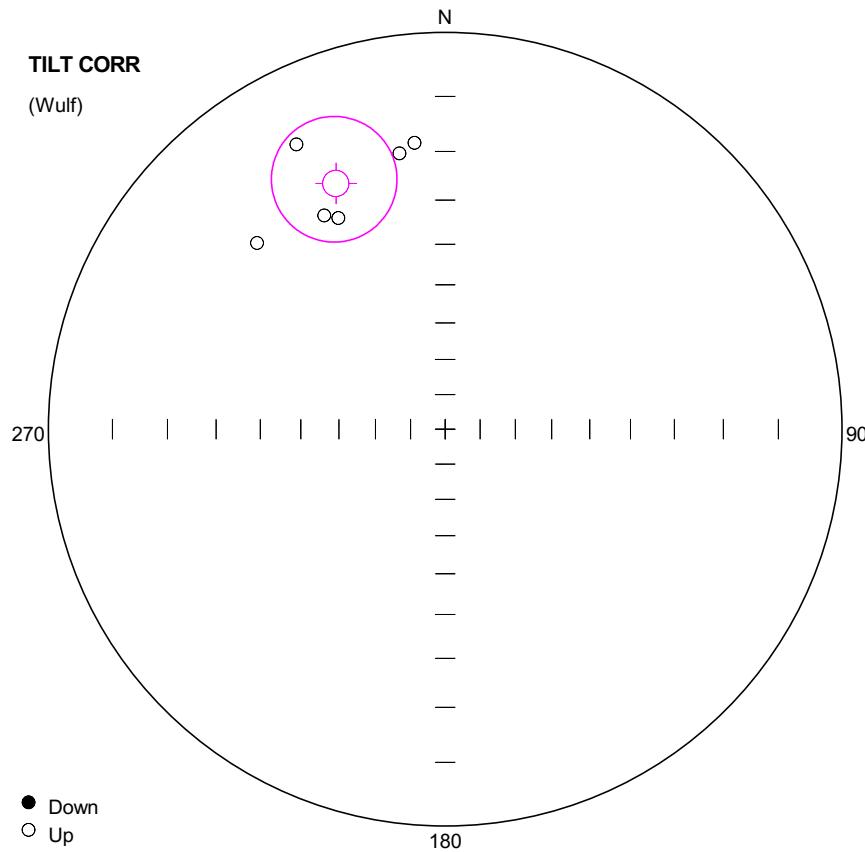
Pole longitude: 165.72 (-14.28)

Paleolatitude: 32.01 (-32.01)

dp: 11.35

dm: 16.72

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6601	Ch	8.2	54.7	4.6	35	68
2	YUNN6602	Ch	11.8	57.0	4.5	40	68
3	YUNN6603	Ch	24.1	71.0	12.8	48	68
4	YUNN6604	Ch	17.8	34.2	8.6	20	56
5	YUNN6605	Ch	35.3	51.7	12.5	30	56
6	YUNN6606	Ch	16.9	37.6	7.9	30	56



**Number of data points: 6**

## Fisher statistics

Mean vector: 336.04 / -21.74

Resultant vector: 5.84

(X= 4.95, Y= -2.2, Z= -2.16)

Estimated precision, k: 30.58

95% Confidence limit: 12.31

95% Confidence limit, approximation: 10.34

## Orientation matrix

1st eigenvalue: 0.95

1st eigenvector: 156.06 / 21.76

2nd eigenvalue: 0.05

2nd eigenvector: 62.75 / 8.24

3rd eigenvalue: 0.01

3rd eigenvector: 313.23 / 66.58

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 11.63 (-11.63)

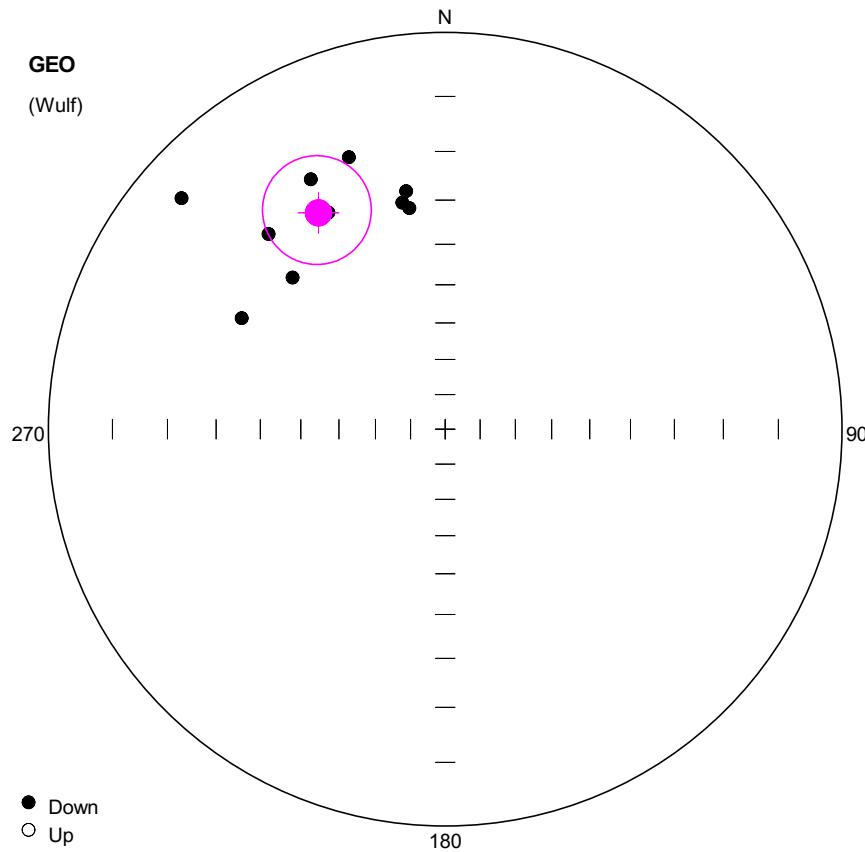
Pole longitude: -141.97 (38.03)

Paleolatitude: -11.28 (11.28)

dp: 6.86

dm: 13

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6601	Ch	333.2	-28.4	4.6	35	68
2	YUNN6602	Ch	330.5	-26.5	4.5	40	68
3	YUNN6603	Ch	314.7	-22.6	12.8	48	68
4	YUNN6604	Ch	353.9	-18.1	8.6	20	56
5	YUNN6605	Ch	332.4	-12.0	12.5	30	56
6	YUNN6606	Ch	350.6	-19.7	7.9	30	56



**Number of data points: 10**

## Fisher statistics

Mean vector: 329.64 / 25.44

Resultant vector: 9.54

(X= 7.44, Y= -4.36, Z= 4.1)

Estimated precision, k: 19.76

95% C

95% Confidence limit, approximation: 9.96

## Orientation matrix

1st eigenvalue: 0.91

1st eigenvector: 329.94 / 25.43

2nd eigenvalue: 0.07

2nd eigenvector: 63.25 / 6.92

3rd eigenvalue: 0.02

3rd eigenvector: 167.35 / 63.51

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 34.37 (-34.37)

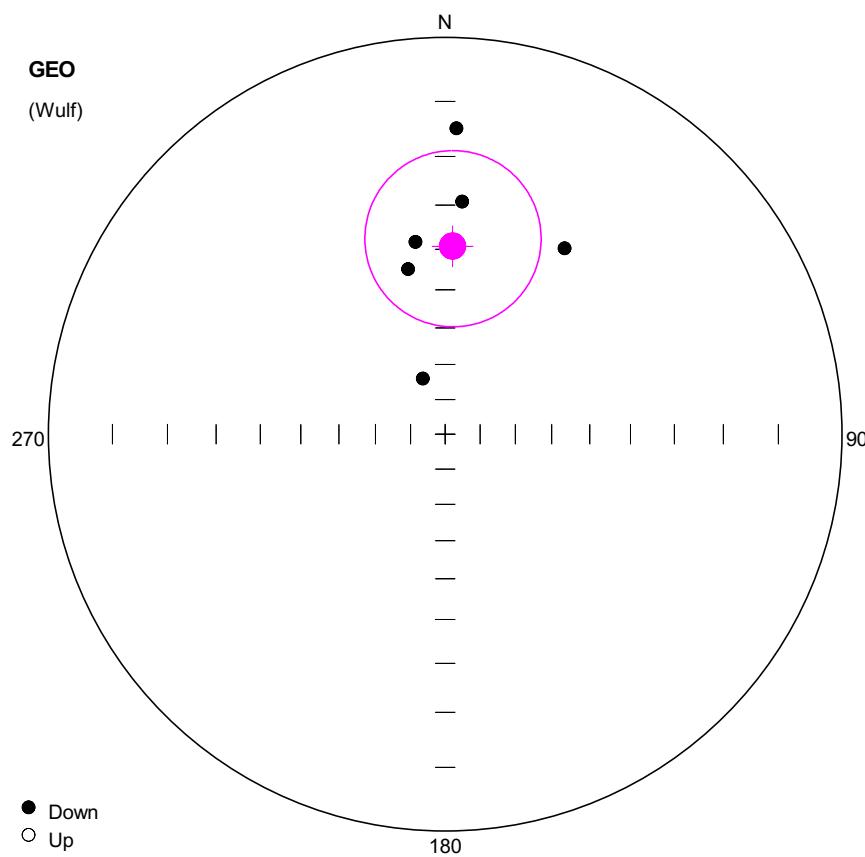
Pole longitude: -129.4 (50.6)

Paleolatitude: 13.38 (-13.38)

dp: 6.47

dm: 12

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6701	Ch	349.3	29.7	3.0	30	64 OA
2	YUNN6702	Ch	311.2	7.1	5.7	30	64 OA
3	YUNN6703	Ch	350.7	27.4	5.1	20	64 OA
4	YUNN6704	Ch	298.6	29.4	6.7	20	64 OA
5	YUNN6705	Ch	317.9	22.9	3.9	20	56 OA
6	YUNN6706	Ch	350.8	31.1	4.0	20	64 OA
7	YUNN6708	Ch	331.7	26.4	3.4	30	64
8	YUNN6709	Ch	340.5	18.0	5.3	35	64 OA
9	YUNN6710	Ch	331.7	18.9	4.3	20	64 OA
10	YUNN6711	Ch	314.8	33.1	5.3	20	68 OA



**Number of data points: 6**

**Fisher statistics**

Mean vector: 2.26 / 39.26  
 Resultant vector: 5.57  
 $(X= 4.31, Y= 0.17, Z= 3.53)$   
 Estimated precision,  $k$ : 11.74  
 95% Confidence limit: 20.39  
 95% Confidence limit, approximation: 16.68

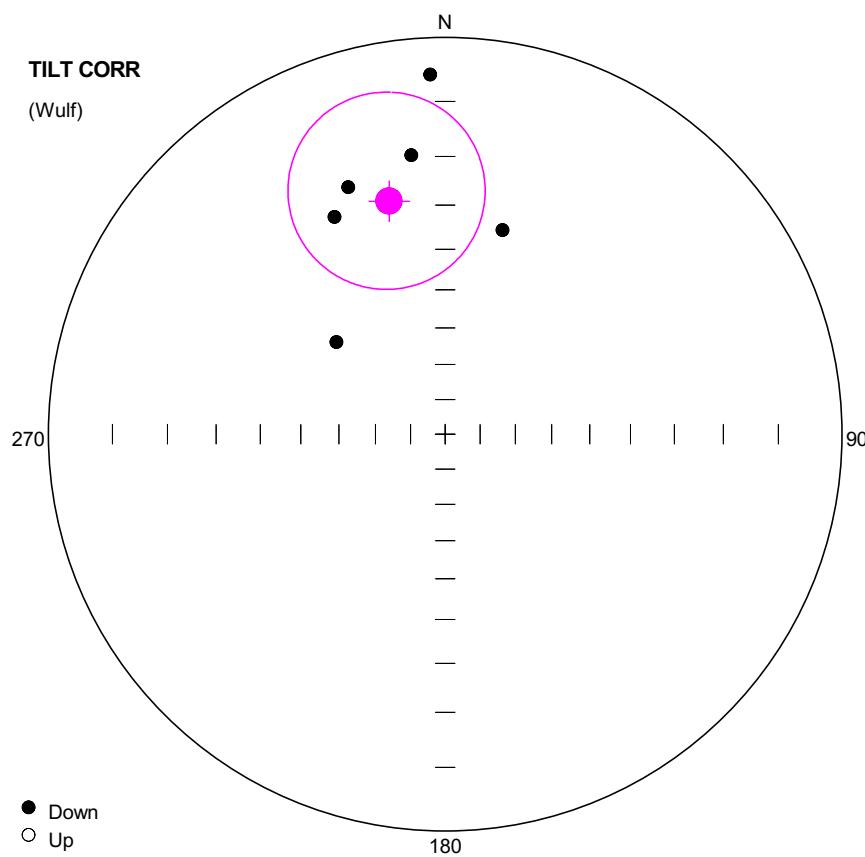
**Orientation matrix**

1st eigenvalue: 0.87  
 1st eigenvector: 2.06 / 38.5  
 2nd eigenvalue: 0.1  
 2nd eigenvector: 211.61 / 47.57  
 3rd eigenvalue: 0.04  
 3rd eigenvector: 104.45 / 15.09

**VGP**

Site latitude: 65  
 Site longitude: 14.04  
 Pole latitude: 47.2 (-47.2)  
 Pole longitude: -169.04 (10.96)  
 Paleolatitude: 22.23 (-22.23)  
 dp: 14.57  
 dm: 24.38

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6801	Ch	2.1	14.7	8.2	20	68 OA
2	YUNN6803	Ch	351.2	37.8	2.2	20	52 OA
3	YUNN6804	Ch	4.2	29.1	5.1	20	64 OA
4	YUNN6805	Ch	347.4	43.8	6.2	20	64 OA
5	YUNN6806	Ch	338.1	72.8	7.3	40	64 OA
6	YUNN6808	Ch	32.7	31.8	10.1	48	64 OA



**Number of data points: 6**

**Fisher statistics**

Mean vector: 346.44 / 27.72  
Resultant vector: 5.57  
(X= 4.8, Y= -1.16, Z= 2.59)  
Estimated precision, k: 11.74  
95% Confidence limit: 20.39  
95% Confidence limit, approximation: 16.68

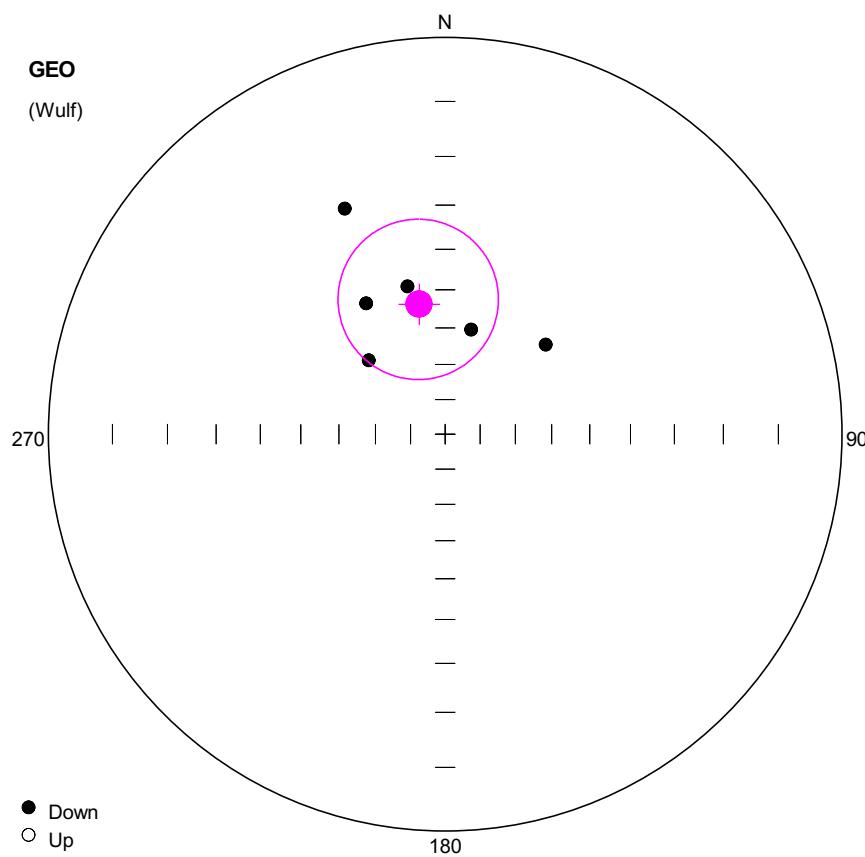
**Orientation matrix**

1st eigenvalue: 0.87  
1st eigenvector: 346.69 / 26.98  
2nd eigenvalue: 0.1  
2nd eigenvector: 233.97 / 37.18  
3rd eigenvalue: 0.04  
3rd eigenvector: 102.87 / 40.91

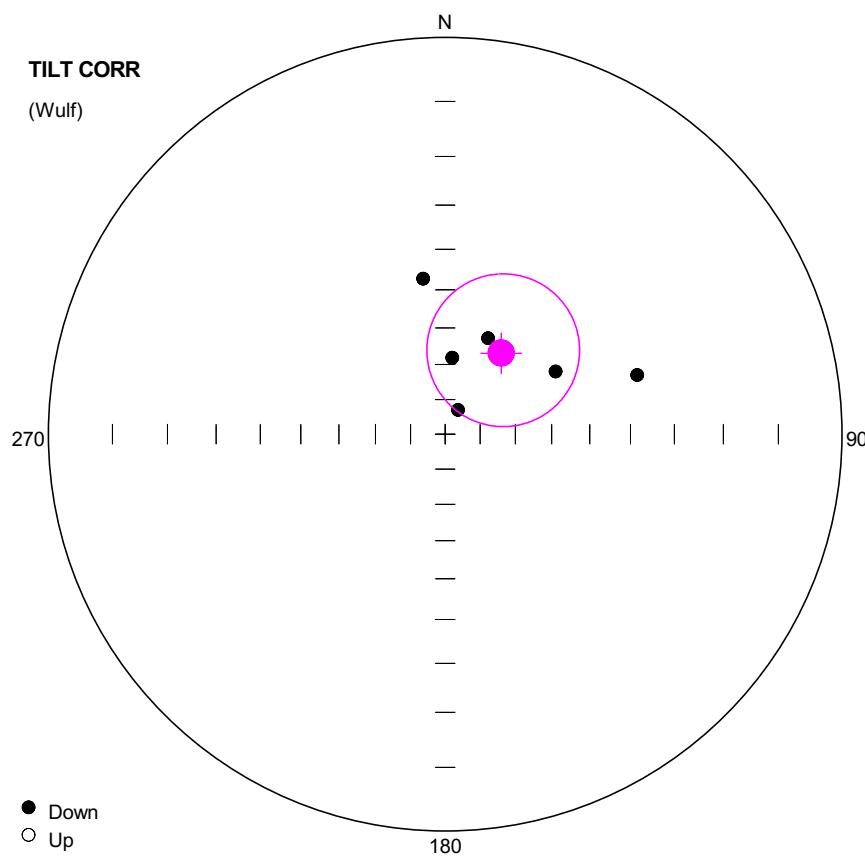
**VGP**

Site latitude: 65  
Site longitude: 14.04  
Pole latitude: 38.88 (-38.88)  
Pole longitude: -149.02 (30.98)  
Paleolatitude: 14.72 (-14.72)  
dp: 12.17  
dm: 22.28

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6801	Ch	357.6	5.6	8.2	20	68 OA
2	YUNN6803	Ch	338.6	22.5	2.2	20	52 OA
3	YUNN6804	Ch	353.1	19.4	5.1	20	64 OA
4	YUNN6805	Ch	333.0	26.9	6.2	20	64 OA
5	YUNN6806	Ch	310.3	50.5	7.3	40	64 OA
6	YUNN6808	Ch	15.7	33.8	10.1	48	64 OA



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6901	Cmt	336.0	26.2	6.1	40	60
2	YUNN6902	Ch	314.0	60.0	7.4	20	68 OA
3	YUNN6903	Ch	14.0	59.6	8.8	20	64 OA
4	YUNN6904	Ch	48.3	52.5	15.8	40	64
5	YUNN6907	Ch	328.9	47.9	4.4	30	64 OA
6	YUNN6908	Ch	345.7	48.0	3.4	20	56 OA



**Number of data points: 6**

**Fisher statistics**

Mean vector: 34.58 / 62.09

Resultant vector: 5.57

(X= 2.15, Y= 1.48, Z= 4.92)

Estimated precision, k: 11.63

95% Confidence limit: 20.5

95% Confidence limit, approximation: 16.76

**Orientation matrix**

1st eigenvalue: 0.87

1st eigenvector: 33.55 / 62.58

2nd eigenvalue: 0.1

2nd eigenvector: 286.23 / 8.78

3rd eigenvalue: 0.03

3rd eigenvector: 191.95 / 25.76

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 61.06 (-61.06)

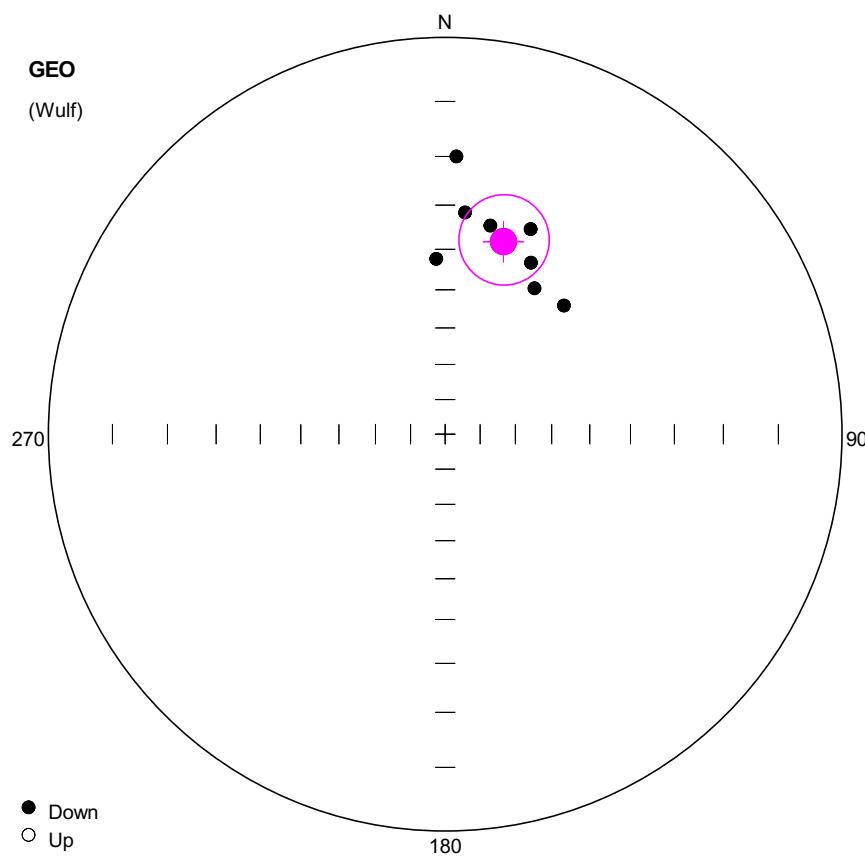
Pole longitude: 135.5 (-44.5)

Paleolatitude: 43.34 (-43.34)

dp: 24.74

dm: 31.85

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN6901	Cmt	351.9	46.8	6.1	40	60
2	YUNN6902	Ch	27.7	82.1	7.4	20	68 OA
3	YUNN6903	Ch	60.4	54.5	8.8	20	64 OA
4	YUNN6904	Ch	72.9	36.3	15.8	40	64
5	YUNN6907	Ch	5.4	68.1	4.4	30	64 OA
6	YUNN6908	Ch	24.1	60.3	3.4	20	56 OA



**Number of data points: 8**

**Fisher statistics**

Mean vector: 16.83 / 36.27

Resultant vector: 7.76

(X= 5.99, Y= 1.81, Z= 4.59)

Estimated precision, k: 29.71

95% Confidence limit: 10.33

95% Confidence limit, approximation: 9.08

**Orientation matrix**

1st eigenvalue: 0.94

1st eigenvector: 16.83 / 36.26

2nd eigenvalue: 0.05

2nd eigenvector: 123.6 / 21.47

3rd eigenvalue: 0.01

3rd eigenvector: 237.57 / 45.93

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 43.78 (-43.78)

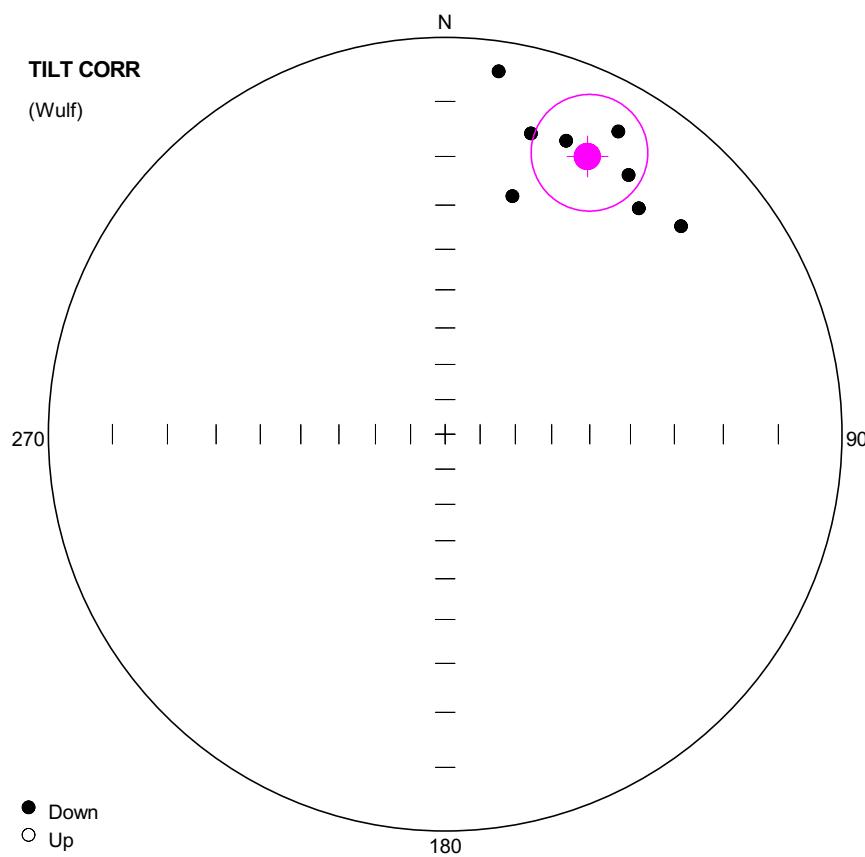
Pole longitude: 171.93 (-8.07)

Paleolatitude: 20.15 (-20.15)

dp: 7

dm: 12.03

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7001	Ch	5.1	31.4	6.8	20	68
2	YUNN7002	Ch	22.6	31.5	4.5	30	64
3	YUNN7003	Ch	26.6	38.4	5.0	35	60
4	YUNN7004	Ch	2.3	20.0	6.9	30	64 OA
5	YUNN7005	Ch	12.2	33.5	5.4	30	64
6	YUNN7006	Ch	42.8	42.4	5.1	40	60
7	YUNN7007	Ch	31.5	43.4	6.7	35	64
8	YUNN7008	Ch	357.1	42.3	12.2	30	64



**Number of data points:** 8

**Fisher statistics**

Mean vector: 27.12 / 13.63

Resultant vector: 7.76

(X= 6.72, Y= 3.44, Z= 1.83)

Estimated precision, k: 29.77

95% Confidence limit: 10.32

95% Confidence limit, approximation: 9.07

**Orientation matrix**

1st eigenvalue: 0.94

1st eigenvector: 27.12 / 13.62

2nd eigenvalue: 0.05

2nd eigenvector: 117.9 / 3.2

3rd eigenvalue: 0.01

3rd eigenvector: 220.86 / 75.99

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 28.85 (-28.85)

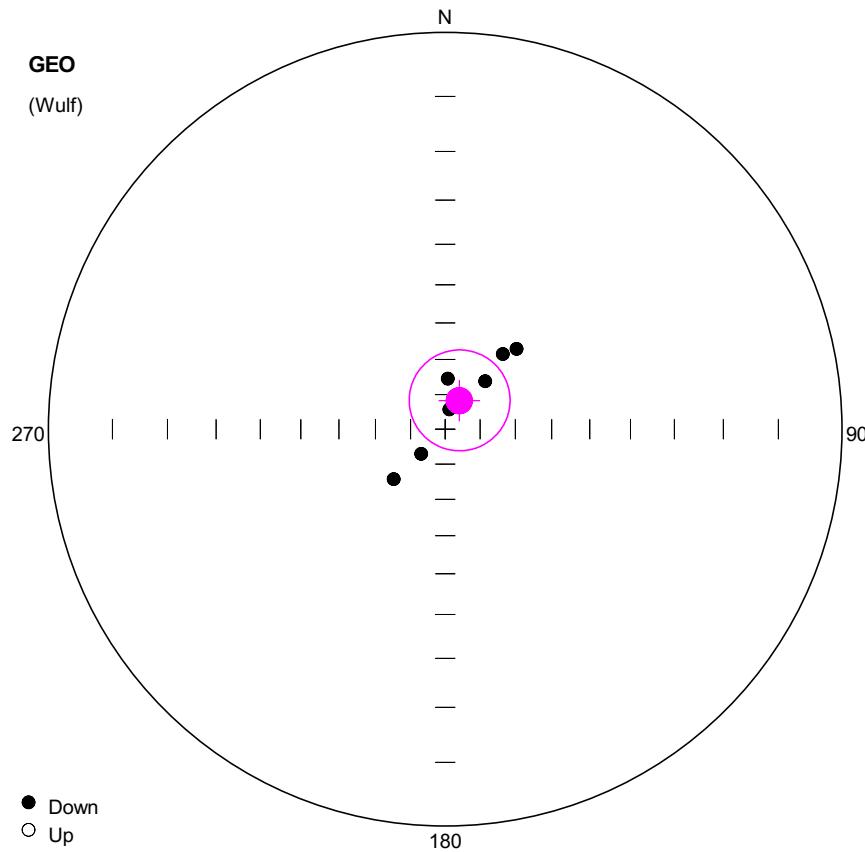
Pole longitude: 162.93 (-17.07)

Paleolatitude: 6.91 (-6.91)

dp: 5.38

dm: 10.54

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7001	Ch	15.9	13.5	6.8	20	68
2	YUNN7002	Ch	29.8	7.4	4.5	30	64
3	YUNN7003	Ch	35.3	12.7	5.0	35	60
4	YUNN7004	Ch	8.4	4.5	6.9	30	64 OA
5	YUNN7005	Ch	22.4	12.7	5.4	30	64
6	YUNN7006	Ch	48.6	13.2	5.1	40	60
7	YUNN7007	Ch	40.6	16.3	6.7	35	64
8	YUNN7008	Ch	15.8	26.1	12.2	30	64



**Number of data points: 7**

## Fisher statistics

Mean vector: 26.34 / 80.84

Resultant vector: 6.68

(X= 0.95, Y= 0.47, Z= 6.59)

Estimated precision, k: 18.56

95% Confidence limit: 14.39

95% Confidence limit, approximation: 12.28

## Orientation matrix

1st eigenvalue: 0.91

1st eigenvector: 26.61 / 80.48

2nd eigenvalue: 0.09

2nd eigenvector: 221.79 / 9.2

3rd eigenvalue: 0

3rd eigenvector: 131.4 / 2.45

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 78.18 (-78.18)

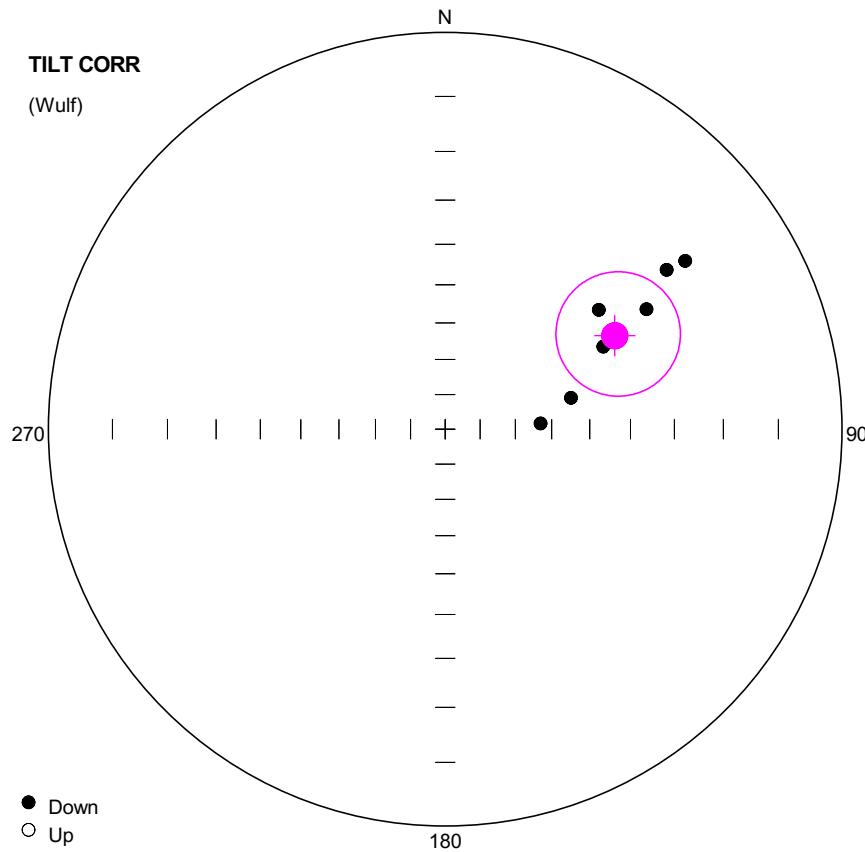
Pole longitude: 55.72 (-124.28)

Paleolatitude: 72.13 (-72.13)

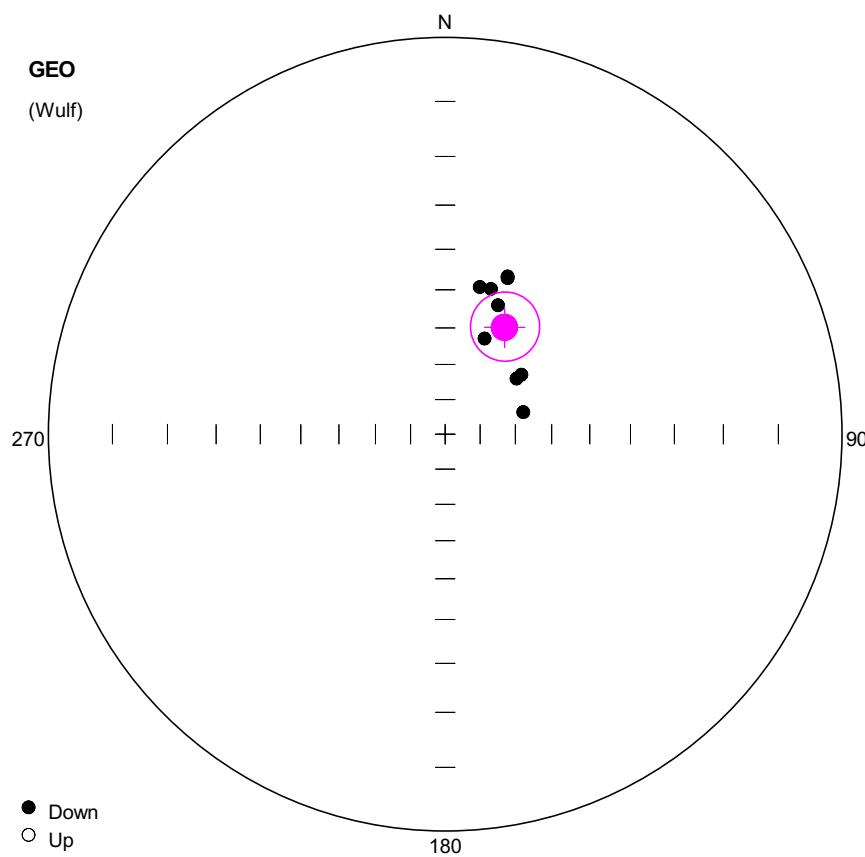
dp: 26.74

dm: 27.74

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7101	Ch	2.9	75.5	22.0	20	56
2	YUNN7102	Ch	37.5	63.1	12.3	20	64
3	YUNNT104	Ch	223.8	80.1	3.6	20	60 OA
4	YUNN7105	Ch	225.9	69.5	7.6	20	68 OA
5	YUNN7106	Ch	12.1	84.2	21.2	56	68 OA
6	YUNN7108	Ch	41.7	59.7	3.2	48	56 OA
7	YUNN7109	Ch	39.9	72.1	4.4	30	56 OA



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7101	Ch	52.2	37.8	22.0	20	56
2	YUNN7102	Ch	54.3	21.0	12.3	20	64
3	YUNN7104	Ch	76.1	53.8	3.6	20	60 OA
4	YUNN7105	Ch	86.5	62.9	7.6	20	68 OA
5	YUNN7106	Ch	62.5	41.6	21.2	56	68 OA
6	YUNN7108	Ch	55.0	17.1	3.2	48	56 OA
7	YUNN7109	Ch	59.2	28.9	4.4	30	56 OA



**Number of data points: 9**

**Fisher statistics**

Mean vector: 29.04 / 55.74

Resultant vector: 8.76

(X= 4.31, Y= 2.39, Z= 7.24)

Estimated precision, k: 33.06

95% Confidence limit: 9.09

95% Confidence limit, approximation: 8.12

**Orientation matrix**

1st eigenvalue: 0.95

1st eigenvector: 28.76 / 55.6

2nd eigenvalue: 0.05

2nd eigenvector: 161.49 / 24.92

3rd eigenvalue: 0

3rd eigenvector: 262.37 / 22.11

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 56.53 (-56.53)

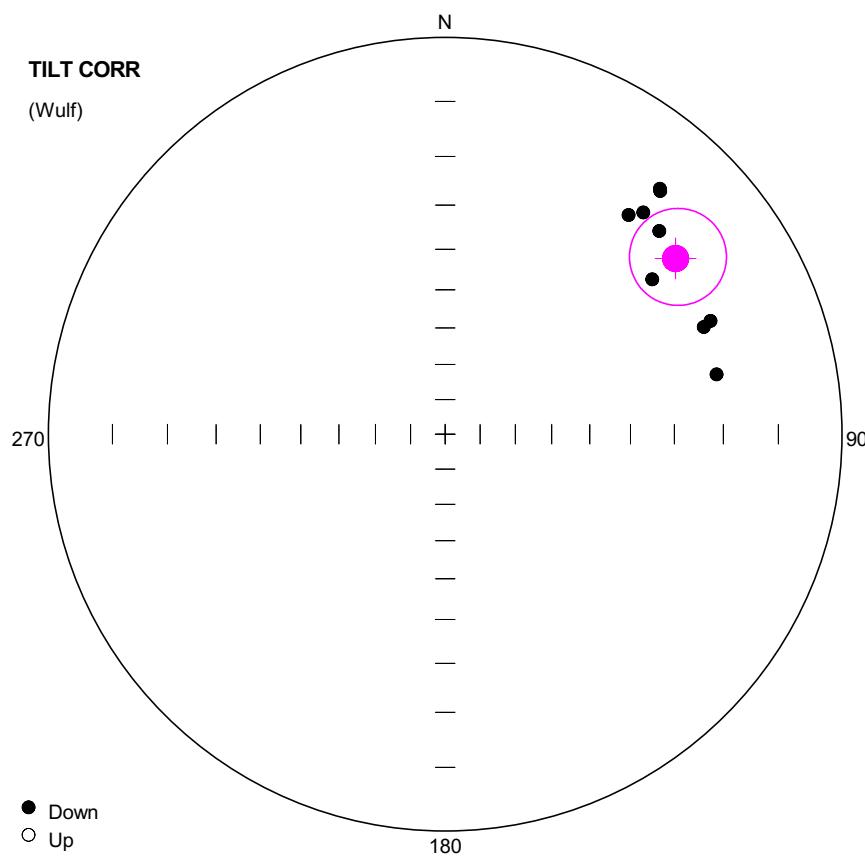
Pole longitude: 148.84 (-31.16)

Paleolatitude: 36.28 (-36.28)

dp: 9.32

dm: 13.01

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7202	Ch	21.8	44.1	7.9	35	64
2	YUNN7203	Ch	21.6	43.7	5.8	30	60
3	YUNN7204	Ch	17.5	48.0	7.4	30	64
4	YUNN7205	Ch	13.3	48.3	4.1	20	64
5	YUNN7206	Ch	74.3	66.9	5.3	35	64
6	YUNN7207	Ch	22.5	60.8	4.4	30	64
7	YUNN7208	Cht	22.3	51.3	3.6	48	64
8	YUNN7209	Ch	52.0	62.6	6.3	30	60
9	YUNN7210	Ch	52.0	64.3	4.3	35	64



**Number of data points:** 9

**Fisher statistics**

Mean vector: 52.68 / 17.73

Resultant vector: 8.76

(X= 5.06, Y= 6.63, Z= 2.67)

Estimated precision, k: 33.12

95% Confidence limit: 9.08

95% Confidence limit, approximation: 8.11

**Orientation matrix**

1st eigenvalue: 0.95

1st eigenvector: 52.46 / 17.71

2nd eigenvalue: 0.05

2nd eigenvector: 146 / 10.96

3rd eigenvalue: 0

3rd eigenvector: 266.25 / 68.98

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 23.33 (-23.33)

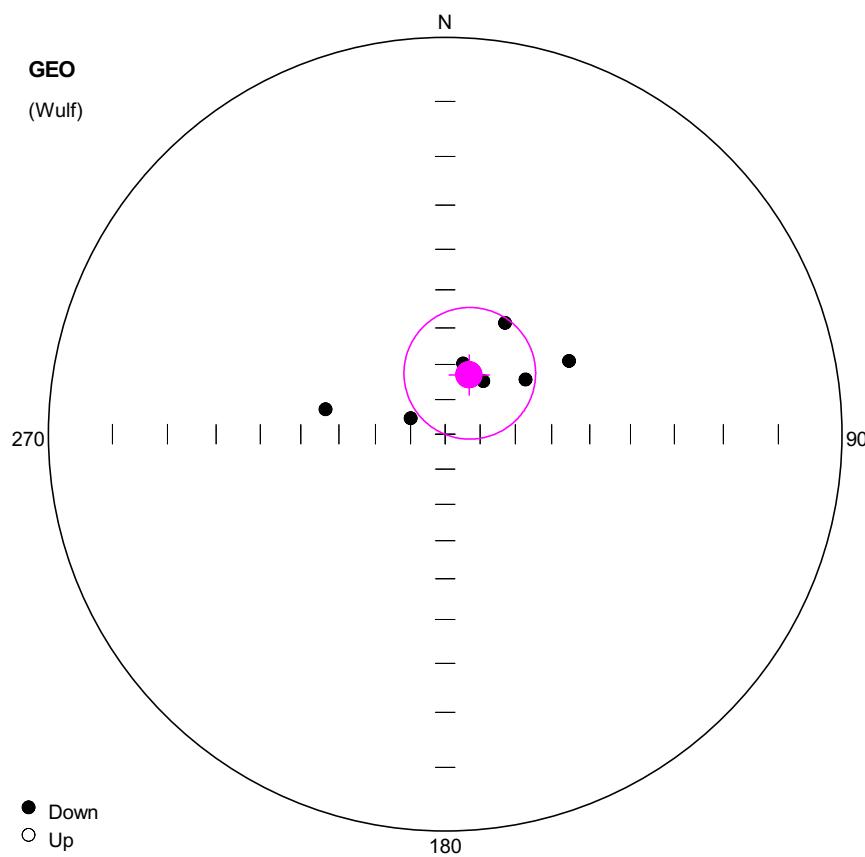
Pole longitude: 135.26 (-44.74)

Paleolatitude: 9.08 (-9.08)

dp: 4.88

dm: 9.41

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7202	Ch	41.5	11.4	7.9	35	64
2	YUNN7203	Ch	41.2	11.2	5.8	30	60
3	YUNN7204	Ch	41.8	16.3	7.4	30	64
4	YUNN7205	Ch	39.9	18.5	4.1	20	64
5	YUNN7206	Ch	77.6	20.0	5.3	35	64
6	YUNN7207	Ch	53.2	23.8	4.4	30	64
7	YUNN7208	Cht	46.5	16.7	3.6	48	64
8	YUNN7209	Ch	66.9	18.0	6.3	30	60
9	YUNN7210	Ch	67.5	19.6	4.3	35	64



**Number of data points:** 7

**Fisher statistics**

Mean vector: 21.9 / 71.68

Resultant vector: 6.49

(X= 1.89, Y= 0.76, Z= 6.16)

Estimated precision, k: 11.77

95% Confidence limit: 18.35

95% Confidence limit, approximation: 15.42

**Orientation matrix**

1st eigenvalue: 0.87

1st eigenvector: 24.99 / 70.95

2nd eigenvalue: 0.12

2nd eigenvector: 256.15 / 12.22

3rd eigenvalue: 0.01

3rd eigenvector: 162.96 / 14.39

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 76.45 (-76.45)

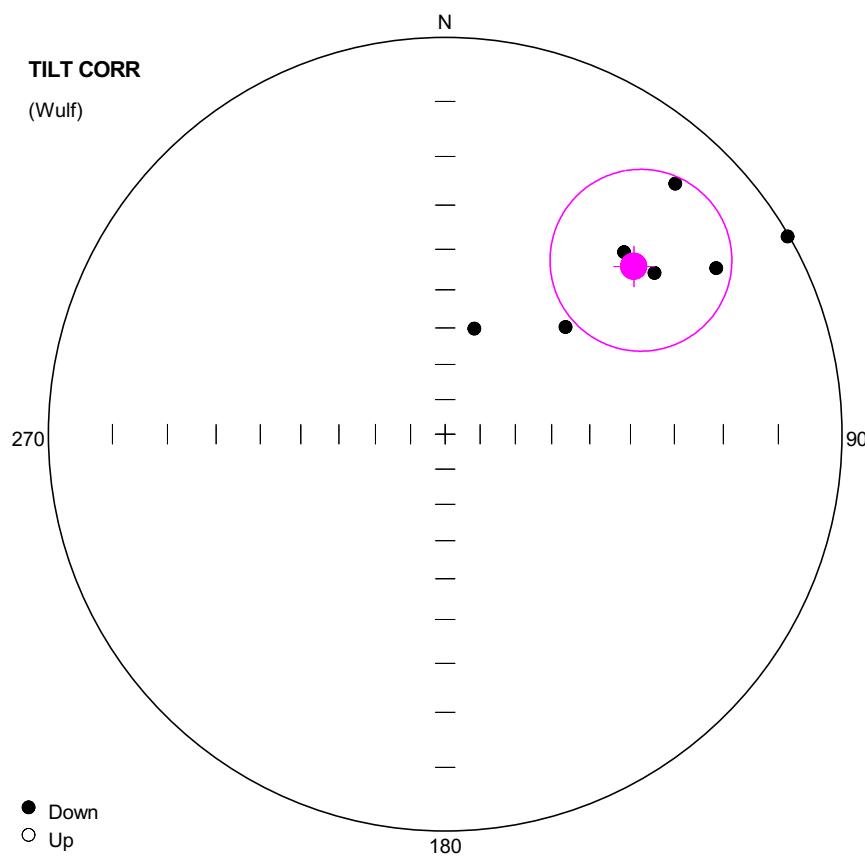
Pole longitude: 132.57 (-47.43)

Paleolatitude: 56.49 (-56.49)

dp: 28.31

dm: 32.23

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7301	Ch	294.9	79.1	11.1	30	64
2	YUNN7303	Ch	281.7	55.8	20.4	30	64
3	YUNN7304	Ch	14.3	69.2	9.2	30	60
4	YUNN7305	Ch	35.7	71.3	4.5	44	60
5	YUNN7306	Ch	59.4	50.1	5.9	44	64
6	YUNN7307	Ch	55.8	62.5	7.2	30	56
7	YUNN7309	Ch	28.3	54.7	8.6	40	60



**Number of data points:** 7

**Fisher statistics**

Mean vector: 48.35 / 25.07

Resultant vector: 6.49

(X= 3.91, Y= 4.39, Z= 2.75)

Estimated precision, k: 11.8

95% Confidence limit: 18.33

95% Confidence limit, approximation: 15.41

**Orientation matrix**

1st eigenvalue: 0.87

1st eigenvector: 48.87 / 23.93

2nd eigenvalue: 0.12

2nd eigenvector: 270.9 / 59.14

3rd eigenvalue: 0.01

3rd eigenvector: 147.3 / 18.29

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 28.68 (-28.68)

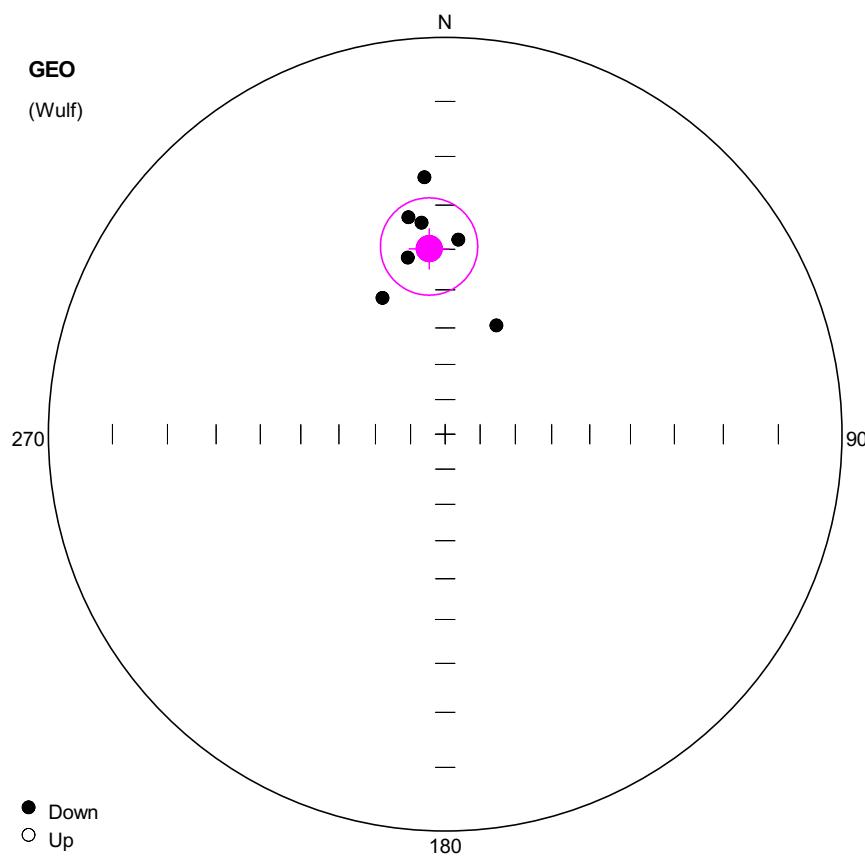
Pole longitude: 138.01 (-41.99)

Paleolatitude: 13.16 (-13.16)

dp: 10.59

dm: 19.7

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7301	Ch	48.3	45.8	11.1	30	64
2	YUNN7303	Ch	15.5	59.1	20.4	30	64
3	YUNN7304	Ch	44.5	24.5	9.2	30	60
4	YUNN7305	Ch	52.4	22.7	4.5	44	60
5	YUNN7306	Ch	60.0	0.2	5.9	44	64
6	YUNN7307	Ch	58.5	12.6	7.2	30	56
7	YUNN7309	Ch	42.6	8.8	8.6	40	60



**Number of data points:** 7

**Fisher statistics**

Mean vector: 355.06 / 39.73

Resultant vector: 6.79

(X= 5.2, Y= -0.45, Z= 4.34)

Estimated precision, k: 28.75

95% Confidence limit: 11.45

95% Confidence limit, approximation: 9.87

**Orientation matrix**

1st eigenvalue: 0.94

1st eigenvector: 354.85 / 39.47

2nd eigenvalue: 0.04

2nd eigenvector: 137.86 / 44.12

3rd eigenvalue: 0.02

3rd eigenvector: 247.91 / 19.48

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 47.44 (-47.44)

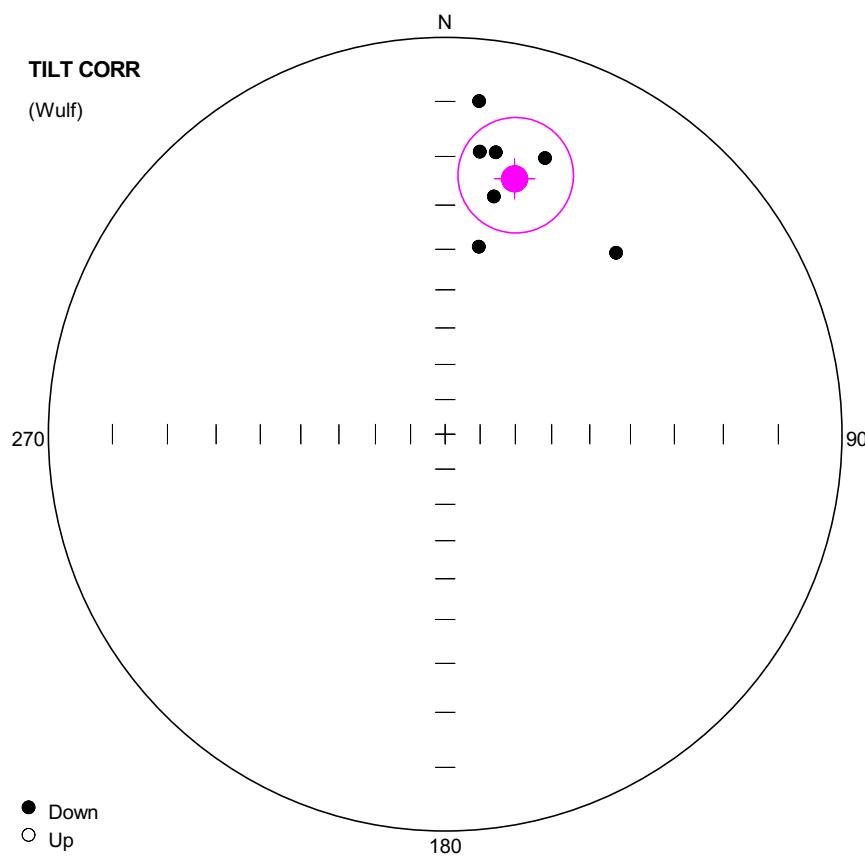
Pole longitude: -159.21 (20.79)

Paleolatitude: 22.57 (-22.57)

dp: 8.25

dm: 13.75

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7404	Ch	348.0	41.1	3.6	30	64
2	YUNN7405	Ch	335.3	48.6	6.7	44	64
3	YUNN7406	Ch	3.9	37.7	5.1	56	64
4	YUNN7407	Cht	350.4	32.0	8.2	52	64
5	YUNN7408	Ch	355.4	24.0	5.2	40	68
6	YUNN7409	Ch	353.6	33.6	7.0	30	64
7	YUNN7410	Ch	25.3	56.3	4.4	44	64



**Number of data points:** 7

**Fisher statistics**

Mean vector: 15.2 / 22.62

Resultant vector: 6.79

(X= 6.05, Y= 1.64, Z= 2.61)

Estimated precision, k: 28.75

95% Confidence limit: 11.45

95% Confidence limit, approximation: 9.87

**Orientation matrix**

1st eigenvalue: 0.94

1st eigenvector: 14.89 / 22.51

2nd eigenvalue: 0.04

2nd eigenvector: 115.98 / 24.89

3rd eigenvalue: 0.02

3rd eigenvector: 248.09 / 55.32

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 35.74 (-35.74)

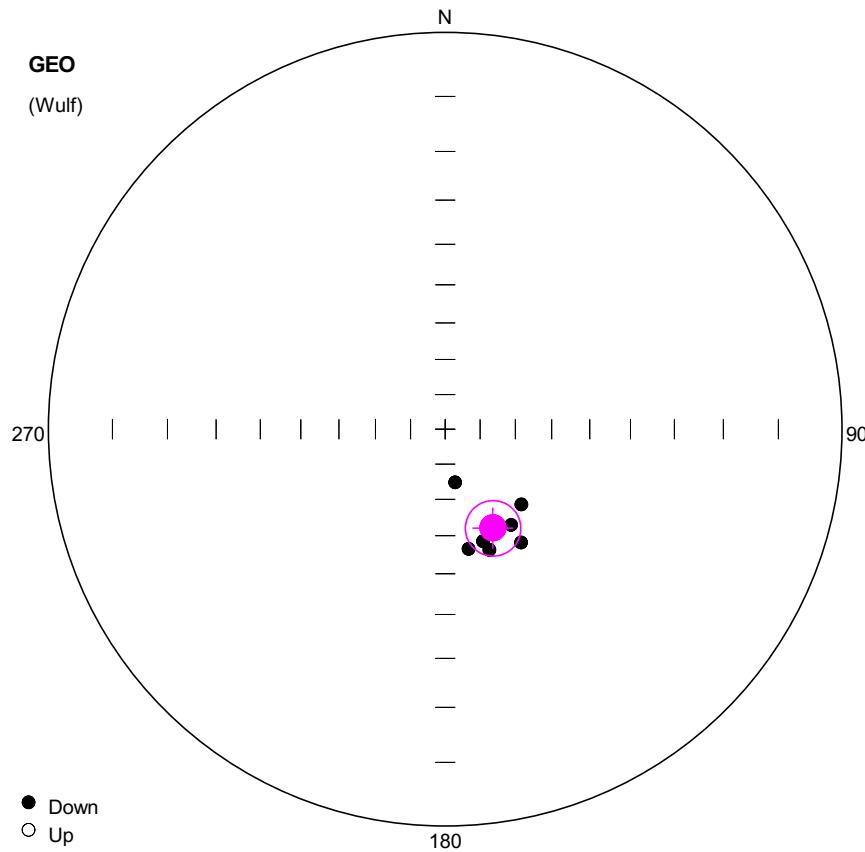
Pole longitude: 175.6 (-4.4)

Paleolatitude: 11.77 (-11.77)

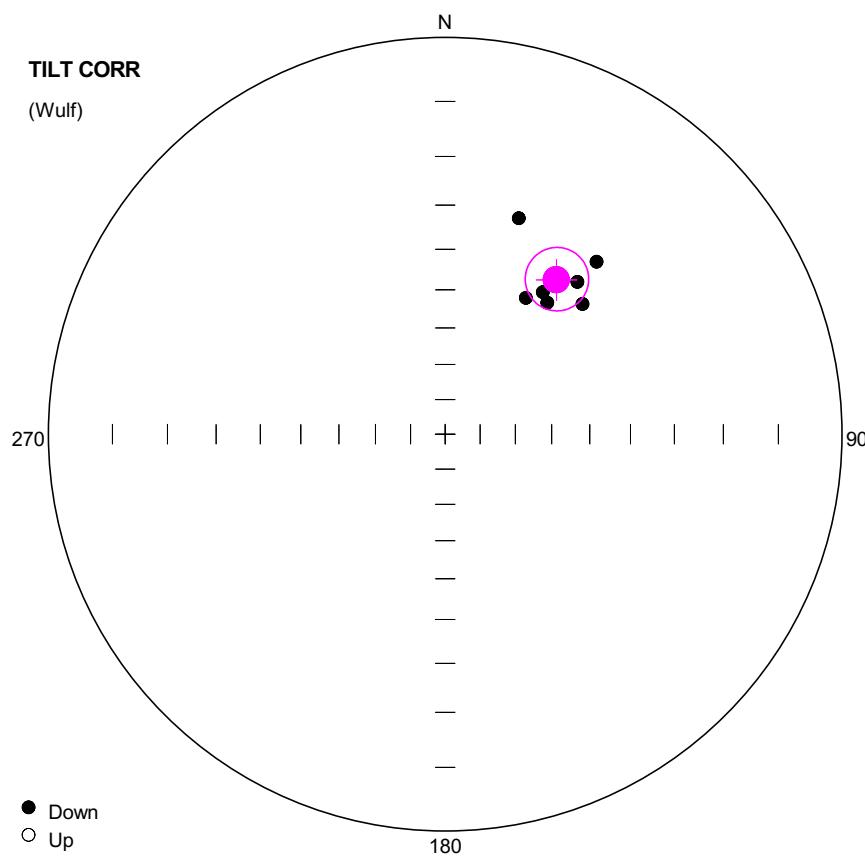
dp: 6.44

dm: 12.14

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7404	Ch	11.6	27.1	3.6	30	64
2	YUNN7405	Ch	10.2	38.7	6.7	44	64
3	YUNN7406	Ch	19.9	17.0	5.1	56	64
4	YUNN7407	Cht	7.0	18.7	8.2	52	64
5	YUNN7408	Ch	5.8	9.7	5.2	40	68
6	YUNN7409	Ch	10.2	18.4	7.0	30	64
7	YUNN7410	Ch	43.3	25.7	4.4	44	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7501	Ch	134.6	59.8	12.6	30	56
2	YUNN7502	Clt	146.2	52.1	10.9	20	52
3	YUNN7503	Cmt	161.3	56.8	6.5	48	56
4	YUNN7505	Ch	145.4	57.3	9.5	35	60
5	YUNN7506	Ch	169.3	74.5	8.3	30	60
6	YUNN7507	Cmt	169.0	55.8	5.1	44	56
7	YUNN7508	Cmt	159.9	54.1	20.0	35	52



**Number of data points:** 7

**Fisher statistics**

Mean vector: 35.75 / 38.79

Resultant vector: 6.91

(X= 4.37, Y= 3.15, Z= 4.33)

Estimated precision, k: 66.99

95% Confidence limit: 7.43

95% Confidence limit, approximation: 6.47

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 35.83 / 38.84

2nd eigenvalue: 0.02

2nd eigenvector: 138.08 / 14.76

3rd eigenvalue: 0.01

3rd eigenvector: 244.72 / 47.4

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 41.01 (-41.01)

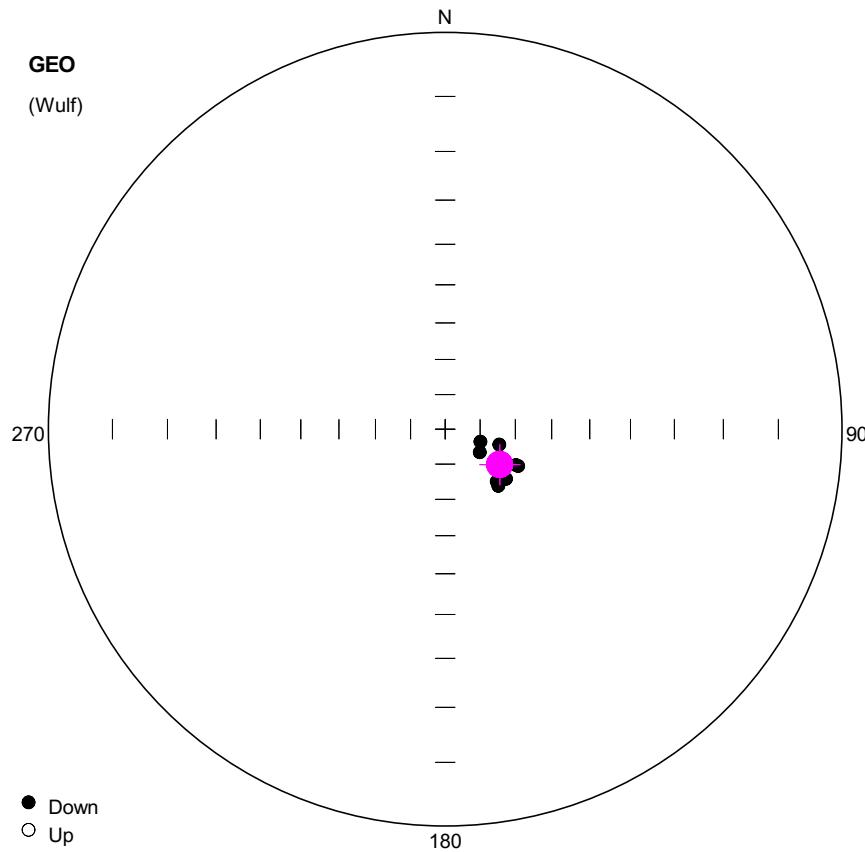
Pole longitude: 148.12 (-31.88)

Paleolatitude: 21.9 (-21.9)

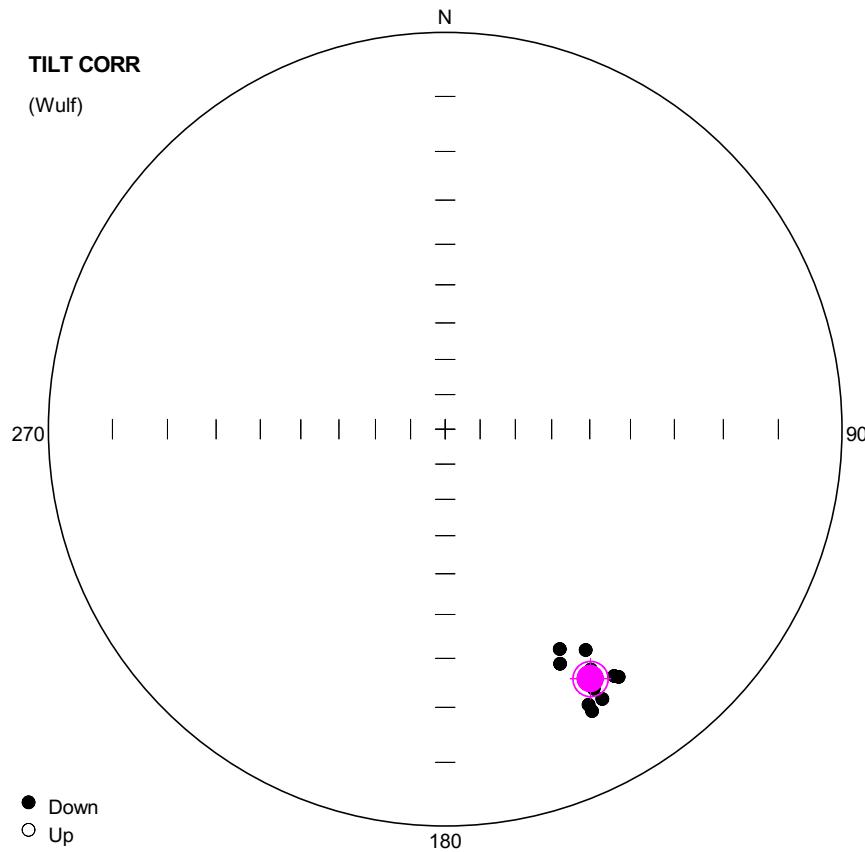
dp: 5.26

dm: 8.84

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7501	Ch	41.3	29.9	12.6	30	56
2	YUNN7502	Clt	46.5	39.0	10.9	20	52
3	YUNN7503	Cmt	34.5	43.1	6.5	48	56
4	YUNN7505	Ch	41.0	36.1	9.5	35	60
5	YUNN7506	Ch	18.8	30.2	8.3	30	60
6	YUNN7507	Cmt	30.6	46.5	5.1	44	56
7	YUNN7508	Cmt	37.9	44.5	20.0	35	52



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7601	Ch	129.2	67.6	1.5	40	64
2	YUNN7602	Ch	109.7	79.2	8.3	30	64
3	YUNN7603	Ch	123.7	78.0	5.7	20	64
4	YUNN7604	Ch	136.8	67.8	3.1	40	64
5	YUNN7605	Ch	135.3	69.0	4.1	30	68
6	YUNN7606	Ch	105.9	73.9	3.5	48	68
7	YUNN7607	Ch	117.9	72.0	5.0	30	64
8	YUNN7608	Ch	117.0	67.4	9.7	40	64
9	YUNN7609	Ch	116.8	66.7	5.1	30	68
10	YUNN7610	Ch	127.2	69.9	5.1	30	68



**Number of data points: 10**

## Fisher statistics

Mean vector: 149.81 / 17.96

Resultant vector: 9.96

(X= -8.19, Y= 4.76, Z= 3.07)

Estimated precision, k: 212.34

95% Confidence limit: 3.32

95% Confidence limit, ap

## Orientation matrix

1st eigenvalue: 0.99

1st eigenvector: 149.81 / 17.95

2nd eigenvalue: 0.01

2nd eigenvector: 301.42 / 69.78

3rd eigenvalue: 0

3rd eigenvector: 5

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -12.45 (12.45)

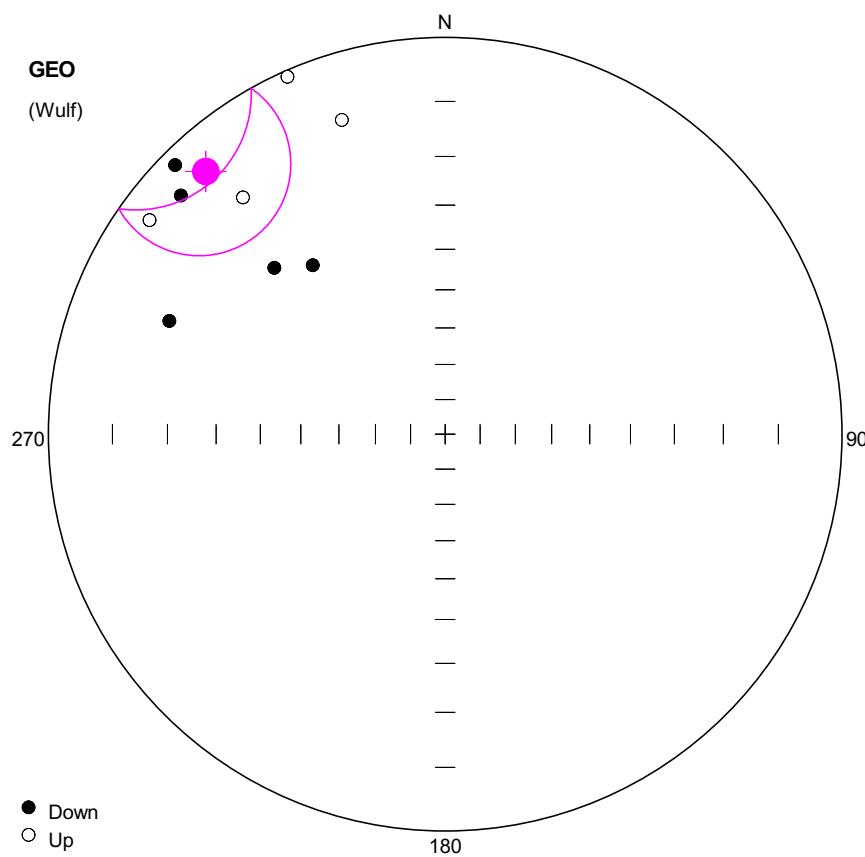
Pole longitude: 44.59 (-135.41)

Paleolatitude: 9.21 (-9.21)

dp: 1.79

dm: 3.45

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7601	Ch	149.8	13.6	1.5	40	64
2	YUNN7602	Ch	152.5	26.0	8.3	30	64
3	YUNN7603	Ch	153.9	23.3	5.7	20	64
4	YUNN7604	Ch	152.5	12.6	3.1	40	64
5	YUNN7605	Ch	152.5	13.9	4.1	30	68
6	YUNN7606	Ch	147.5	23.2	3.5	48	68
7	YUNN7607	Ch	148.8	19.4	5.0	30	64
8	YUNN7608	Ch	145.6	16.0	9.7	40	64
9	YUNN7609	Ch	145.0	15.4	5.1	30	68
10	YUNN7610	Ch	150.2	16.0	5.1	30	68



**Number of data points:** 9

#### Fisher statistics

Mean vector: 317.67 / 6.32  
 Resultant vector: 8.41  
 $(X= 6.18, Y= -5.63, Z= 0.93)$   
 Estimated precision,  $k$ : 13.65  
 95% Confidence limit: 14.45  
 95% Confidence limit, approximation: 12.63

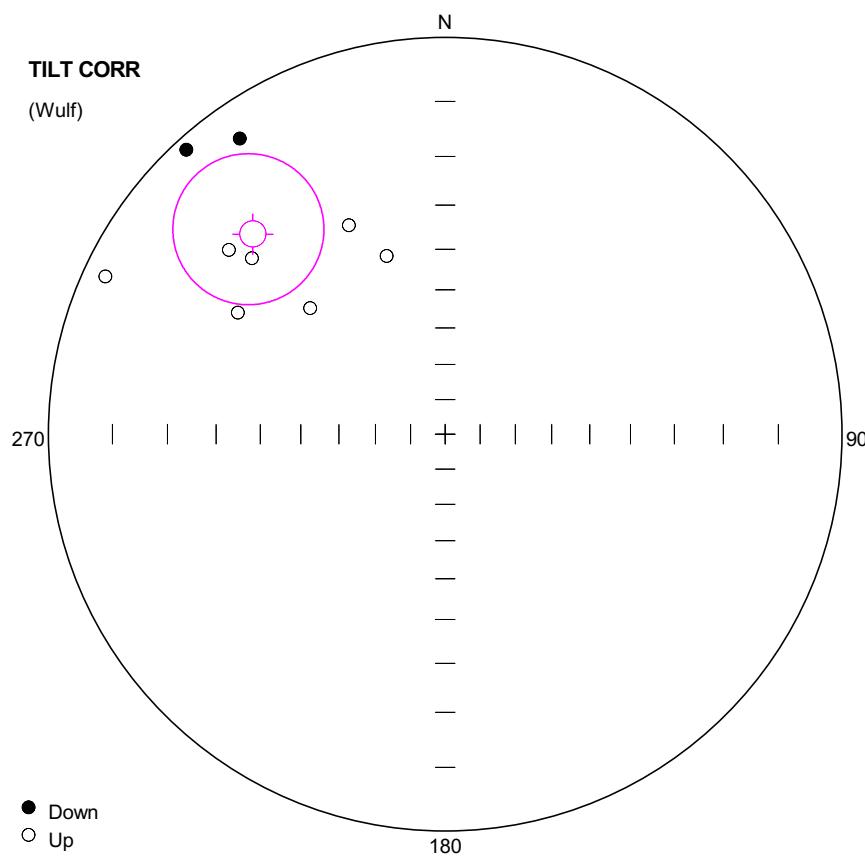
#### Orientation matrix

1st eigenvalue: 0.88  
 1st eigenvector: 317.5 / 6.13  
 2nd eigenvalue: 0.08  
 2nd eigenvector: 218.87 / 54.4  
 3rd eigenvalue: 0.04  
 3rd eigenvector: 51.79 / 34.9

#### VGP

Site latitude: 65  
 Site longitude: 14.04  
 Pole latitude: 21.22 (-21.22)  
 Pole longitude: -119.8 (60.2)  
 Paleolatitude: 3.17 (-3.17)  
 dp: 7.29  
 dm: 14.52

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7701	Ch	319.5	-13.8	6.1	44	60
2	YUNN7702	Cmt	336.2	-0.9	2.3	40	56 OA
3	YUNN7703	Ch	305.9	-4.8	17.4	40	60
4	YUNN7704	Ch	321.9	33.2	9.3	35	64
5	YUNN7705	Ch	314.9	2.3	9.0	35	56
6	YUNN7706	Ch	292.3	16.2	7.9	35	56
7	YUNN7707	Ch	341.8	-10.4	19.1	30	64
8	YUNN7708	Ch	312.1	6.2	11.4	35	60
9	YUNN7709	Ch	314.2	28.0	17.7	35	60



**Number of data points: 9**

**Fisher statistics**

Mean vector: 316.16 / -20.06

Resultant vector: 8.41

(X= 5.7, Y= -5.47, Z= -2.89)

Estimated precision, k: 13.65

95% Confidence limit: 14.45

95% Confidence limit, approximation: 12.63

**Orientation matrix**

1st eigenvalue: 0.88

1st eigenvector: 135.94 / 20.2

2nd eigenvalue: 0.08

2nd eigenvector: 265.39 / 59.92

3rd eigenvalue: 0.04

3rd eigenvector: 37.7 / 21.3

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 7.88 (-7.88)

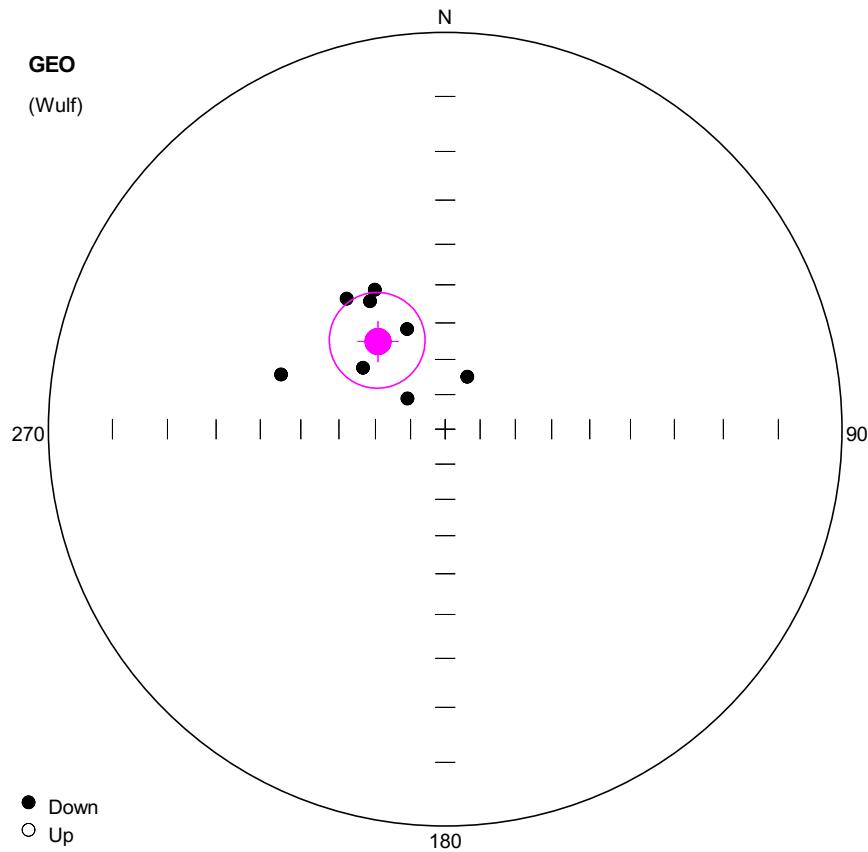
Pole longitude: -122.5 (57.5)

Paleolatitude: -10.34 (10.34)

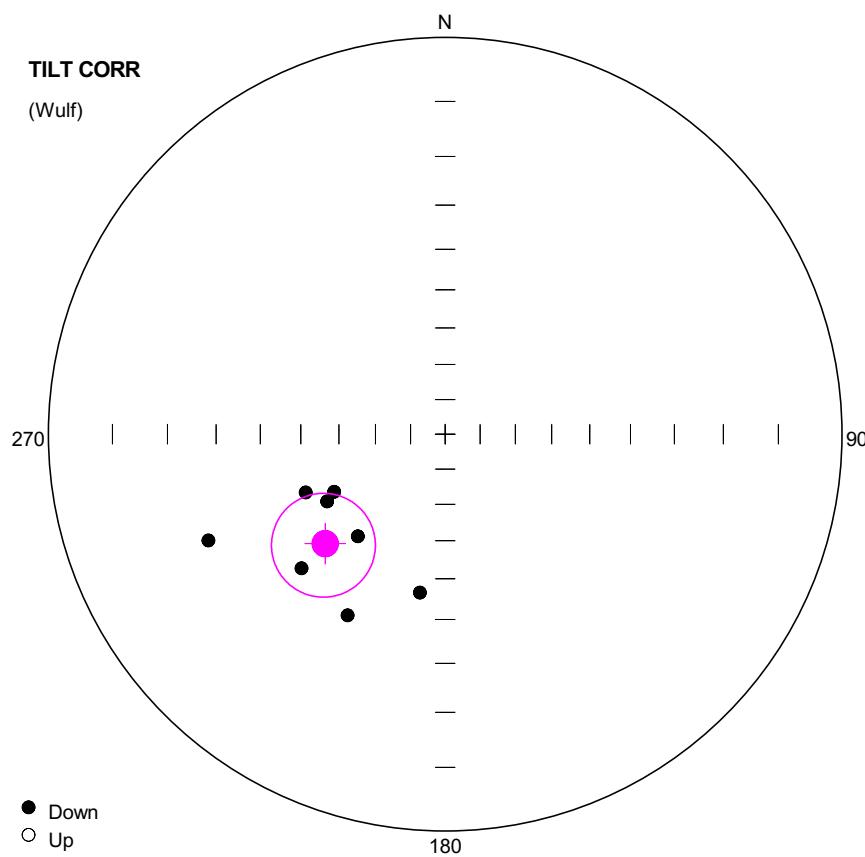
dp: 7.92

dm: 15.13

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7701	Ch	313.0	-40.1	6.1	44	60
2	YUNN7702	Cmt	335.3	-29.8	2.3	40	56 OA
3	YUNN7703	Ch	300.4	-27.6	17.4	40	60
4	YUNN7704	Ch	325.2	5.6	9.3	35	64
5	YUNN7705	Ch	312.3	-23.3	9.0	35	56
6	YUNN7706	Ch	294.9	-3.3	7.9	35	56
7	YUNN7707	Ch	341.8	-39.4	19.1	30	64
8	YUNN7708	Ch	310.4	-18.8	11.4	35	60
9	YUNN7709	Ch	317.7	1.8	17.7	35	60



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7801	Ch	306.7	61.0	11.9	30	56
2	YUNN7802	Ch	333.2	47.0	14.1	30	64
3	YUNN7803	Ch	339.2	59.8	16.4	20	64
4	YUNN7804	Ch	288.4	42.9	25.7	20	60
5	YUNN7805	Clt	329.6	49.0	9.6	20	56
6	YUNN7808	Ch	323.0	45.2	16.5	30	60
7	YUNN7809	Ch	22.9	73.7	25.9	30	56
8	YUNN7810	Ch	309.1	76.1	24.8	20	64



**Number of data points: 8**

**Fisher statistics**

Mean vector: 227.68 / 45.49

Resultant vector: 7.65

(X= -3.61, Y= -3.96, Z= 5.45)

Estimated precision, k: 19.8

95% Confidence limit: 12.76

95% Confidence limit, approximation: 11.12

**Orientation matrix**

1st eigenvalue: 0.91

1st eigenvector: 227.9 / 45.65

2nd eigenvalue: 0.06

2nd eigenvector: 127.07 / 10.41

3rd eigenvalue: 0.03

3rd eigenvector: 27.39 / 42.48

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 9.05 (-9.05)

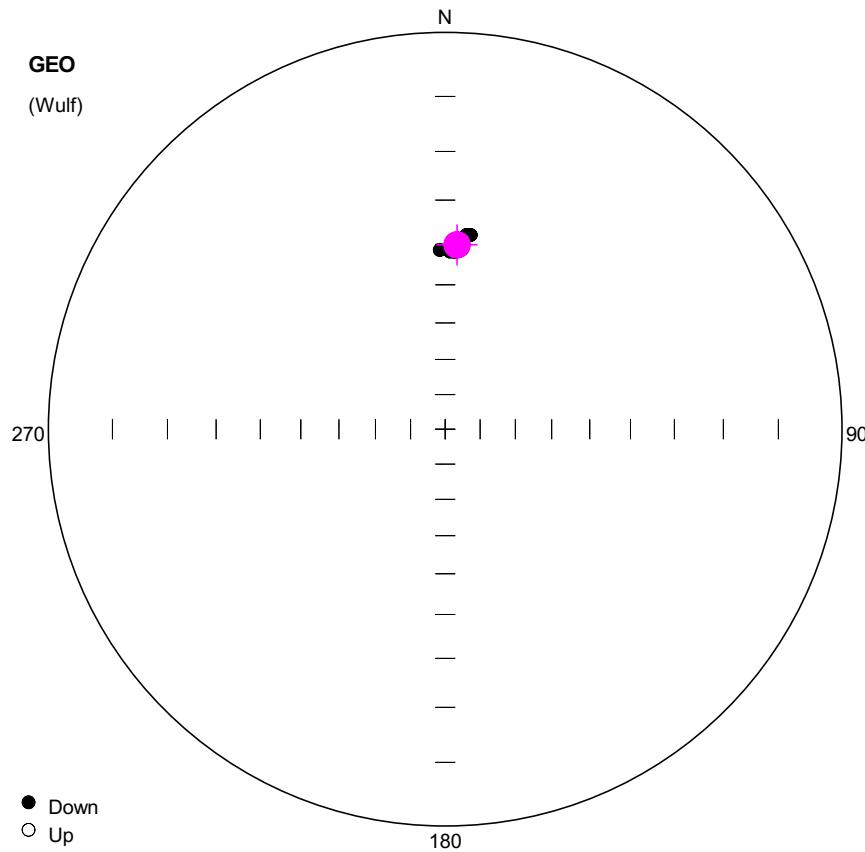
Pole longitude: -27.83 (152.17)

Paleolatitude: 26.96 (-26.96)

dp: 10.31

dm: 16.22

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7801	Ch	227.0	37.3	11.9	30	56
2	YUNN7802	Ch	242.6	55.0	14.1	30	64
3	YUNN7803	Ch	220.5	52.6	16.4	20	64
4	YUNN7804	Ch	245.8	23.6	25.7	20	60
5	YUNN7805	Clt	240.4	52.2	9.6	20	56
6	YUNN7808	Ch	247.3	48.3	16.5	30	60
7	YUNN7809	Ch	189.0	46.0	25.9	30	56
8	YUNN7810	Ch	208.3	35.2	24.8	20	64



**Number of data points: 8**

## Fisher statistics

Mean vector: 3.68 / 40.07

Resultant vector: 7.99

(X= 6.1, Y= 0.39, Z= 5.14)

Estimated precision, k: 814.74

95% Confidence limit: 1.94

95% Confidence limit, approximation: 1.73

## Orientation matrix

1st eigenvalue: 1

1st eigenvector: 3.69 / 40.07

2nd eigenvalue: 0

2nd eigenvector: 247.81 / 27.42

3rd eigenvalue: 0

3rd eigenvector: 134.19 / 37.68

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 47.74 (-47.74)

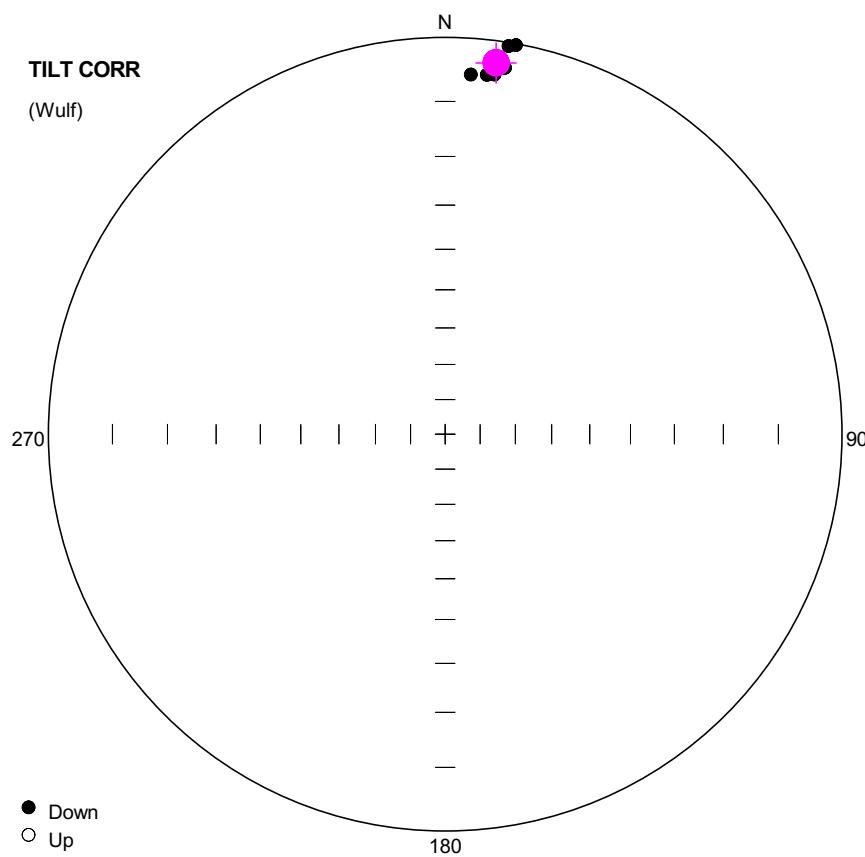
Pole longitude: -171.01 (8.99)

Paleolatitude: 22.81 (-22.81)

dp: 1.41

dm: 2.34

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7901	Ch	358.3	41.4	2.5	35	68
2	YUNN7904	Ch	5.5	40.9	4.0	30	64
3	YUNN7905	Ch	6.2	37.6	8.6	35	60
4	YUNN7906	Ch	3.6	40.1	5.3	20	64
5	YUNN7907	Ch	1.6	41.8	4.9	35	68
6	YUNN7908	Ch	3.2	41.9	4.1	20	68
7	YUNN7909	Ch	7.5	37.5	3.1	20	68
8	YUNN7910	Ch	3.2	39.1	4.9	20	68



**Number of data points: 8**

**Fisher statistics**

Mean vector: 7.8 / 3.28

Resultant vector: 7.99

(X= 7.9, Y= 1.08, Z= 0.46)

Estimated precision, k: 811.91

95% Confidence limit: 1.95

95% Confidence limit, approximation: 1.74

**Orientation matrix**

1st eigenvalue: 1

1st eigenvector: 7.8 / 3.28

2nd eigenvalue: 0

2nd eigenvector: 274.16 / 47.99

3rd eigenvalue: 0

3rd eigenvector: 100.74 / 41.82

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 26.39 (-26.39)

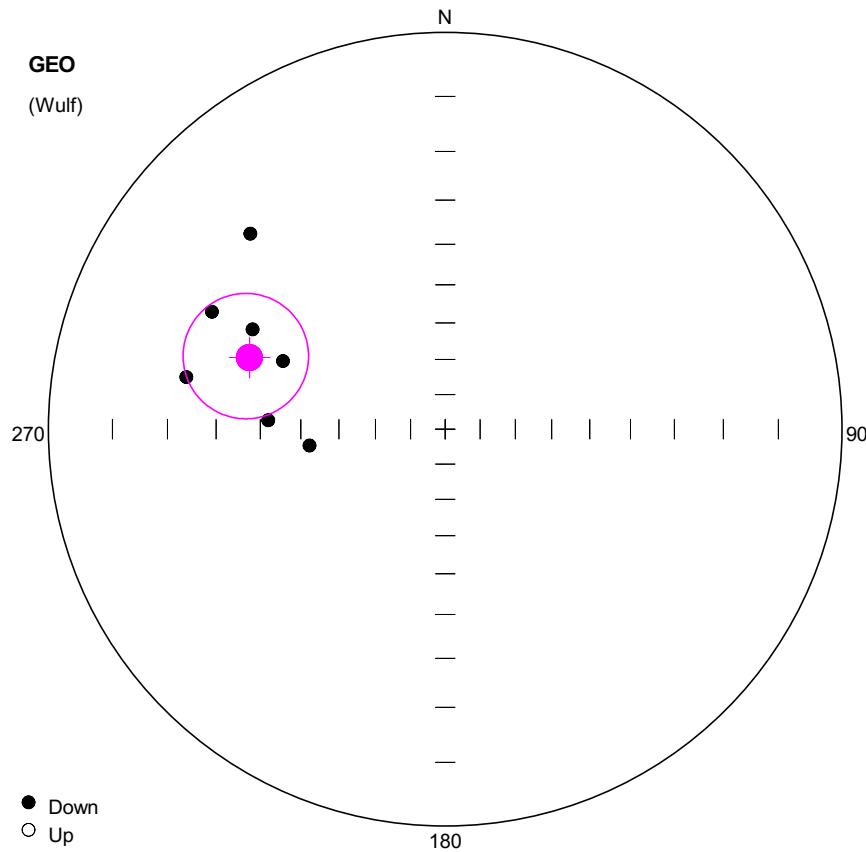
Pole longitude: -174.67 (5.33)

Paleolatitude: 1.64 (-1.64)

dp: 0.97

dm: 1.95

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN7901	Ch	4.1	5.5	2.5	35	68
2	YUNN7904	Ch	9.3	3.8	4.0	30	64
3	YUNN7905	Ch	9.3	0.5	8.6	35	60
4	YUNN7906	Ch	7.7	3.4	5.3	20	64
5	YUNN7907	Ch	6.6	5.3	4.9	35	68
6	YUNN7908	Ch	7.8	5.1	4.1	20	68
7	YUNN7909	Ch	10.3	0.2	3.1	20	68
8	YUNN7910	Ch	7.3	2.4	4.9	20	68



**Number of data points: 7**

## Fisher statistics

Mean vector: 290.15 / 34.59

Resultant vector: 6.69

(X= 1.9, Y= -5.17, Z= 3.8)

Estimated precision, k: 19.38

95% Confidence Int.

95% Confidence limit, approximation: 12.02

## Orientation matrix

1st eigenvalue: 0.91

1st eigenvector: 290.11 / 34.5

2nd eigenvalue: 0.07

2nd eigenvector: 174.08 / 32.56

3rd eigenvalue: 0.01

3rd eigenvector: 53.41 / 38.62

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 25.66 (-25.66)

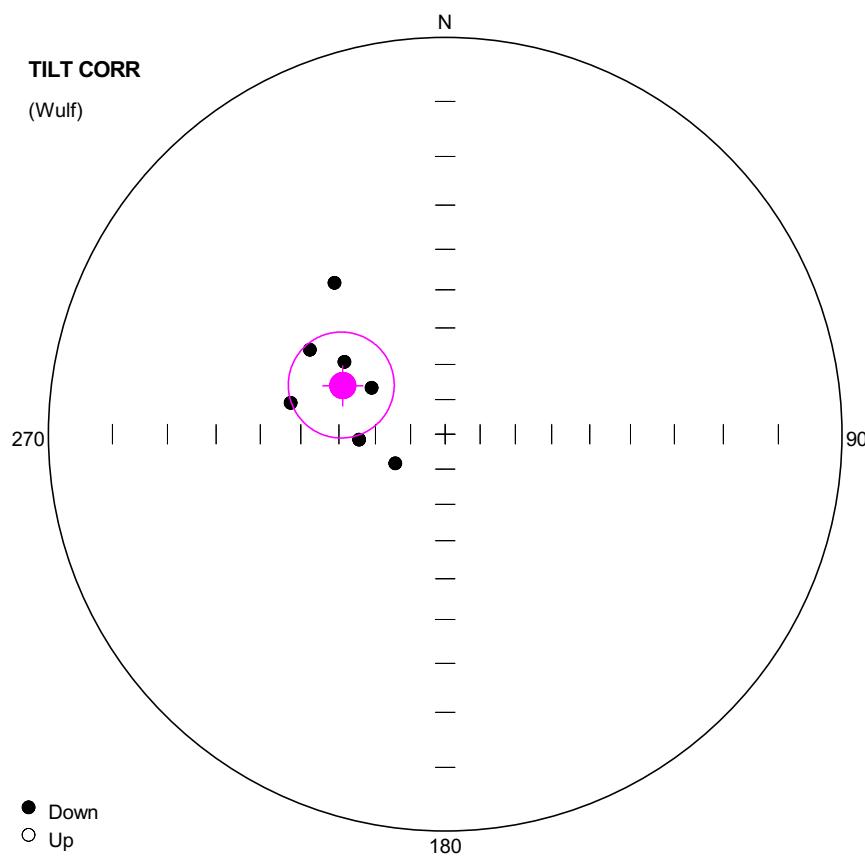
Pole longitude: -86.01 (93.99)

Paleolatitude: 19.02 (-19.02)

dp: 9.27

dm: 16.15

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8001	Ch	296.7	23.3	6.3	20	64
2	YUNN8002	Ch	272.9	41.9	4.5	20	64
3	YUNN8003	Ch	297.4	32.7	18.2	20	60
4	YUNN8005	Ch	263.1	52.0	9.2	20	64
5	YUNN8006	Ch	315.1	20.4	9.6	20	60
6	YUNN8008	Ch	292.8	42.2	30.9	20	64
7	YUNN8009	Ch	281.4	22.7	7.9	20	60



**Number of data points:** 7

**Fisher statistics**

Mean vector: 295.34 / 58.14  
Resultant vector: 6.69  
(X= 1.51, Y= -3.19, Z= 5.68)  
Estimated precision, k: 19.42  
95% Confidence limit: 14.05  
95% Confidence limit, approximation: 12.01

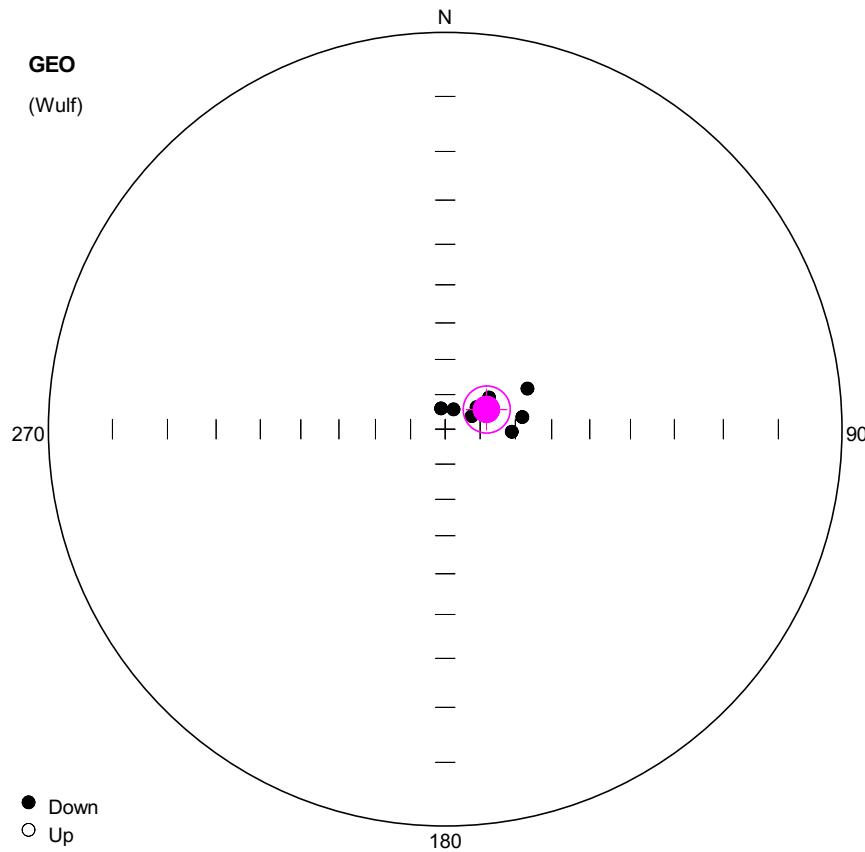
**Orientation matrix**

1st eigenvalue: 0.92  
1st eigenvector: 295.26 / 58.06  
2nd eigenvalue: 0.07  
2nd eigenvector: 162.19 / 23.06  
3rd eigenvalue: 0.01  
3rd eigenvector: 62.87 / 20.83

**VGP**

Site latitude: 65  
Site longitude: 14.04  
Pole latitude: 45.16 (-45.16)  
Pole longitude: -78.98 (101.02)  
Paleolatitude: 38.82 (-38.82)  
dp: 15.3  
dm: 20.74

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8001	Ch	302.0	46.2	6.3	20	64
2	YUNN8002	Ch	266.4	65.5	4.5	20	64
3	YUNN8003	Ch	305.6	55.3	18.2	20	60
4	YUNN8005	Ch	239.6	73.4	9.2	20	64
5	YUNN8006	Ch	323.8	39.4	9.6	20	60
6	YUNN8008	Ch	302.3	65.3	30.9	20	64
7	YUNN8009	Ch	281.5	46.7	7.9	20	60



**Number of data points: 8**

## Fisher statistics

Mean vector: 64.51 / 76.84

Resultant vector: 7.9

(X= 0.77, Y= 1.62, Z= 7.69)

Estimated precision, k: 69.45

95% C

95% Confidence limit, approximation: 5.94

## Orientation matrix

1st eigenvalue: 0.98

1st eigenvector: 64.51 / 76.85

2nd eigenvalue: 0.02

2nd eigenvector: 270.2 / 11.89

3rd eigenvalue: 0

3rd eigenvector: 179.03 / 5.54

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 63.9 (-63.9)

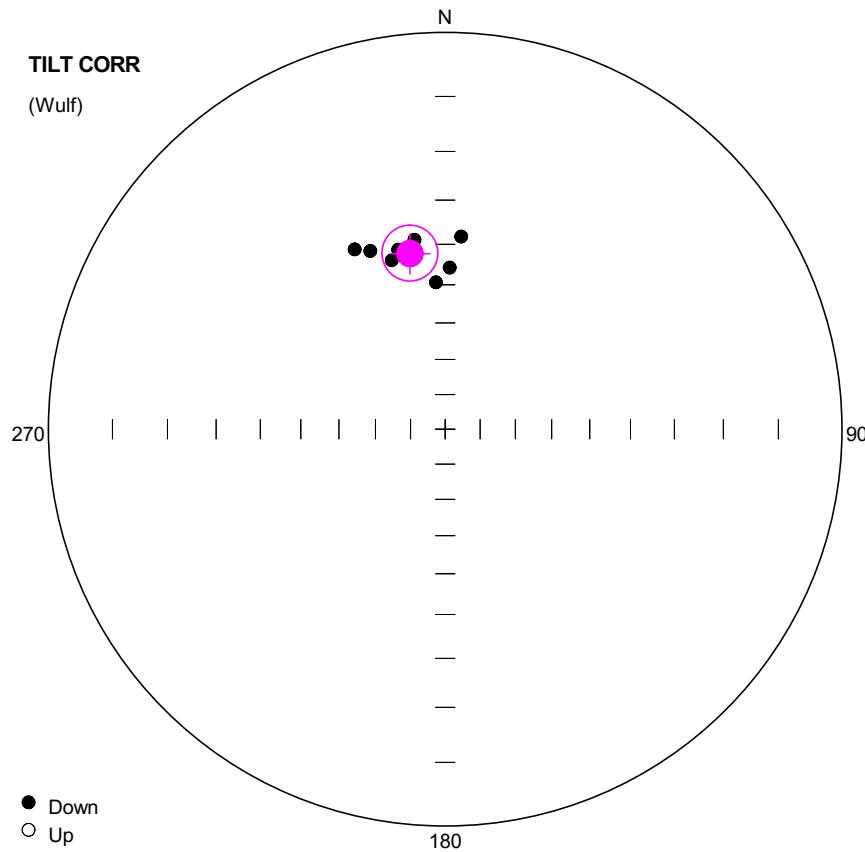
Pole longitude: 74.4 (-105.6)

Paleolatitude: 64.94 (-64.94)

dp: 11.58

dm: 12.45

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8101	Ch	55.0	79.0	6.8	40	64
2	YUNN8102	Ch	81.1	67.7	6.0	35	64
3	YUNN8103	Ch	63.8	64.0	7.0	40	64
4	YUNN8104	Ch	63.9	81.5	6.5	30	64
5	YUNN8105	Ch	349.0	83.9	12.2	44	64
6	YUNN8106	Ch	54.2	74.5	13.2	30	64
7	YUNN8107	Ch	92.2	70.9	7.8	20	64
8	YUNN8109	Ch	23.4	83.8	4.5	44	64



**Number of data points: 8**

## Fisher statistics

Mean vector: 348.63 / 41.45

Resultant vector: 7.9

(X= 5.81, Y= -1.17, Z= 5.23)

Estimated precision, k: 69.64

95% 0

95% Confidence limit, approximation: 5.93

## Orientation matrix

1st eigenvalue: 0.98

1st eigenvector: 348.62 / 41.45

2nd eigenvalue: 0.02

2nd eigenvector: 89.82 / 12.4

3rd eigenvalue: 0

3rd eigenvector: 192.93 / 45.9

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 48.17 (-48.17)

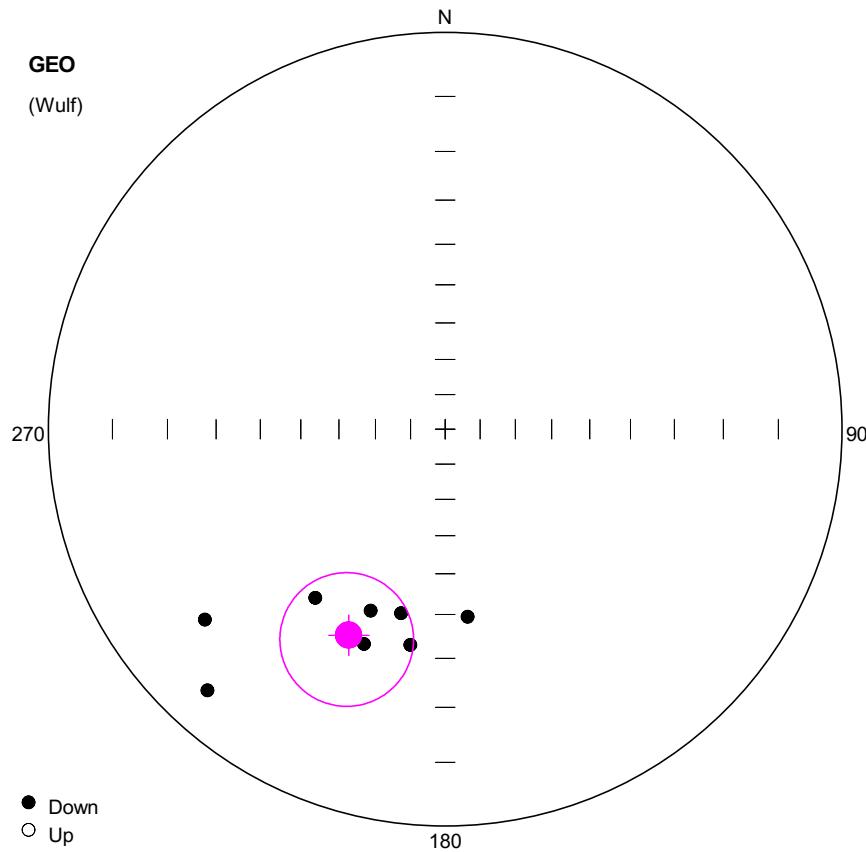
Pole longitude: -150.28 (29.72)

Paleolatitude: 23.82 (-23.82)

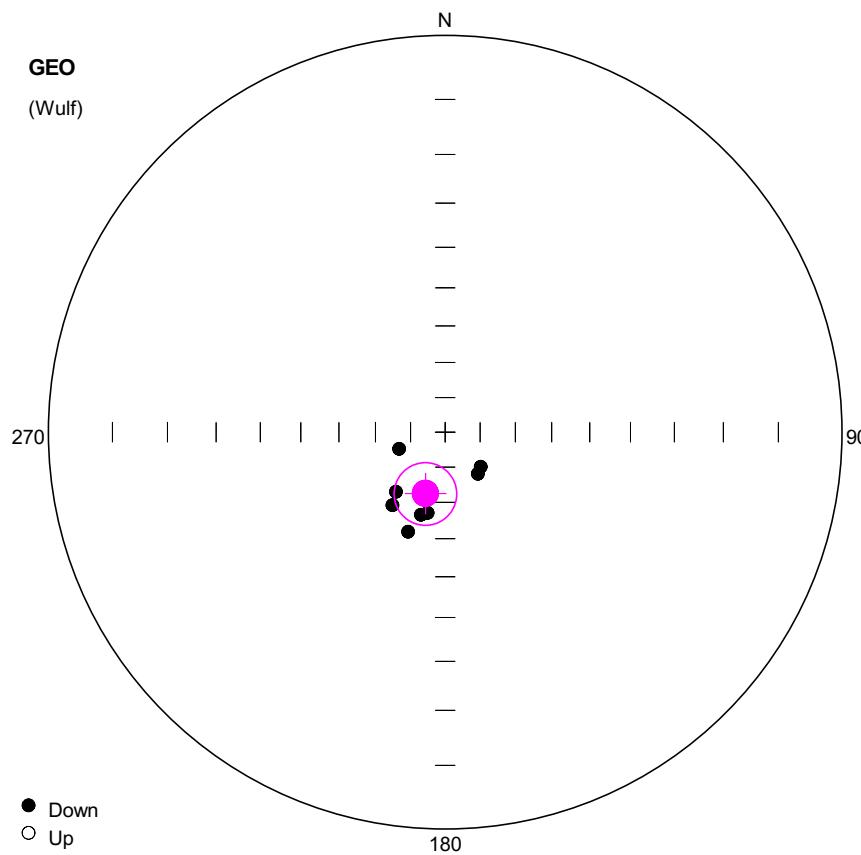
dp: 4.98

dm: 8.16

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8101	Ch	345.3	39.9	6.8	40	64
2	YUNN8102	Ch	1.6	45.7	6.0	35	64
3	YUNN8103	Ch	4.8	38.1	7.0	40	64
4	YUNN8104	Ch	342.5	41.9	6.5	30	64
5	YUNN8105	Ch	333.3	36.2	12.2	44	64
6	YUNN8106	Ch	350.8	38.4	13.2	30	64
7	YUNN8107	Ch	356.4	49.3	7.8	20	64
8	YUNN8109	Ch	337.2	38.1	4.5	44	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8201	Ch	189.2	32.3	6.8	44	56
2	YUNN8202	Ch	217.6	33.6	5.1	30	64
3	YUNN8203	Ch	231.6	14.6	8.0	20	64
4	YUNN8204	Ch	202.3	37.4	9.7	40	64
5	YUNN8205	Ch	200.7	29.9	7.8	40	64
6	YUNN8206	Ch	222.3	6.7	5.4	35	64
7	YUNN8207	Ch	173.2	39.1	7.6	35	64
8	YUNN8208	Ch	193.5	39.0	20.8	20	60



**Number of data points: 8**

**Fisher statistics**

Mean vector: 197.85 / 71.56

Resultant vector: 7.83

(X= -2.36, Y= -0.76, Z= 7.43)

Estimated precision, k: 40.7

95% Confidence limit: 8.79

95% Confidence limit, approximation: 7.76

**Orientation matrix**

1st eigenvalue: 0.96

1st eigenvector: 198.06 / 71.45

2nd eigenvalue: 0.03

2nd eigenvector: 65.64 / 12.75

3rd eigenvalue: 0.01

3rd eigenvector: 332.59 / 13.24

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 32.06 (-32.06)

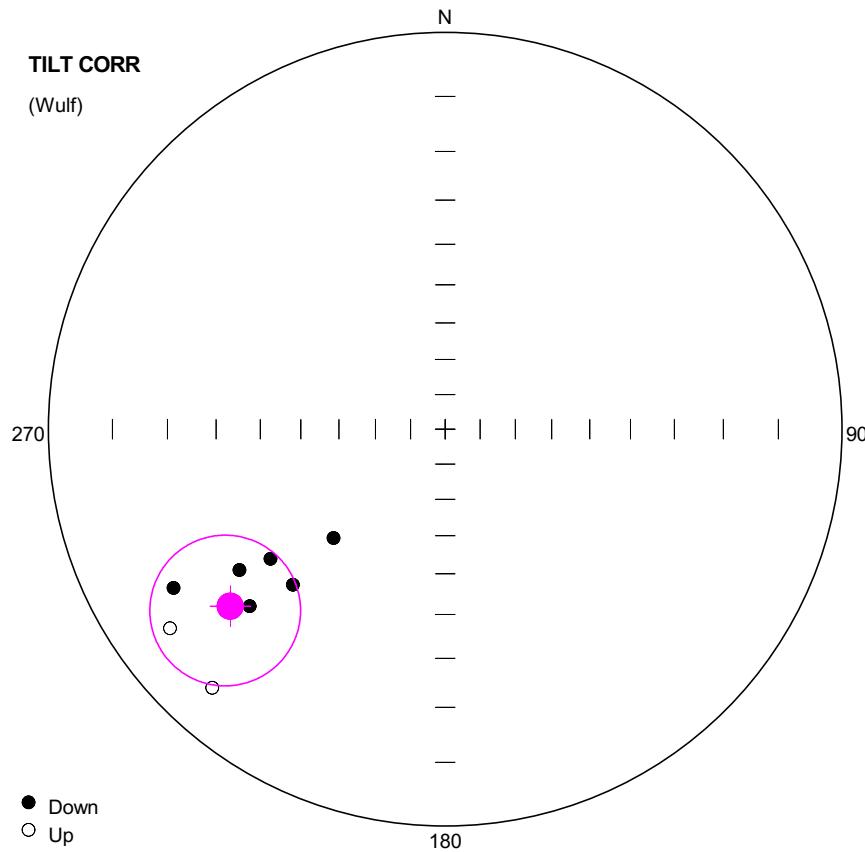
Pole longitude: 2.46 (-177.54)

Paleolatitude: 56.3 (-56.3)

dp: 13.52

dm: 15.41

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8201	CHT	192.3	66.5	1.0	64	64 OA
2	YUNN8202	CHT	196.4	65.5	1.0	64	64 OA
3	YUNN8203	CHT	219.6	68.0	1.0	64	64 OA
4	YUNN8204	CHT	200.4	60.0	1.0	64	64 OA
5	YUNN8205	CHT	250.0	75.9	1.0	64	64 OA
6	YUNN8206	CHT	215.9	64.4	1.0	64	64 OA
7	YUNN8207	CHT	134.3	75.7	1.0	64	64 OA
8	YUNN8208	CHT	142.0	74.8	1.0	64	64 OA



**Number of data points: 8**

## Fisher statistics

Mean vector: 230.53 / 19.9

Resultant vector: 7.56

(X= -4.52, Y= -5.49, Z= 2.57)

Estimated precision, k: 15.77

95% Confidence limit: 14.39

95% Confidence limit, approximation: 12.46

## Orientation matrix

1st eigenvalue: 0.89

1st eigenvector: 230.66 / 20.38

2nd eigenvalue: 0.1

2nd eigenvector: 53.1 / 69.6

3rd eigenvalue: 0.01

3rd eigenvector: 320.96 / 0.8

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -5.91 (5.91)

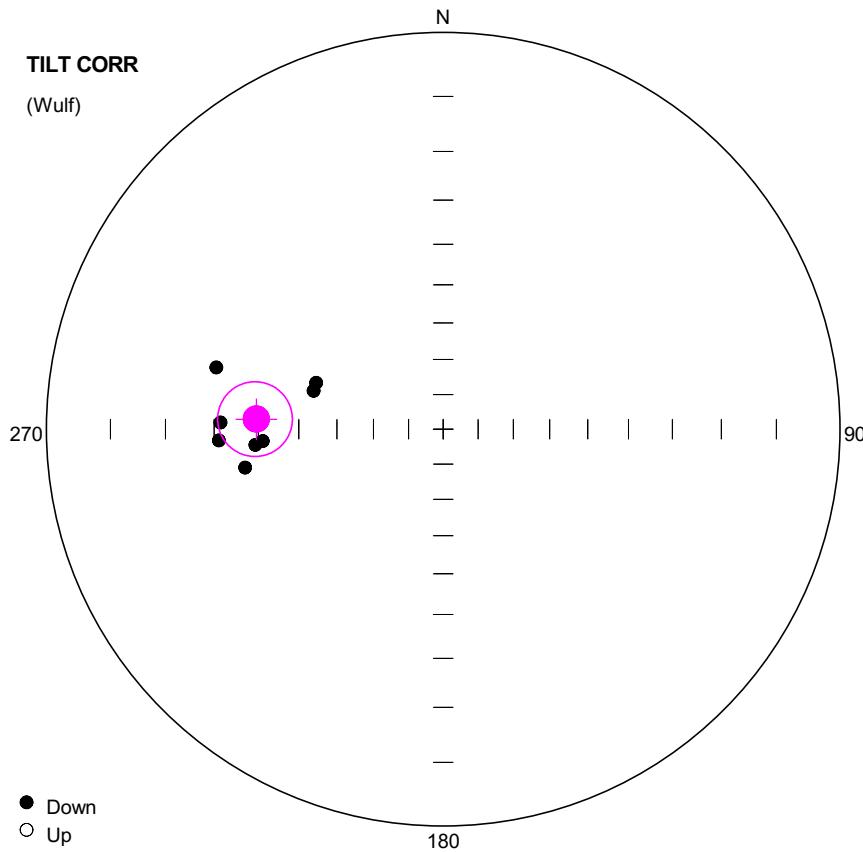
Pole longitude: -35.75 (144.25)

Paleolatitude: 10.26 (-10.26)

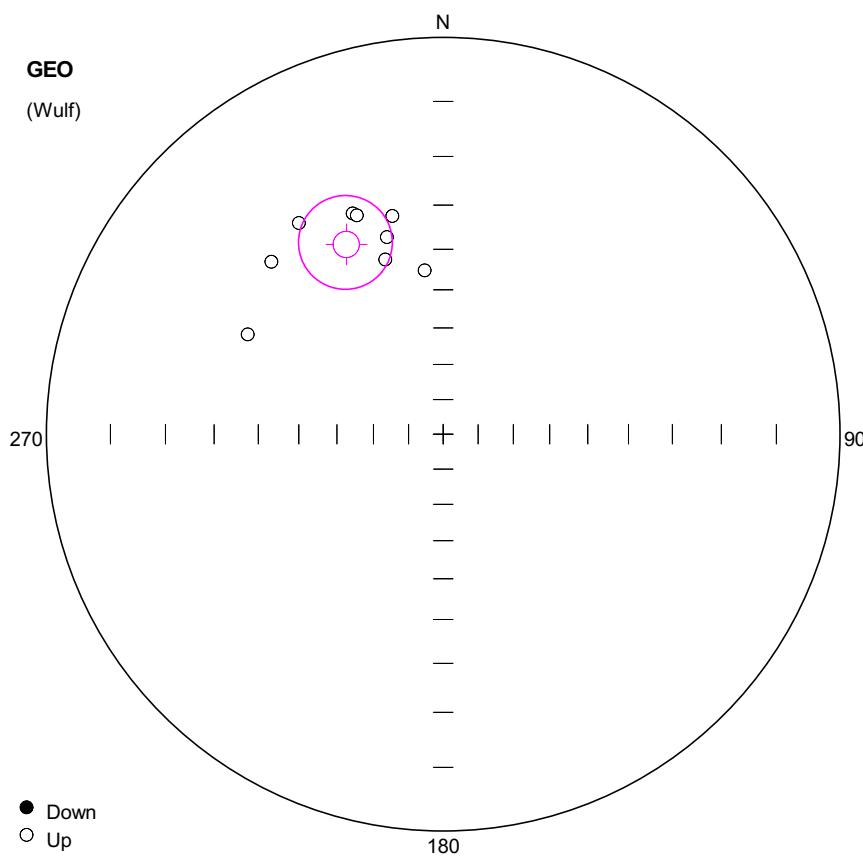
dp: 7.88

dm: 15.06

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8201	Ch	224.4	32.5	6.8	44	56
2	YUNN8202	Ch	239.7	13.2	5.1	30	64
3	YUNN8203	Ch	234.1	-8.9	8.0	20	64
4	YUNN8204	Ch	235.6	25.7	9.7	40	64
5	YUNN8205	Ch	227.8	22.8	7.8	40	64
6	YUNN8206	Ch	222.0	-7.5	5.4	35	64
7	YUNN8207	Ch	225.7	47.1	7.6	35	64
8	YUNN8208	Ch	233.4	32.5	20.8	20	60



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8201	CHT	266.2	41.0	1.0	64	64 OA
2	YUNN8202	CHT	265.2	39.2	1.0	64	64 OA
3	YUNN8203	CHT	271.7	31.4	1.0	64	64 OA
4	YUNN8204	CHT	259.0	36.1	1.0	64	64 OA
5	YUNN8205	CHT	285.2	28.7	1.0	64	64 OA
6	YUNN8206	CHT	267.1	31.0	1.0	64	64 OA
7	YUNN8207	CHT	290.1	52.4	1.0	64	64 OA
8	YUNN8208	CHT	286.5	52.4	1.0	64	64 OA



**Number of data points: 9**

**Fisher statistics**

Mean vector: 332.97 / -33.53

Resultant vector: 8.68

(X= 6.45, Y= -3.29, Z= -4.8)

Estimated precision, k: 25.22

95% Confidence limit: 10.45

95% Confidence limit, approximation: 9.29

**Orientation matrix**

1st eigenvalue: 0.93

1st eigenvector: 153.47 / 33.43

2nd eigenvalue: 0.06

2nd eigenvector: 248.81 / 8.03

3rd eigenvalue: 0.01

3rd eigenvector: 350.59 / 55.36

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 4.14 (-4.14)

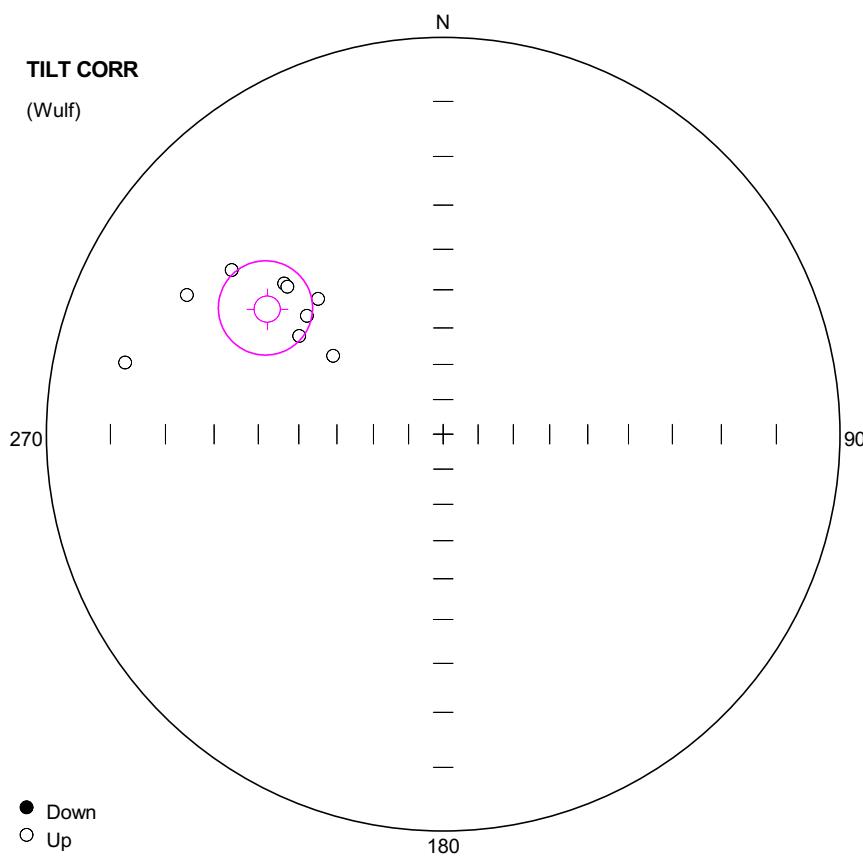
Pole longitude: -140.33 (39.67)

Paleolatitude: -18.33 (18.33)

dp: 6.78

dm: 11.9

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8301	Cht	341.7	-40.3	2.1	60	68
2	YUNN8302	Cht	346.9	-31.1	8.1	60	68
3	YUNN8303	Cht	337.7	-28.0	4.0	60	68
4	YUNN8305	Cht	315.1	-27.0	2.0	60	68
5	YUNN8306	Cht	353.6	-44.9	0.0	64	68
6	YUNN8307	Cht	344.1	-35.4	9.5	60	68
7	YUNN8308	Cht	297.0	-32.1	1.3	60	68
8	YUNN8309	Cht	338.5	-28.7	6.7	60	68
9	YUNN8310	Cht	325.7	-24.4	0.0	64	68



**Number of data points: 9**

#### Fisher statistics

Mean vector: 305.39 / -32.93

Resultant vector: 8.68

(X= 4.22, Y= -5.94, Z= -4.72)

Estimated precision, k: 25.28

95% Confidence limit: 10.44

95% Confidence limit, approximation: 9.28

#### Orientation matrix

1st eigenvalue: 0.93

1st eigenvector: 125.81 / 33.17

2nd eigenvalue: 0.06

2nd eigenvector: 257.01 / 45.22

3rd eigenvalue: 0.01

3rd eigenvector: 16.93 / 26.34

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -2.66 (2.66)

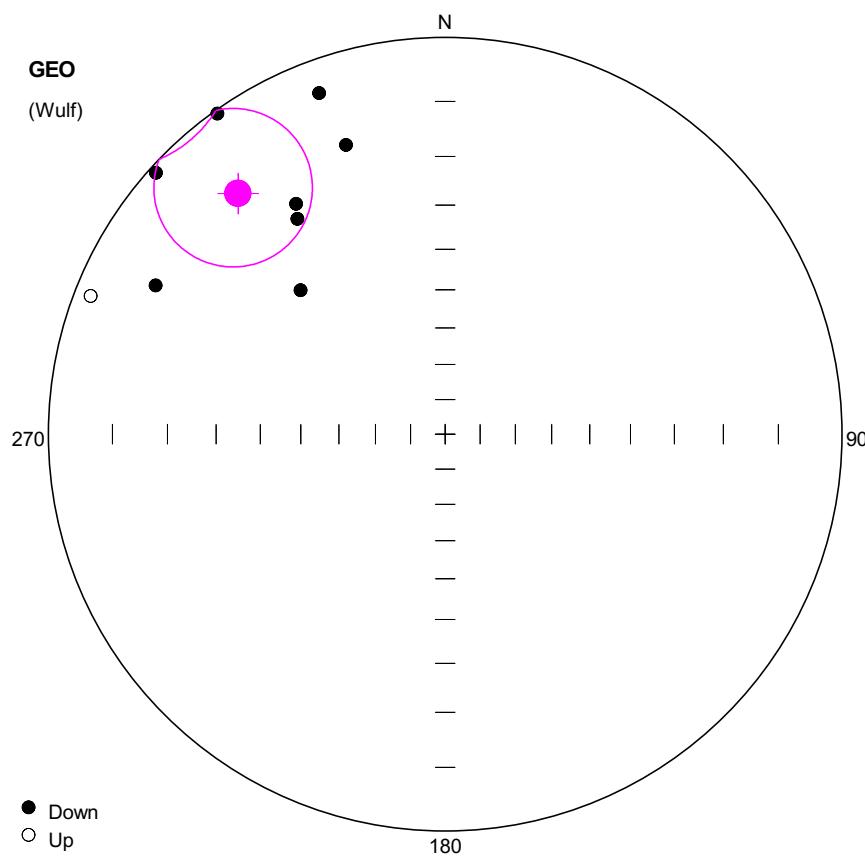
Pole longitude: -115.03 (64.97)

Paleolatitude: -17.94 (17.94)

dp: 6.71

dm: 11.83

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8301	Cht	304.3	-42.6	2.1	60	68
2	YUNN8302	Cht	317.3	-40.2	8.1	60	68
3	YUNN8303	Cht	313.5	-32.2	4.0	60	68
4	YUNN8305	Cht	298.5	-17.4	2.0	60	68
5	YUNN8306	Cht	305.5	-52.4	0.0	64	68
6	YUNN8307	Cht	311.0	-41.1	9.5	60	68
7	YUNN8308	Cht	282.7	-11.2	1.3	60	68
8	YUNN8309	Cht	313.4	-33.2	6.7	60	68
9	YUNN8310	Cht	307.8	-22.0	0.0	64	68



**Number of data points: 9**

**Fisher statistics**

Mean vector: 319.29 / 12.65  
 Resultant vector: 8.47  
 $(X= 6.26, Y= -5.39, Z= 1.85)$   
 Estimated precision,  $k$ : 14.98  
 95% Confidence limit: 13.75  
 95% Confidence limit, approximation: 12.06

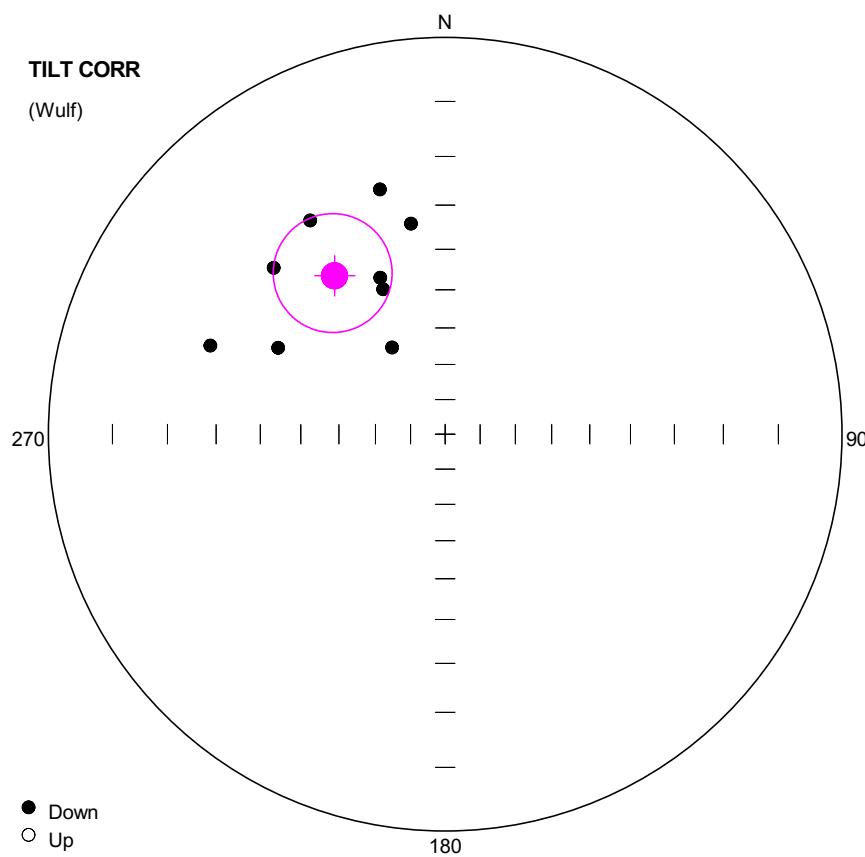
**Orientation matrix**

1st eigenvalue: 0.89  
 1st eigenvector: 319.68 / 12.75  
 2nd eigenvalue: 0.08  
 2nd eigenvector: 53.77 / 17.48  
 3rd eigenvalue: 0.04  
 3rd eigenvector: 195.4 / 68.11

**VGP**

Site latitude: 65  
 Site longitude: 14.04  
 Pole latitude: 24.8 (-24.8)  
 Pole longitude: -120.4 (59.6)  
 Paleolatitude: 6.4 (-6.4)  
 dp: 7.13  
 dm: 14

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8401	Cmt	339.7	5.0	6.9	30	64
2	YUNN8402	Cmt	324.6	0.5	7.0	20	64
3	YUNN8403	Cmt	314.9	35.6	8.3	40	64
4	YUNN8405	Cmt	341.1	14.8	6.3	20	64
5	YUNN8406	Cmt	327.1	20.7	8.0	35	64
6	YUNN8407	Cmt	325.5	23.3	12.3	40	60
7	YUNN8408	Cmt	297.2	11.3	8.3	44	64
8	YUNN8409	Cmt	291.3	-2.4	18.6	35	64
9	YUNN8410	Cmt	312.1	1.0	13.9	20	64



**Number of data points: 9**

**Fisher statistics**

Mean vector: 325.06 / 38.09  
Resultant vector: 8.47  
(X= 5.46, Y= -3.82, Z= 5.22)  
Estimated precision, k: 14.97  
95% Confidence limit: 13.75  
95% Confidence limit, approximation: 12.06

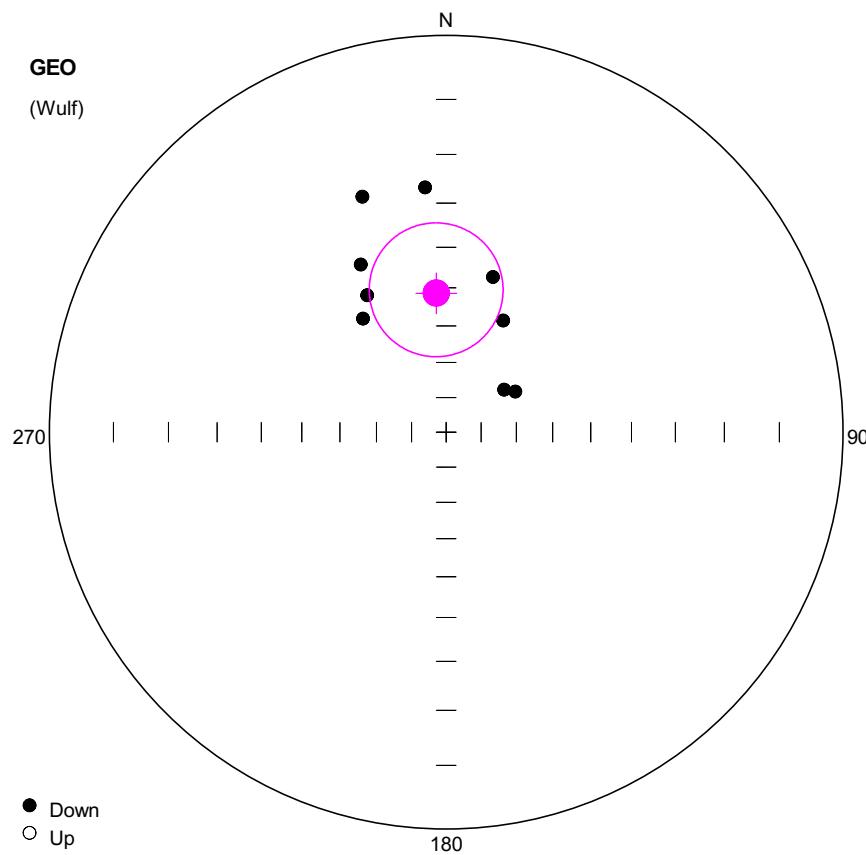
**Orientation matrix**

1st eigenvalue: 0.89  
1st eigenvector: 325.56 / 38.11  
2nd eigenvalue: 0.08  
2nd eigenvector: 58.61 / 3.88  
3rd eigenvalue: 0.04  
3rd eigenvector: 153.52 / 51.63

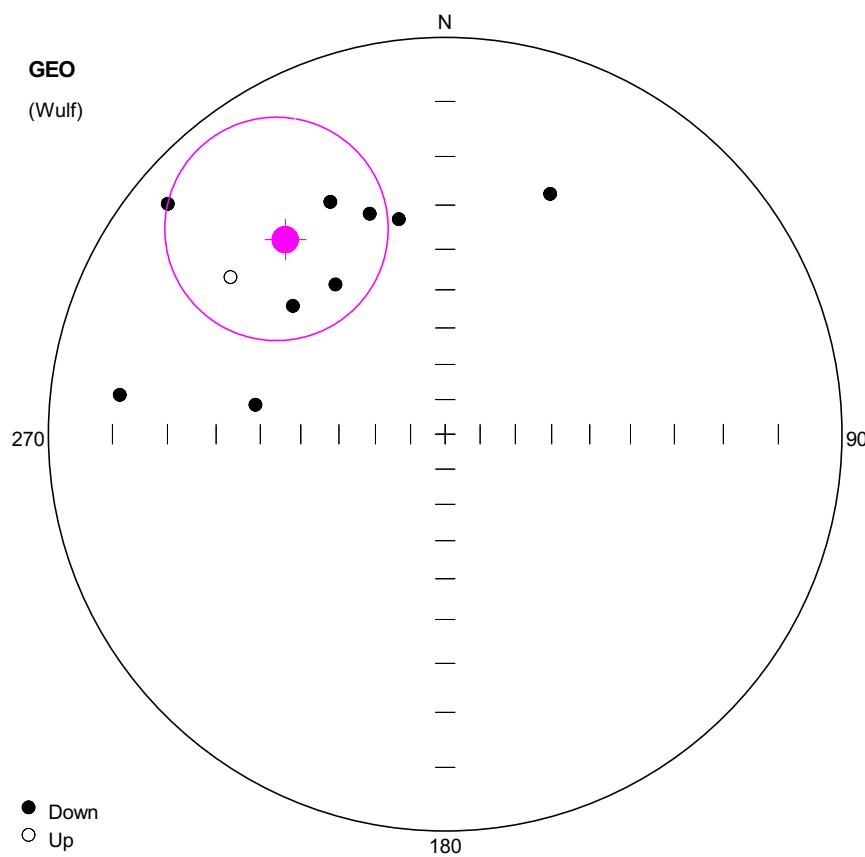
**VGP**

Site latitude: 65  
Site longitude: 14.04  
Pole latitude: 40.79 (-40.79)  
Pole longitude: -121.19 (58.81)  
Paleolatitude: 21.4 (-21.4)  
dp: 9.62  
dm: 16.27

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8401	Cmt	345.1	24.9	6.9	30	64
2	YUNN8402	Cmt	327.7	25.0	7.0	20	64
3	YUNN8403	Cmt	328.5	61.3	8.3	40	64
4	YUNN8405	Cmt	350.8	33.5	6.3	20	64
5	YUNN8406	Cmt	337.5	43.8	8.0	35	64
6	YUNN8407	Cmt	336.8	46.7	12.3	40	60
7	YUNN8408	Cmt	297.3	39.3	8.3	44	64
8	YUNN8409	Cmt	290.7	25.4	18.6	35	64
9	YUNN8410	Cmt	314.1	27.9	13.9	20	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8502	Cmt	53.8	69.5	5.8	30	64
2	YUNN8503	Cmt	16.8	45.6	6.1	30	64
3	YUNN8504	Cmt	323.8	51.0	7.9	40	64
4	YUNN8505	Cmt	27.1	54.9	14.0	40	56
5	YUNN8506	Cmt	330.0	46.6	15.8	30	64
6	YUNN8507	Cmt	355.1	26.5	6.4	20	64
7	YUNN8508	Cmt	59.6	67.2	8.1	35	52
8	YUNN8509	Cmt	340.4	25.6	13.8	30	60
9	YUNN8510	Cmt	333.0	39.3	14.4	30	60



**Number of data points: 10**

**Fisher statistics**

Mean vector: 320.59 / 25.2

Resultant vector: 8.4

(X= 5.87, Y= -4.83, Z= 3.58)

Estimated precision, k: 5.64

95% Confidence limit: 22.34

95% Confidence limit, approximation: 18.65

**Orientation matrix**

1st eigenvalue: 0.73

1st eigenvector: 320.26 / 26.4

2nd eigenvalue: 0.19

2nd eigenvector: 59.46 / 17.86

3rd eigenvalue: 0.08

3rd eigenvector: 179.59 / 57.31

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 31.69 (-31.69)

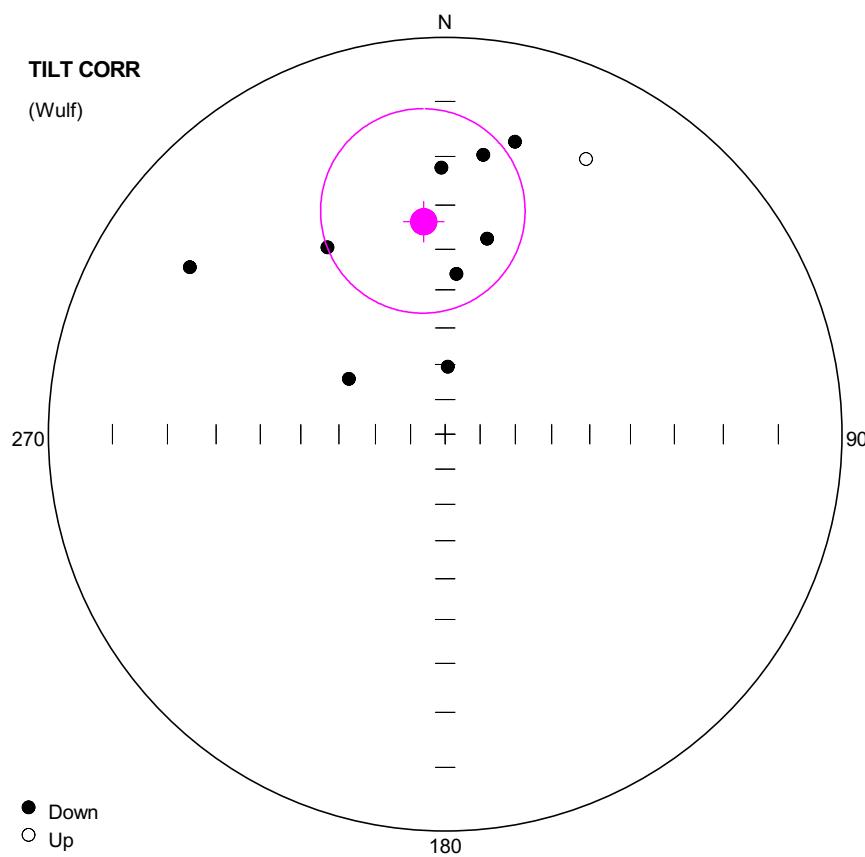
Pole longitude: -119.38 (60.62)

Paleolatitude: 13.24 (-13.24)

dp: 12.93

dm: 24.03

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8501	CHT	333.7	23.7	1.0	64	64 OA
2	YUNN8502	CHT	347.8	32.0	1.0	64	64 OA
3	YUNN8503	CHT	341.1	29.2	1.0	64	64 OA
4	YUNN8504	CHT	310.1	36.7	1.0	64	64 OA
5	YUNN8505	CHT	278.8	38.4	1.0	64	64 OA
6	YUNN8506	CHT	323.8	39.9	1.0	64	64 OA
7	YUNN8507	CHT	276.9	10.9	1.0	64	64 OA
8	YUNN8508	CHT	23.6	23.1	1.0	64	64 OA
9	YUNN8509	CHT	309.7	5.5	1.0	64	64 OA
10	YUNN8510	CHT	306.2	-22.3	1.0	64	64 OA



Number of data points: 10

#### Fisher statistics

Mean vector: 354.25 / 33.41

Resultant vector: 8.4

(X= 6.98, Y= -0.7, Z= 4.63)

Estimated precision, k: 5.63

95% Confidence limit: 22.34

95% Confidence limit, approximation: 18.65

#### Orientation matrix

1st eigenvalue: 0.73

1st eigenvector: 355.41 / 34.21

2nd eigenvalue: 0.19

2nd eigenvector: 236.86 / 35.1

3rd eigenvalue: 0.08

3rd eigenvector: 115.57 / 36.47

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 43.1 (-43.1)

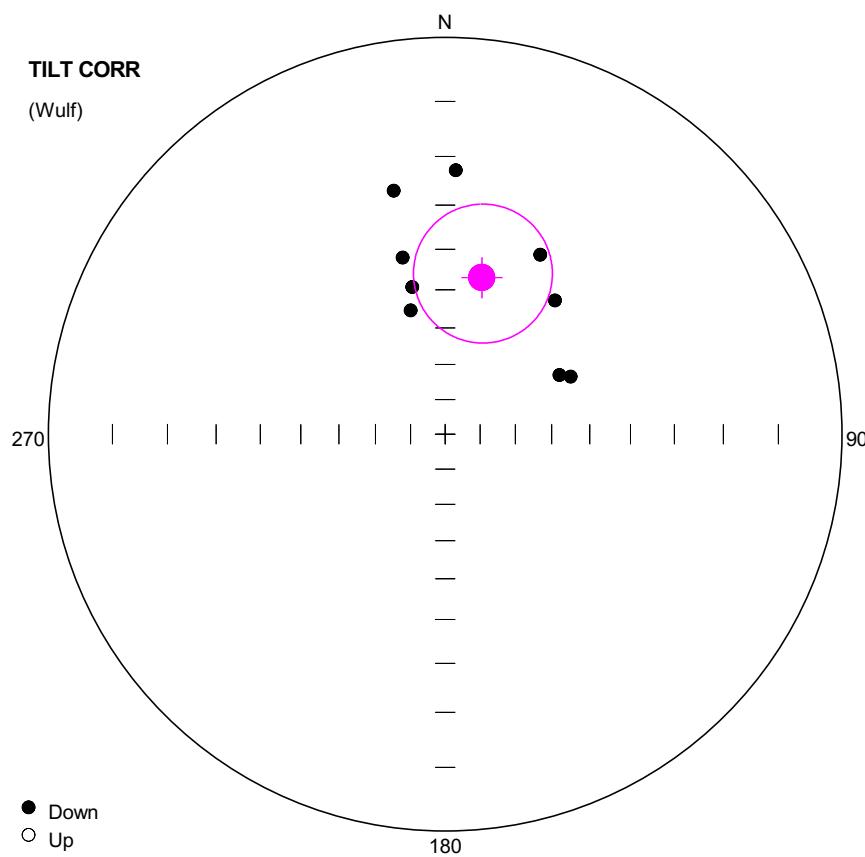
Pole longitude: -158.48 (21.52)

Paleolatitude: 18.26 (-18.26)

dp: 14.46

dm: 25.42

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8501	CHT	359.2	22.2	1.0	64	64 OA
2	YUNN8502	CHT	13.4	15.7	1.0	64	64 OA
3	YUNN8503	CHT	7.8	19.2	1.0	64	64 OA
4	YUNN8504	CHT	4.0	45.9	1.0	64	64 OA
5	YUNN8505	CHT	2.2	70.7	1.0	64	64 OA
6	YUNN8506	CHT	12.1	36.5	1.0	64	64 OA
7	YUNN8507	CHT	299.9	58.7	1.0	64	64 OA
8	YUNN8508	CHT	27.1	-14.2	1.0	64	64 OA
9	YUNN8509	CHT	327.8	31.8	1.0	64	64 OA
10	YUNN8510	CHT	303.2	14.9	1.0	64	64 OA



**Number of data points: 9**

**Fisher statistics**

Mean vector: 13.18 / 45.88

Resultant vector: 8.21

(X= 5.56, Y= 1.3, Z= 5.89)

Estimated precision, k: 10.09

95% Confidence limit: 17.03

95% Confidence limit, approximation: 14.69

**Orientation matrix**

1st eigenvalue: 0.83

1st eigenvector: 12.35 / 45.59

2nd eigenvalue: 0.13

2nd eigenvector: 120.79 / 17.22

3rd eigenvalue: 0.03

3rd eigenvector: 225.51 / 39.35

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 51.36 (-51.36)

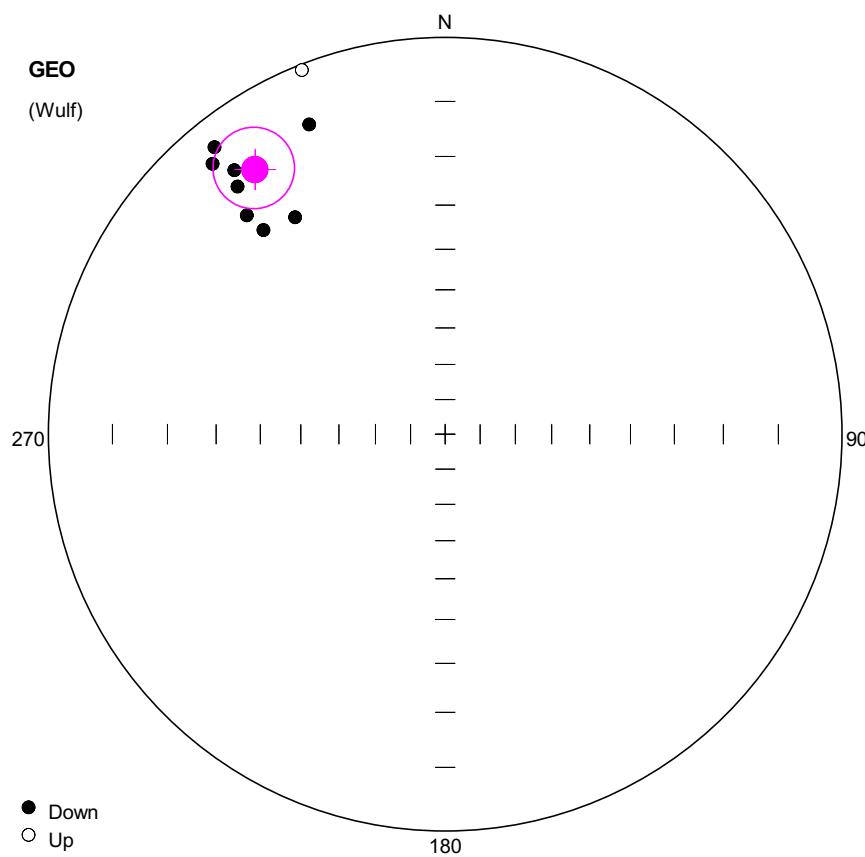
Pole longitude: 175.1 (-4.9)

Paleolatitude: 27.28 (-27.28)

dp: 13.88

dm: 21.75

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8502	Cmt	62.6	54.1	5.8	30	64
2	YUNN8503	Cmt	27.9	35.8	6.1	30	64
3	YUNN8504	Cmt	344.5	54.1	7.9	40	64
4	YUNN8505	Cmt	39.4	42.9	14.0	40	56
5	YUNN8506	Cmt	347.4	48.4	15.8	30	64
6	YUNN8507	Cmt	2.3	22.7	6.4	20	64
7	YUNN8508	Cmt	65.4	51.6	8.1	35	52
8	YUNN8509	Cmt	348.1	25.8	13.8	30	60
9	YUNN8510	Cmt	346.5	40.8	14.4	30	60



**Number of data points: 9**

**Fisher statistics**

Mean vector: 324.28 / 11.2

Resultant vector: 8.85

(X= 7.05, Y= -5.07, Z= 1.72)

Estimated precision, k: 55.13

95% Confidence limit: 7

95% Confidence limit, approximation: 6.29

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 324.19 / 11.23

2nd eigenvalue: 0.02

2nd eigenvector: 223.03 / 44.26

3rd eigenvalue: 0.01

3rd eigenvector: 65.07 / 43.57

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 25.51 (-25.51)

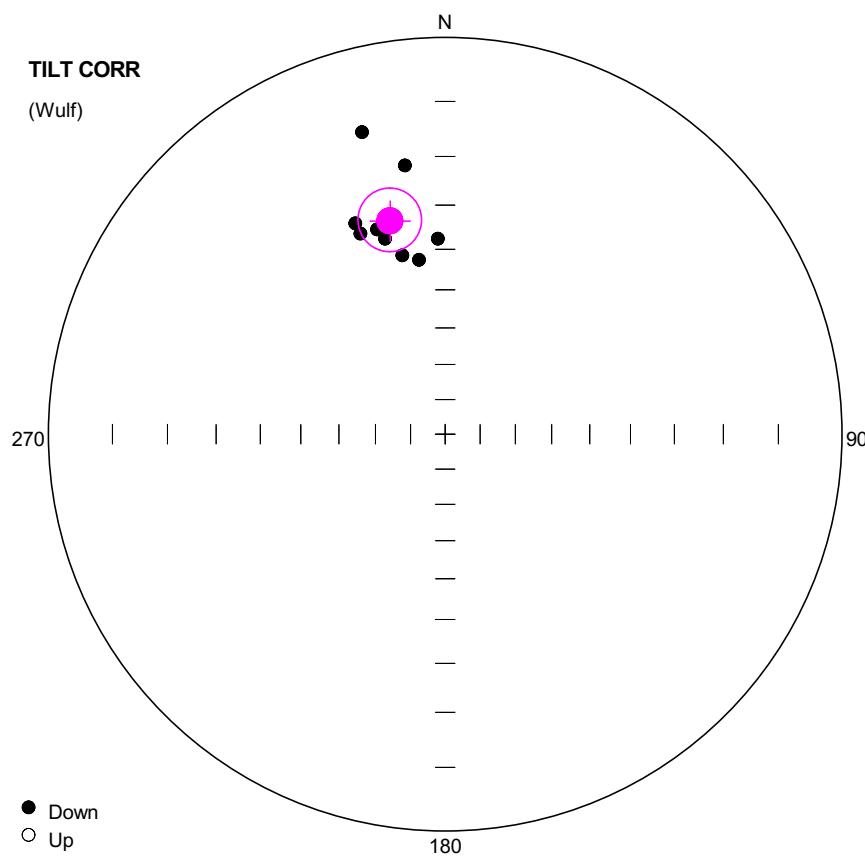
Pole longitude: -125.88 (54.12)

Paleolatitude: 5.65 (-5.65)

dp: 3.6

dm: 7.1

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8601	Cmt	319.3	6.1	2.2	30	56 OA
2	YUNN8602	Cmt	320.0	11.7	4.6	30	56 OA
3	YUNN8603	Cmt	336.3	9.1	3.5	20	56 OA
4	YUNN8604	Cmt	325.3	22.8	27.2	20	64
5	YUNN8605	Cmt	321.2	4.3	3.0	20	56 OA
6	YUNN8606	Cmt	317.8	16.7	4.2	30	60 OA
7	YUNN8607	Cmt	321.4	9.2	4.0	20	60 OA
8	YUNN8608	Cmt	318.3	20.9	6.9	20	64 OA
9	YUNN8609	Cmt	338.5	-0.8	3.6	20	52 OA



**Number of data points: 9**

**Fisher statistics**

Mean vector: 345.47 / 31.91

Resultant vector: 8.85

(X= 7.28, Y= -1.89, Z= 4.68)

Estimated precision, k: 55.13

95% Confidence limit: 7

95% Confidence limit, approximation: 6.28

**Orientation matrix**

1st eigenvalue: 0.97

1st eigenvector: 345.44 / 32

2nd eigenvalue: 0.02

2nd eigenvector: 149.19 / 56.94

3rd eigenvalue: 0.01

3rd eigenvector: 250.76 / 7.44

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 41.3 (-41.3)

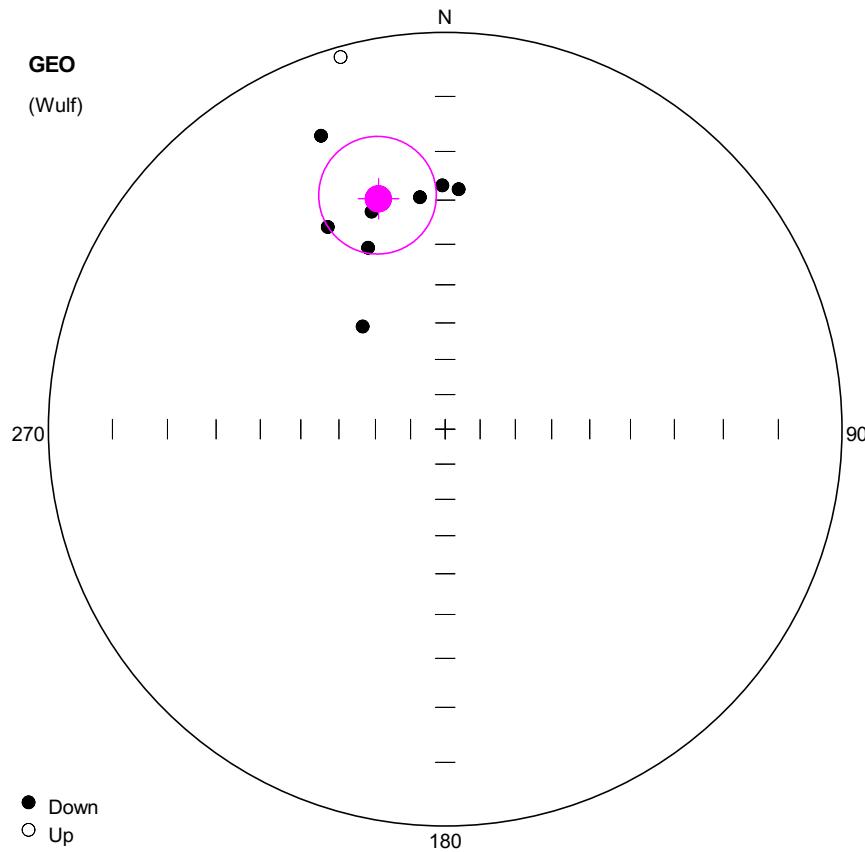
Pole longitude: -147.37 (32.63)

Paleolatitude: 17.29 (-17.29)

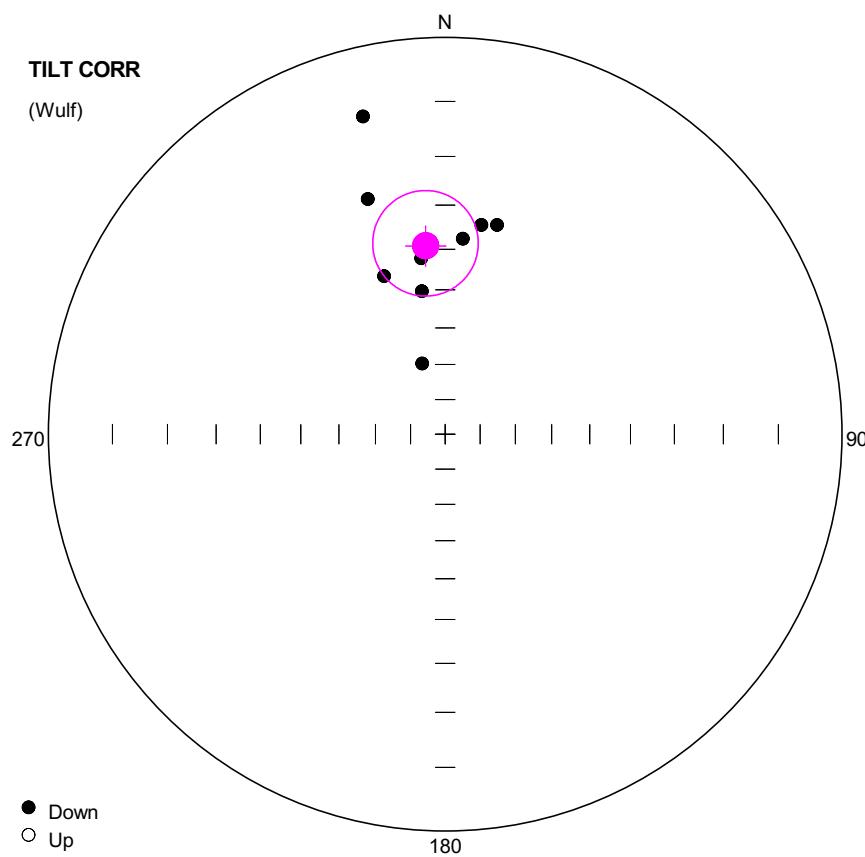
dp: 4.42

dm: 7.87

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8601	Cmt	337.1	32.5	2.2	30	56 OA
2	YUNN8602	Cmt	342.9	35.5	4.6	30	56 OA
3	YUNN8603	Cmt	351.5	21.2	3.5	20	56 OA
4	YUNN8604	Cmt	357.9	37.5	27.2	20	64
5	YUNN8605	Cmt	336.9	30.0	3.0	20	56 OA
6	YUNN8606	Cmt	346.5	40.3	4.2	30	60 OA
7	YUNN8607	Cmt	341.6	32.9	4.0	20	60 OA
8	YUNN8608	Cmt	351.5	42.1	6.9	20	64 OA
9	YUNN8609	Cmt	344.6	13.4	3.6	20	52 OA



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8701	Cmt	329.9	29.0	8.2	30	56
2	YUNN8702	Cmt	337.0	37.2	12.5	20	52
3	YUNN8703	Cmt	321.2	53.3	9.8	30	56
4	YUNN8704	Cmt	344.3	-1.5	11.3	35	60
5	YUNN8705	Cmt	359.3	26.9	17.7	30	60
6	YUNN8706	Cmt	337.1	12.5	23.7	30	60
7	YUNN8707	Cmt	341.3	29.9	6.2	35	56
8	YUNN8708	Cmt	3.2	27.6	14.6	44	60
9	YUNN8709	Cmt	353.8	29.1	6.6	44	60



**Number of data points: 9**

**Fisher statistics**

Mean vector: 354.08 / 39  
 Resultant vector: 8.57  
 $(X= 6.62, Y= -0.69, Z= 5.39)$   
 Estimated precision, k: 18.4  
 95% Confidence limit: 12.33  
 95% Confidence limit, approximation: 10.88

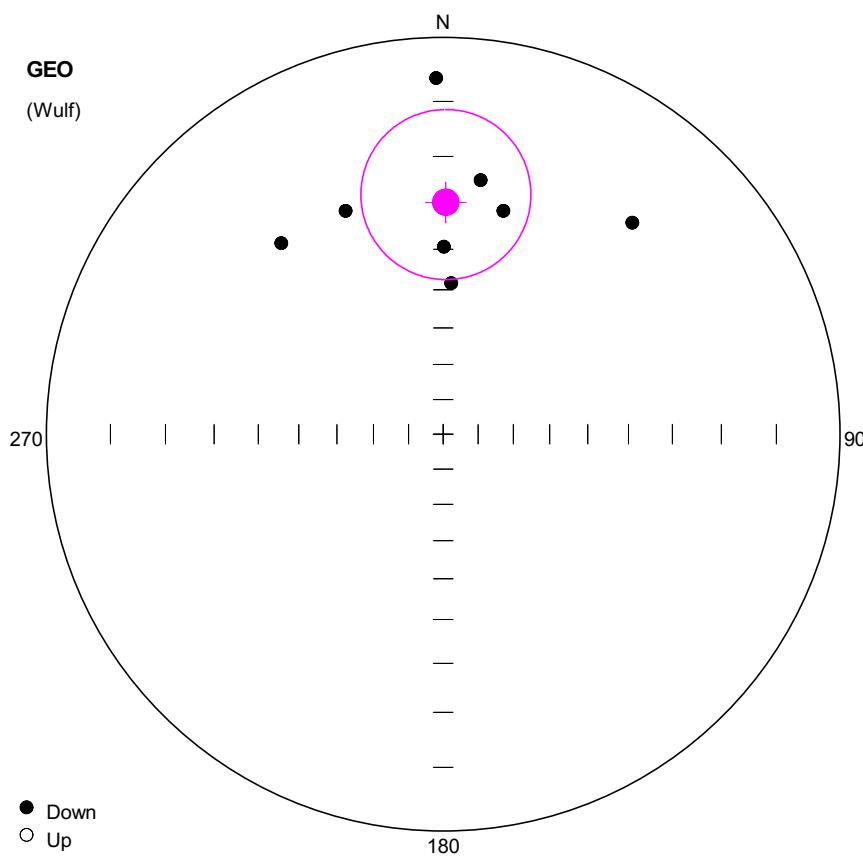
**Orientation matrix**

1st eigenvalue: 0.91  
 1st eigenvector: 354.27 / 38.98  
 2nd eigenvalue: 0.06  
 2nd eigenvector: 173.12 / 51.01  
 3rd eigenvalue: 0.03  
 3rd eigenvector: 263.82 / 0.56

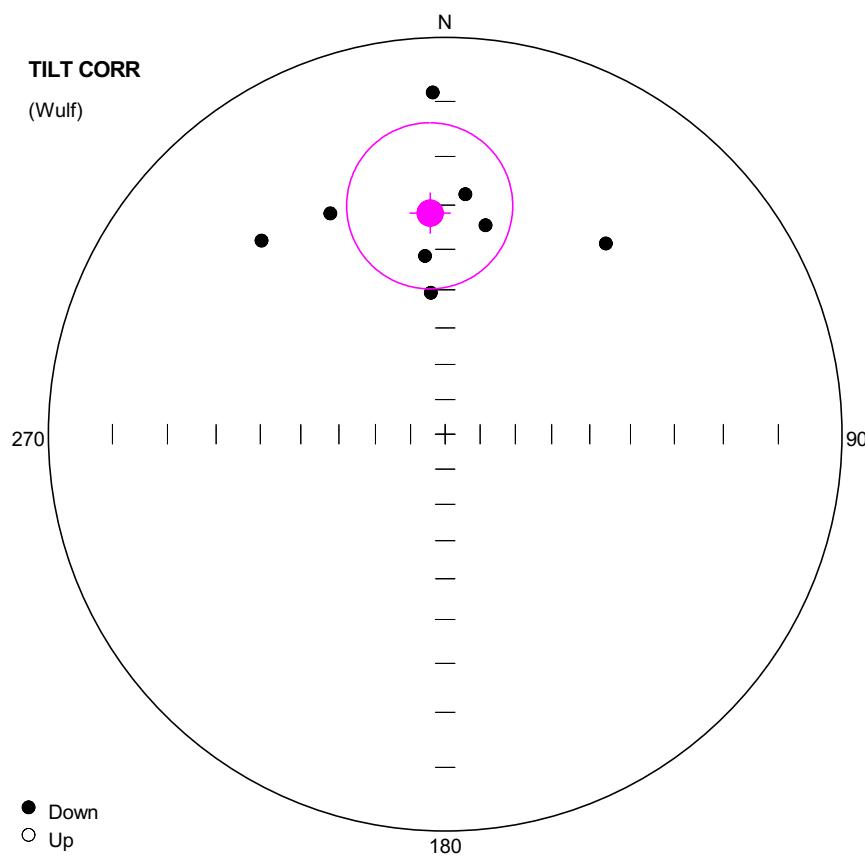
**VGP**

Site latitude: 65  
 Site longitude: 14.04  
 Pole latitude: 46.87 (-46.87)  
 Pole longitude: -157.92 (22.08)  
 Paleolatitude: 22.04 (-22.04)  
 dp: 8.77  
 dm: 14.7

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8701	Cmt	338.9	43.8	8.2	30	56
2	YUNN8702	Cmt	350.8	49.9	12.5	20	52
3	YUNN8703	Cmt	342.0	68.8	9.8	30	56
4	YUNN8704	Cmt	345.5	10.8	11.3	35	60
5	YUNN8705	Cmt	9.8	33.7	17.7	30	60
6	YUNN8706	Cmt	341.8	26.1	23.7	30	60
7	YUNN8707	Cmt	352.2	41.8	6.2	35	56
8	YUNN8708	Cmt	13.9	33.0	14.6	44	60
9	YUNN8709	Cmt	5.2	37.4	6.6	44	60



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8801	Cmt	0.2	39.4	8.8	44	64
2	YUNN8802	Cmt	8.4	24.2	7.9	40	64
3	YUNN8803	Cmt	336.4	26.9	11.5	35	60
4	YUNN8805	Cmt	358.9	6.2	12.2	35	60
5	YUNN8806	Cmt	319.7	25.5	10.9	30	60
6	YUNN8807	Cmt	15.1	29.5	11.6	30	64
7	YUNN8808	Cmt	3.1	48.3	13.5	20	48
8	YUNN8809	Cmt	41.8	18.9	5.4	40	64



**Number of data points: 8**

**Fisher statistics**

Mean vector: 356.1 / 31.63  
Resultant vector: 7.33  
(X= 6.23, Y= -0.42, Z= 3.84)  
Estimated precision, k: 10.44  
95% Confidence limit: 17.98  
95% Confidence limit, approximation: 15.32

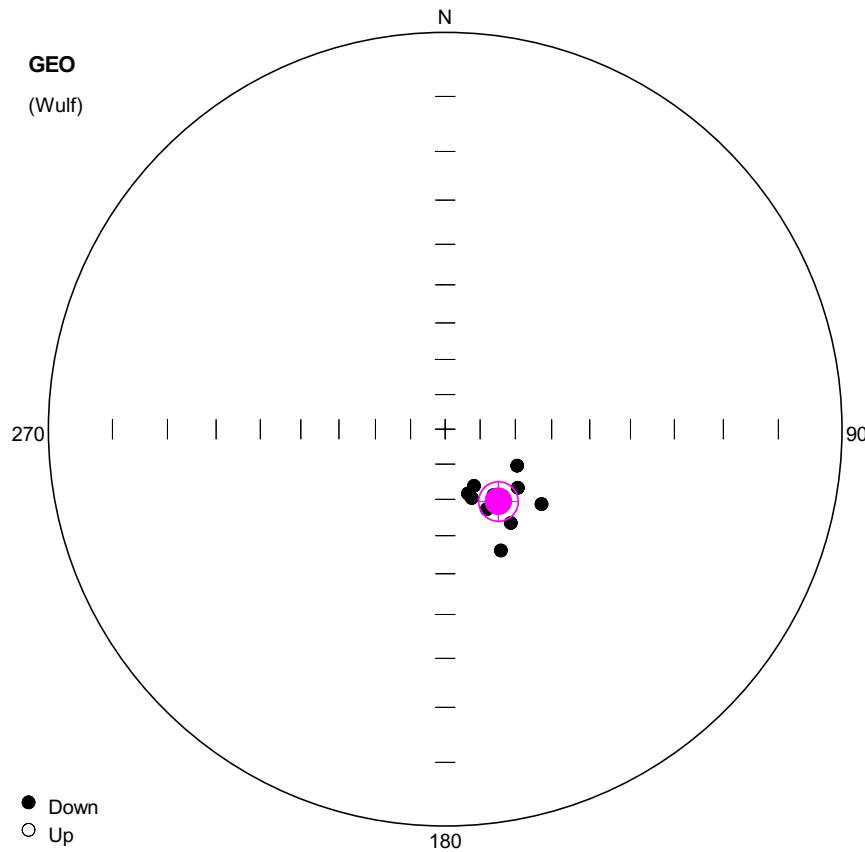
**Orientation matrix**

1st eigenvalue: 0.85  
1st eigenvector: 356.03 / 31.8  
2nd eigenvalue: 0.12  
2nd eigenvector: 88.35 / 3.74  
3rd eigenvalue: 0.04  
3rd eigenvector: 184.33 / 57.93

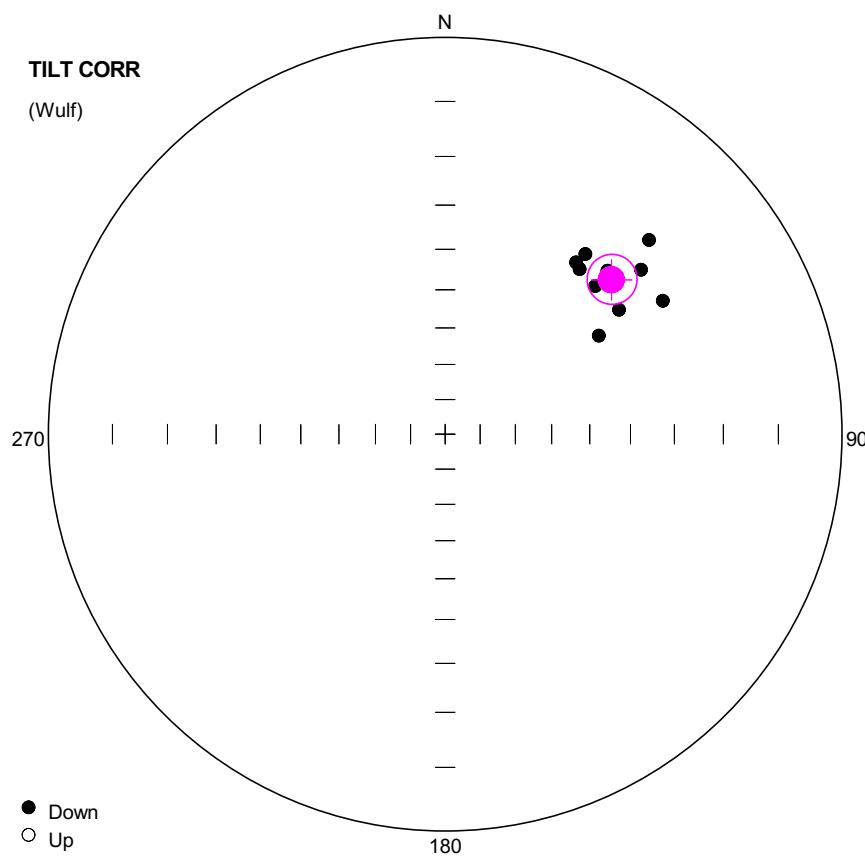
**VGP**

Site latitude: 65  
Site longitude: 14.04  
Pole latitude: 42.04 (-42.04)  
Pole longitude: -160.93 (19.07)  
Paleolatitude: 17.12 (-17.12)  
dp: 11.33  
dm: 20.18

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8801	Cmt	353.6	41.4	8.8	44	64
2	YUNN8802	Cmt	4.8	27.5	7.9	40	64
3	YUNN8803	Cmt	332.5	25.8	11.5	35	60
4	YUNN8805	Cmt	357.9	8.5	12.2	35	60
5	YUNN8806	Cmt	316.5	22.2	10.9	30	60
6	YUNN8807	Cmt	10.9	33.6	11.6	30	64
7	YUNN8808	Cmt	354.2	50.6	13.5	20	48
8	YUNN8809	Cmt	40.1	25.7	5.4	40	64



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8901	Cht	159.0	68.9	0.0	64	68
2	YUNN8902	Cht	152.5	64.4	0.0	64	68
3	YUNN8903	Cht	128.8	63.5	0.0	64	68
4	YUNN8904	Cht	116.8	67.0	0.0	64	68
5	YUNN8905	Cht	145.0	57.8	0.0	64	68
6	YUNN8906	Cht	127.9	55.8	0.0	64	68
7	YUNN8907	Cht	155.3	52.8	0.0	64	68
8	YUNN8908	Cht	152.9	71.8	0.0	64	68
9	YUNN8909	Cht	160.4	70.5	0.0	64	68
10	YUNN8910	Cht	143.7	66.7	0.0	64	68



Number of data points: 10

#### Fisher statistics

Mean vector: 47.11 / 30.47

Resultant vector: 9.89

(X= 5.8, Y= 6.24, Z= 5.01)

Estimated precision, k: 80.32

95% Confidence limit: 5.42

95% Confidence limit, approximation: 4.94

#### Orientation matrix

1st eigenvalue: 0.98

1st eigenvector: 47.08 / 30.47

2nd eigenvalue: 0.01

2nd eigenvector: 309.93 / 11.95

3rd eigenvalue: 0.01

3rd eigenvector: 201.06 / 56.79

#### VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: 32.12 (-32.12)

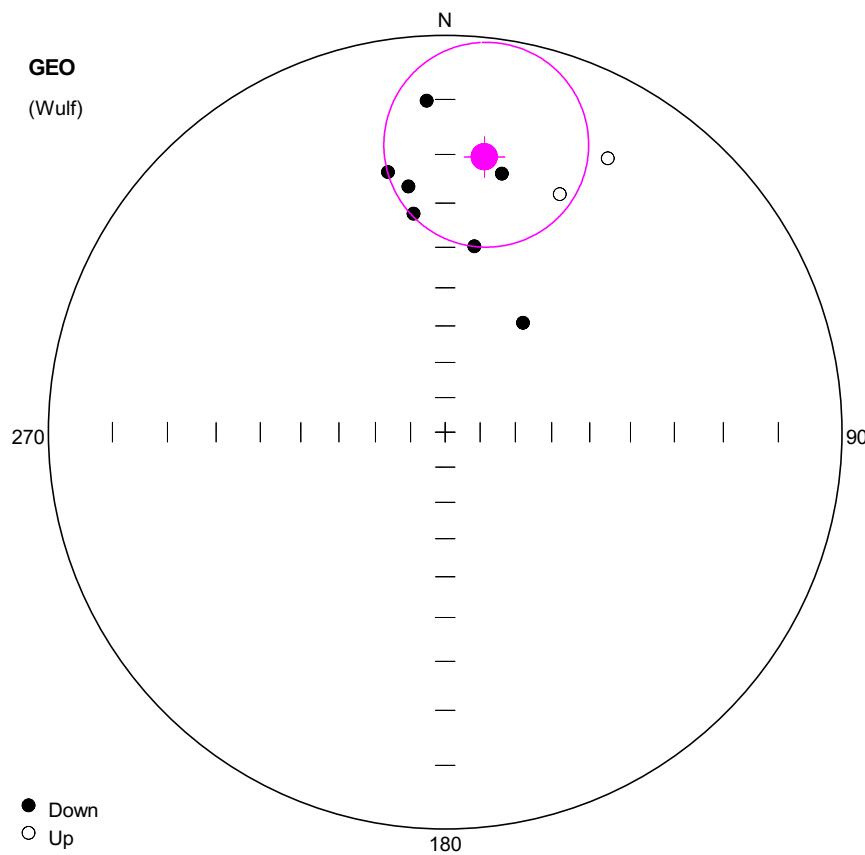
Pole longitude: 137.95 (-42.05)

Paleolatitude: 16.39 (-16.39)

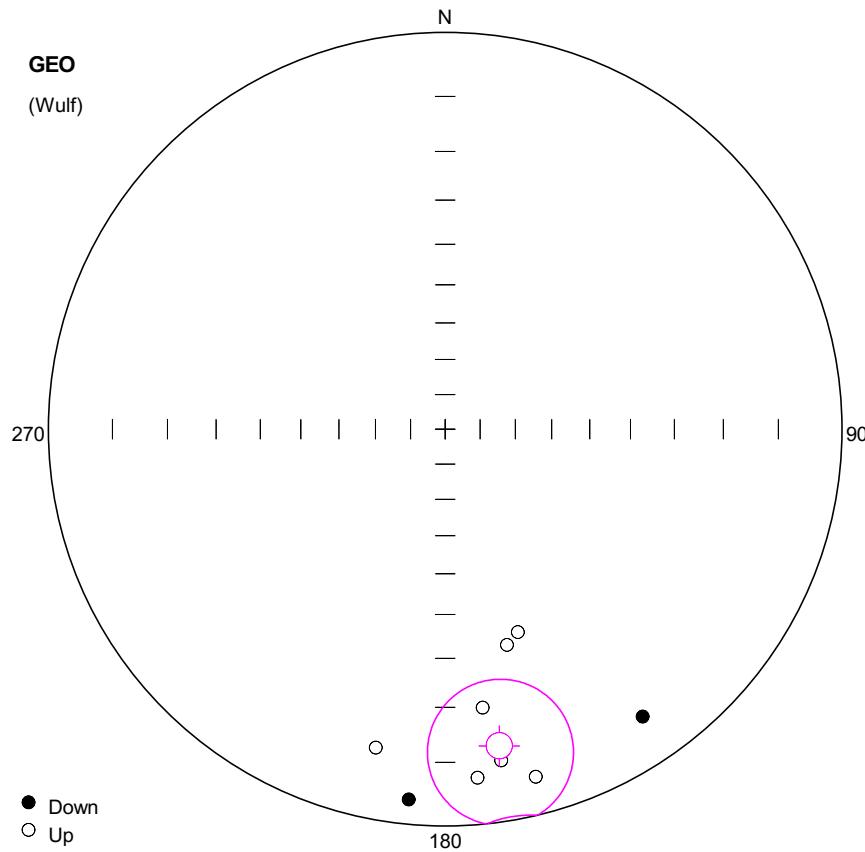
dp: 3.36

dm: 6.04

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN8901	Cht	39.2	33.6	0.0	64	68
2	YUNN8902	Cht	45.4	34.0	0.0	64	68
3	YUNN8903	Cht	50.0	24.4	0.0	64	68
4	YUNN8904	Cht	46.4	19.3	0.0	64	68
5	YUNN8905	Cht	54.4	33.4	0.0	64	68
6	YUNN8906	Cht	58.5	24.5	0.0	64	68
7	YUNN8907	Cht	57.4	40.6	0.0	64	68
8	YUNN8908	Cht	37.9	30.2	0.0	64	68
9	YUNN8909	Cht	37.3	32.9	0.0	64	68
10	YUNN8910	Cht	44.8	29.7	0.0	64	68



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9101	Cmt	12.4	22.6	9.9	30	52
2	YUNN9102	Cmt	25.7	-22.7	4.1	30	48
3	YUNN9103	Cmt	351.8	31.8	7.8	44	56
4	YUNN9104	Cmt	8.9	39.3	6.8	40	60
5	YUNN9105	Cmt	30.7	-12.5	5.6	35	52
6	YUNN9106	Cmt	35.4	52.6	11.6	20	60
7	YUNN9109	Cmt	356.8	10.2	9.2	35	60
8	YUNN9110	Cmt	351.5	25.9	3.2	35	64
9	YUNN9111	Cmt	347.6	22.3	7.9	40	56



**Number of data points: 9**

## Fisher statistics

Mean vector: 170.31 / -12.04

Resultant vector: 8.55

(X= -8.24, Y= 1.41, Z= -1.78)

Estimated precision, k: 17.69

95% Confidence limit: 12.5

95% Confidence limit, approximation: 11.1

## Orientation matrix

1st eigenvalue: 0.9

1st eigenvector: 350.56 / 12.17

2nd eigenvalue: 0.05

2nd eigenvector: 238.92 / 59.68

3rd eigenvalue: 0.05

3rd eigenvector: 86.95 / 27.3

VGP

Site latitude: 65

Site longitude: 14.04

Pole latitude: -30.68 (30.68)

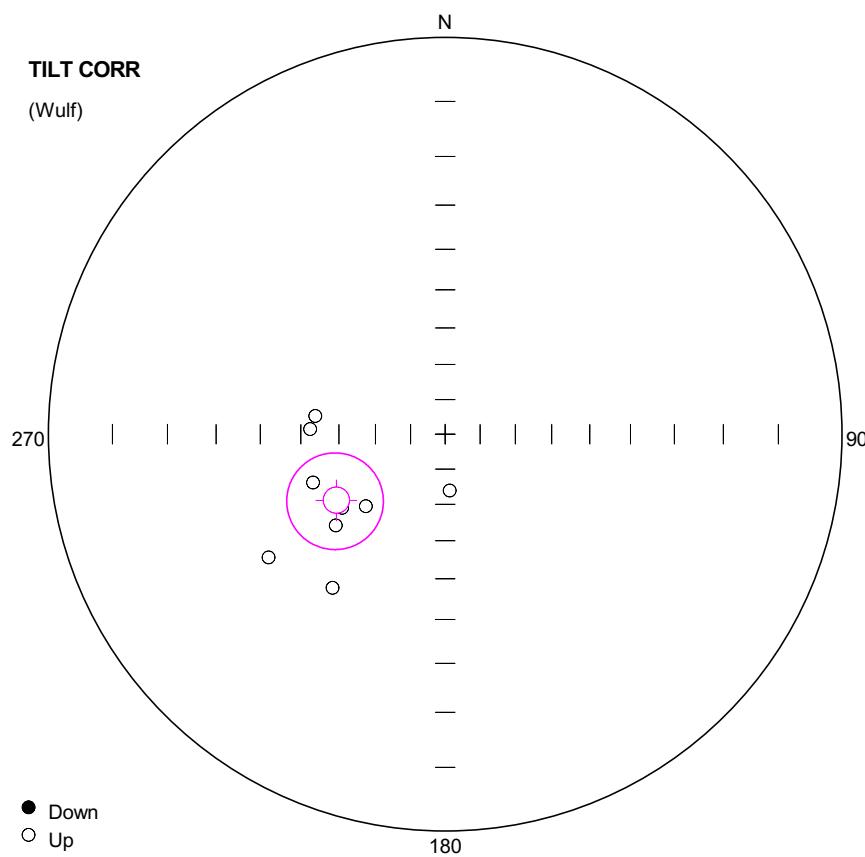
Pole longitude: 25.27 (-154.73)

Paleolatitude: -6.09 (6.09)

dp: 6.51

dm: 12.8

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9101	Cht	165.4	-5.7	6.1	56	64 OA
2	YUNN9102	Cht	164.0	-31.0	7.0	52	60 OA
3	YUNN9103	Cht	174.7	-7.2	10.9	52	60 OA
4	YUNN9104	Cht	192.3	-11.2	9.7	60	64 OA
5	YUNN9105	Cht	160.3	-33.0	2.0	56	60 OA
6	YUNN9106	Cht	172.3	-19.4	4.5	60	64 OA
7	YUNN9109	Cht	145.5	7.4	1.7	60	64 OA
8	YUNN9110	Cht	170.4	-9.5	1.4	60	64 OA
9	YUNN9111	Cht	185.6	3.7	19.2	60	64 OA



**Number of data points: 9**

**Fisher statistics**

Mean vector: 238.73 / -54.44  
Resultant vector: 8.55  
(X= -2.58, Y= -4.25, Z= -6.95)  
Estimated precision, k: 17.7  
95% Confidence limit: 12.58  
95% Confidence limit, approximation: 11.09

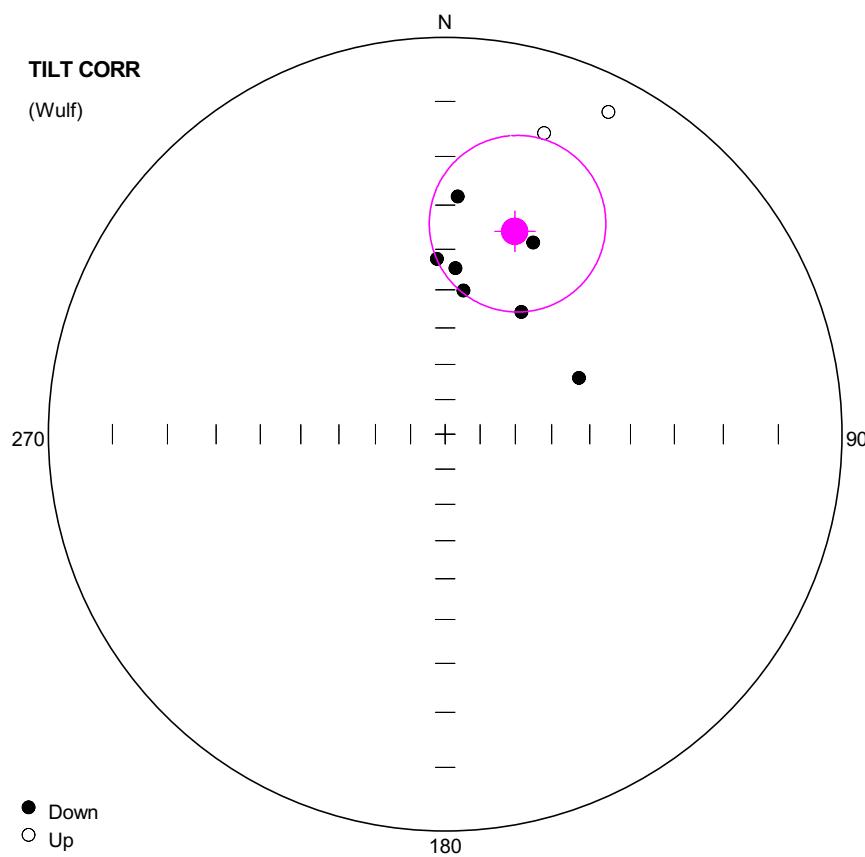
**Orientation matrix**

1st eigenvalue: 0.9  
1st eigenvector: 58.89 / 54.17  
2nd eigenvalue: 0.05  
2nd eigenvector: 166.27 / 12.17  
3rd eigenvalue: 0.05  
3rd eigenvector: 264.35 / 33.1

**VGP**

Site latitude: 65  
Site longitude: 14.04  
Pole latitude: -44.36 (44.36)  
Pole longitude: -64.39 (115.61)  
Paleolatitude: -34.96 (34.96)  
dp: 12.49  
dm: 17.73

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9101	Cht	227.8	-59.8	6.1	56	64 OA
2	YUNN9102	Cht	272.1	-52.4	7.0	52	60 OA
3	YUNN9103	Cht	230.2	-50.5	10.9	52	60 OA
4	YUNN9104	Cht	235.1	-33.0	9.7	60	64 OA
5	YUNN9105	Cht	278.0	-53.4	2.0	56	60 OA
6	YUNN9106	Cht	249.9	-51.0	4.5	60	64 OA
7	YUNN9109	Cht	175.3	-73.8	1.7	60	64 OA
8	YUNN9110	Cht	234.4	-54.6	1.4	60	64 OA
9	YUNN9111	Cht	216.2	-38.7	19.2	60	64 OA



**Number of data points: 9**

**Fisher statistics**

Mean vector: 18.94 / 33.22

Resultant vector: 8

(X= 6.33, Y= 2.17, Z= 4.38)

Estimated precision, k: 8.02

95% Confidence limit: 19.37

95% Confidence limit, approximation: 16.47

**Orientation matrix**

1st eigenvalue: 0.8

1st eigenvector: 17.87 / 35.35

2nd eigenvalue: 0.15

2nd eigenvector: 192.61 / 54.53

3rd eigenvalue: 0.05

3rd eigenvector: 286.11 / 2.49

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 41.45 (-41.45)

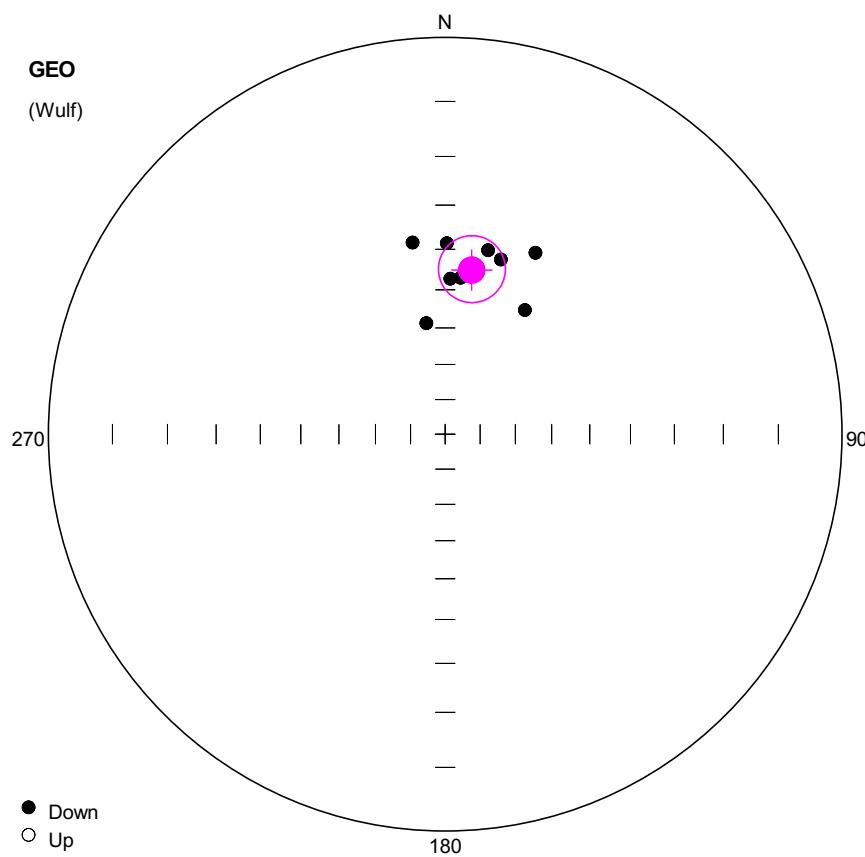
Pole longitude: 169.74 (-10.26)

Paleolatitude: 18.13 (-18.13)

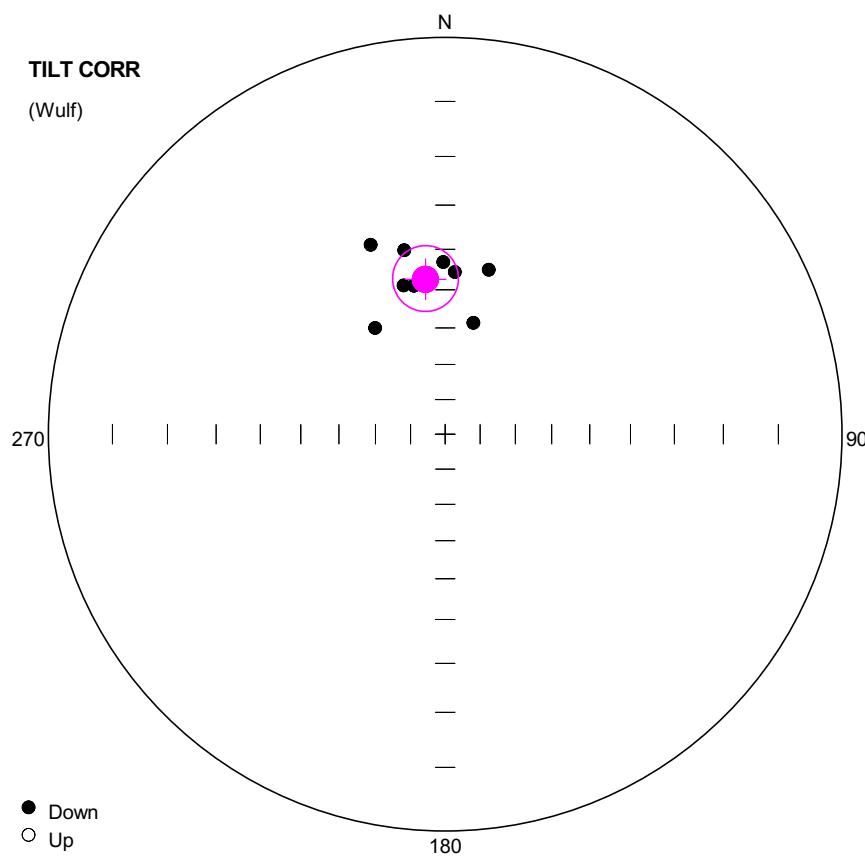
dp: 12.5

dm: 22

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9101	Cmt	24.7	34.0	9.9	30	52
2	YUNN9102	Cmt	18.2	-12.8	4.1	30	48
3	YUNN9103	Cmt	7.3	49.9	7.8	44	56
4	YUNN9104	Cmt	31.9	50.1	6.8	40	60
5	YUNN9105	Cmt	26.9	-5.4	5.6	35	52
6	YUNN9106	Cmt	67.2	49.8	11.6	20	60
7	YUNN9109	Cmt	3.0	28.1	9.2	35	60
8	YUNN9110	Cmt	3.6	44.5	3.2	35	64
9	YUNN9111	Cmt	357.3	42.3	7.9	40	56



#	Name	State	Dec	Inc	MAD	Limit1	Limit2	
1	YUNN9201	Cmt	17.7	40.4	12.6	20	64	
2	YUNN9202	Cmt	13.1	39.1	10.0	20	64	
3	YUNN9204	Cmt	5.6	46.8	5.6	20	64	
4	YUNN9205	Cmt	350.4	58.3	5.8	20	64	
5	YUNN9206	Cmt	26.5	35.9	6.0	20	64	
6	YUNN9207	Cmt	0.5	38.6	7.4	30	60	
7	YUNN9208	Ch	350.4	37.8	3.7	30	68	
8	YUNN9209	Ch		1.9	47.2	7.0	30	68
9	YUNN9210	Cmt		32.7	49.2	19.1	30	60



**Number of data points: 9**

**Fisher statistics**

Mean vector: 352.77 / 47.06

Resultant vector: 8.8

(X= 5.95, Y= -0.75, Z= 6.44)

Estimated precision, k: 40.33

95% Confidence limit: 8.21

95% Confidence limit, approximation: 7.35

**Orientation matrix**

1st eigenvalue: 0.96

1st eigenvector: 352.79 / 47.01

2nd eigenvalue: 0.03

2nd eigenvector: 90.9 / 7.5

3rd eigenvalue: 0.01

3rd eigenvector: 187.71 / 42.01

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 52.97 (-52.97)

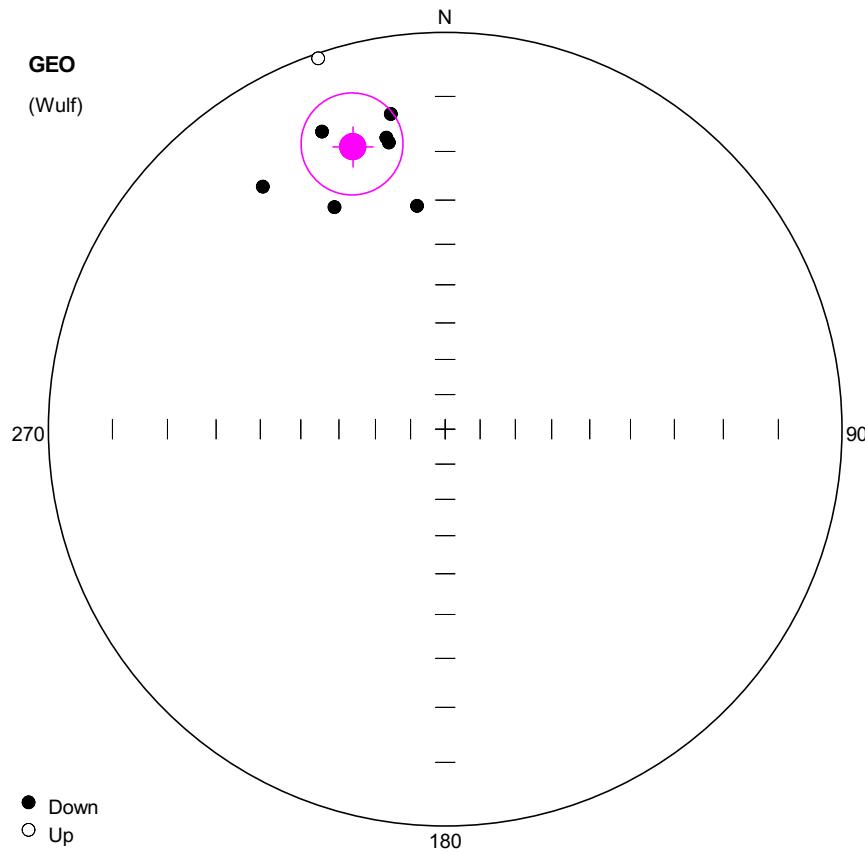
Pole longitude: -155.36 (24.64)

Paleolatitude: 28.25 (-28.25)

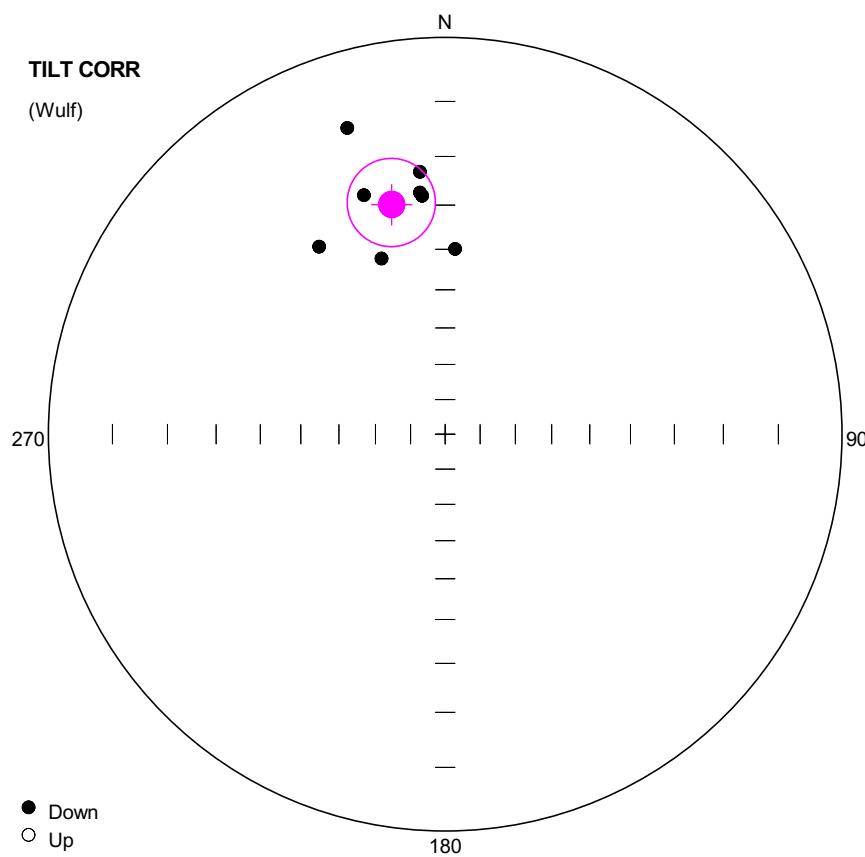
dp: 6.86

dm: 10.61

#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9201	Cmt	3.4	45.5	12.6	20	64
2	YUNN9202	Cmt	359.4	43.1	10.0	20	64
3	YUNN9204	Cmt	348.1	48.2	5.6	20	64
4	YUNN9205	Cmt	326.6	54.5	5.8	20	64
5	YUNN9206	Cmt	14.8	43.6	6.0	20	64
6	YUNN9207	Cmt	347.5	39.2	7.4	30	60
7	YUNN9208	Ch	338.5	35.7	3.7	30	68
8	YUNN9209	Ch	344.4	47.5	7.0	30	68
9	YUNN9210	Cmt	14.2	57.7	19.1	30	60



#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9301	Cmt	337.5	11.9	12.5	30	60
2	YUNN9302	Cmt	352.8	30.9	8.3	20	64
3	YUNN9304	Ch	323.1	15.2	10.3	30	64
4	YUNN9305	Ch	333.5	26.0	6.0	30	64
5	YUNN9306	Ch	341.1	-0.7	5.1	30	60
6	YUNN9308	Ch	350.2	12.3	8.7	30	60
7	YUNN9309	Cmt	348.6	16.3	8.2	44	68
8	YUNN9310	Cmt	348.9	17.3	9.0	30	64



**Number of data points: 8**

**Fisher statistics**

Mean vector: 346.88 / 28.56

Resultant vector: 7.81

(X= 6.68, Y= -1.56, Z= 3.73)

Estimated precision, k: 35.92

95% Confidence limit: 9.37

95% Confidence limit, approximation: 8.26

**Orientation matrix**

1st eigenvalue: 0.95

1st eigenvector: 346.96 / 28.53

2nd eigenvalue: 0.03

2nd eigenvector: 92.66 / 26.47

3rd eigenvalue: 0.02

3rd eigenvector: 217.91 / 49.21

**VGP**

Site latitude: 65

Site longitude: 14.04

Pole latitude: 39.43 (-39.43)

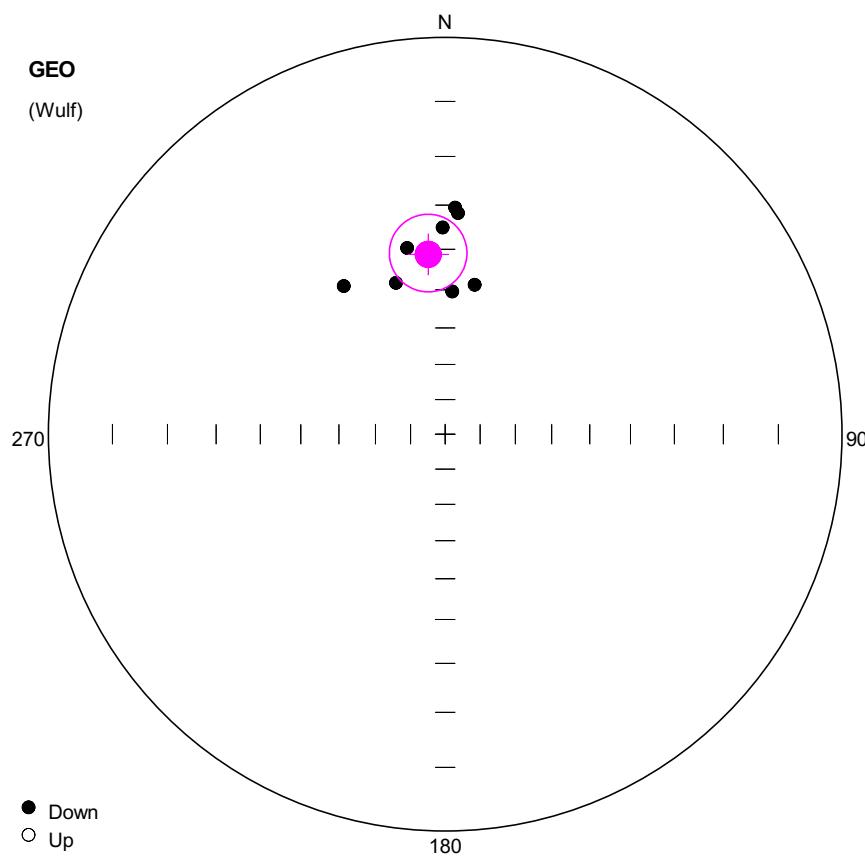
Pole longitude: -149.49 (30.51)

Paleolatitude: 15.22 (-15.22)

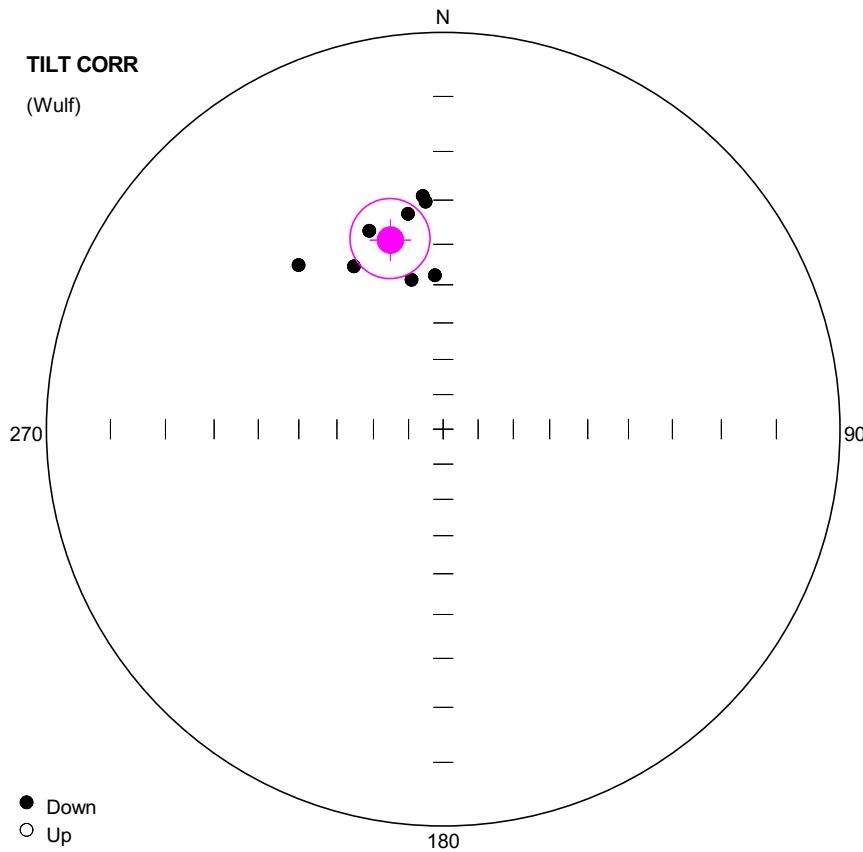
dp: 5.65

dm: 10.29

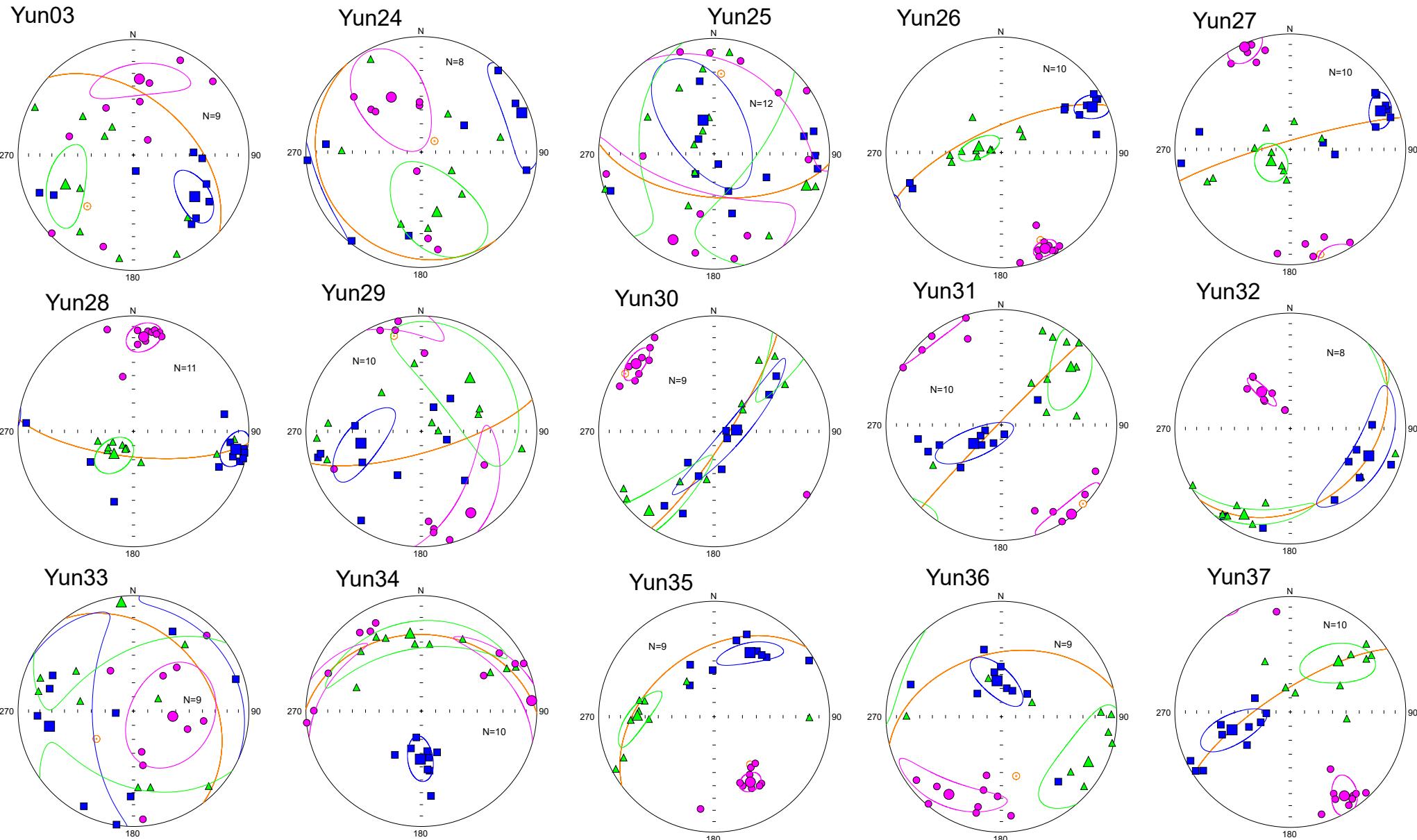
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9301	Cmt	341.2	25.1	12.5	30	60
2	YUNN9302	Cmt	3.1	39.9	8.3	20	64
3	YUNN9304	Ch	326.1	30.7	10.3	30	64
4	YUNN9305	Ch	340.1	39.6	6.0	30	64
5	YUNN9306	Ch	342.3	12.0	5.1	30	60
6	YUNN9308	Ch	354.5	22.8	8.7	30	60
7	YUNN9309	Cmt	354.0	27.1	8.2	44	68
8	YUNN9310	Cmt	354.5	27.8	9.0	30	64

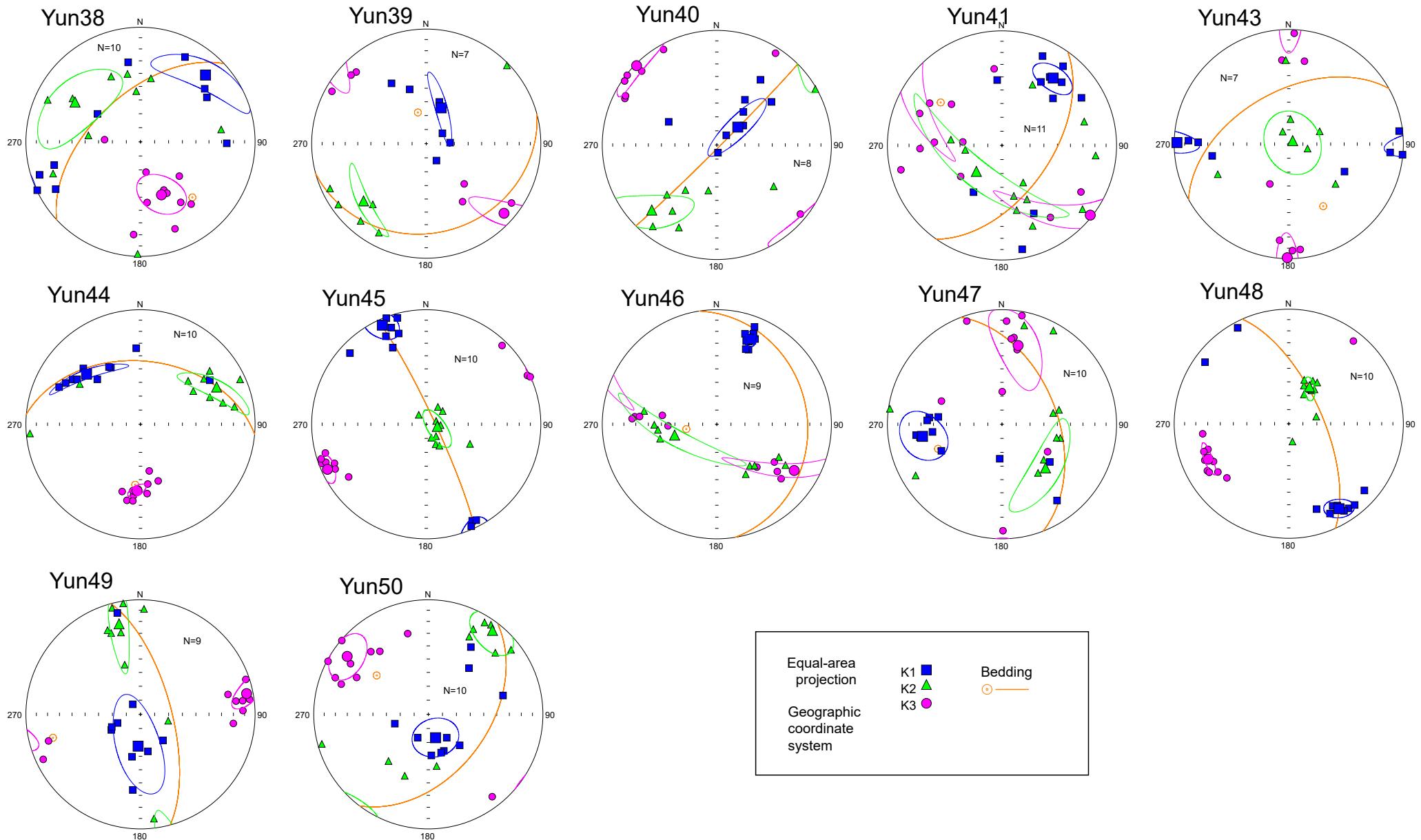


#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9401	Ch	325.6	41.3	5.9	30	64
2	YUNN9402	Ch	359.3	35.0	19.5	30	64
3	YUNN9403	Ch	342.0	46.3	3.5	20	60 OA
4	YUNN9404	Ch	348.5	38.8	5.7	20	64 OA
5	YUNN9405	Ch	2.9	50.4	14.7	20	64
6	YUNN9407	Ch	2.5	30.5	16.3	20	60
7	YUNN9408	Ch	3.3	31.7	11.4	20	68 OA
8	YUNN9409	Ch	11.2	48.0	9.0	20	52



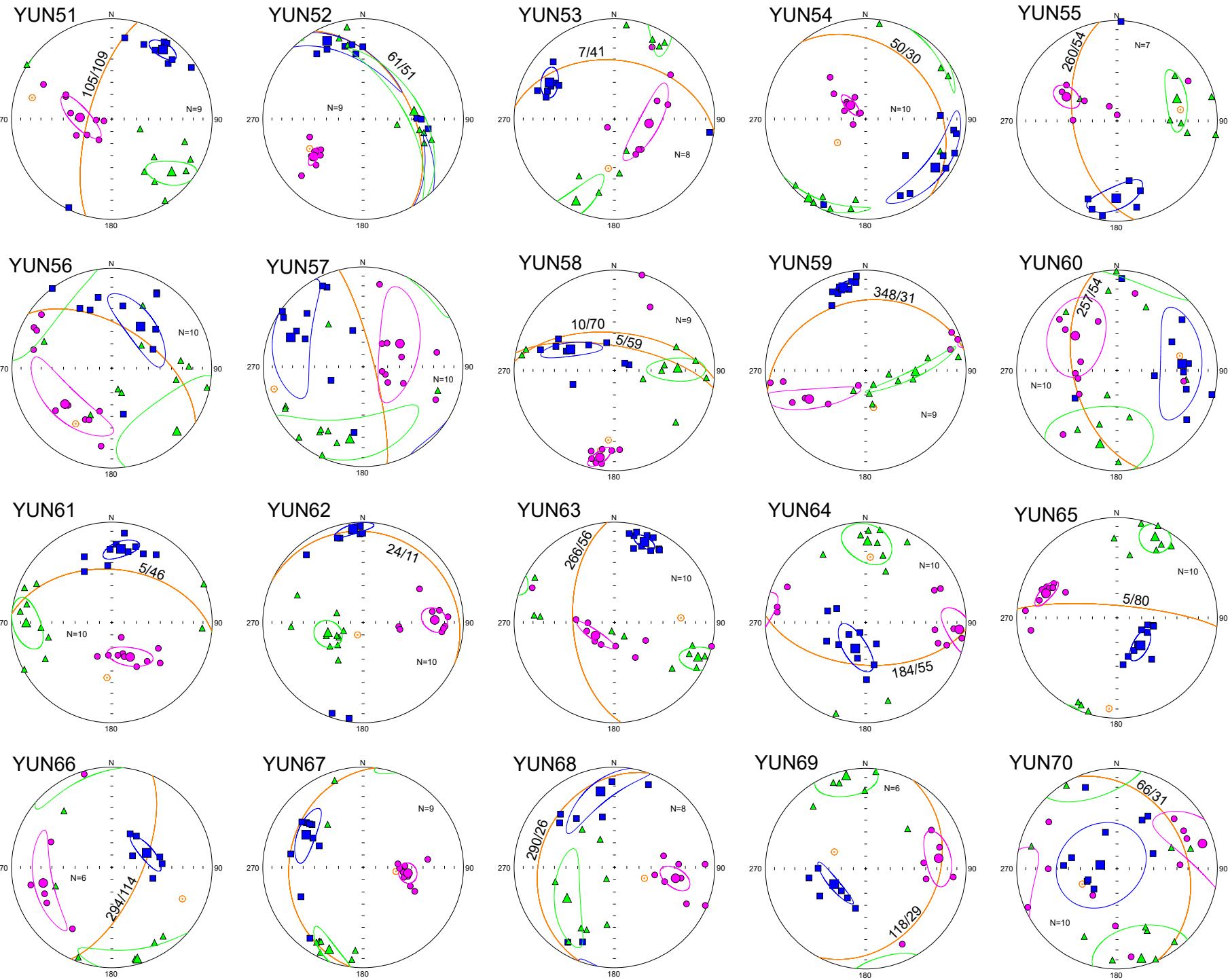
#	Name	State	Dec	Inc	MAD	Limit1	Limit2
1	YUNN9401	Ch	318.6	32.3	5.9	30	64
2	YUNN9402	Ch	350.7	32.4	19.5	30	64
3	YUNN9403	Ch	331.3	39.9	3.5	20	60 OA
4	YUNN9404	Ch	339.6	33.9	5.7	20	64 OA
5	YUNN9405	Ch	348.0	47.9	14.7	20	64
6	YUNN9407	Ch	355.0	28.9	16.3	20	60
7	YUNN9408	Ch	355.6	30.2	11.4	20	68 OA
8	YUNN9409	Ch	356.9	47.6	9.0	20	52

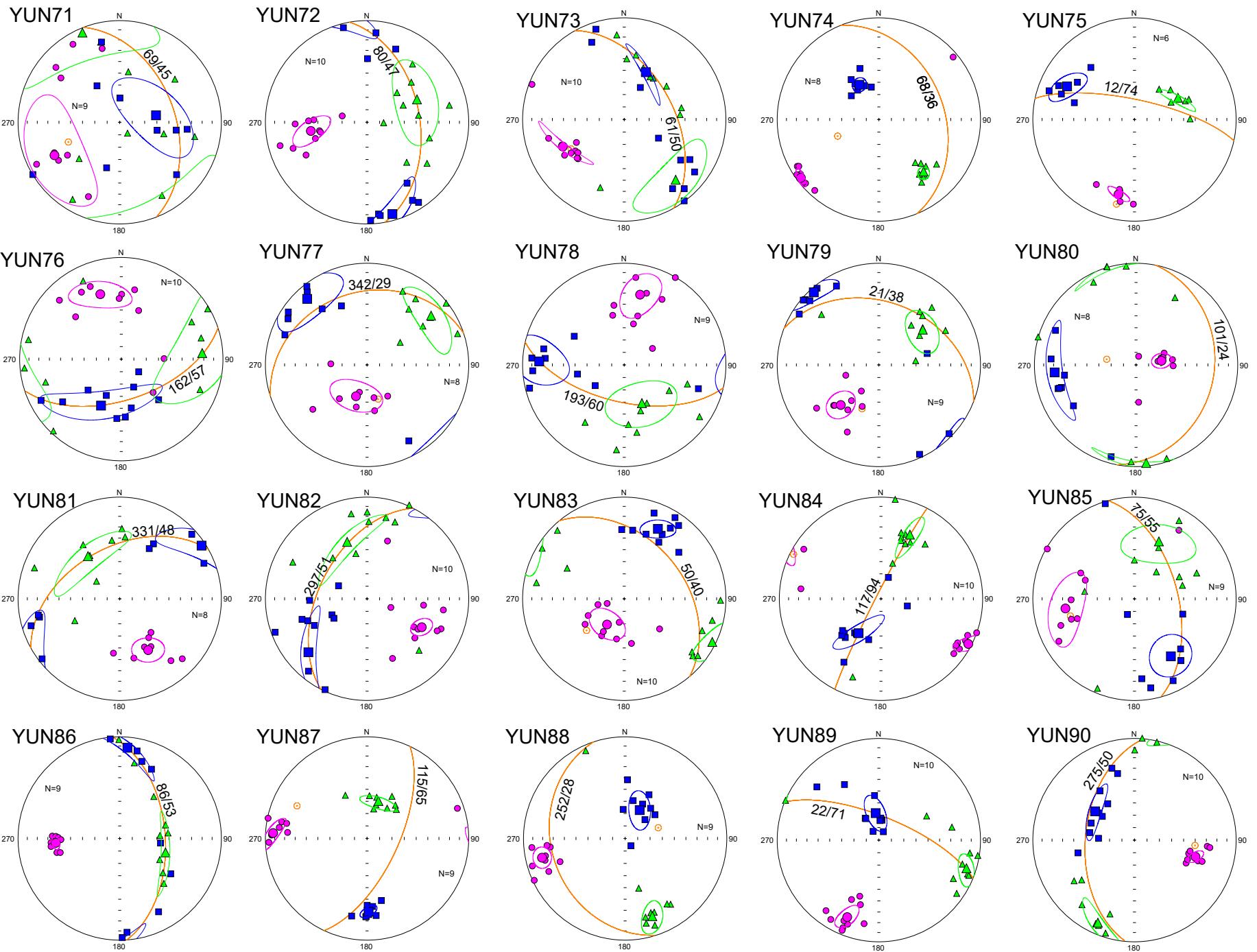


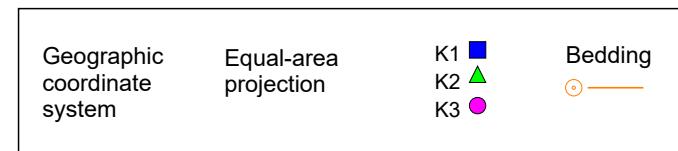
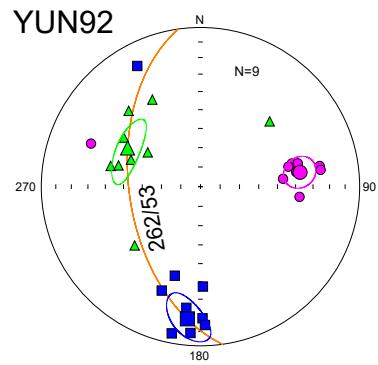
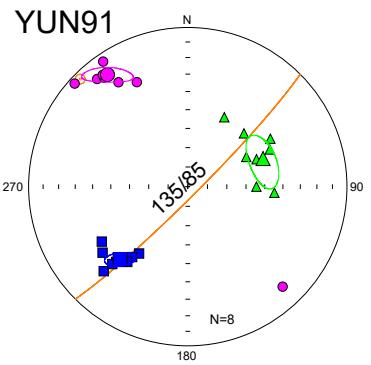


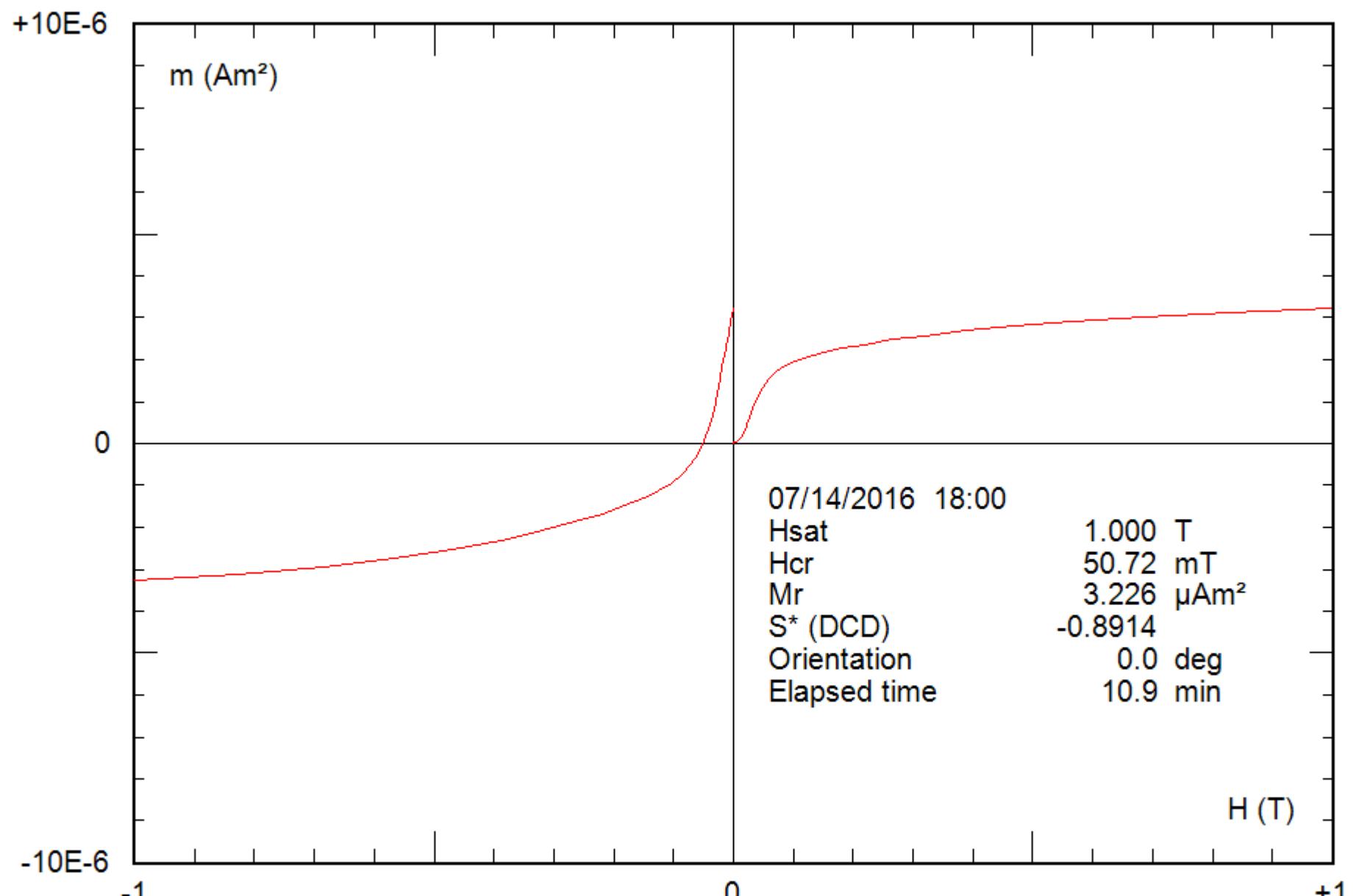
Equal-area  
projection  
Geographic  
coordinate  
system

K1    K2    K3    Bedding



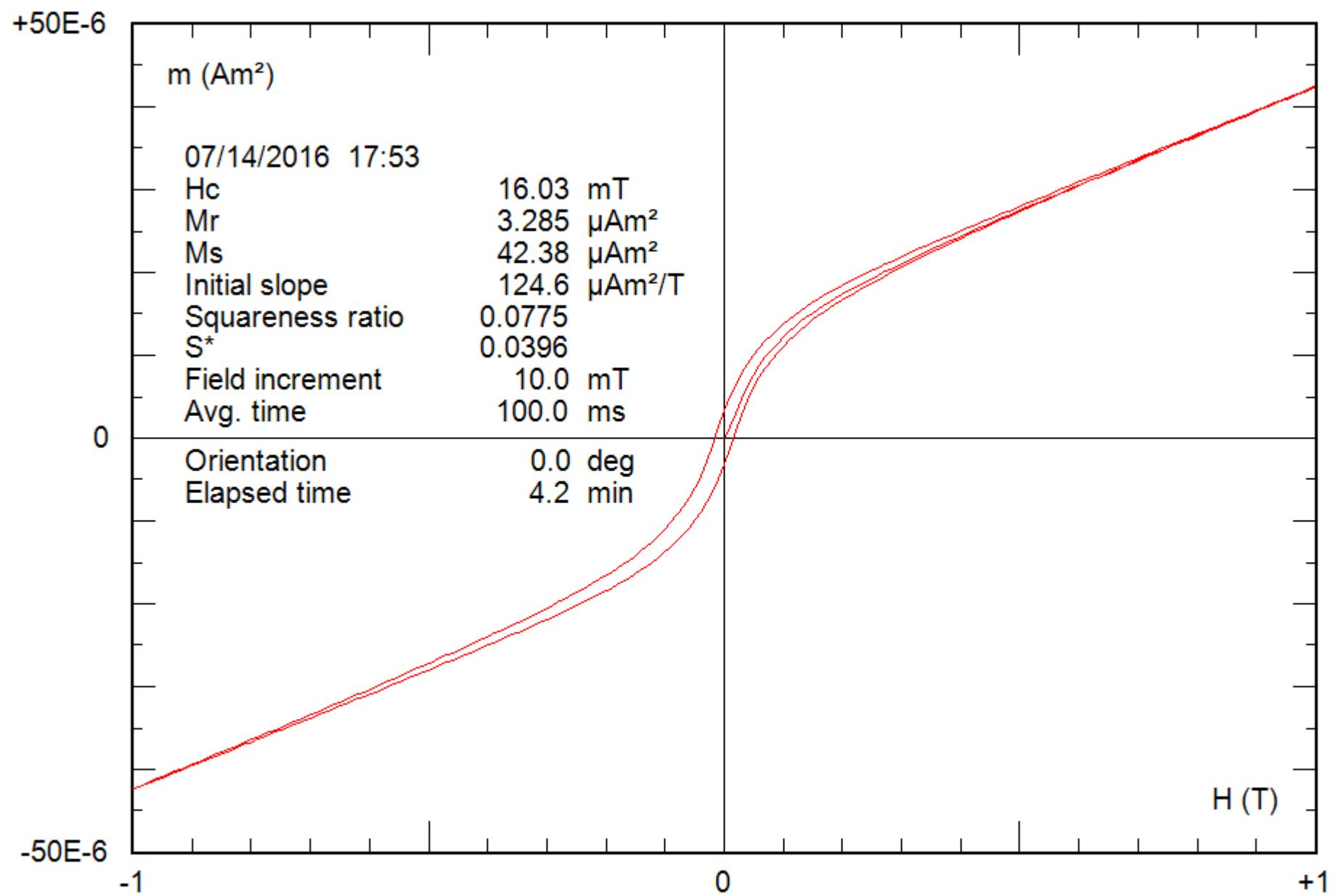






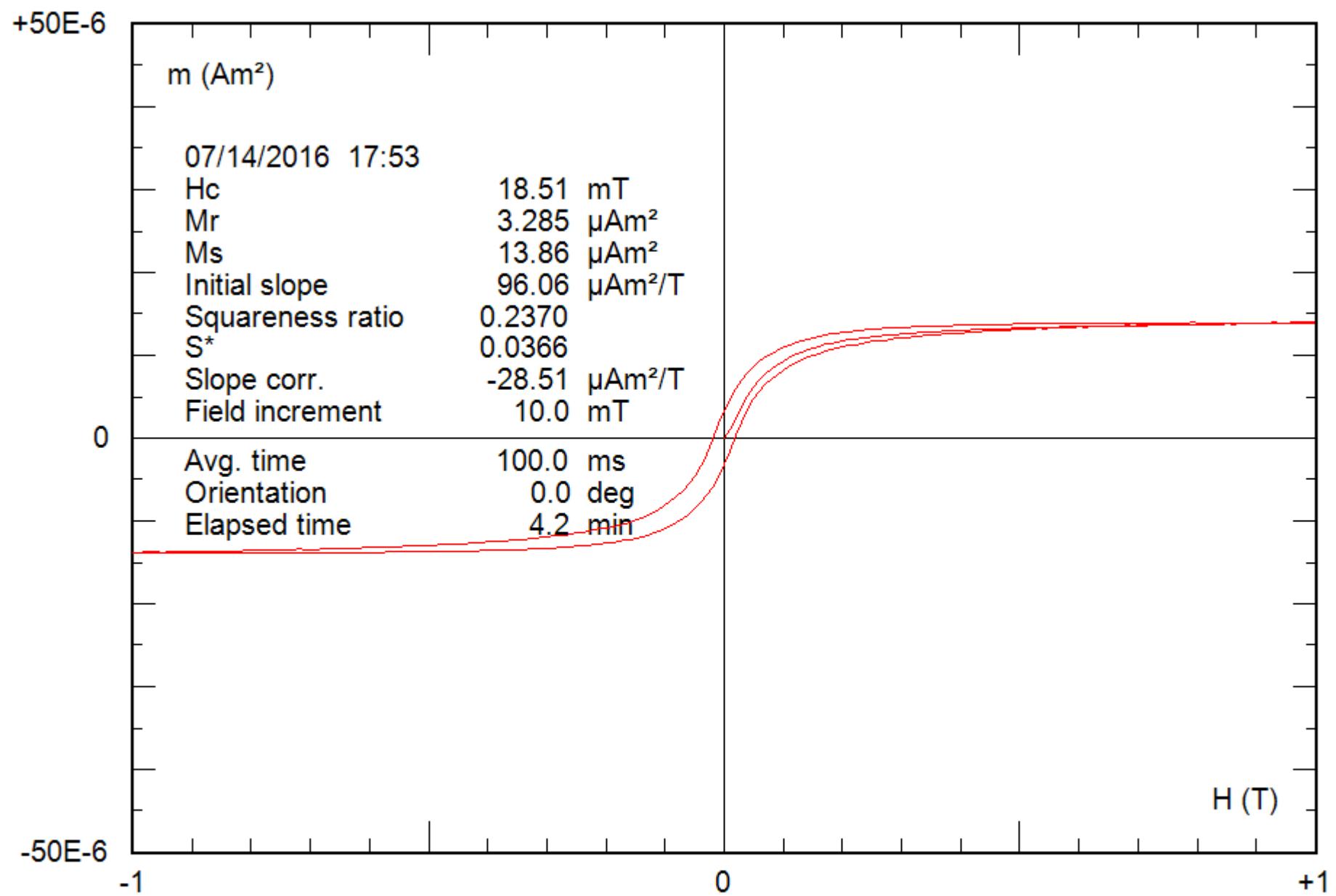
0.27 g

File: 01-03b\_bcr



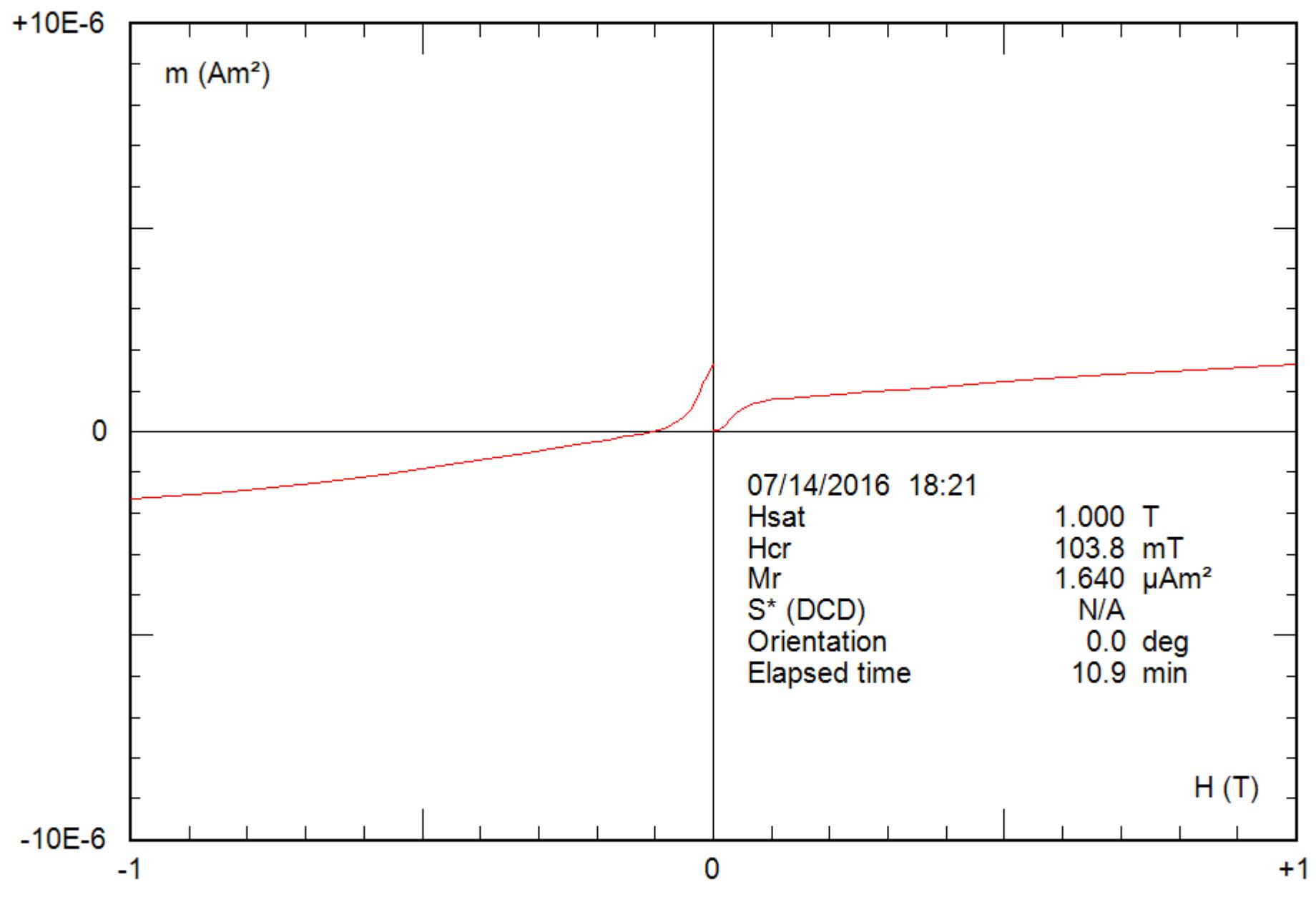
0.27 g

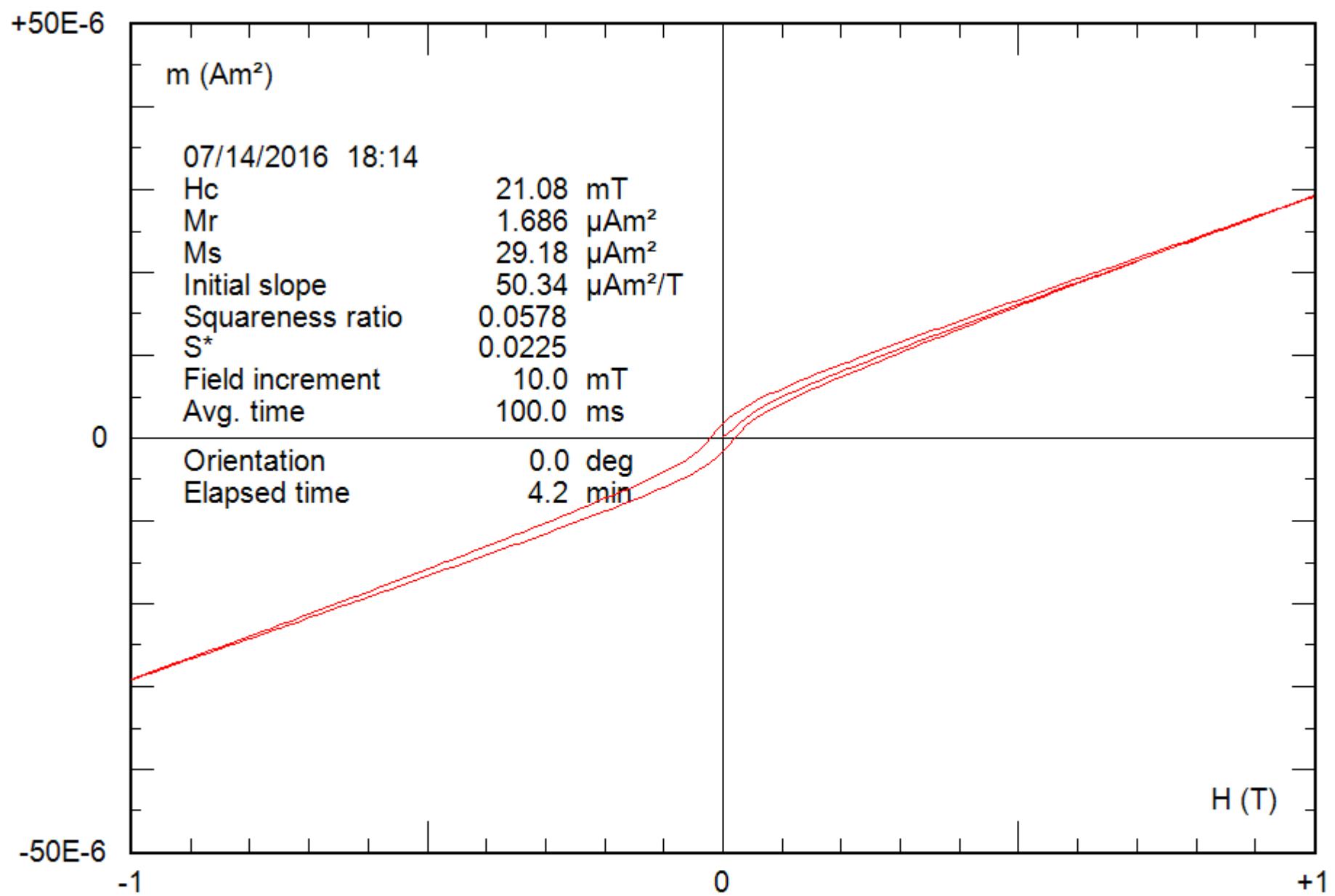
File: 01-03b\_hyst

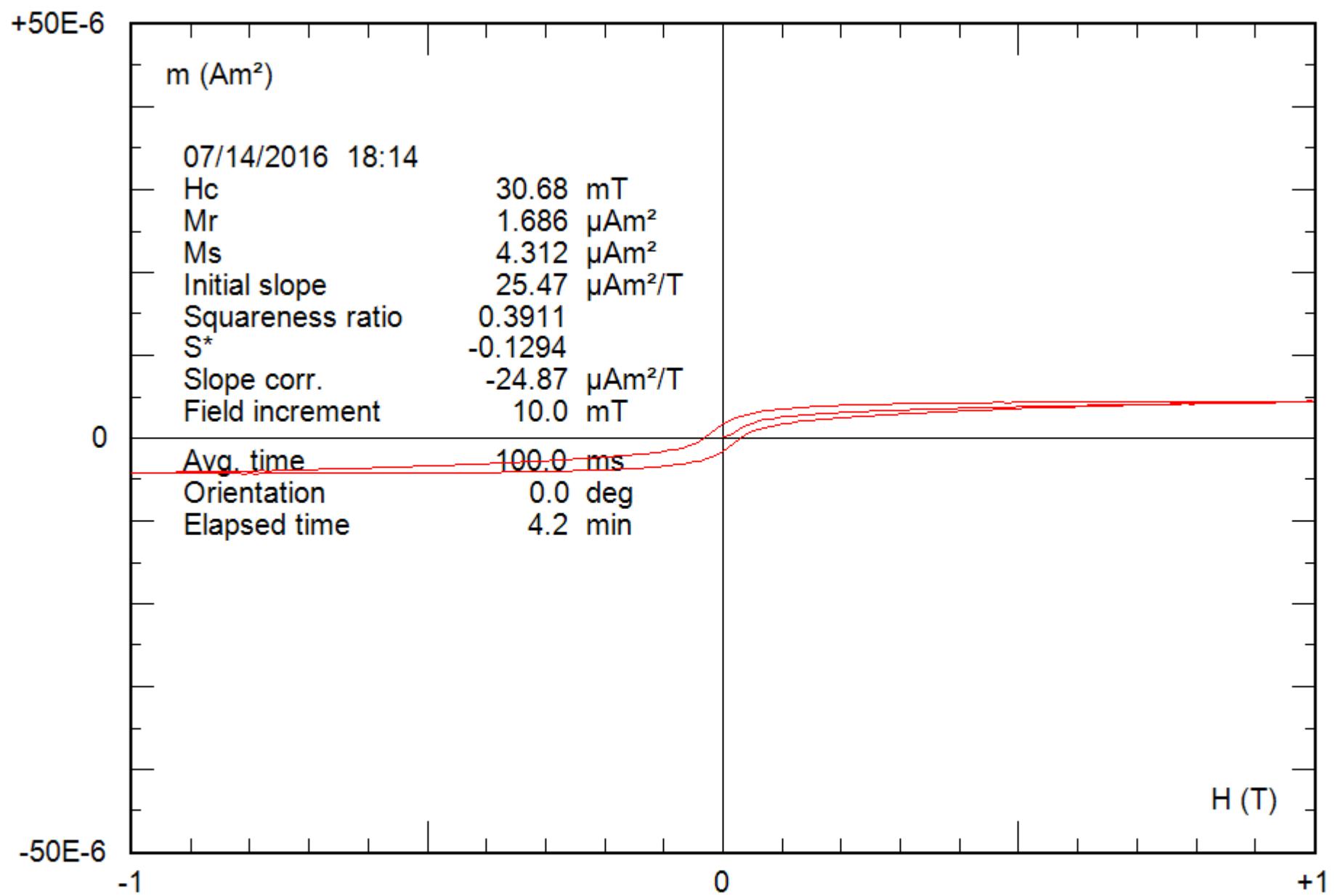


0.27 g

File: 01-03b\_hyst\_corr

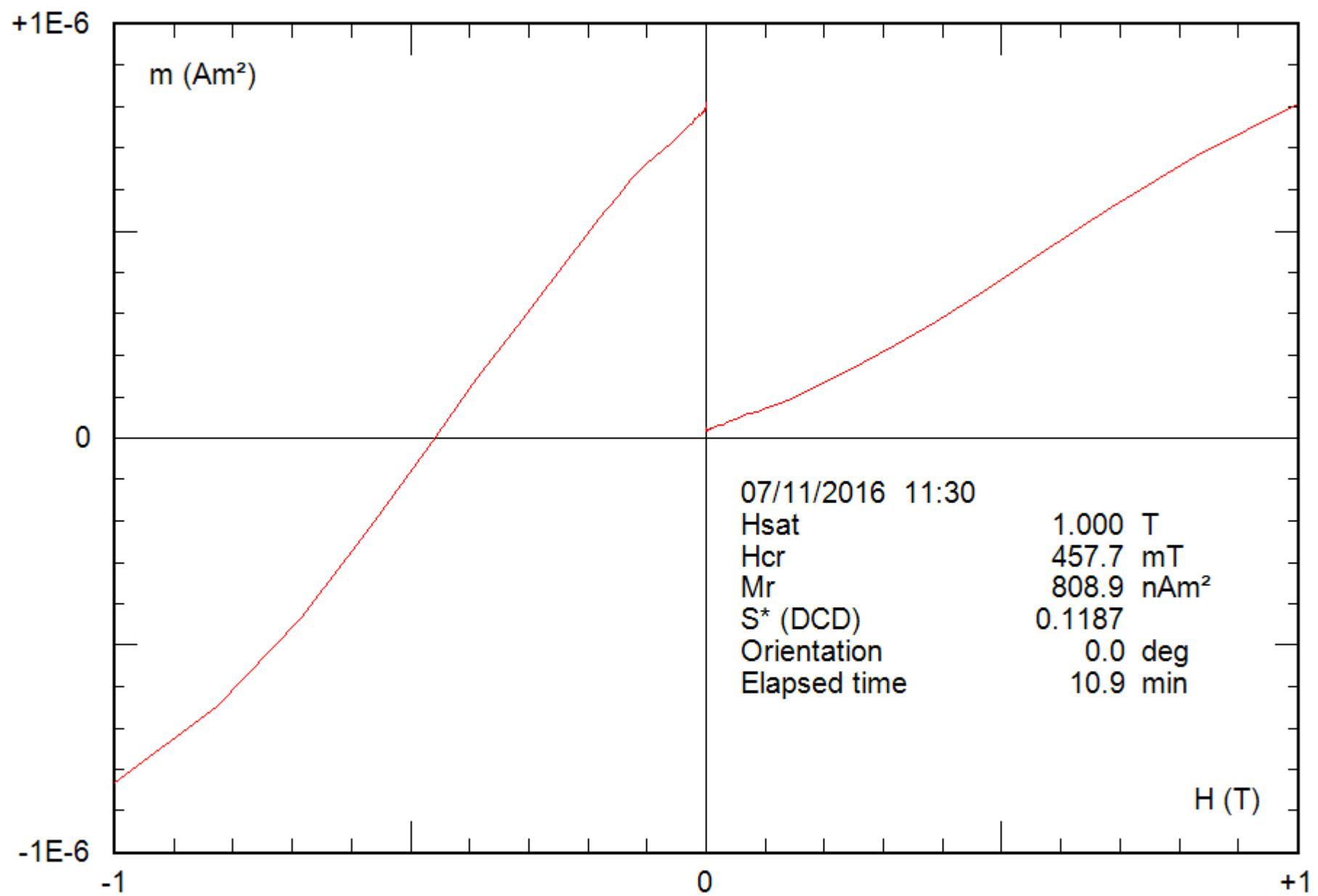






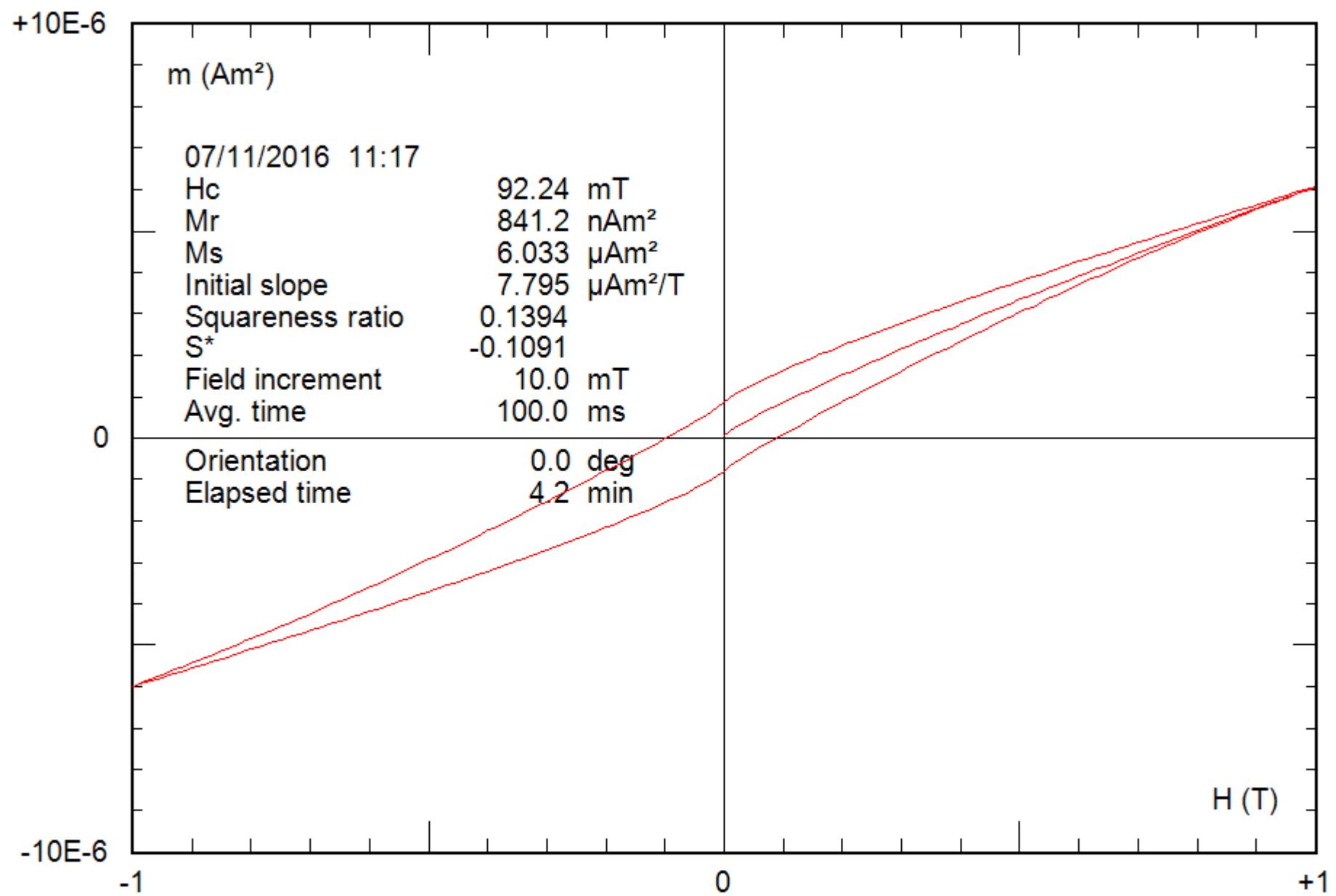
0.29 g

File: 02-08b\_hyst\_corr



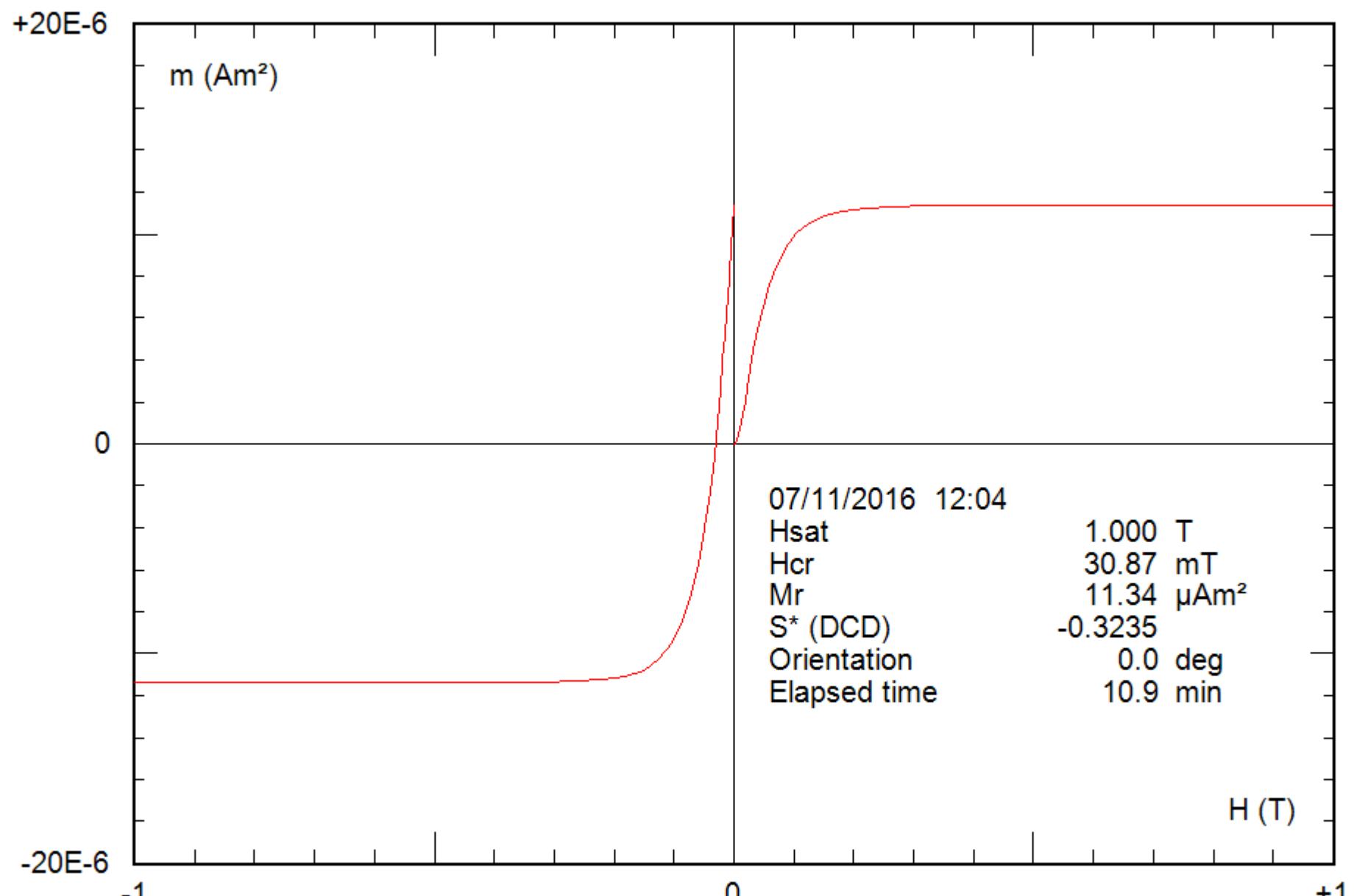
0.24 g

File: 03-10b\_bcr



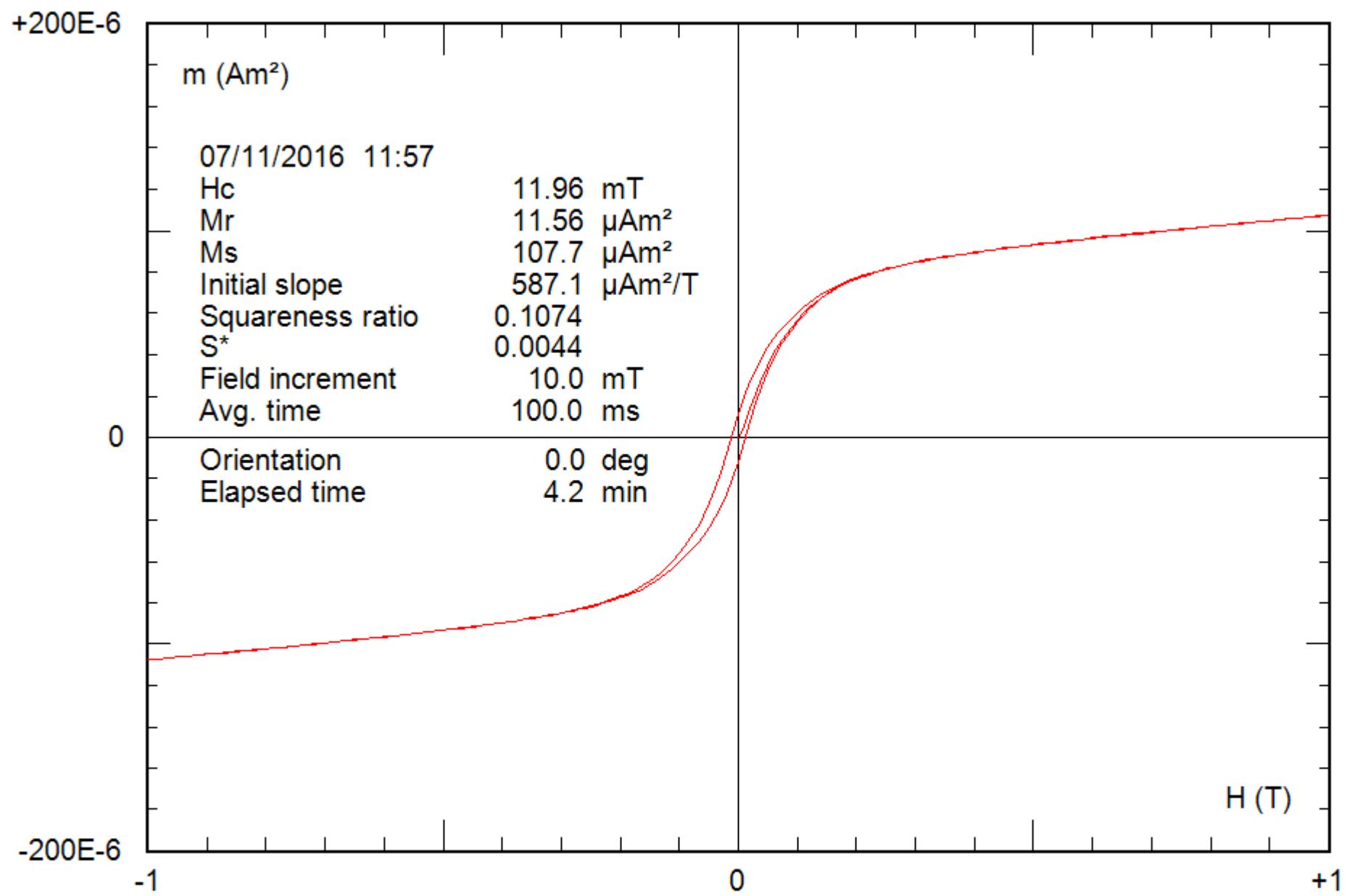
0.24 g

File: 03-10b\_hyst



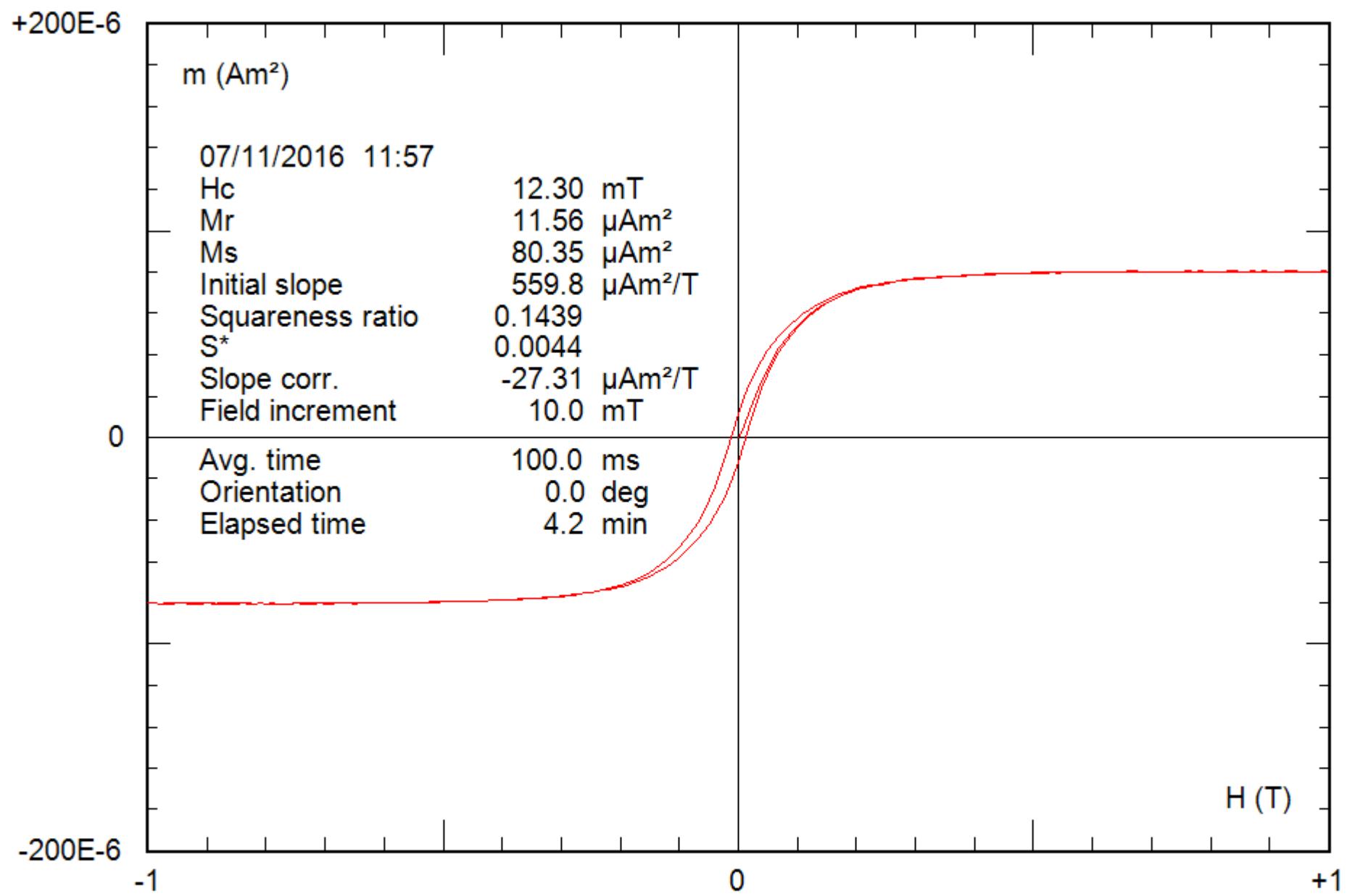
0.16 g

File: 04-4a\_bcr



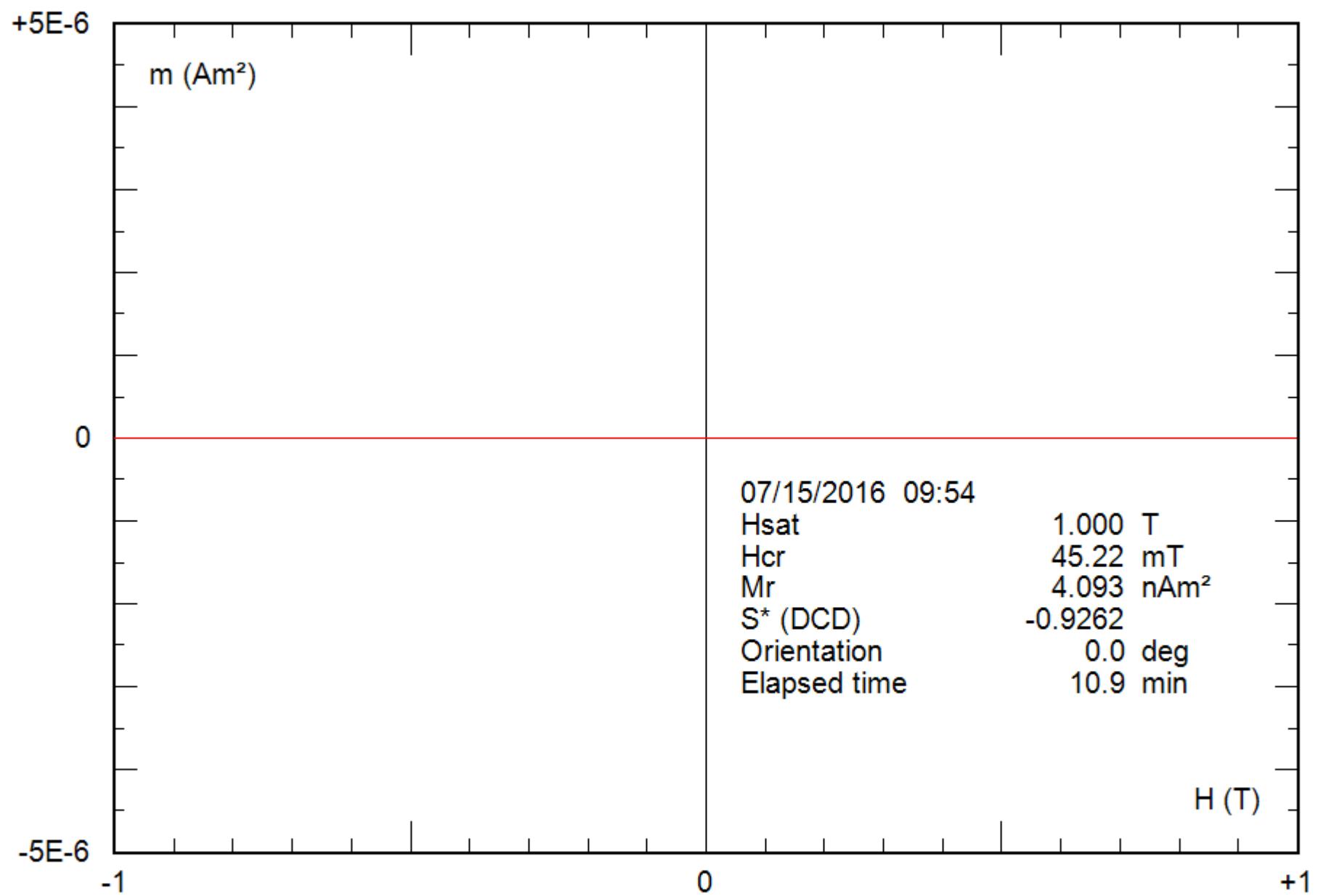
0.16 g

File: 04-4a\_hyst



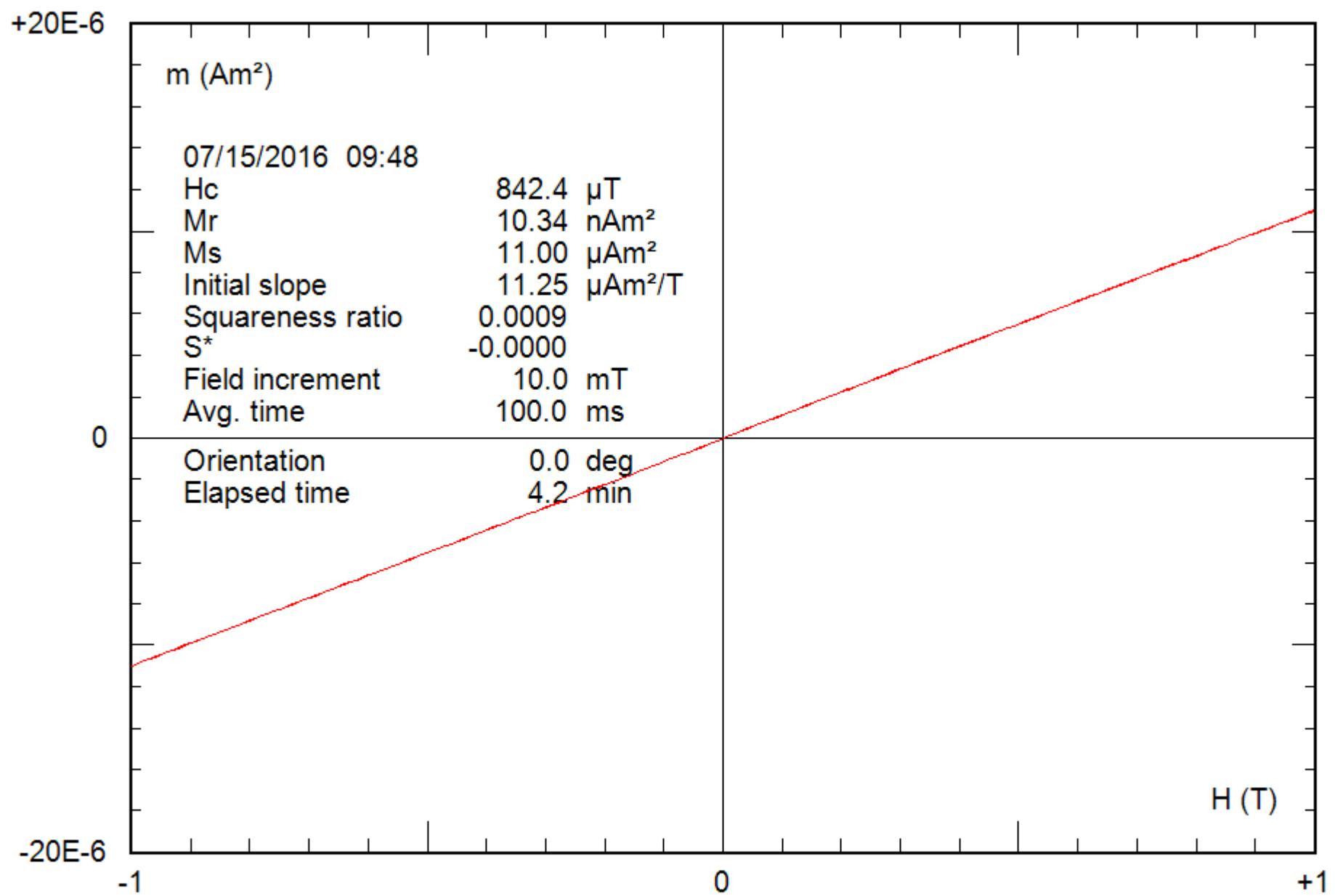
0.16 g

File: 04-4a\_hyst\_corr



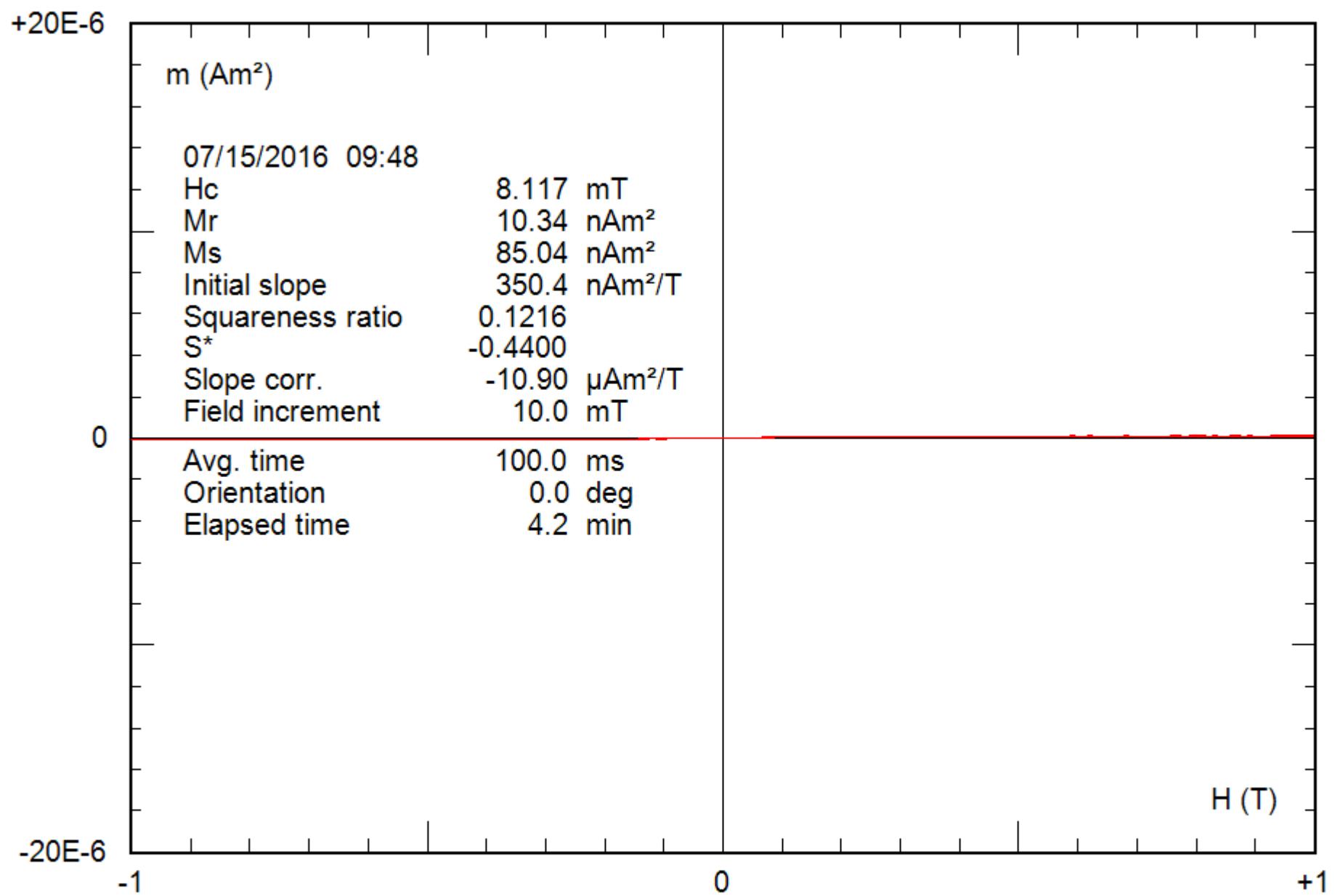
0.21 g

File: 05-01a\_bcr



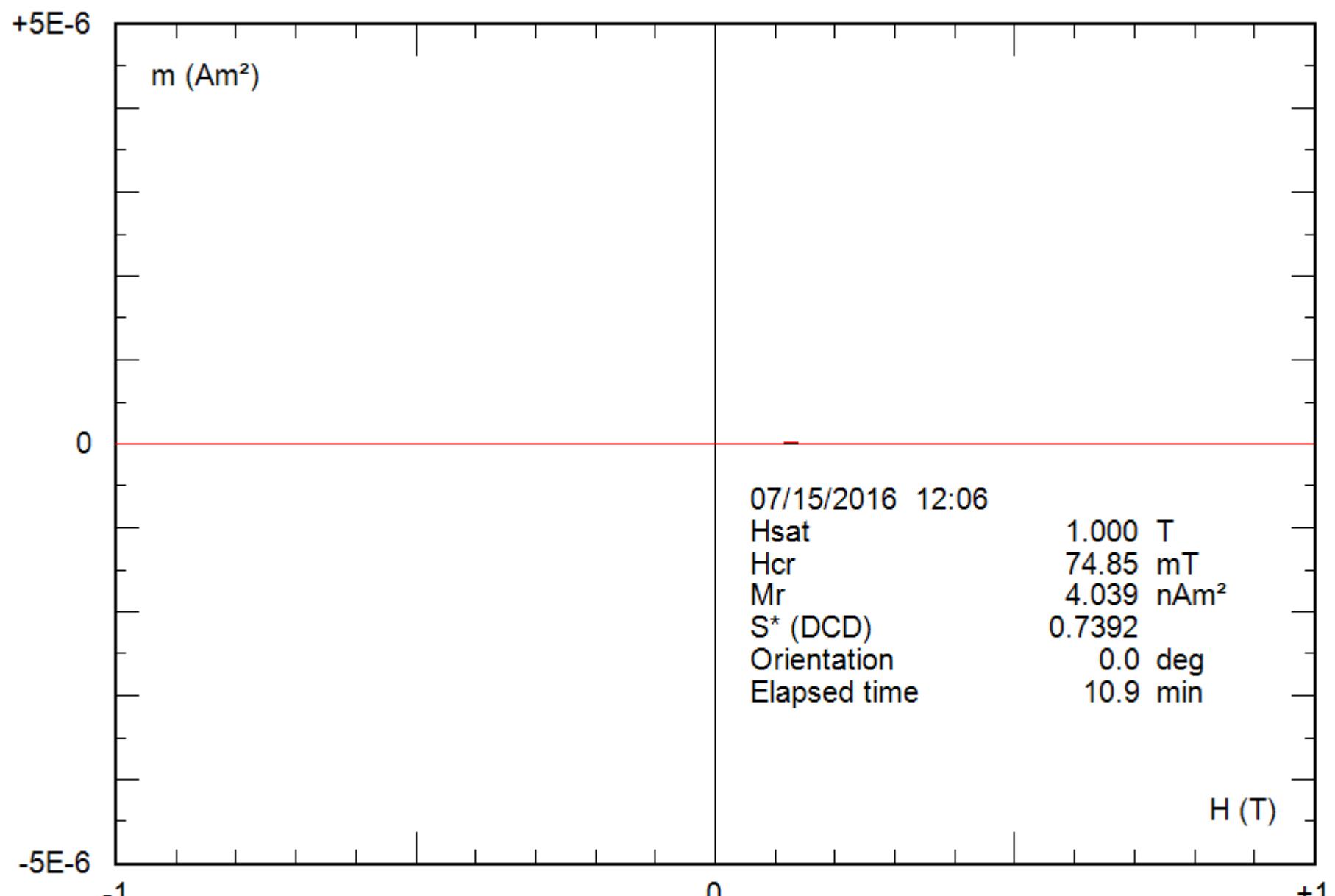
0.21 g

File: 05-01a\_hyst



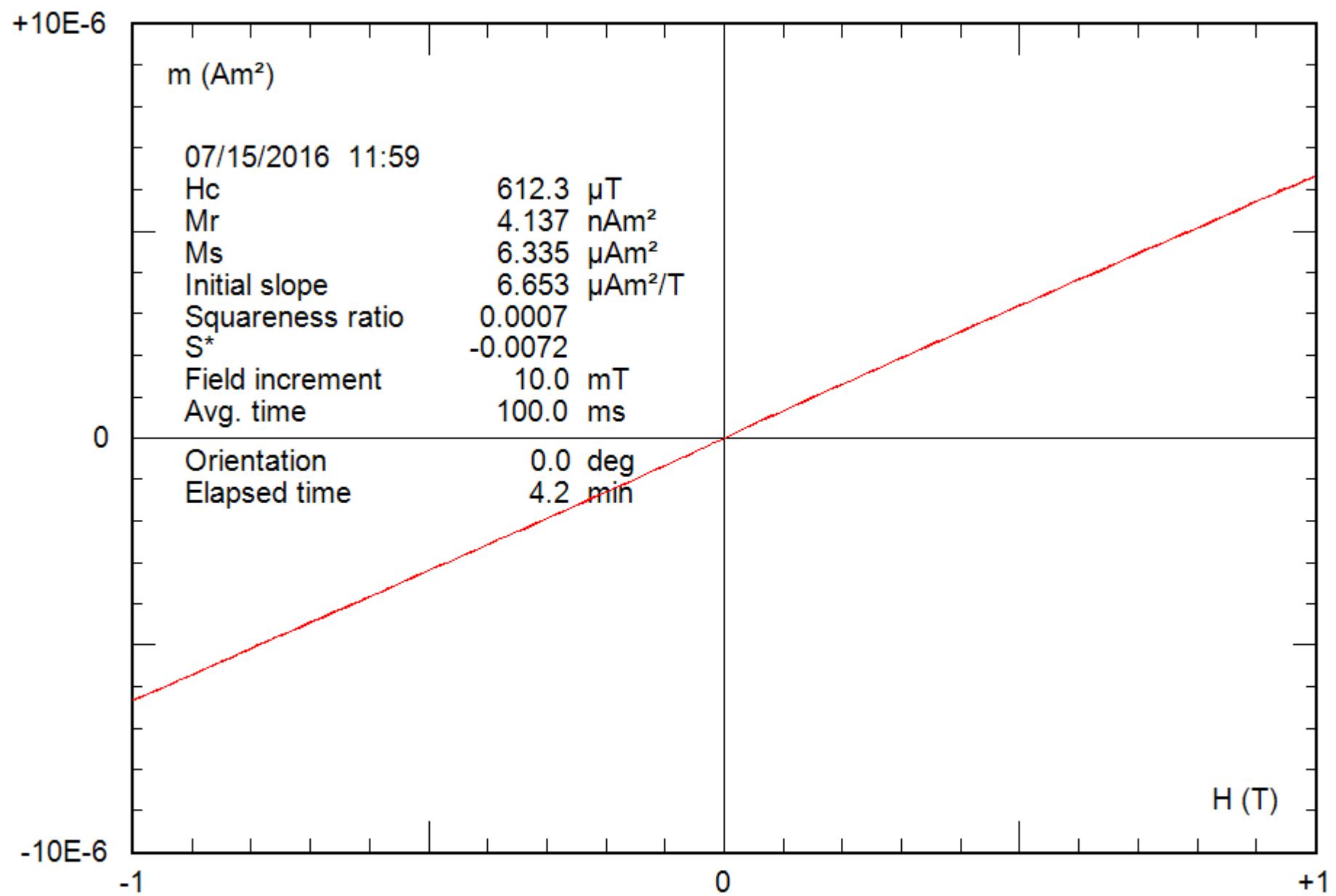
0.21 g

File: 05-01a\_hyst\_corr



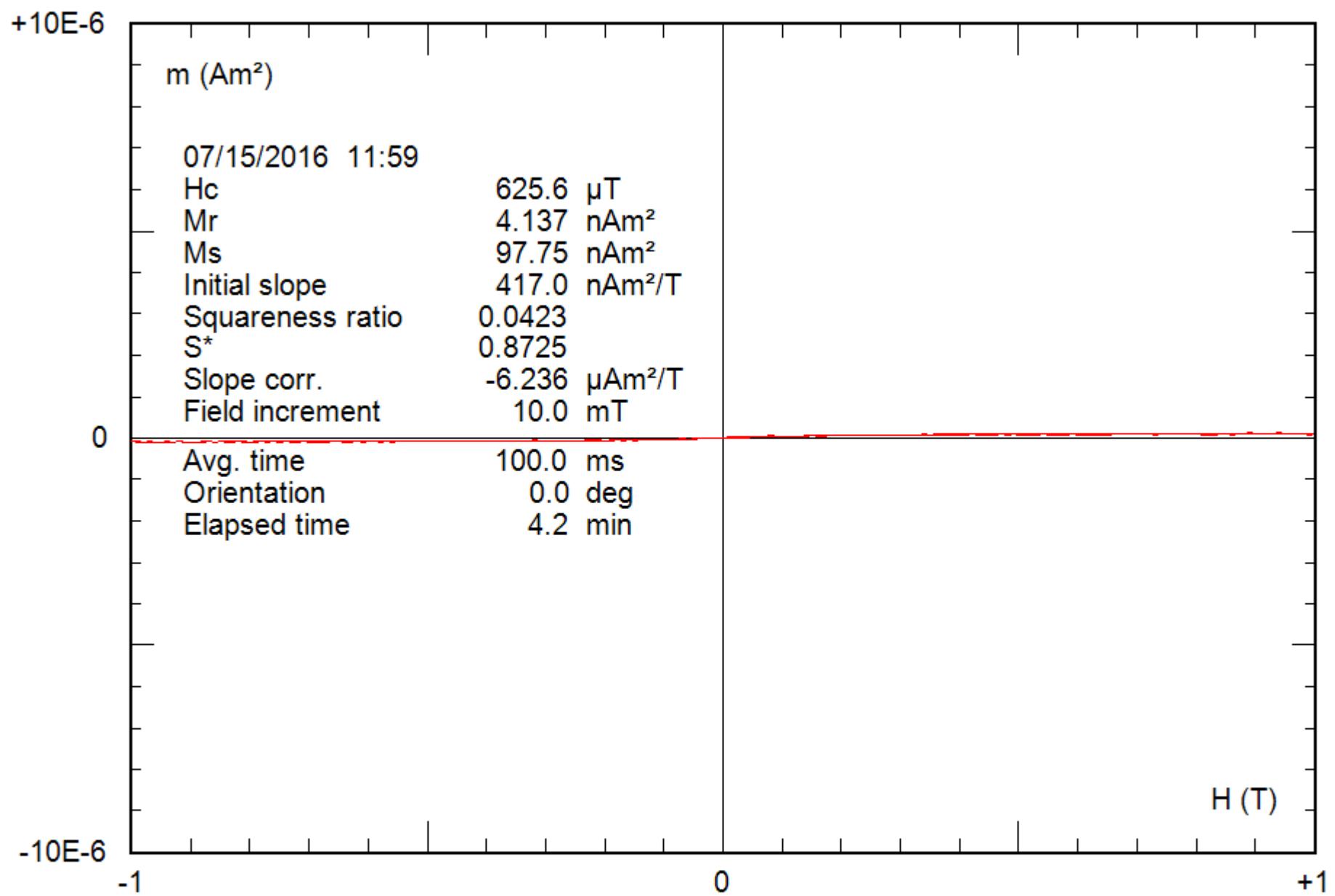
0.19 g

File: 06-07a\_bcr



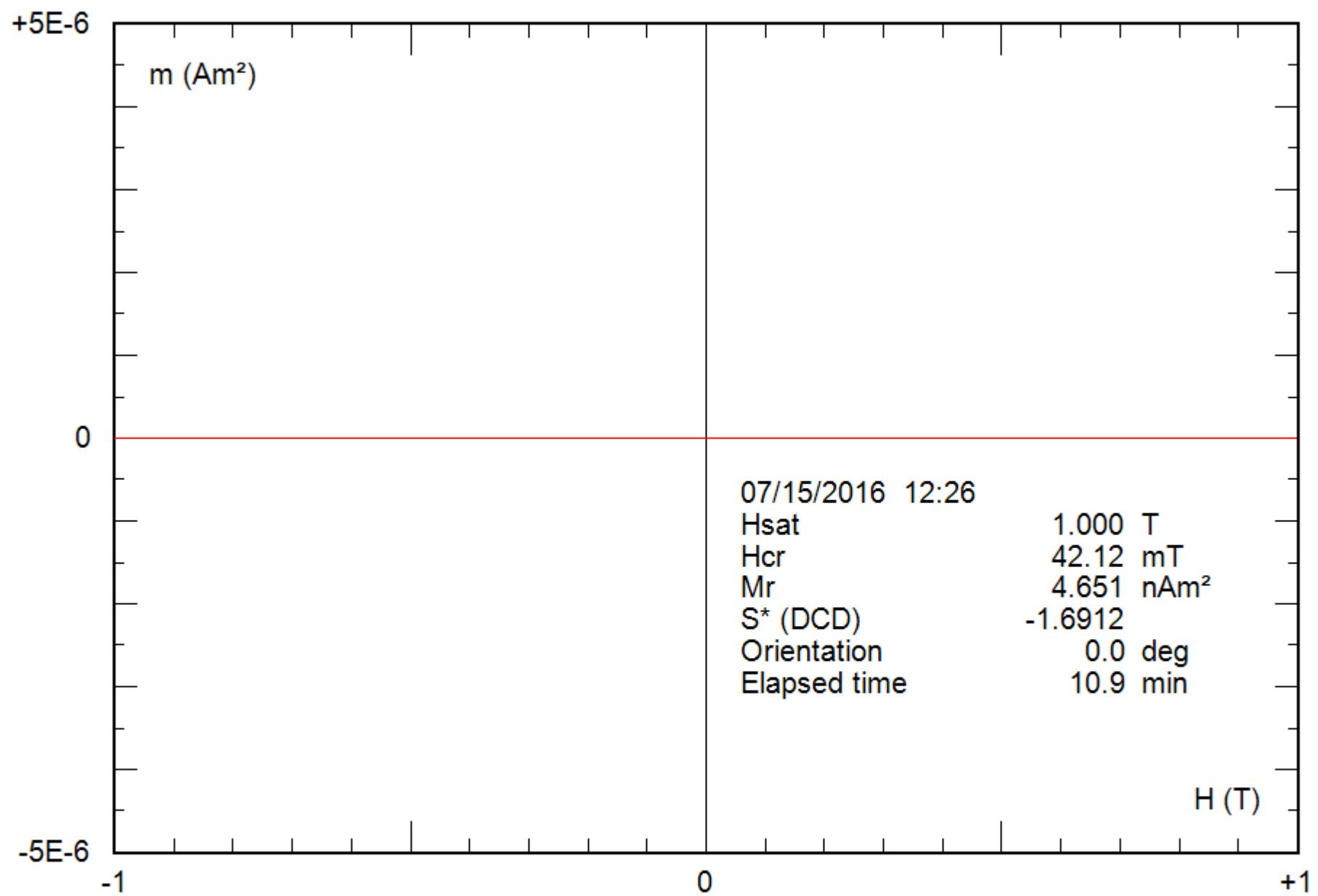
0.19 g

File: 06-07a\_hyst



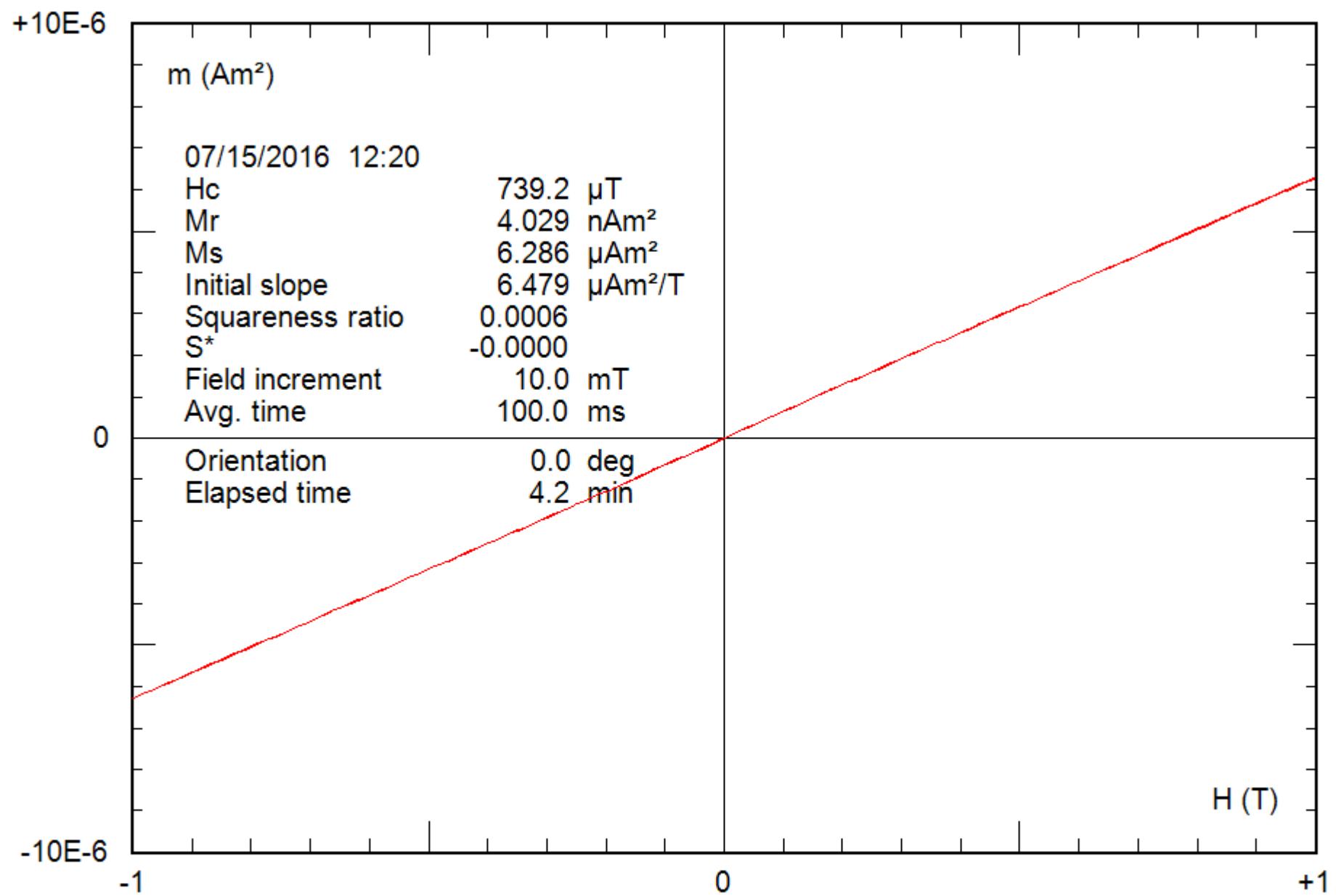
0.19 g

File: 06-07a\_hyst\_corr



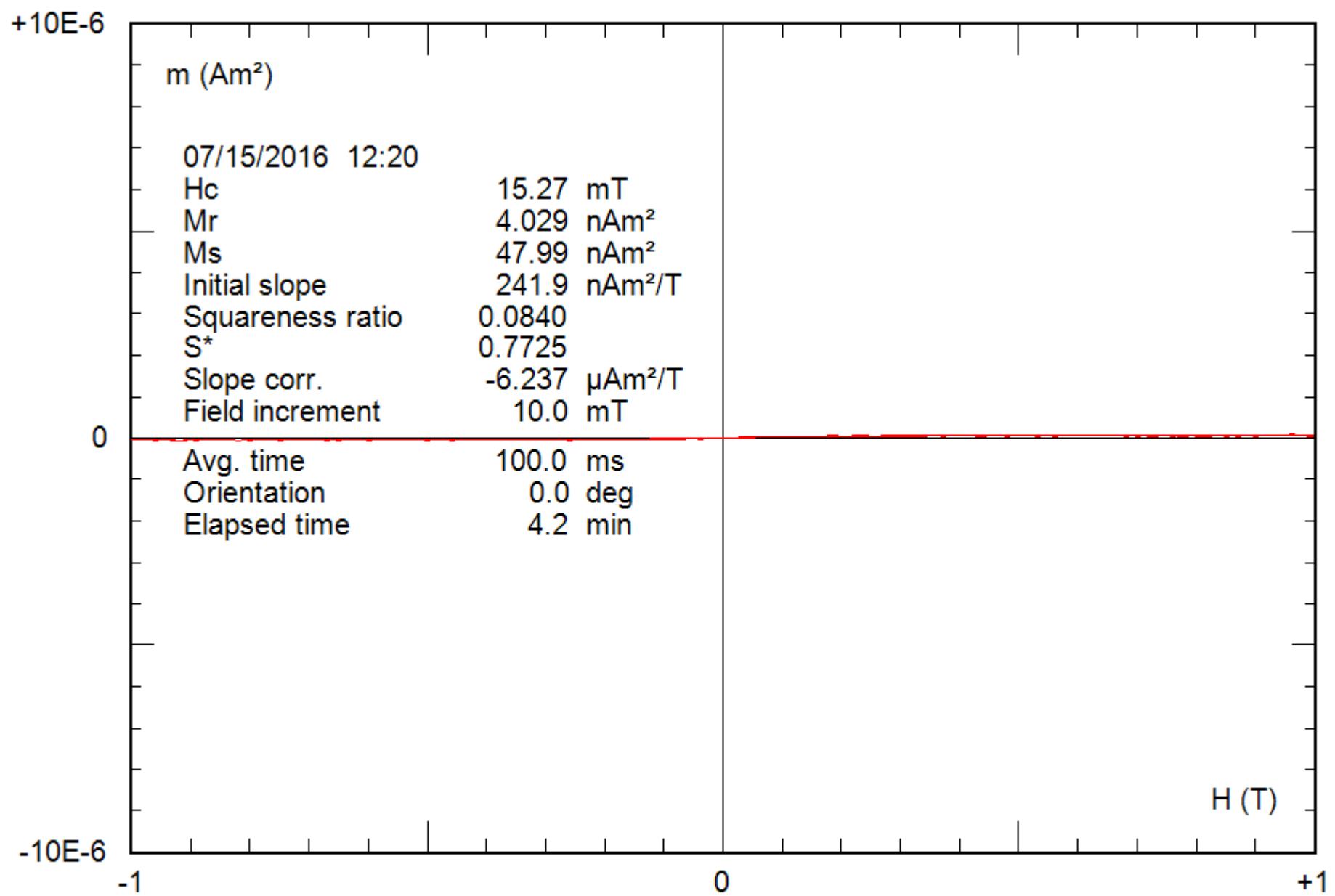
0.22 g

File: 07-01a\_bcr



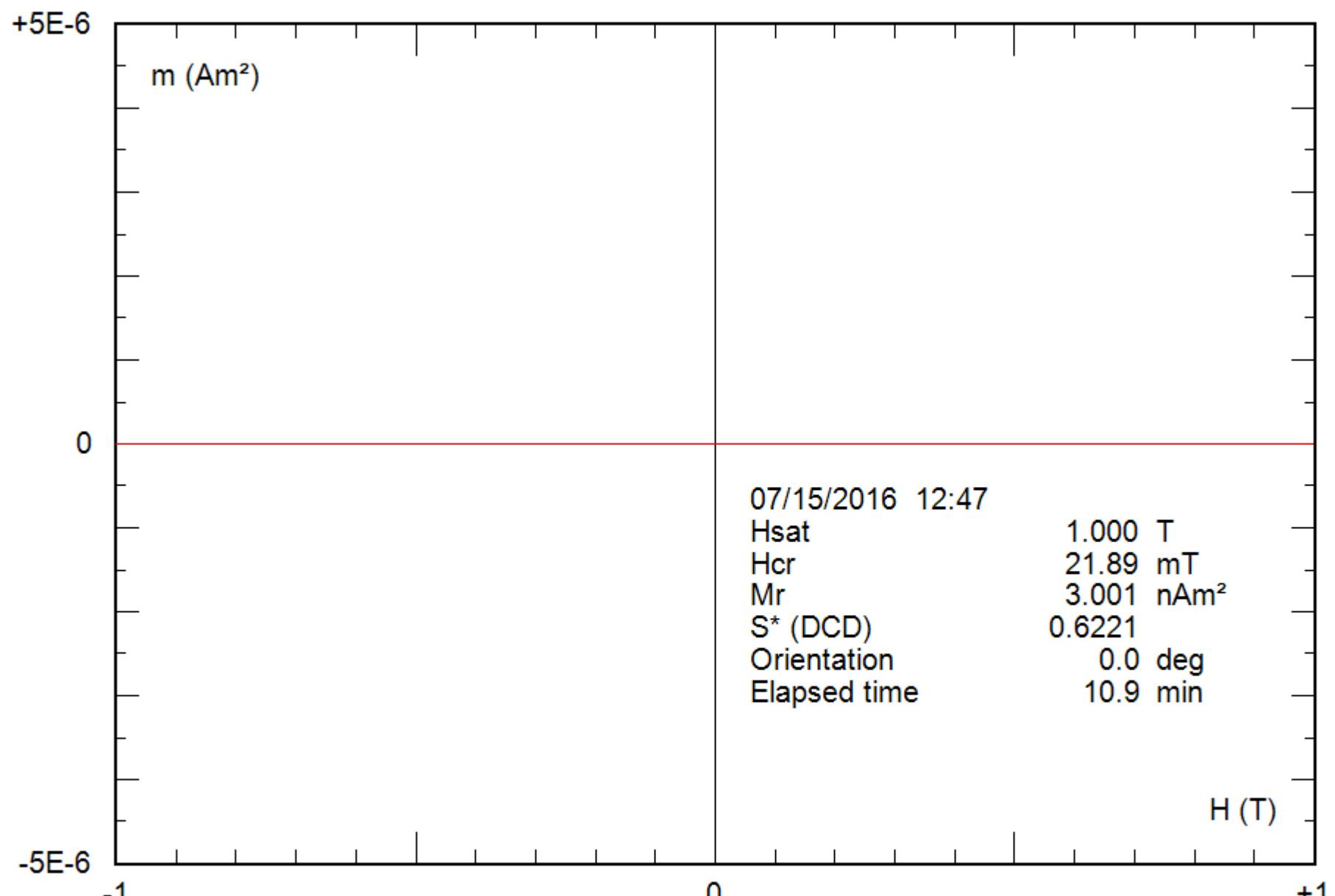
0.22 g

File: 07-01a\_hyst



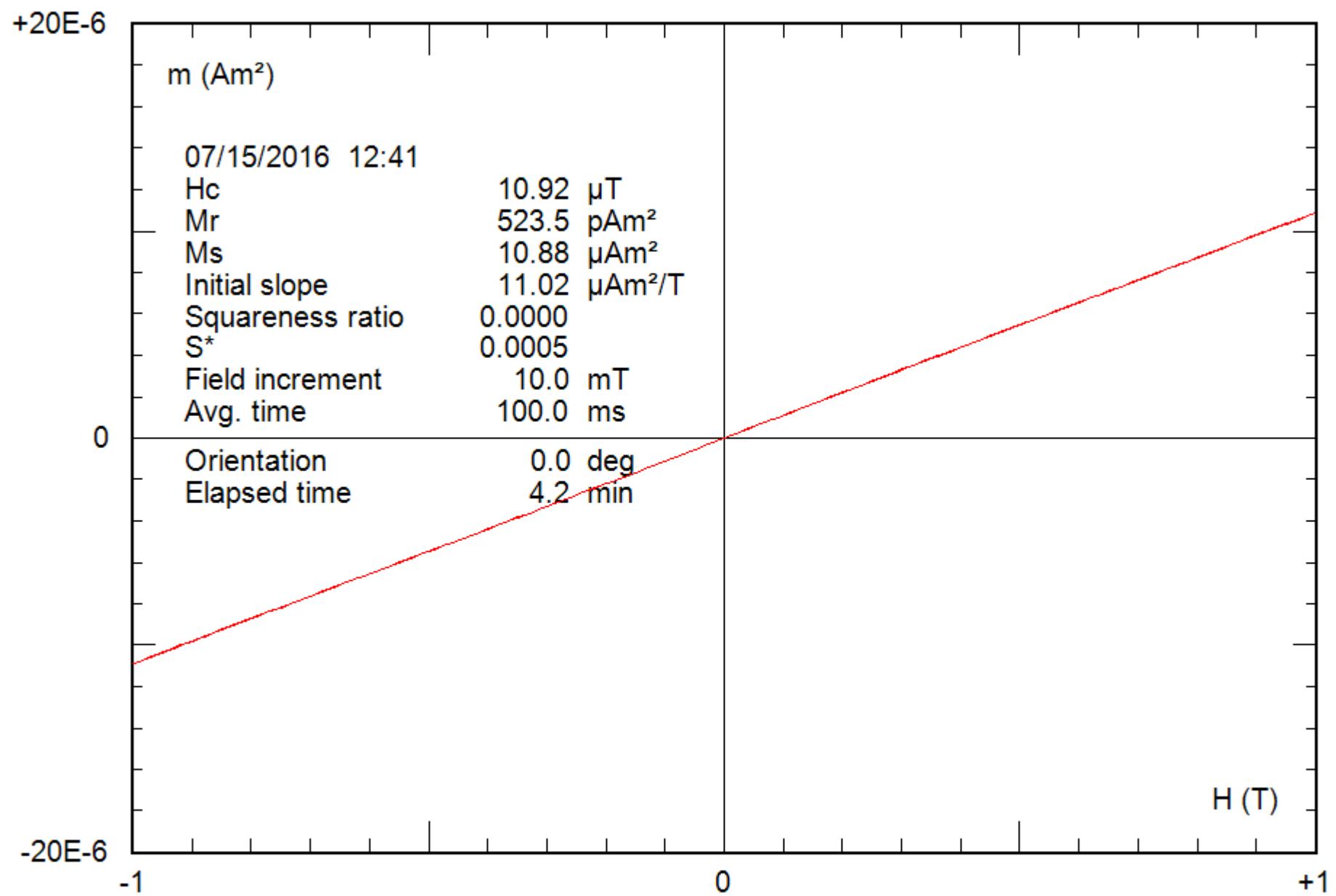
0.22 g

File: 07-01a\_hyst\_corr



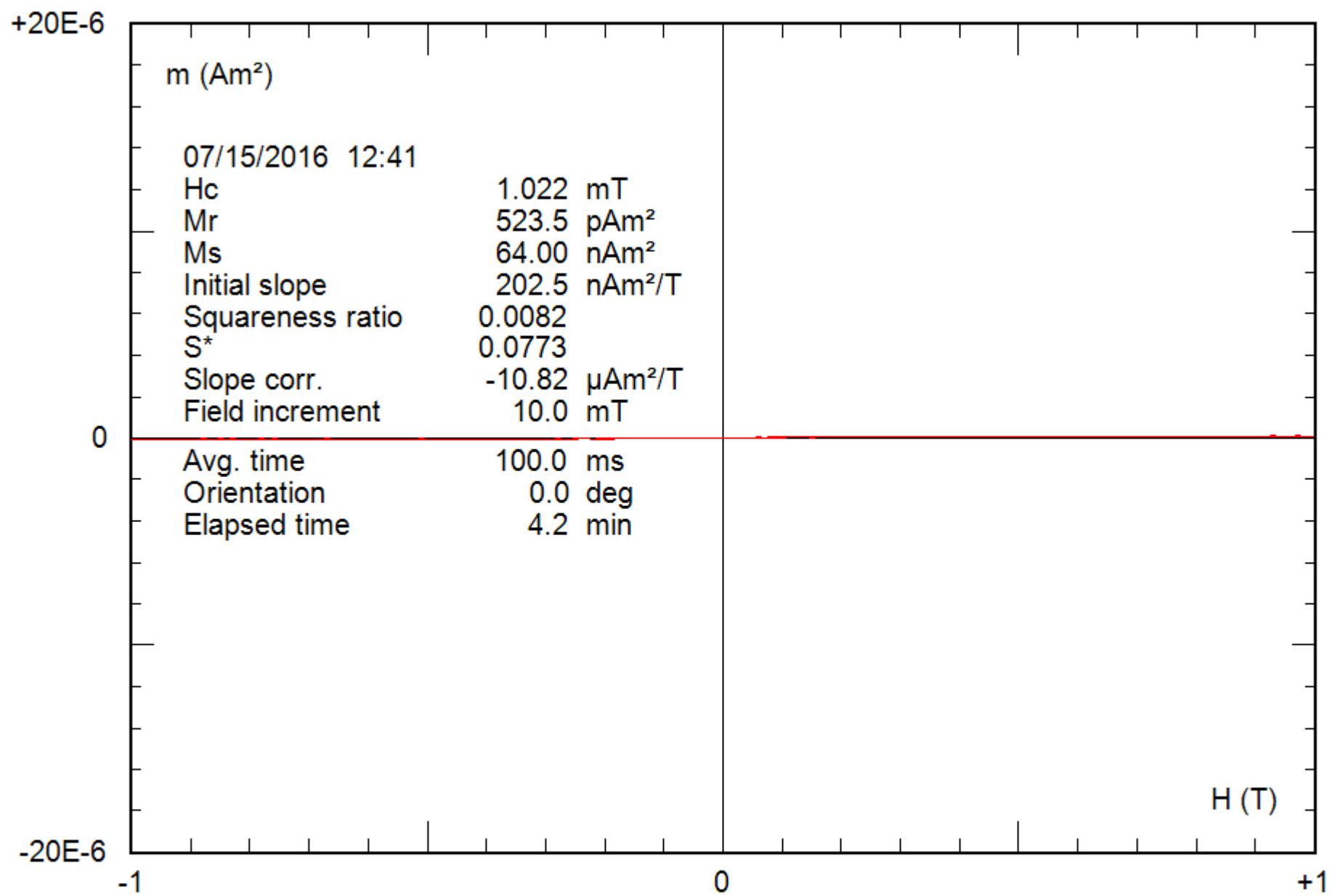
0.20 g

File: 08-04a\_bcr



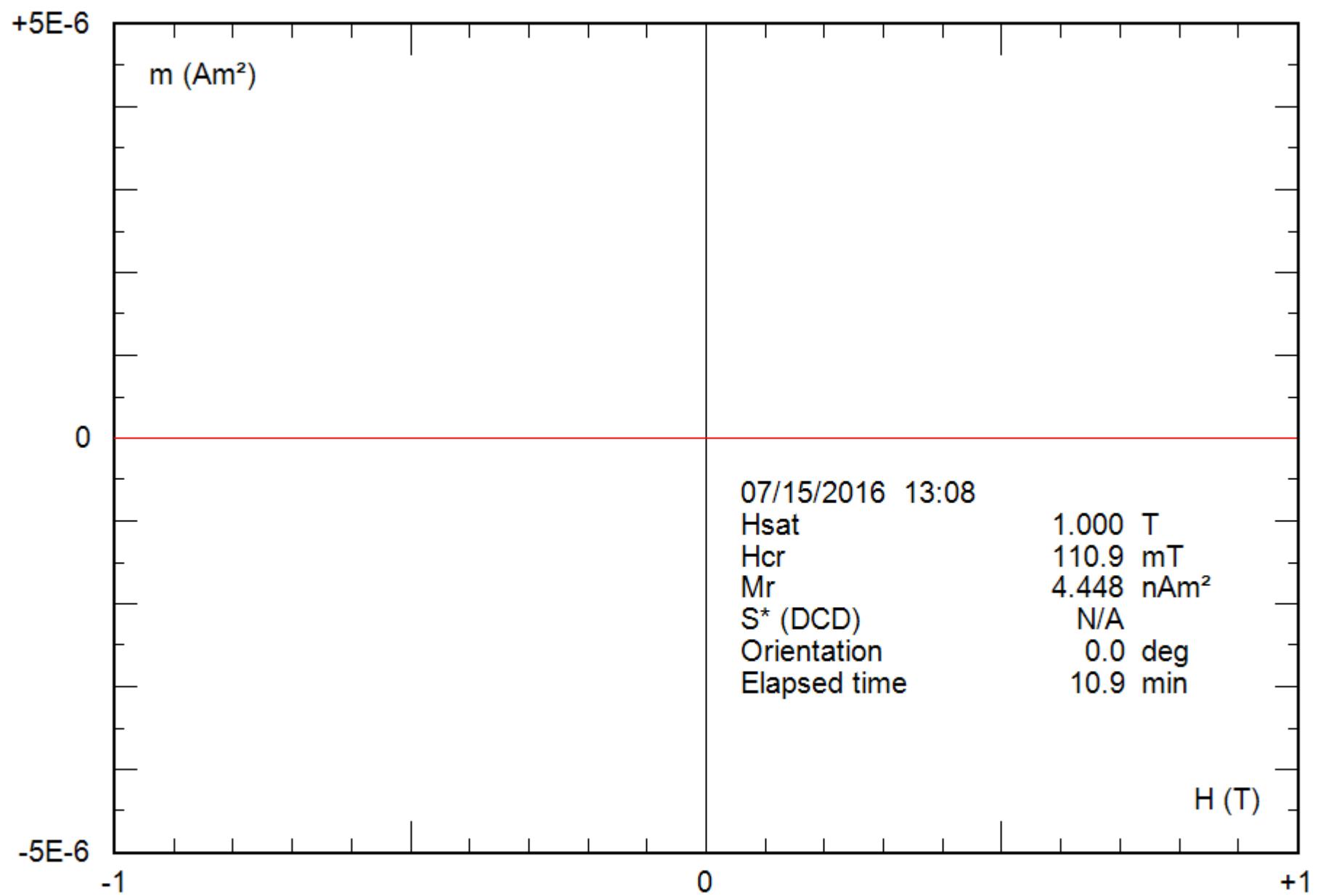
0.20 g

File: 08-04a\_hyst



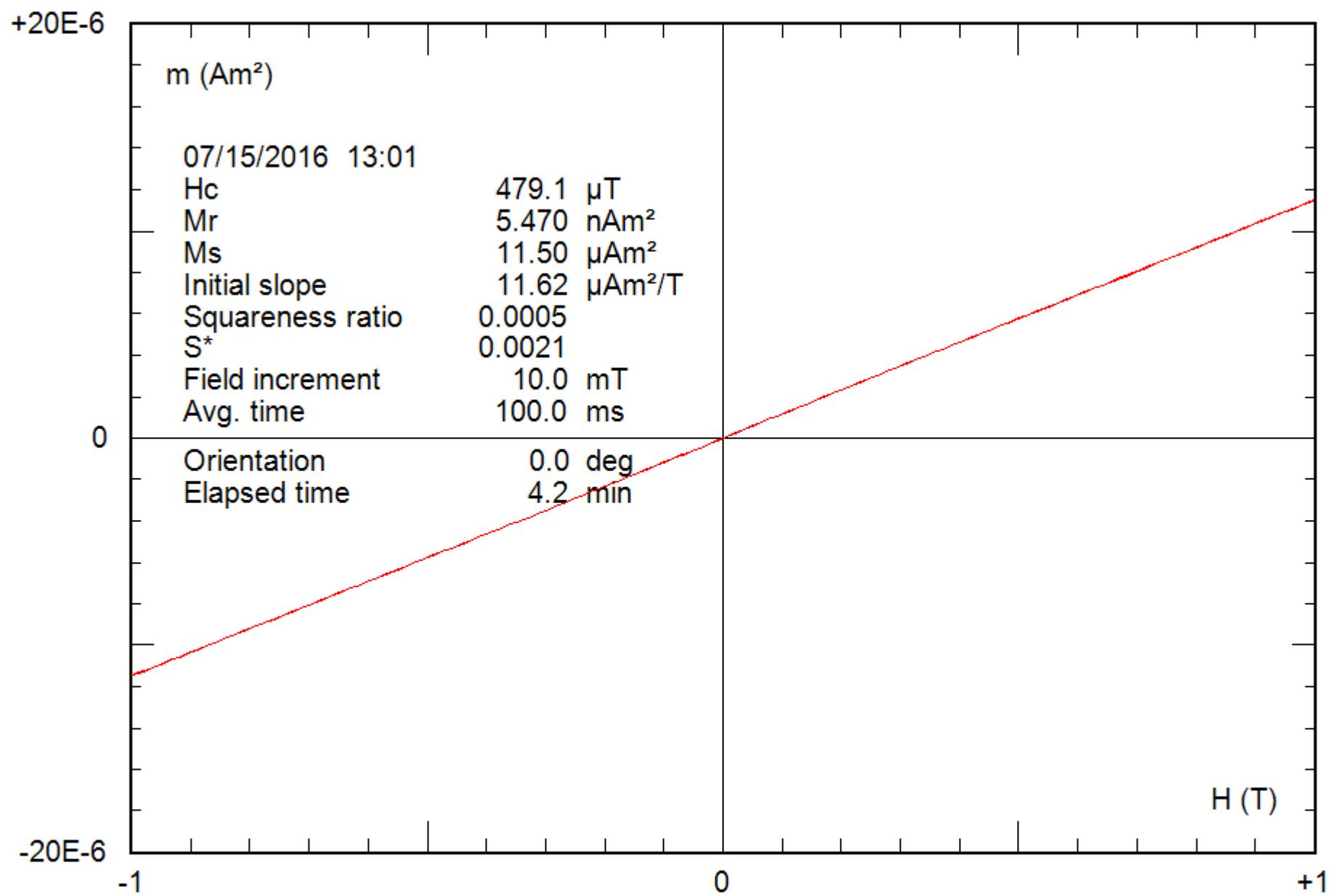
0.20 g

File: 08-04a\_hyst\_corr



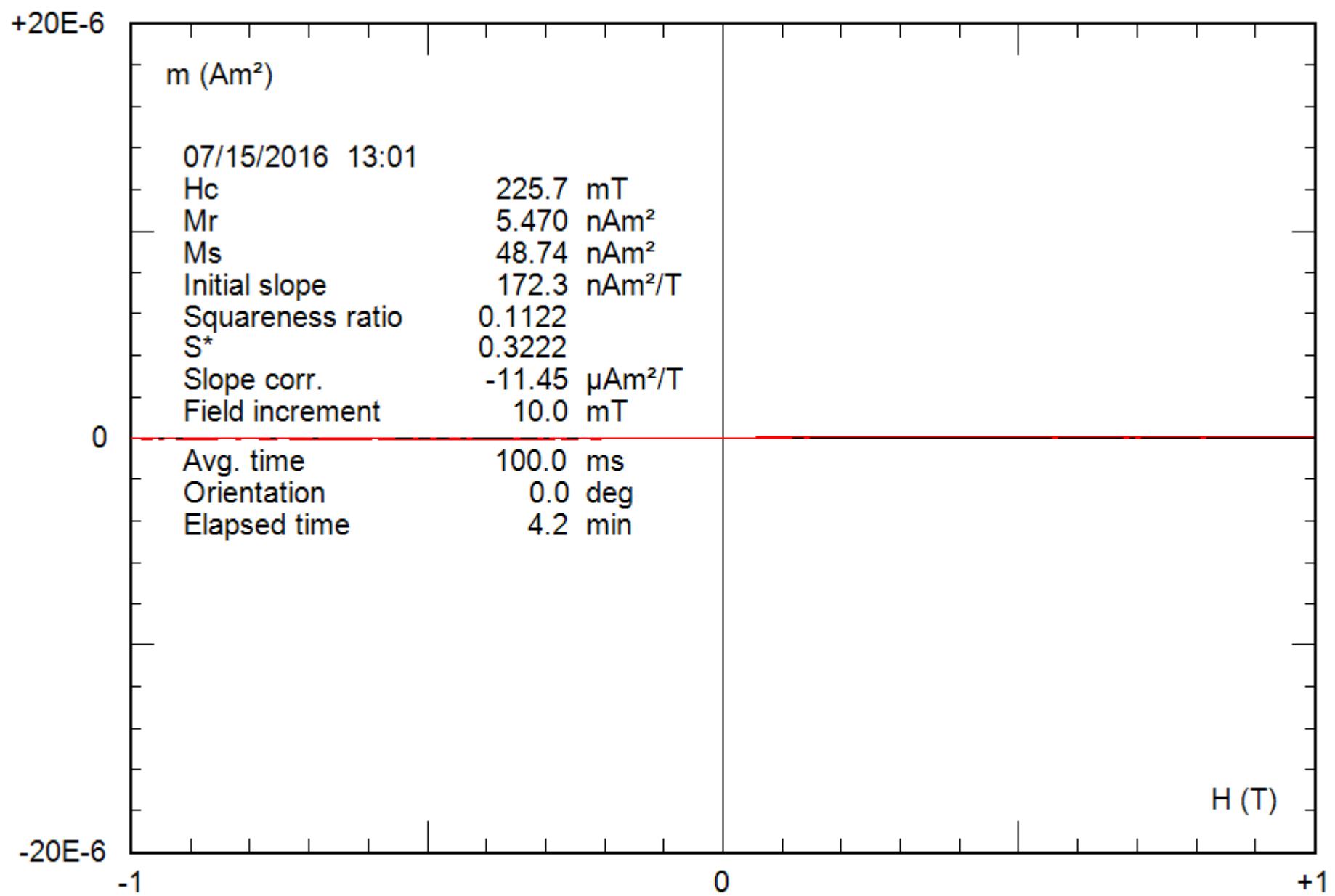
0.19 g

File: 09-04a\_bcr



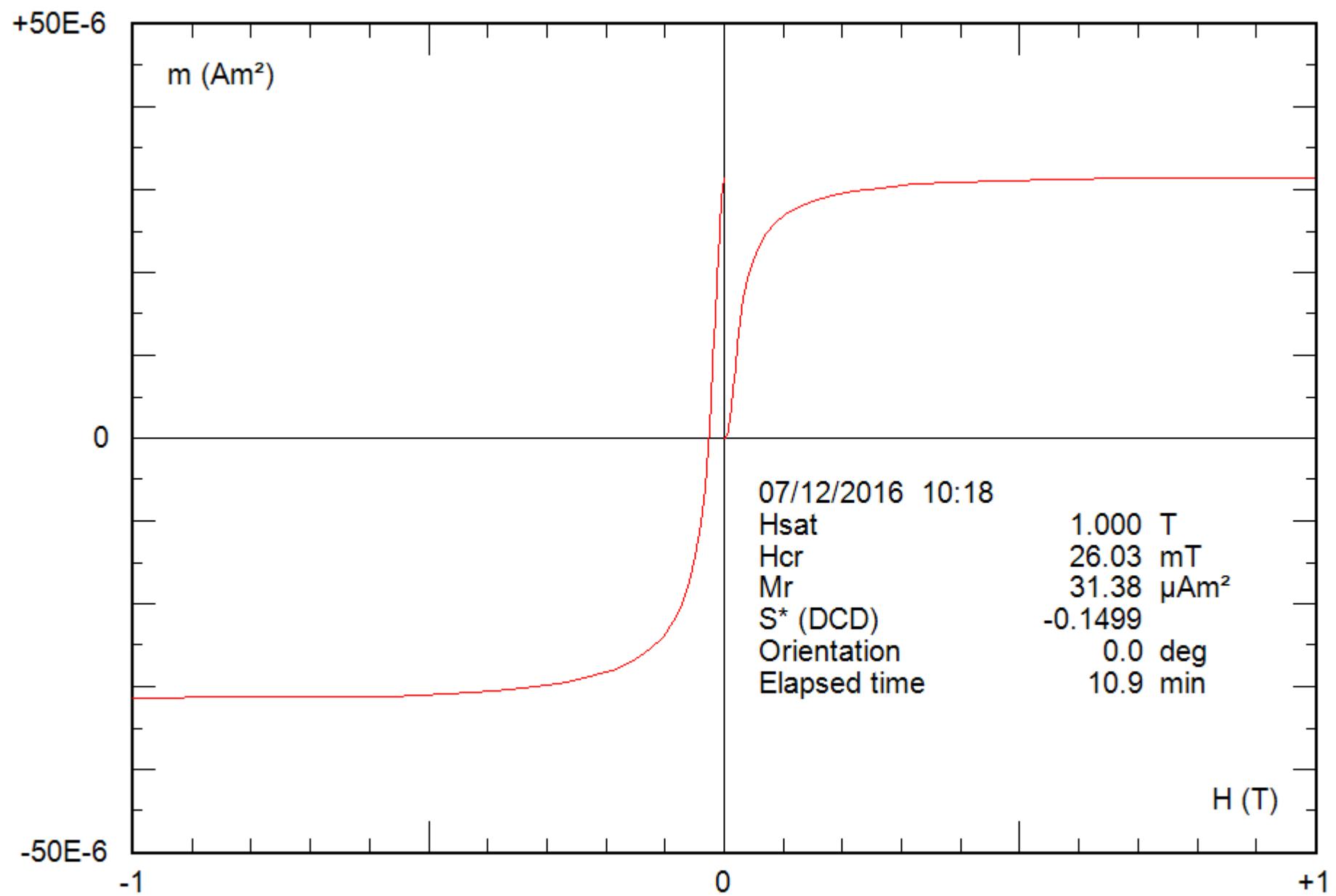
0.19 g

File: 09-04a\_hyst



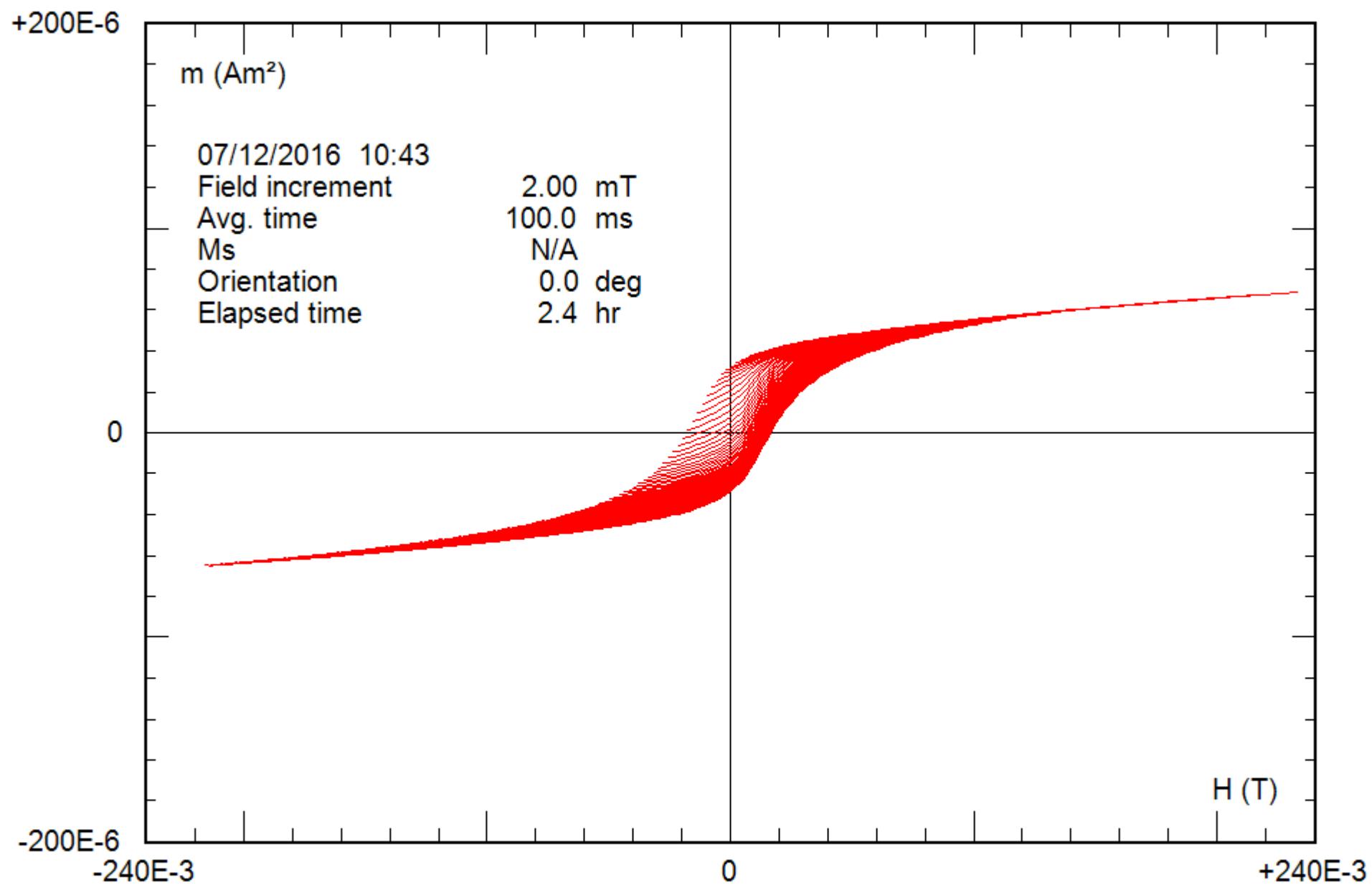
0.19 g

File: 09-04a\_hyst\_corr

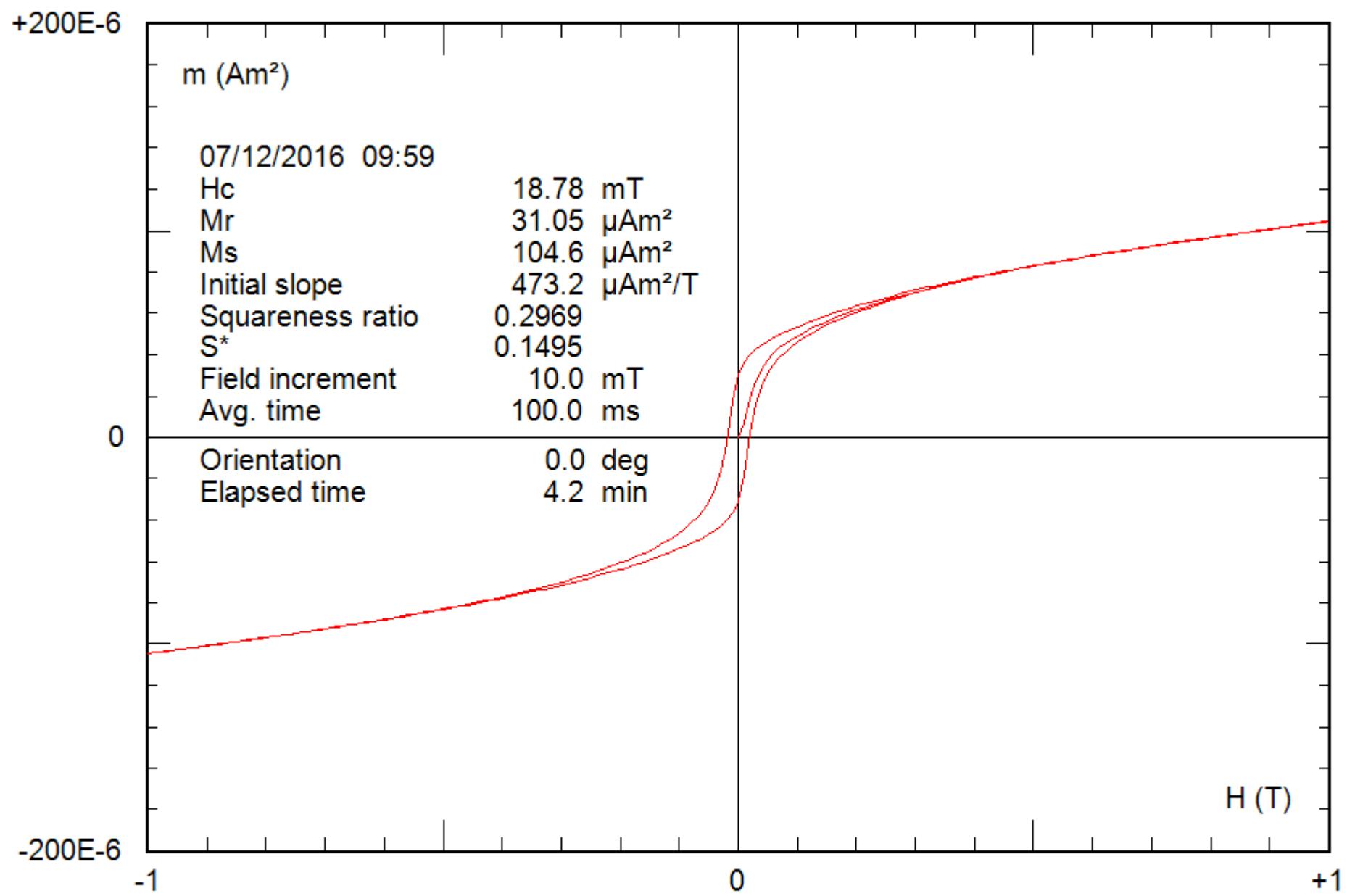


Description: [Not assigned]

File: 10-04a\_bcr

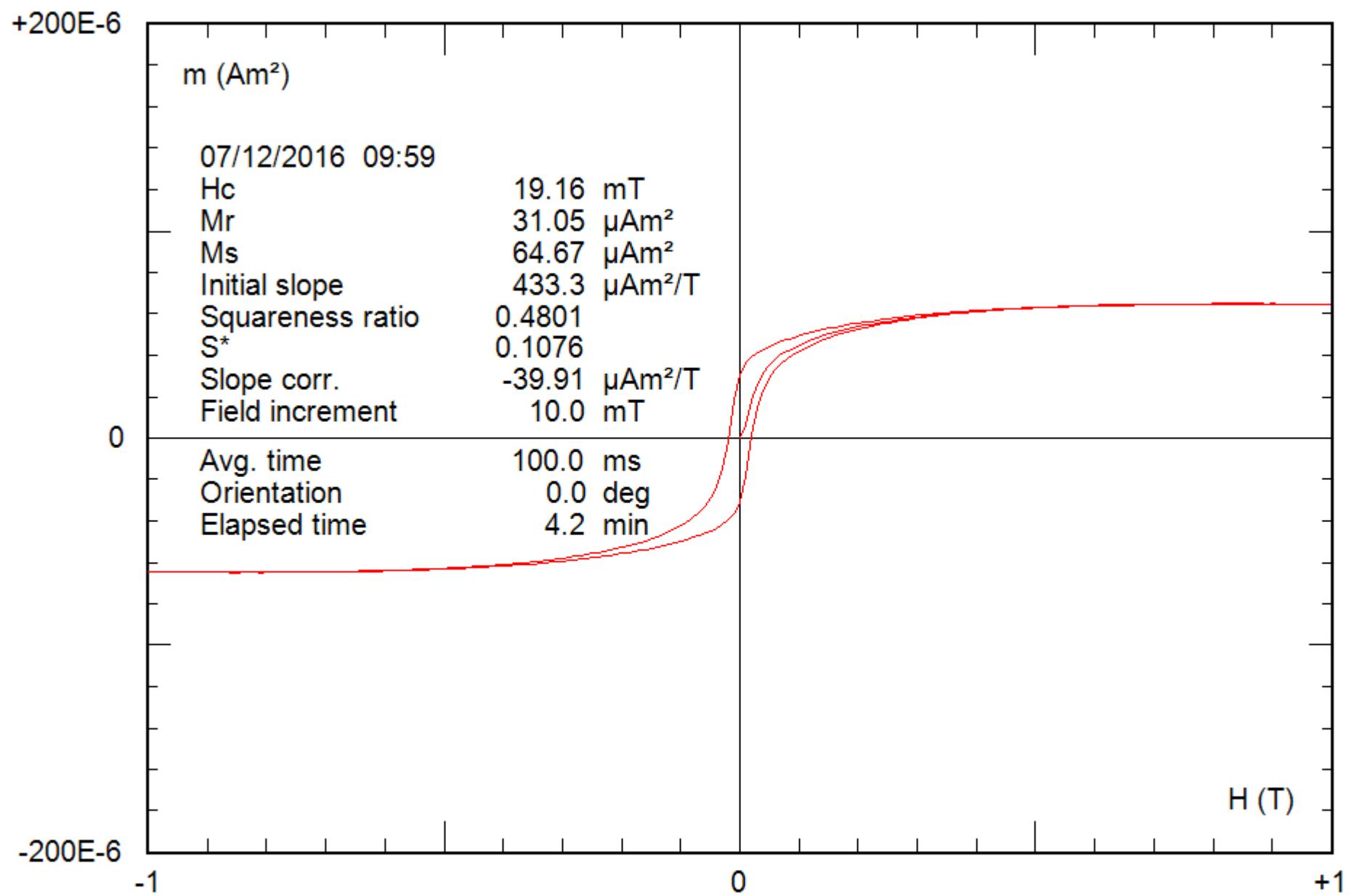


0.25 g  
File: 10-04a\_forc



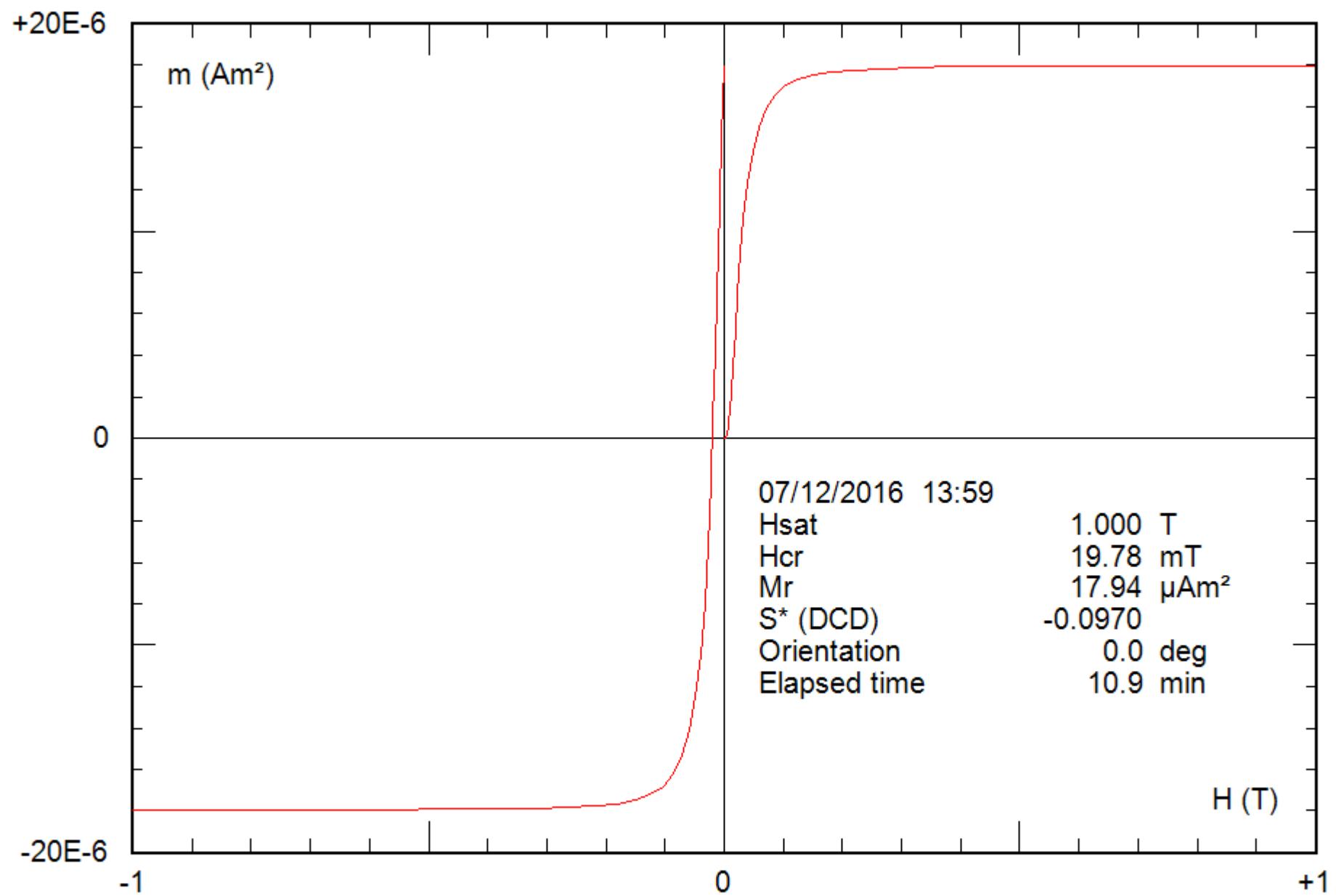
0.25 g

File: 10-04a\_hyst



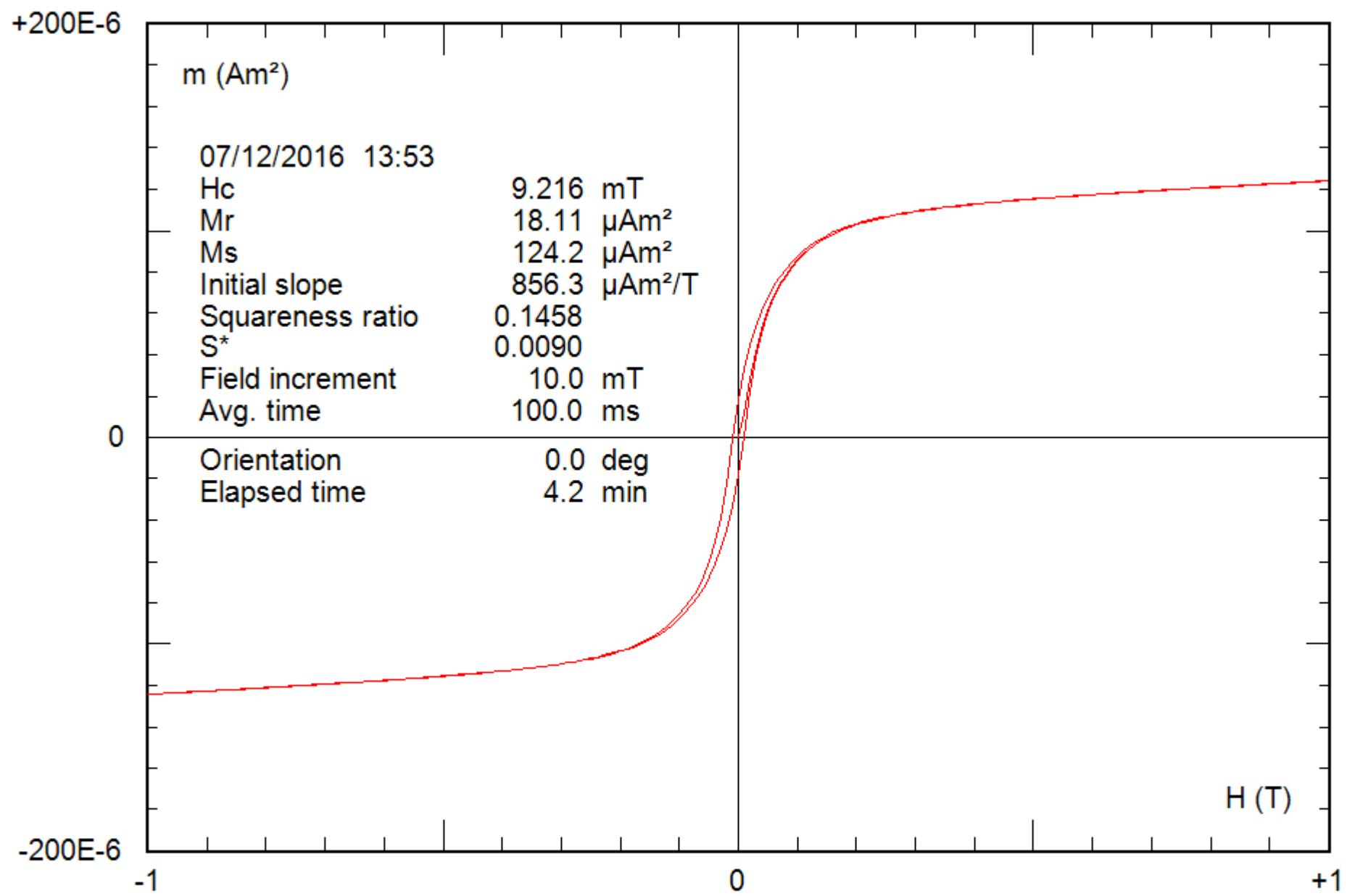
0.25 g

File: 10-04a\_hyst\_corr



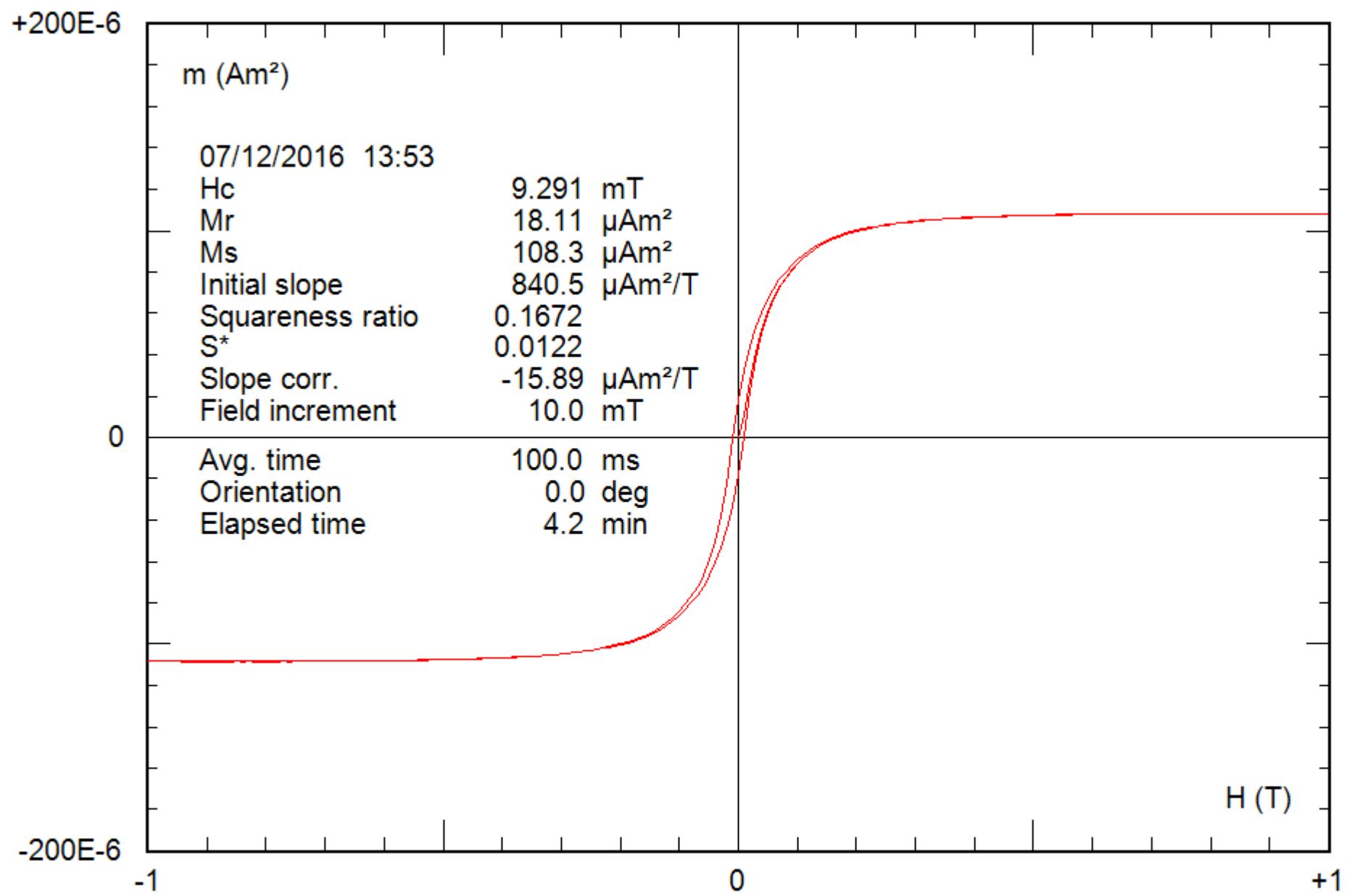
0.26 g

File: 11-04a\_bcr



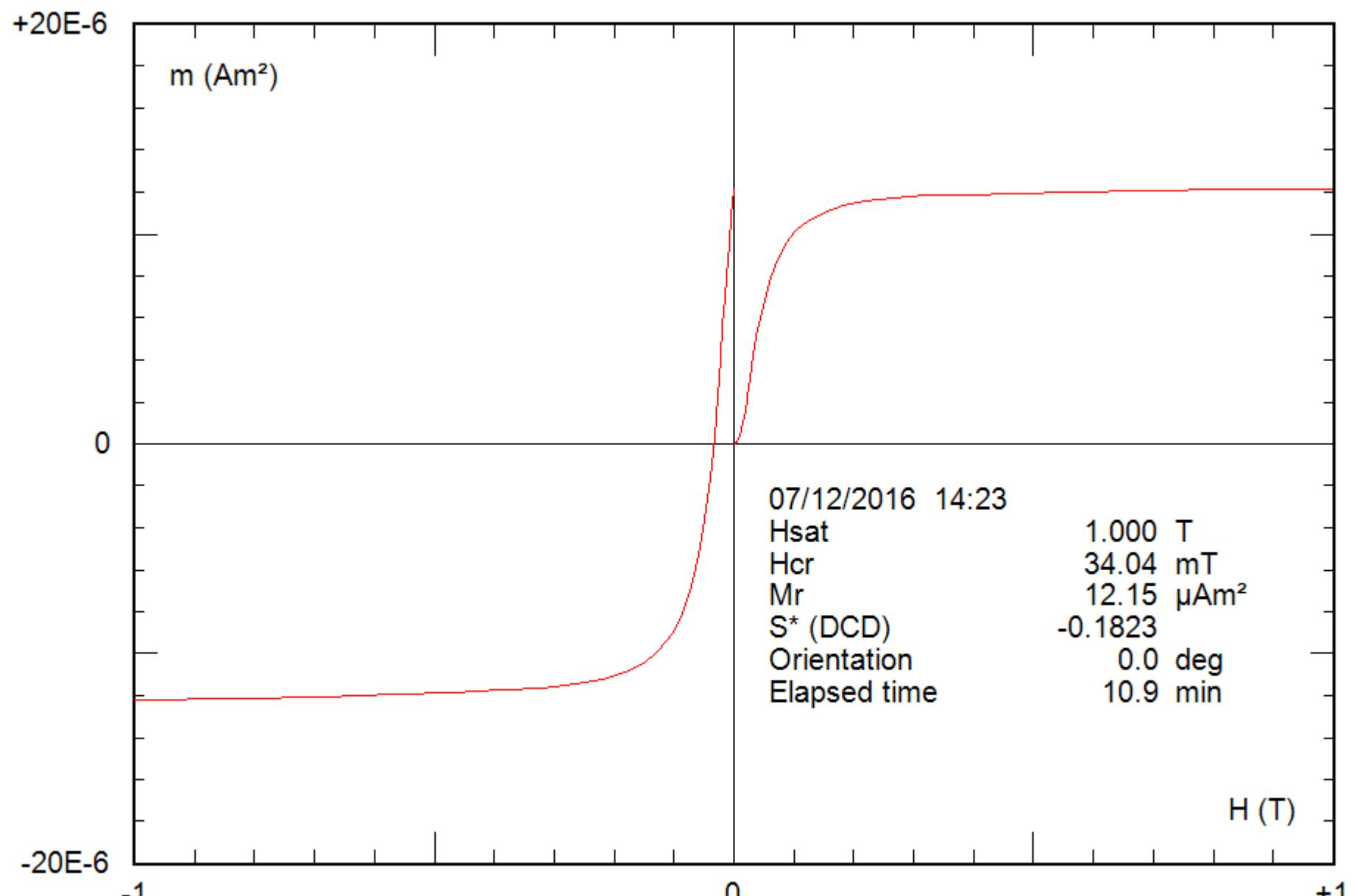
0.26 g

File: 11-04a\_hyst



0.26 g

File: 11-04a\_hyst



0.22 g

File: 12-04a\_bcr

+100E-6

$m$  ( $\text{Am}^2$ )

07/12/2016 14:15

$H_c$  16.28 mT

$M_r$  12.24  $\mu\text{Am}^2$

$M_s$  72.50  $\mu\text{Am}^2$

Initial slope 419.1  $\mu\text{Am}^2/\text{T}$

Squareness ratio 0.1688

$S^*$  0.0758

Field increment 10.0 mT

Avg. time 100.0 ms

Orientation 0.0 deg

Elapsed time 4.2 min

0

-100E-6

-1

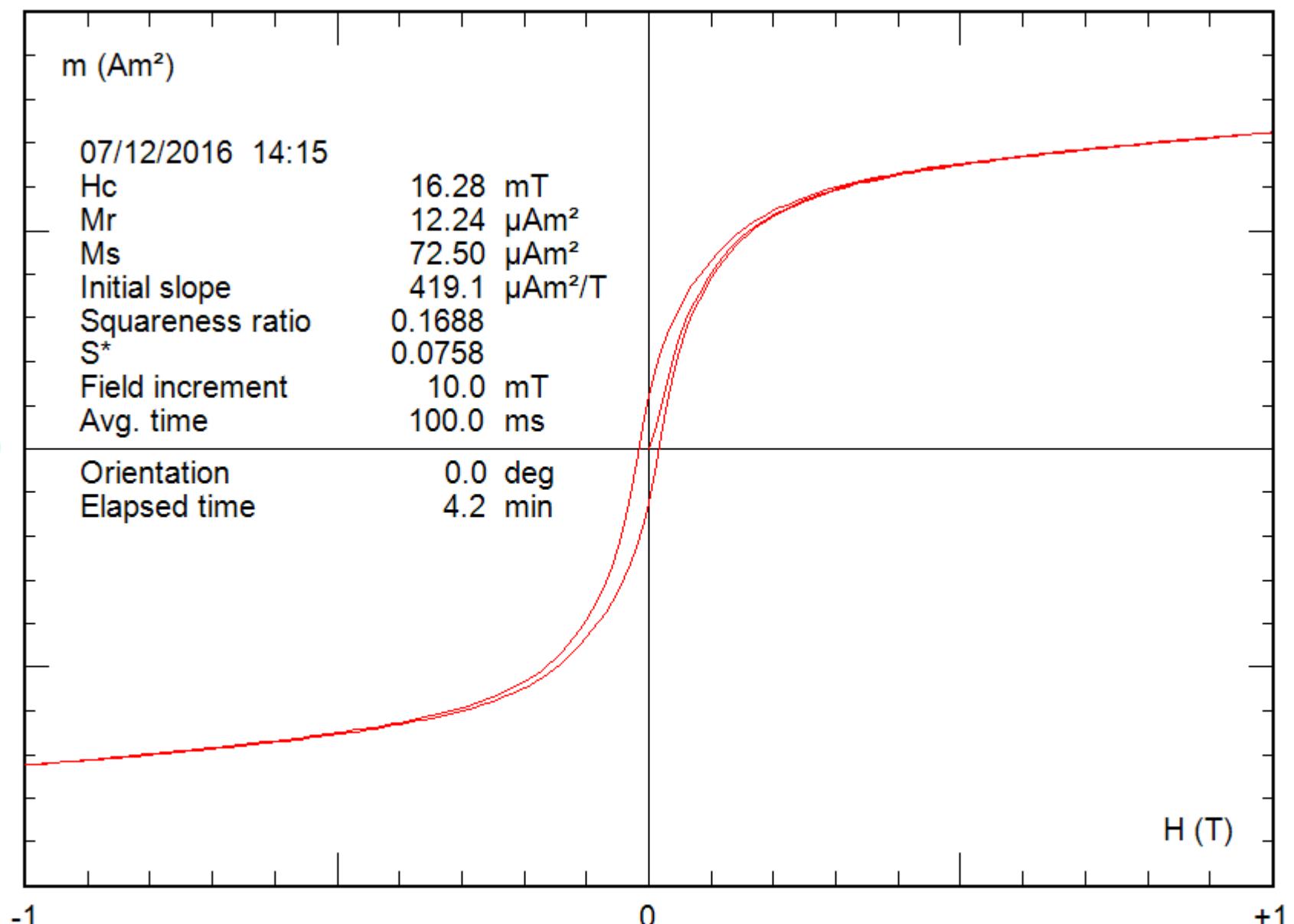
0

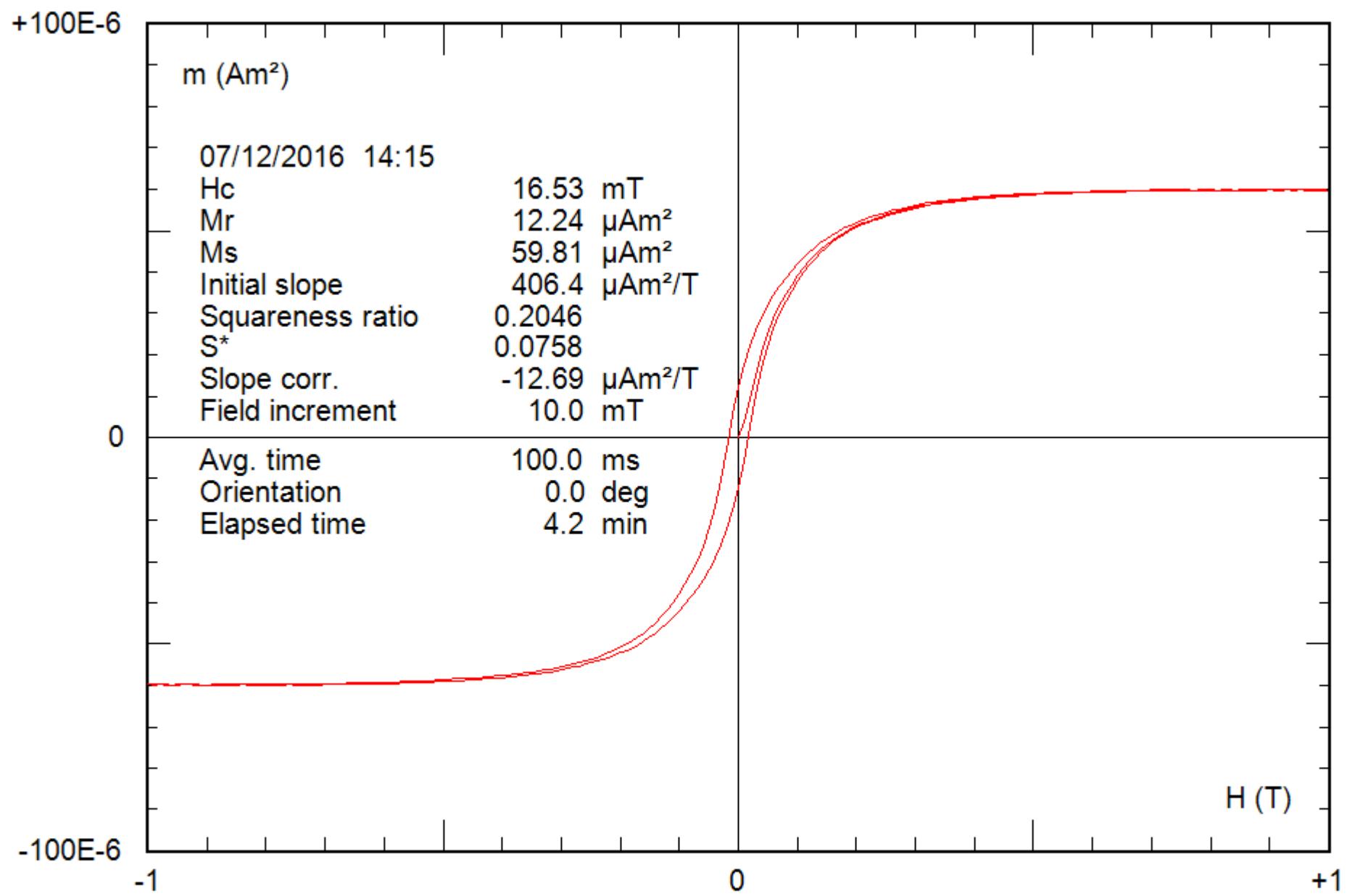
+1

$H$  (T)

0.22 g

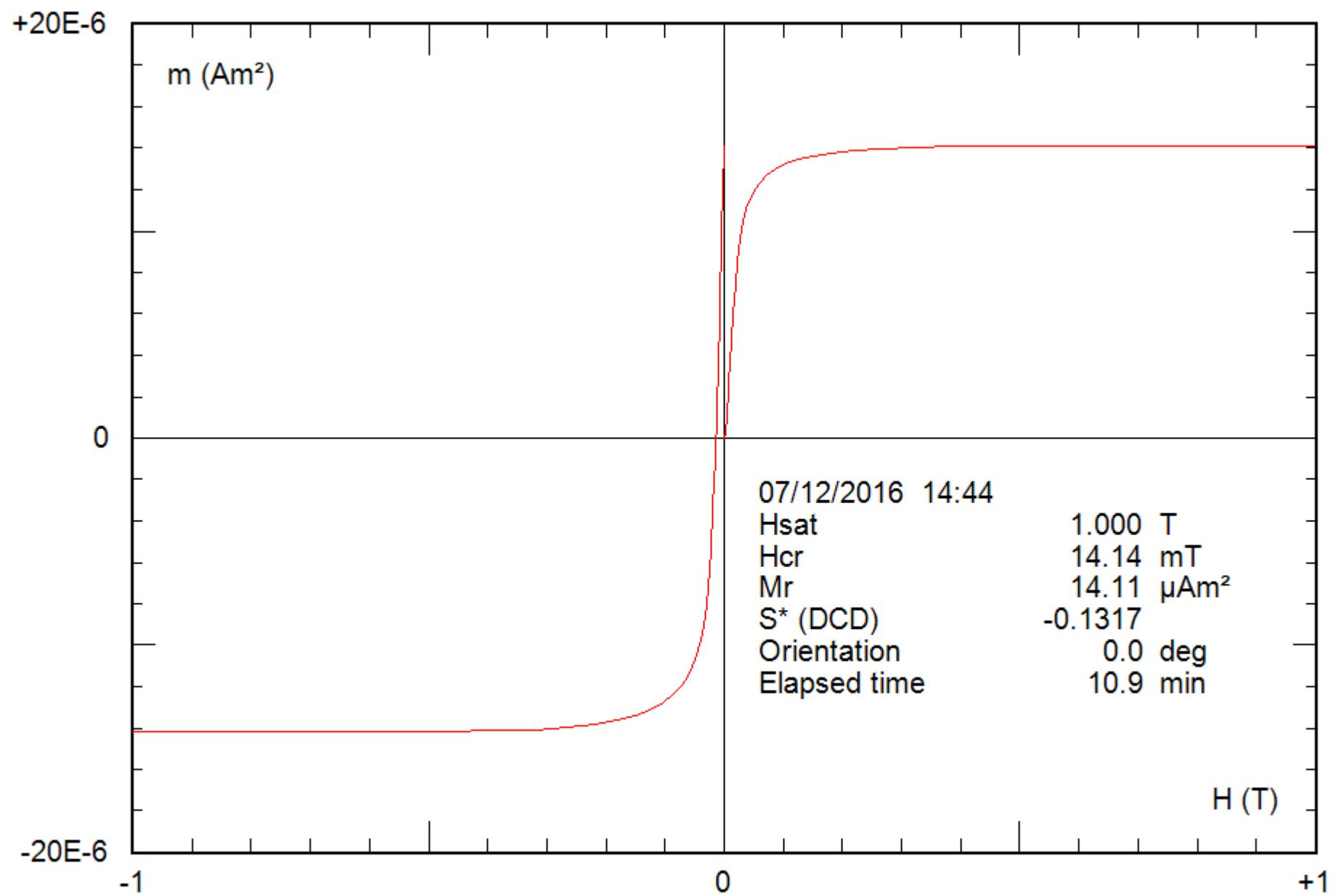
File: 12-04a\_hyst





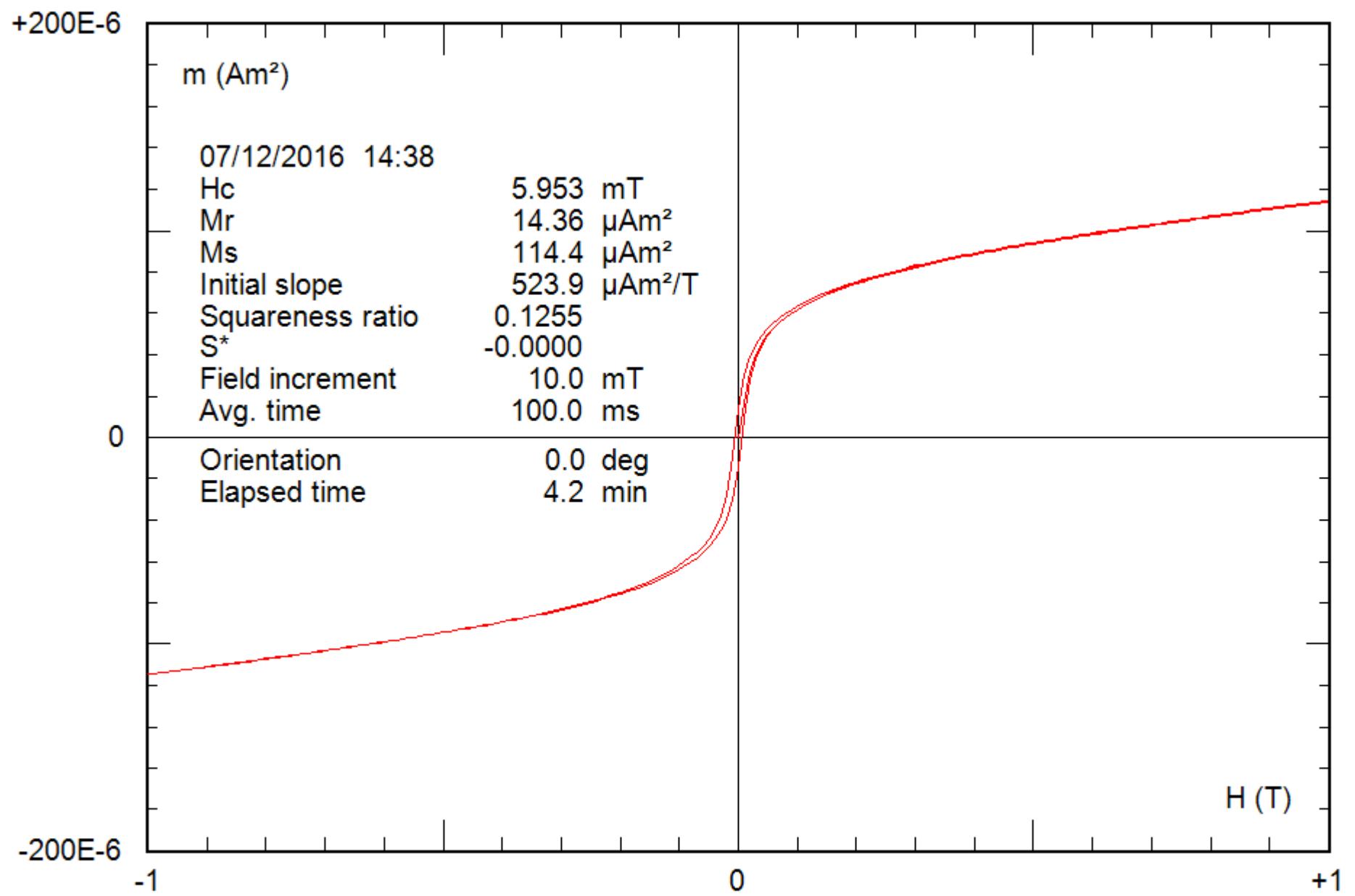
0.22 g

File: 12-04a\_hyst\_corr



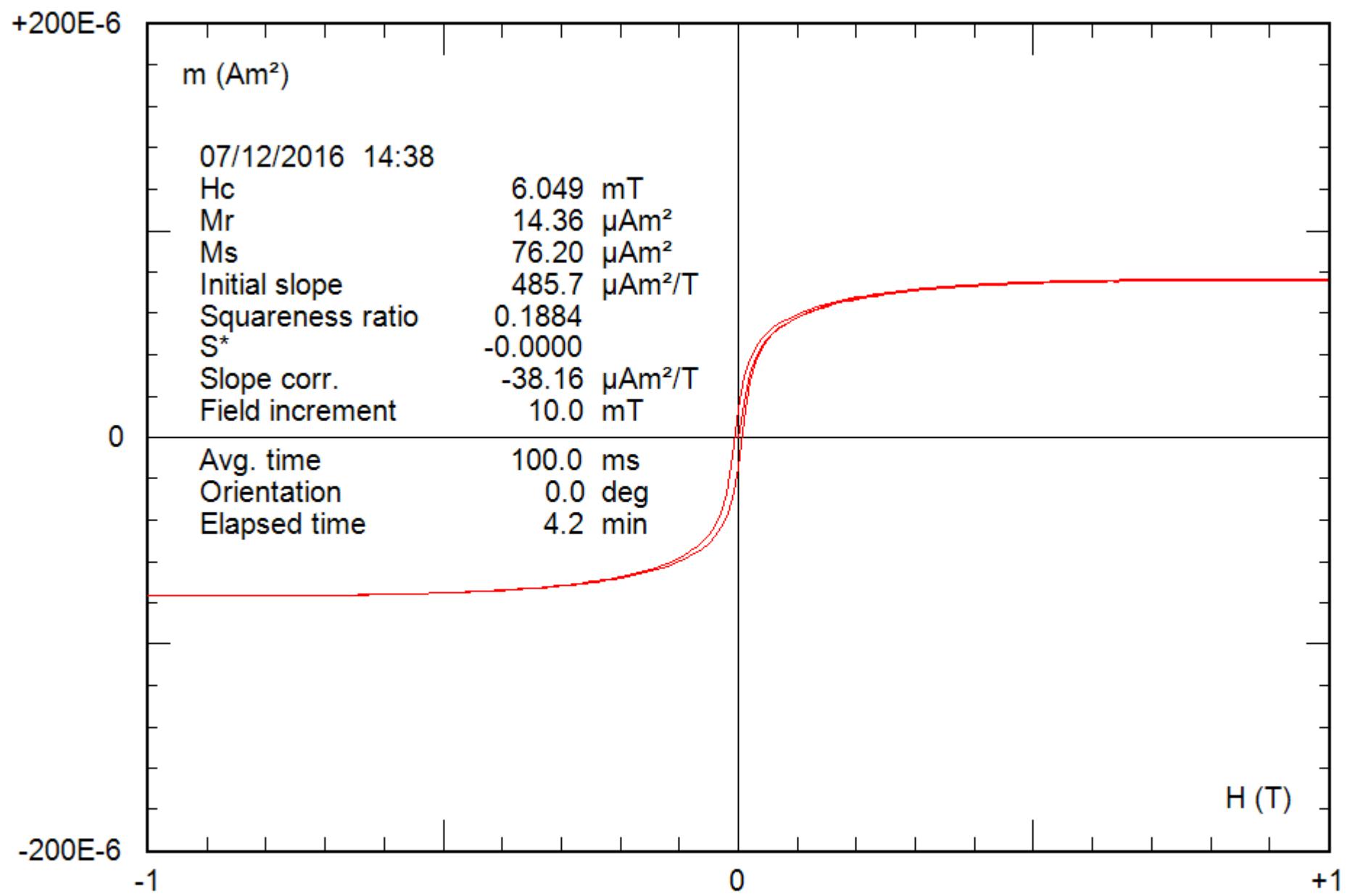
0.25 g

File: 13-04a\_bcr



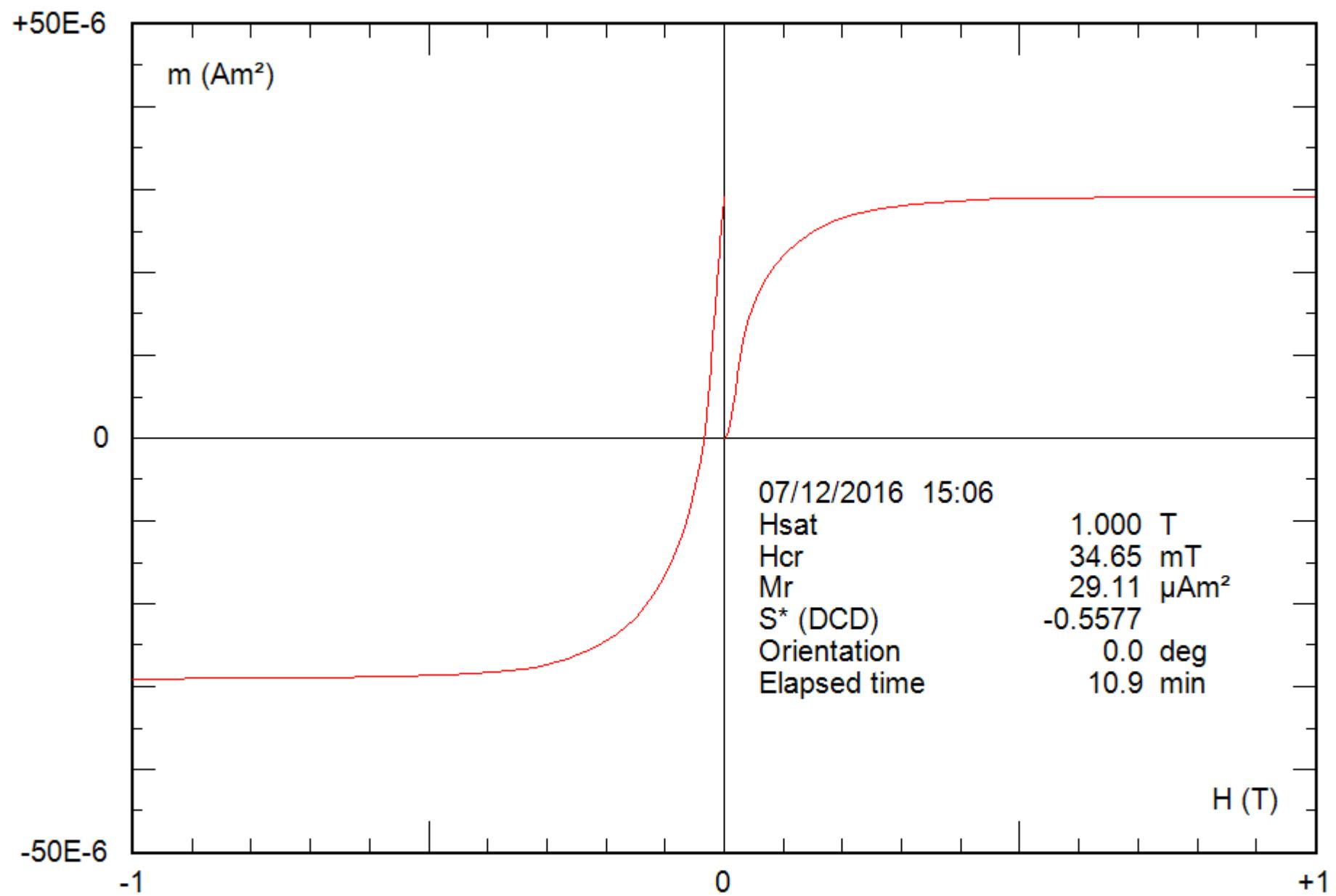
0.25 g

File: 13-04a\_hyst



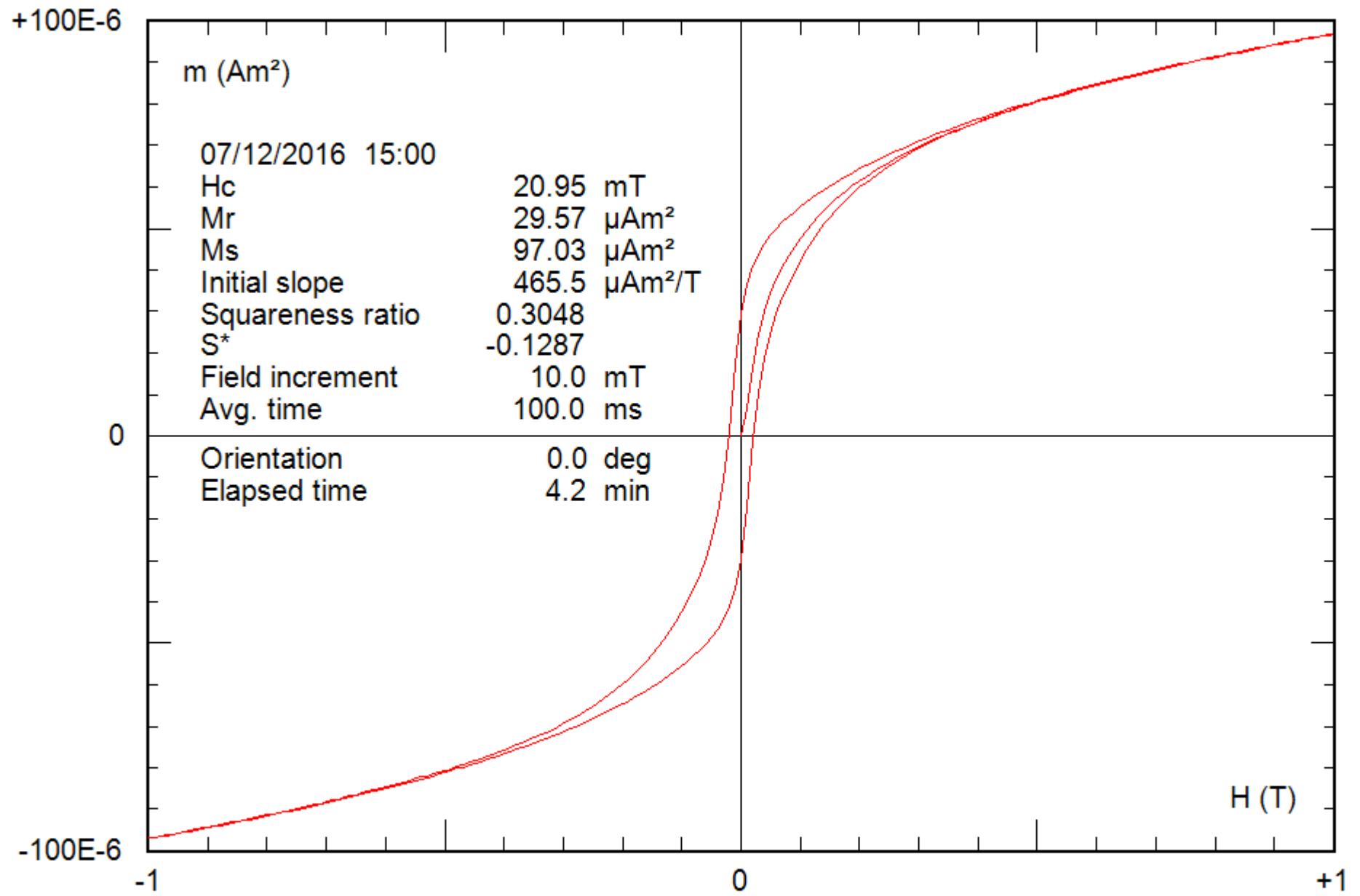
0.25 g

File: 13-04a\_hyst\_corr



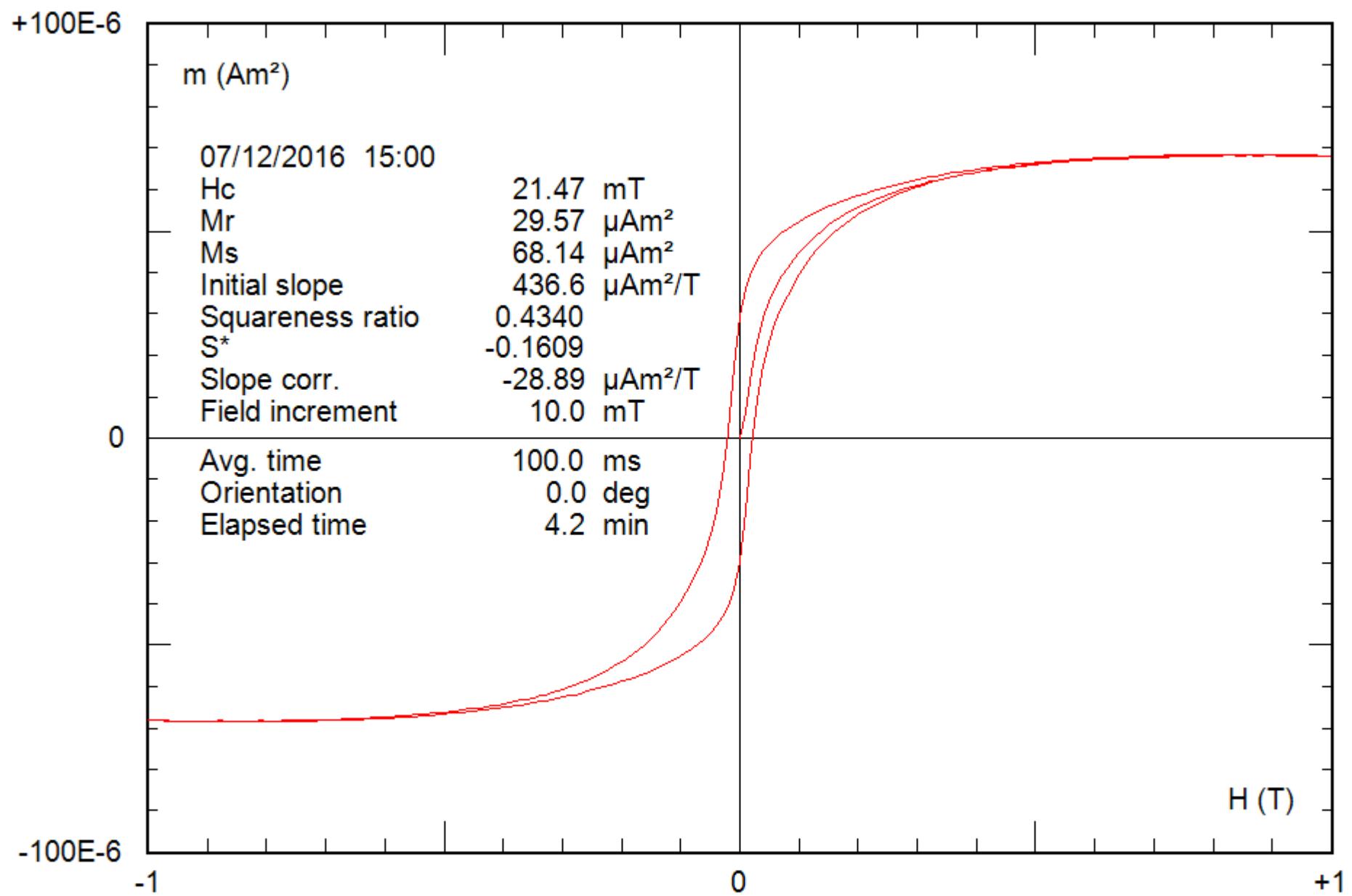
0.21 g

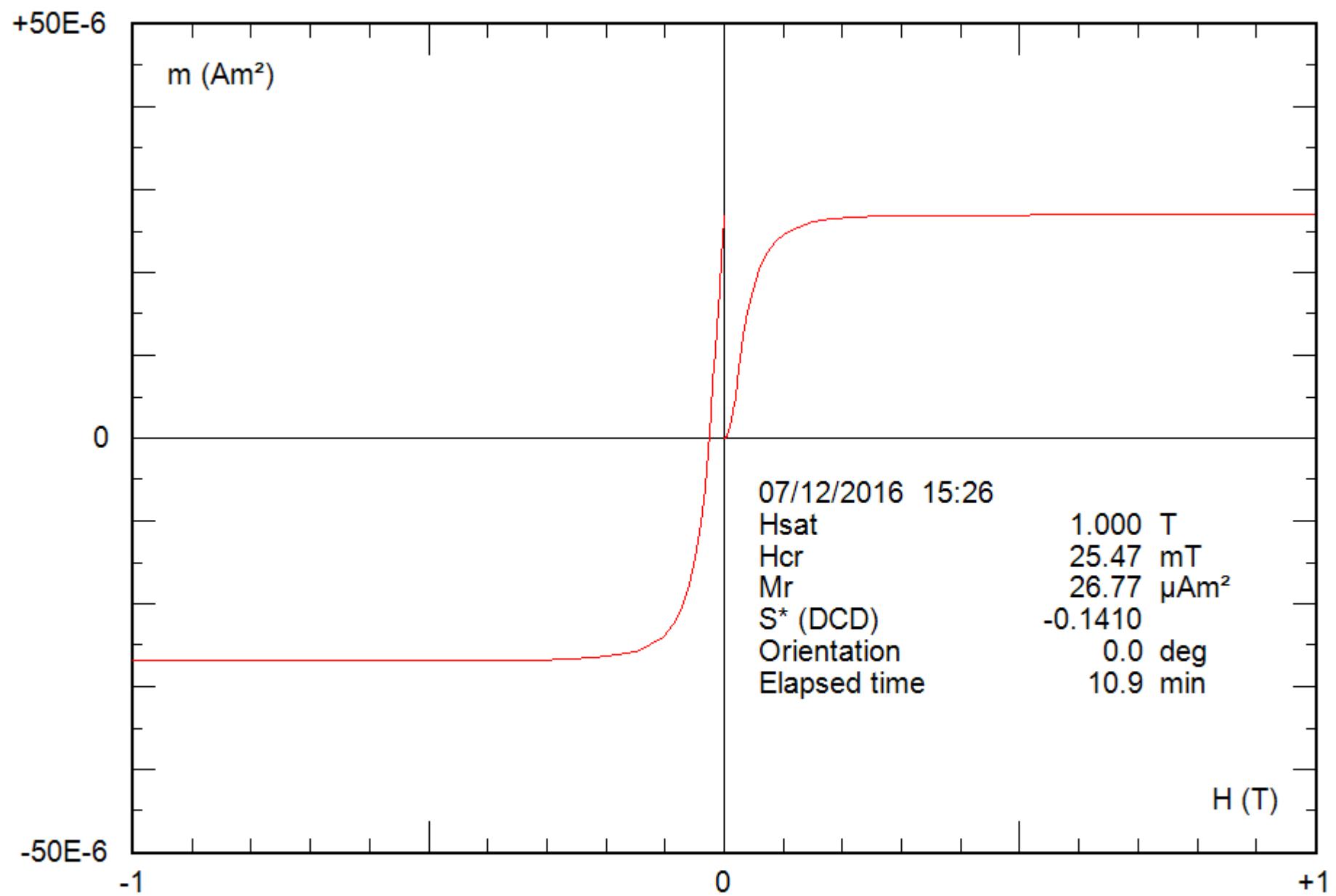
File: 14-04a\_bcr



0.21 g

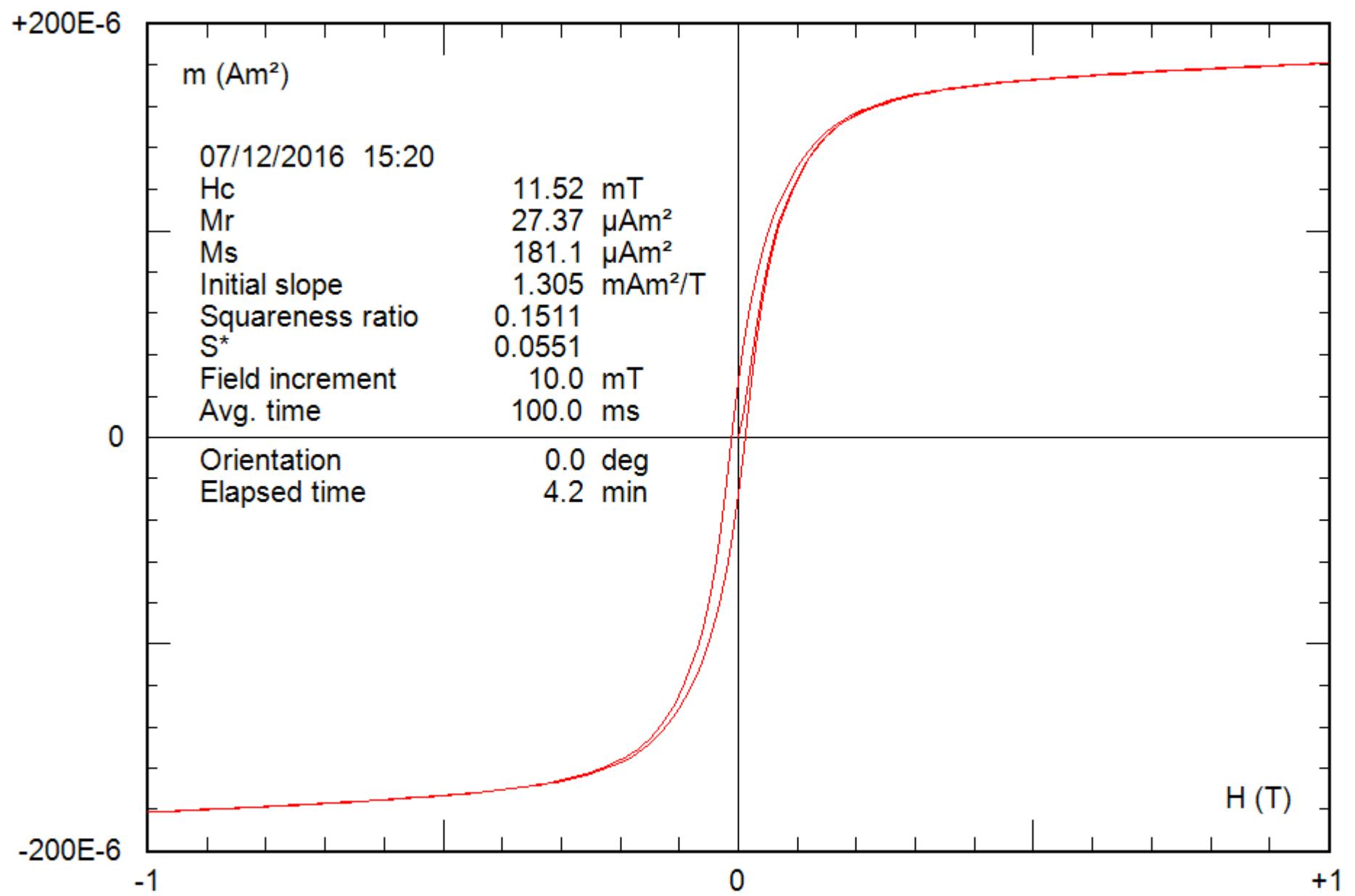
File: 14-04a\_hyst

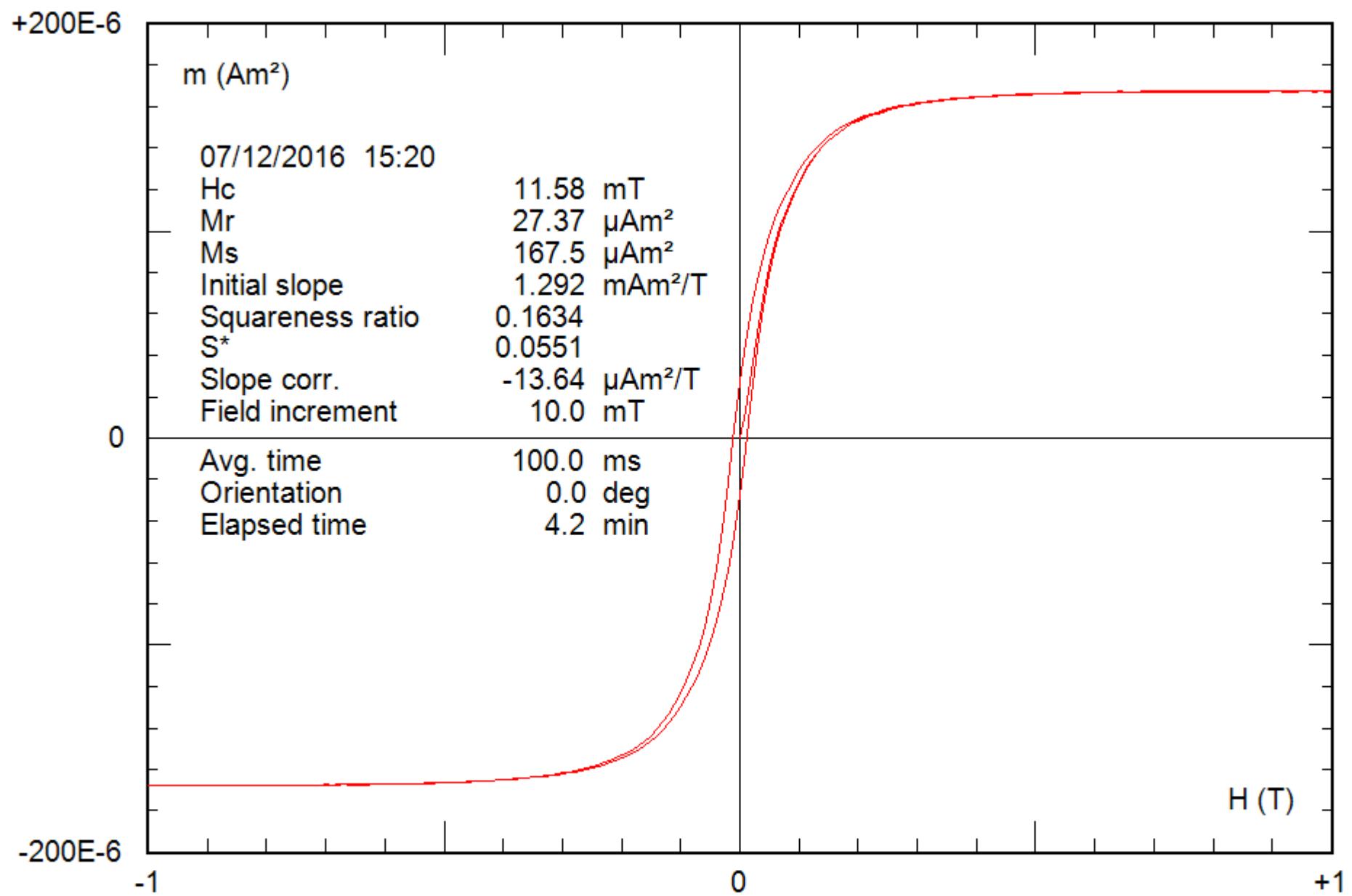




0.25 g

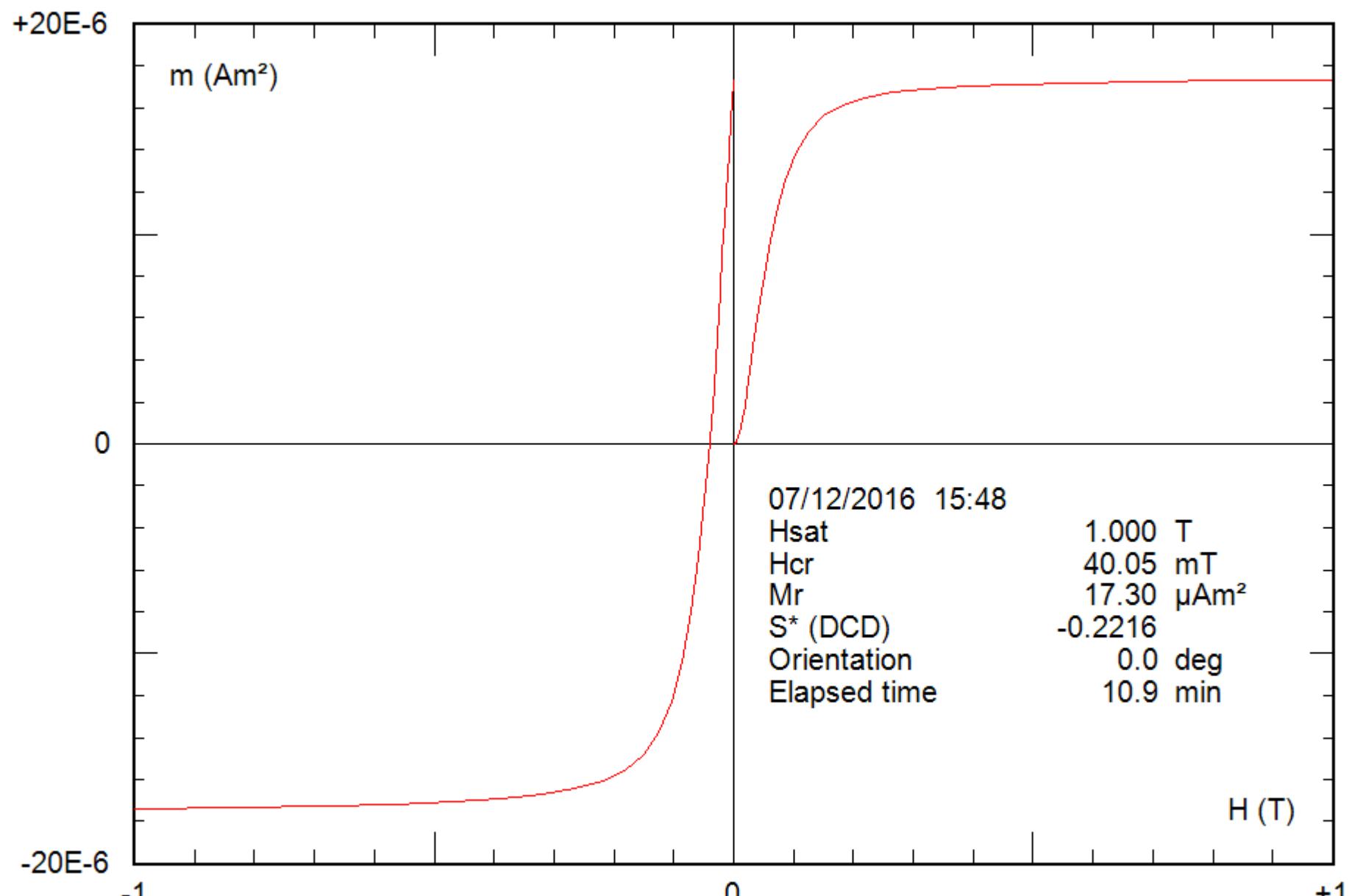
File: 15-04a\_bcr





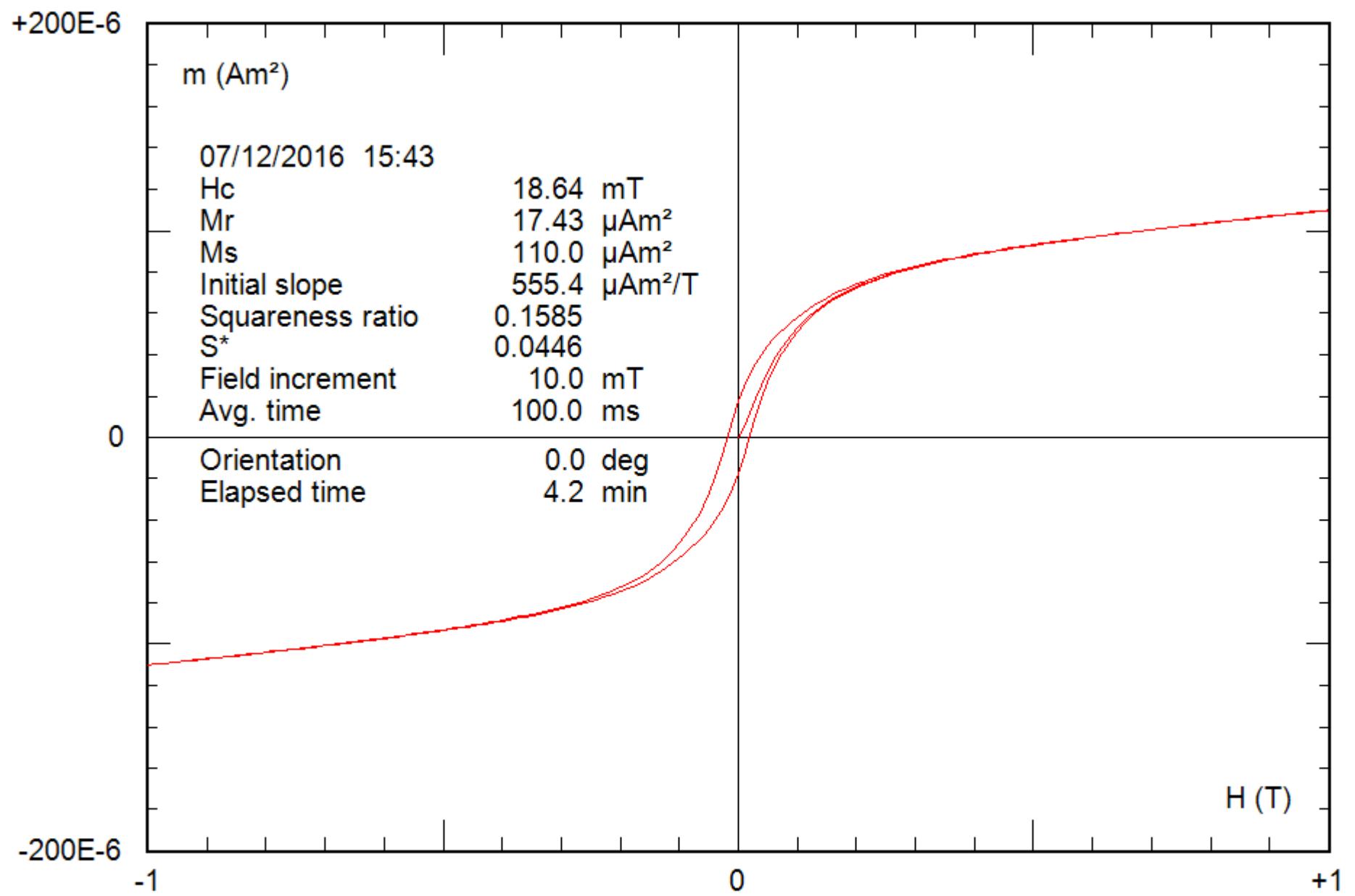
0.25 g

File: 15-04a\_hyst\_corr



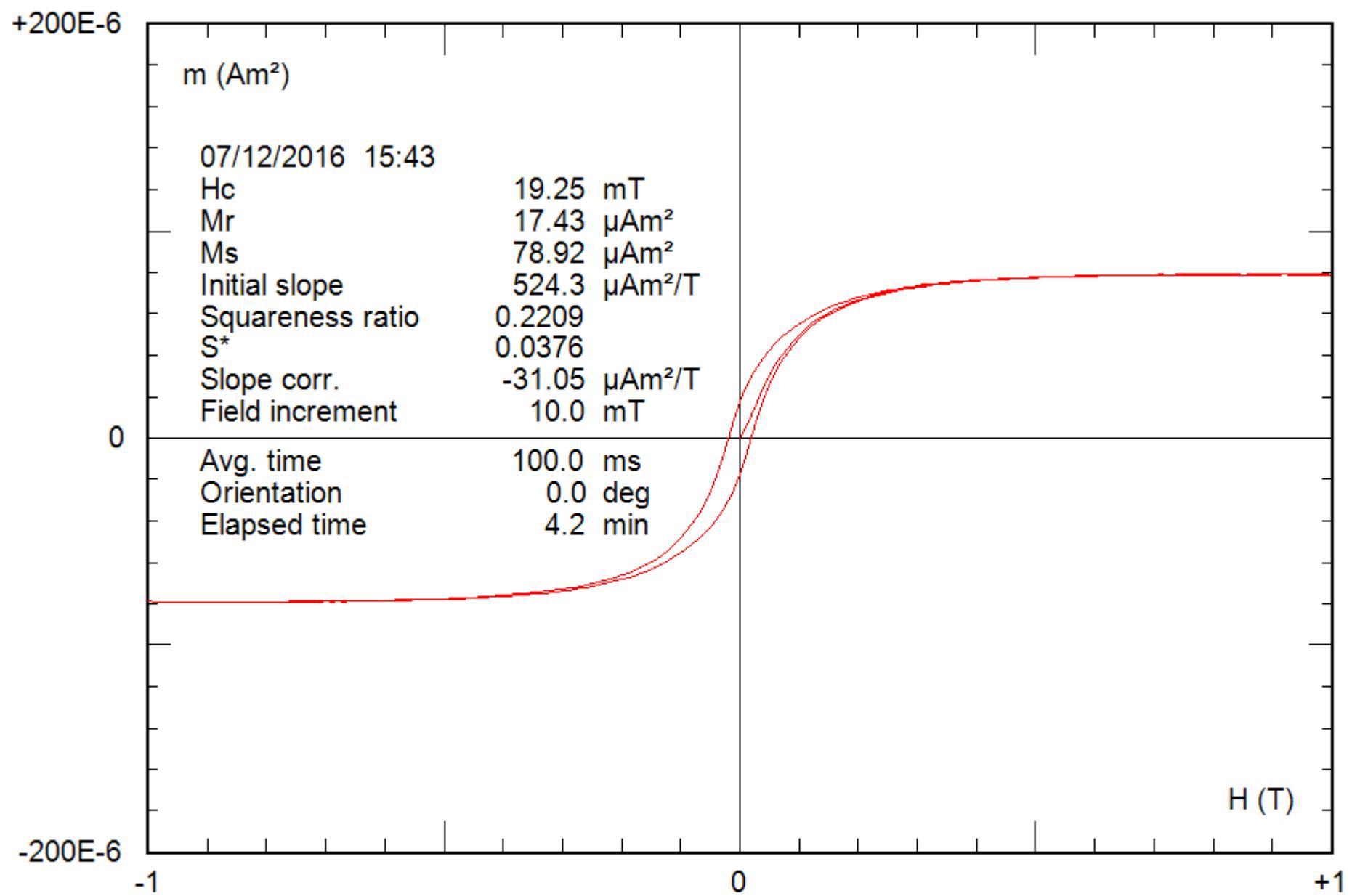
0.28 g

File: 16-04a\_bcr



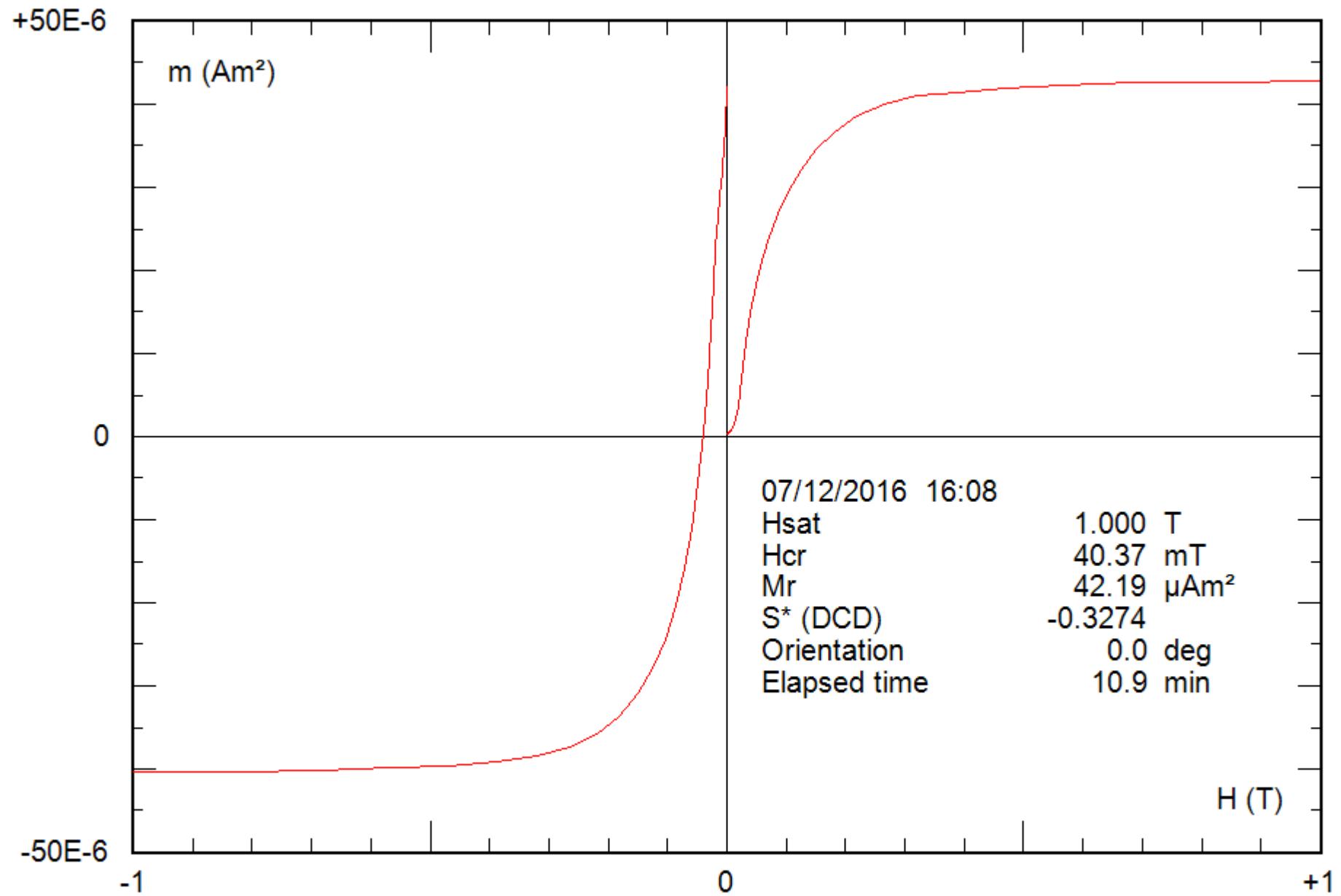
0.28 g

File: 16-04a\_hyst



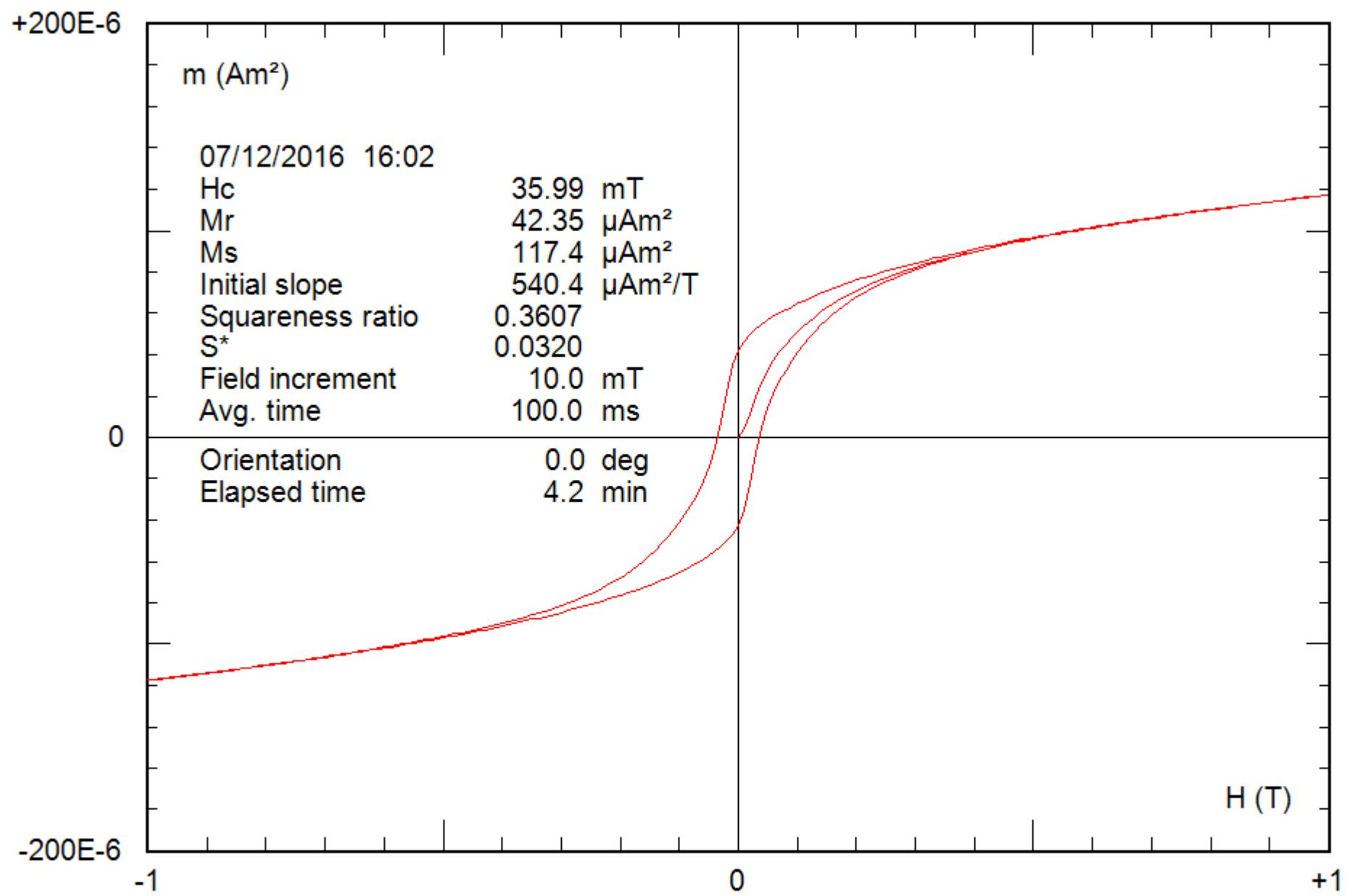
0.28 g

File: 16-04a\_hyst\_corr



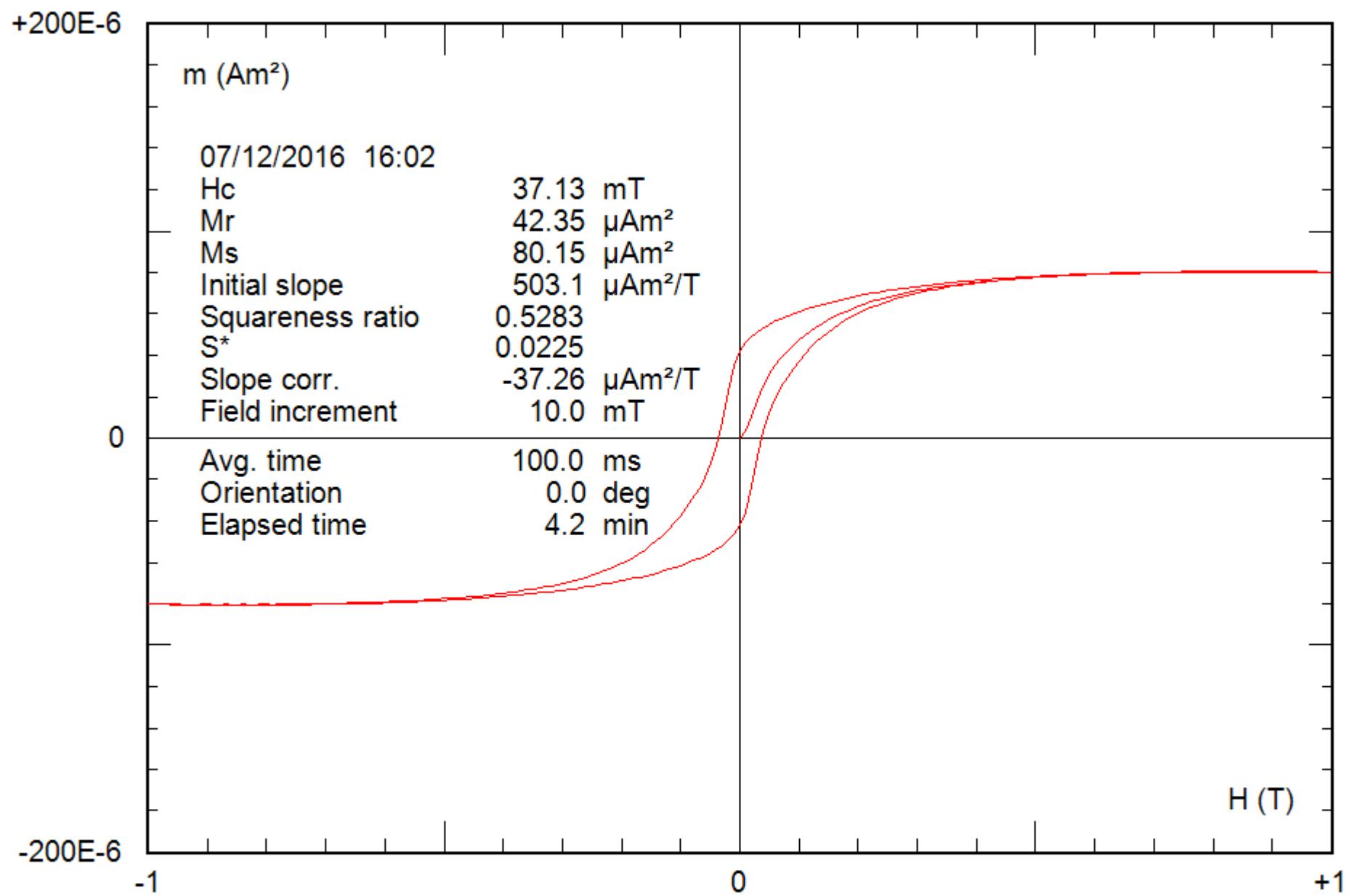
0.25 g

File: 17-04a\_bcr



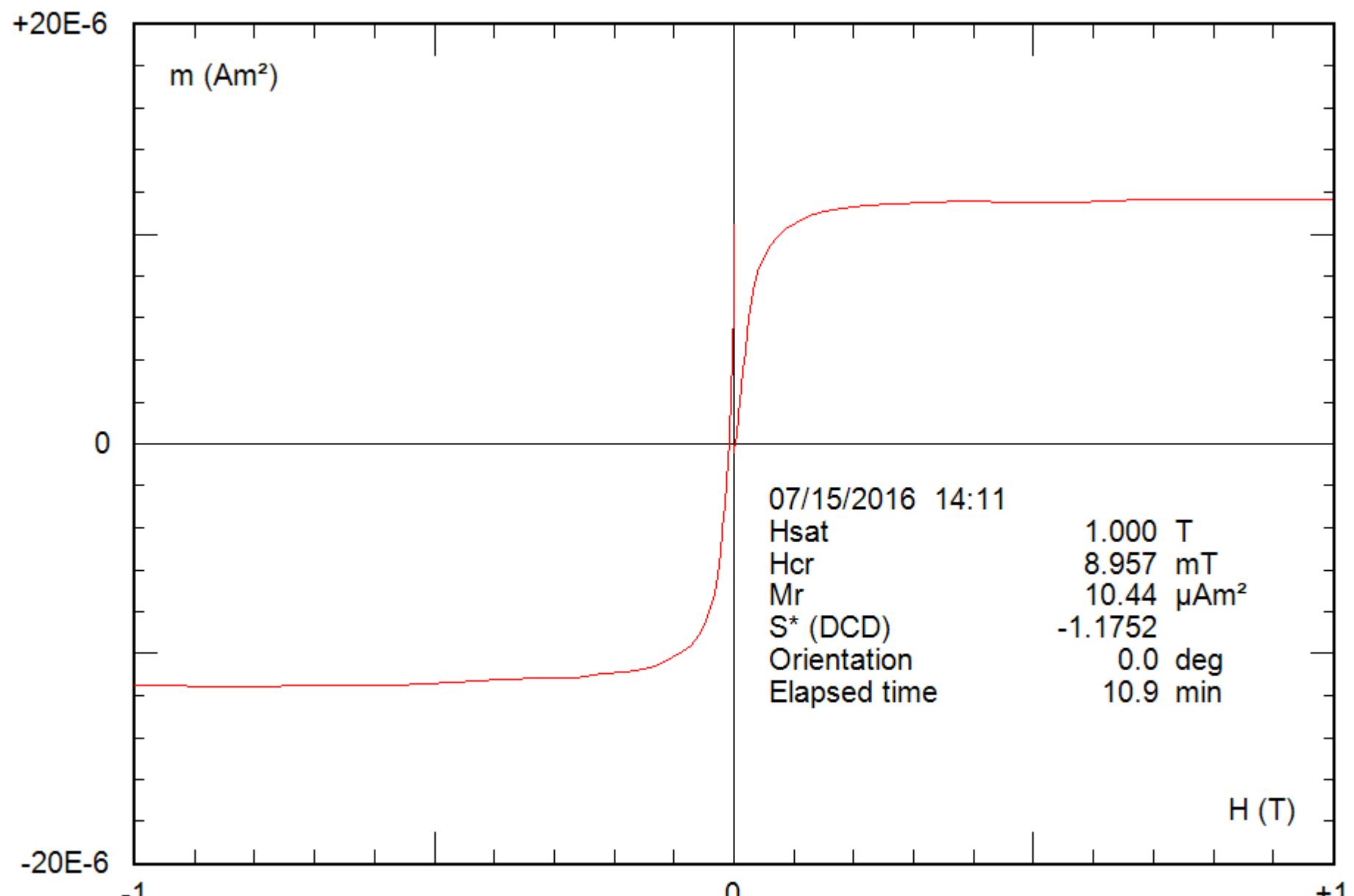
0.25 g

File: 17-04a\_hyst



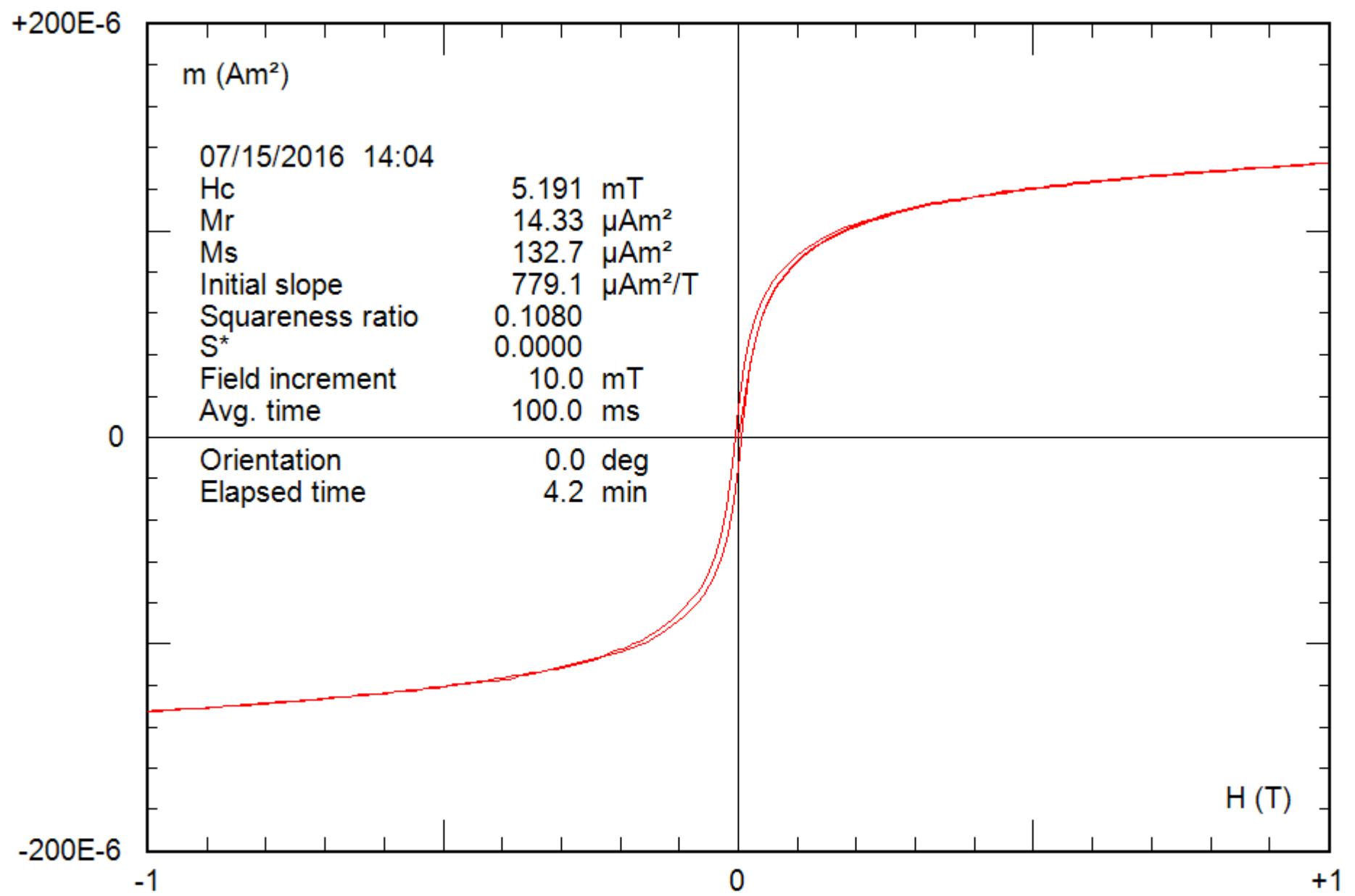
0.25 g

File: 17-04a\_hyst\_corr



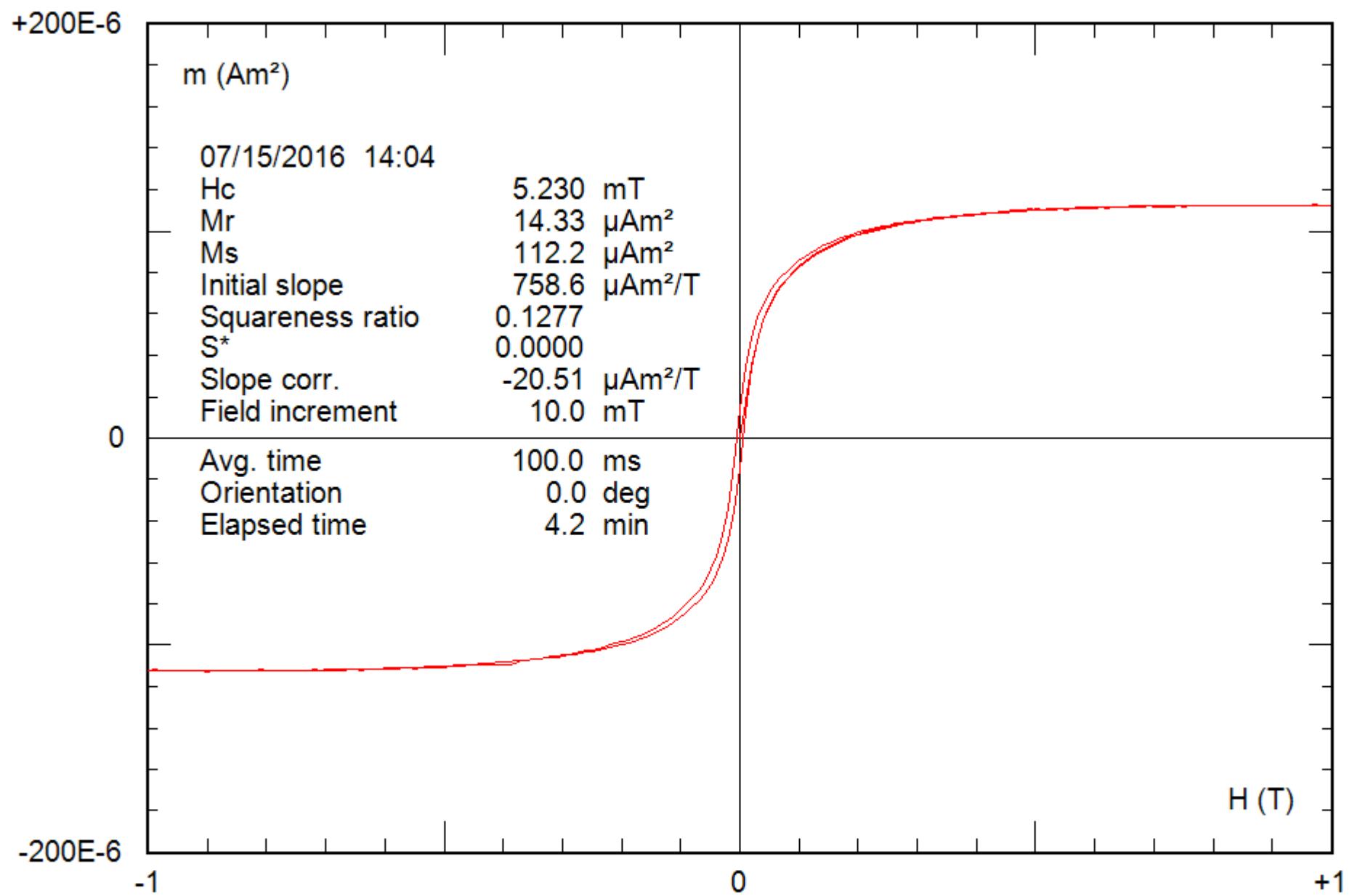
0.19 g

File: 18-04a\_bcr\_bis



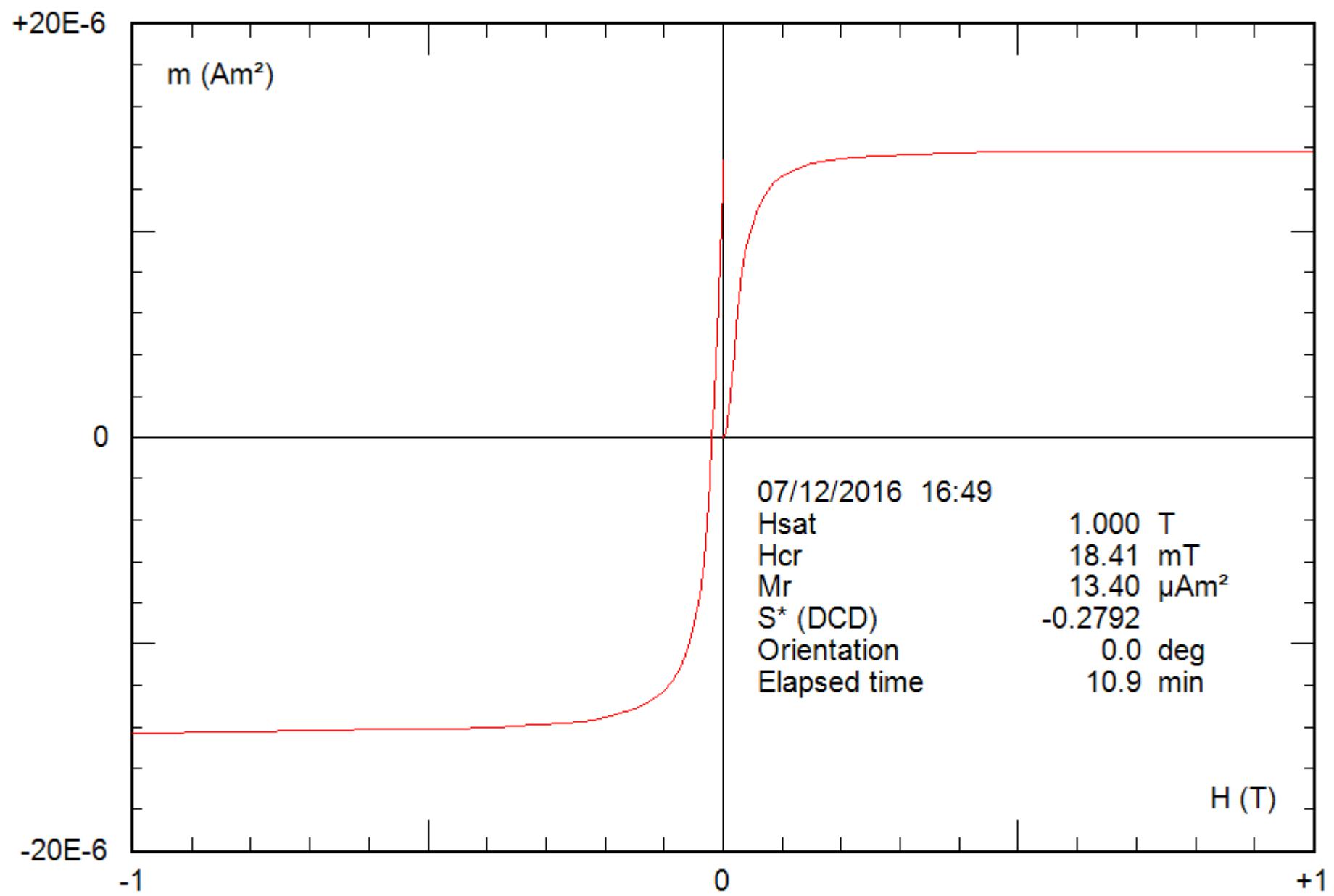
0.19 g

File: 18-04a\_hyst\_bis



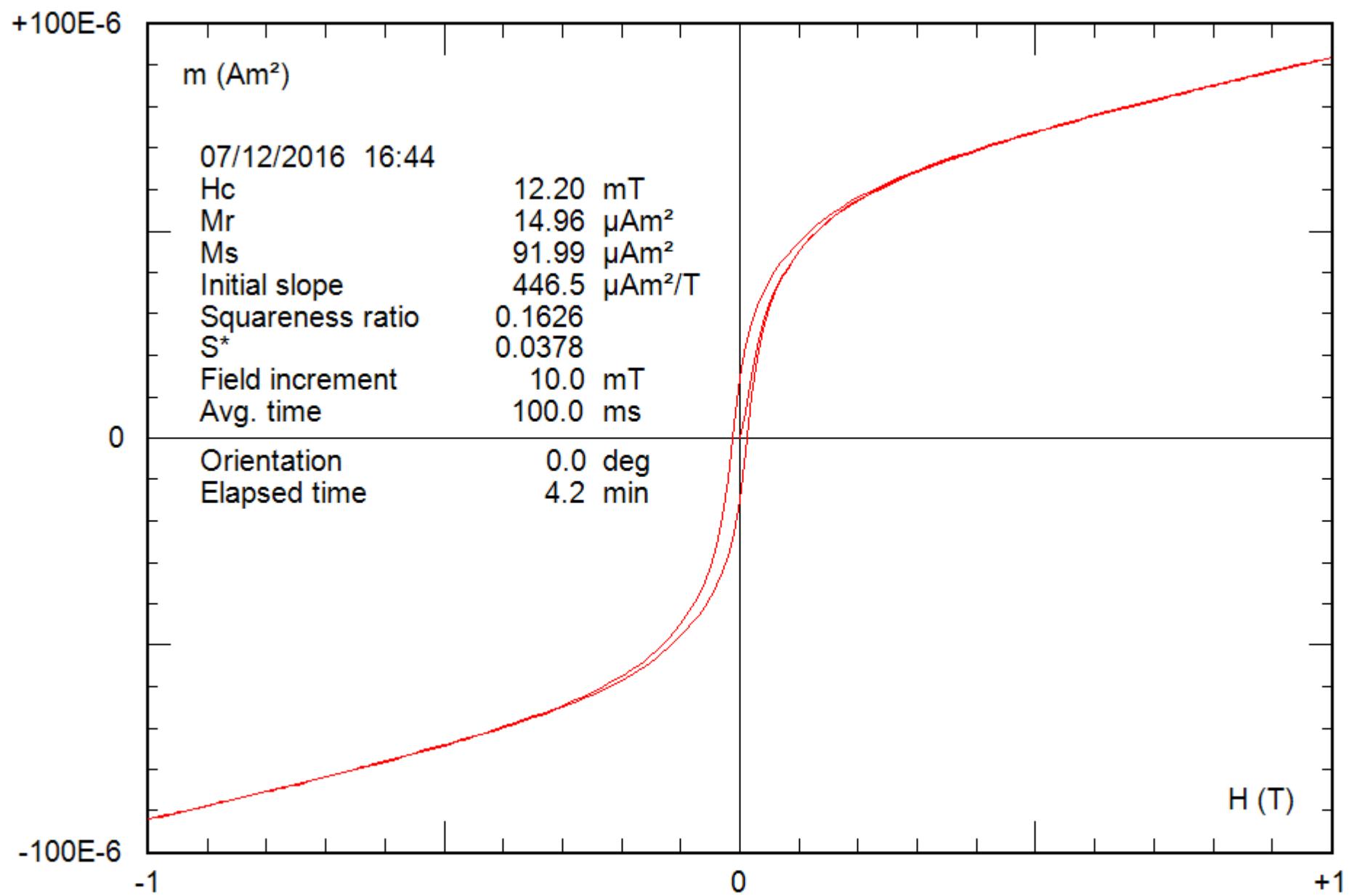
0.19 g

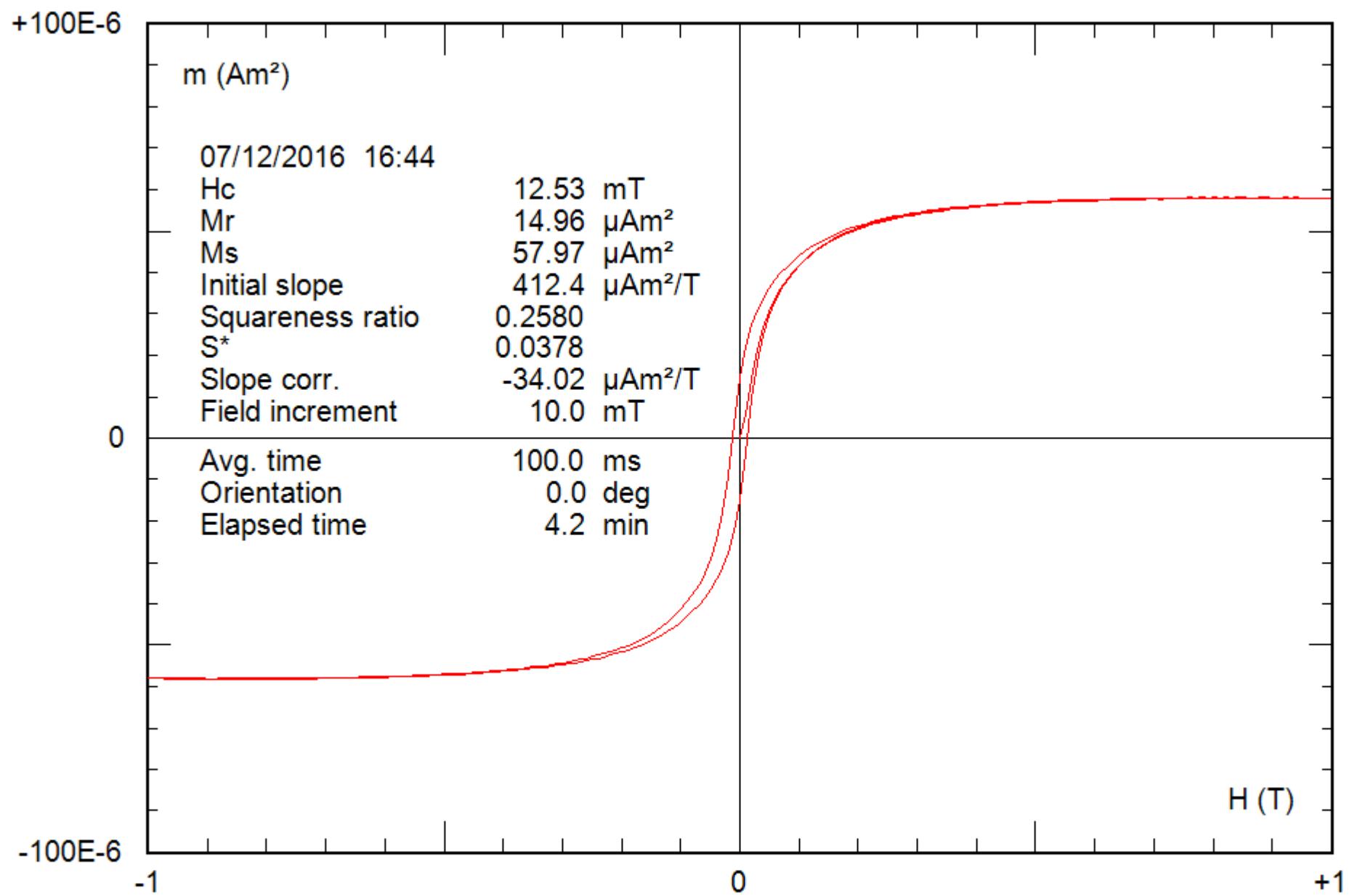
File: 18-04a\_hyst\_corr\_bis



0.25 g

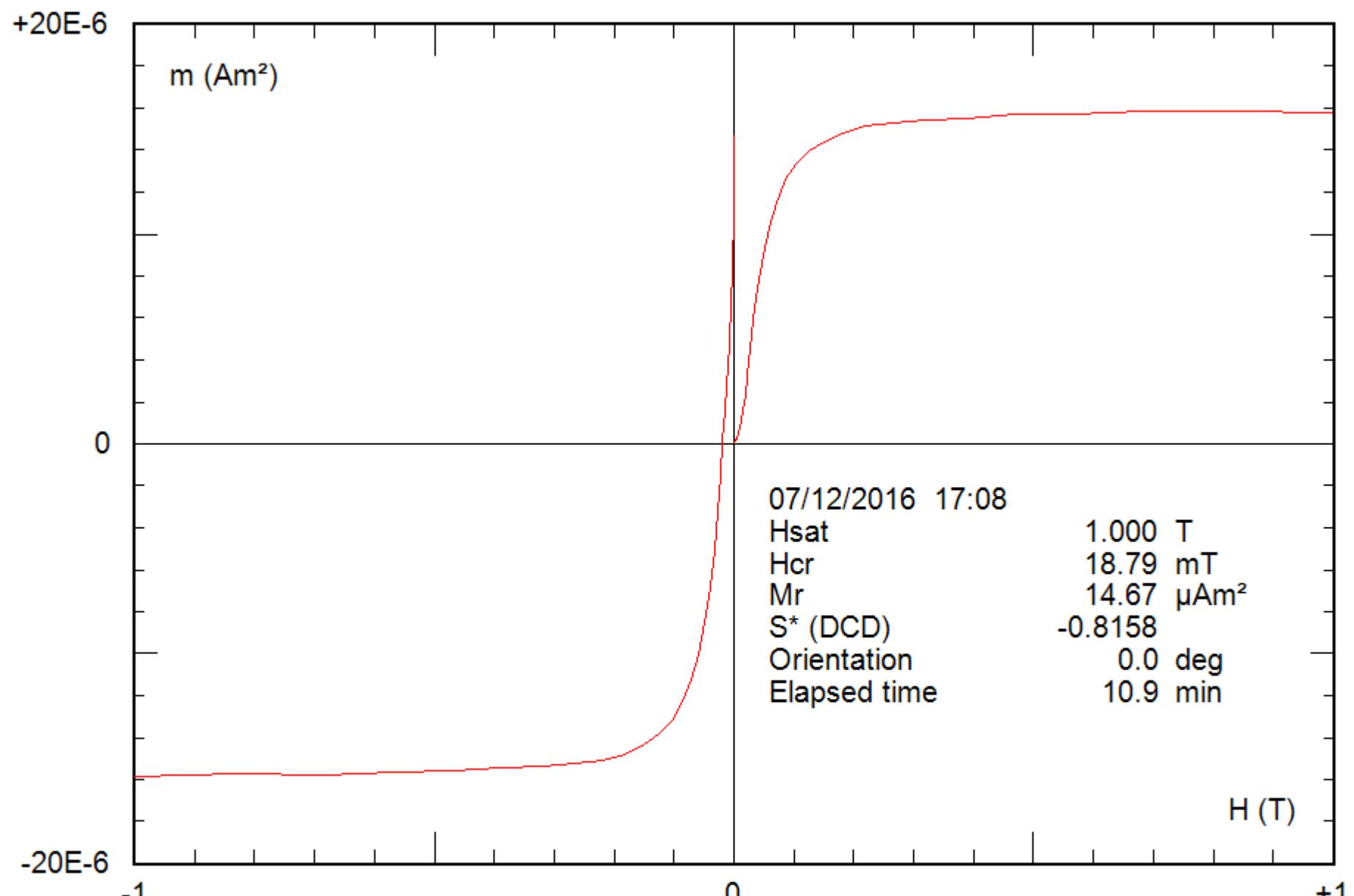
File: 19-04a\_bcr





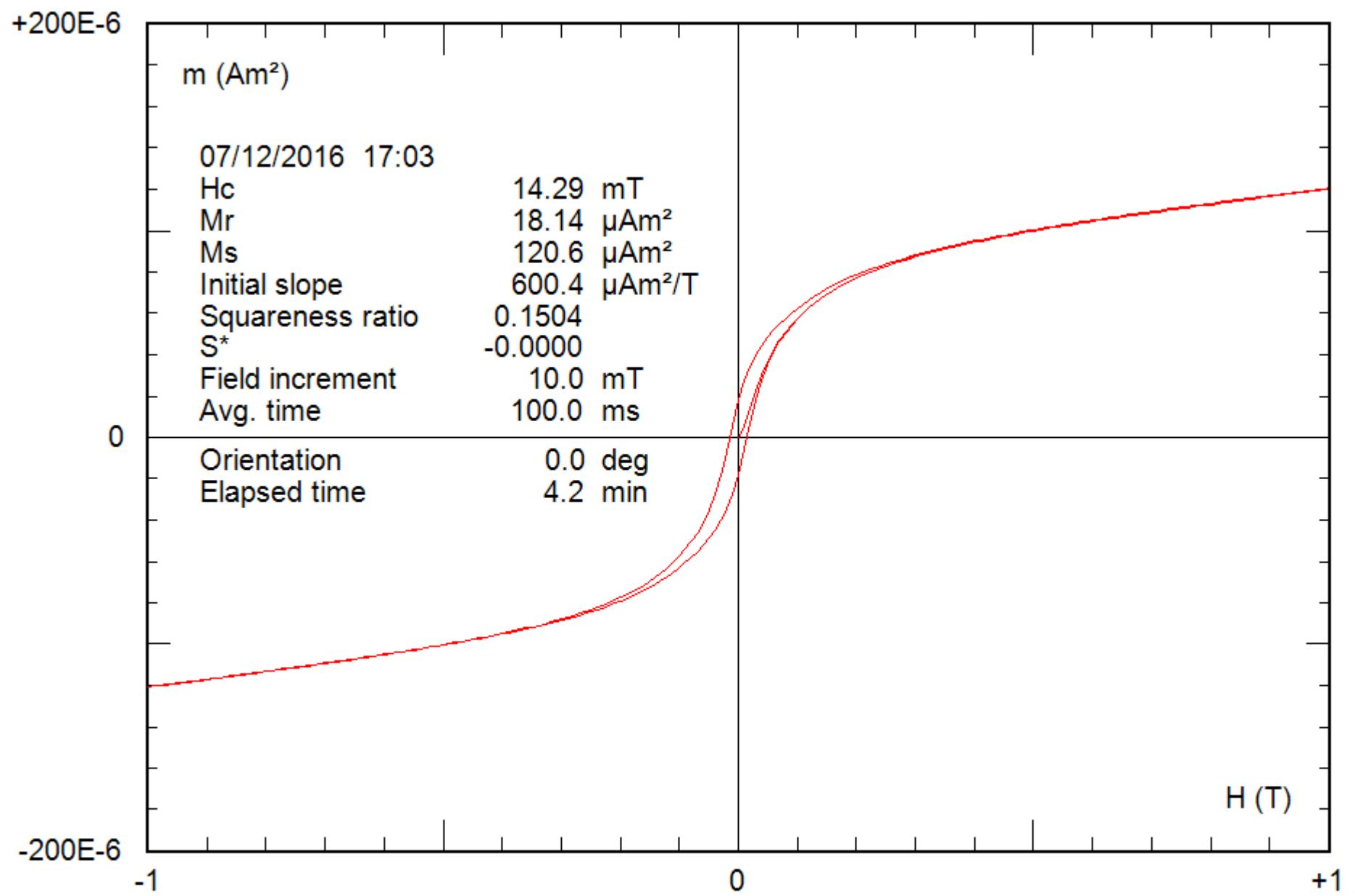
0.25 g

File: 19-04a\_hyst\_corr



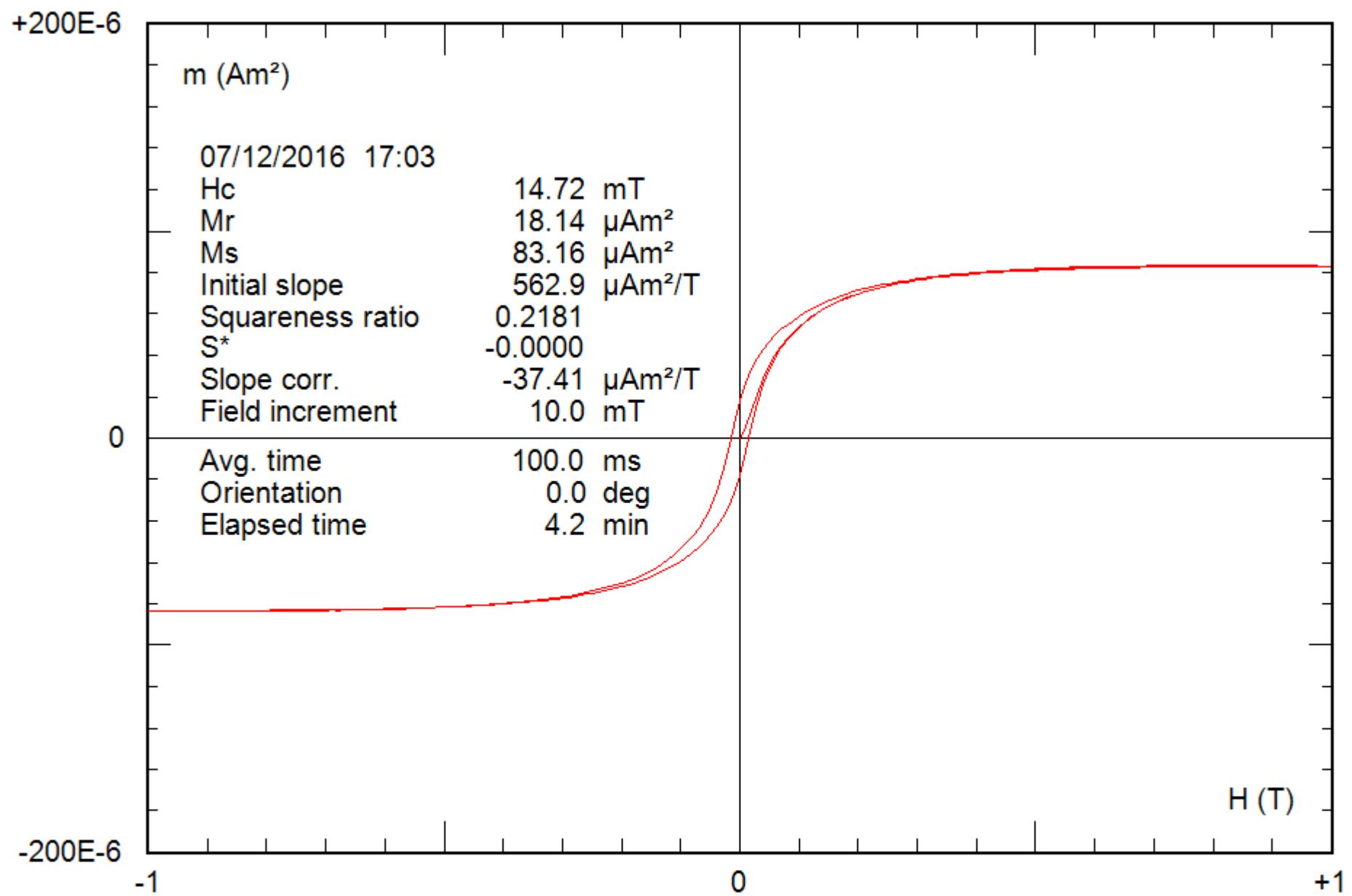
0.25 g

File: 20-04a\_bcr



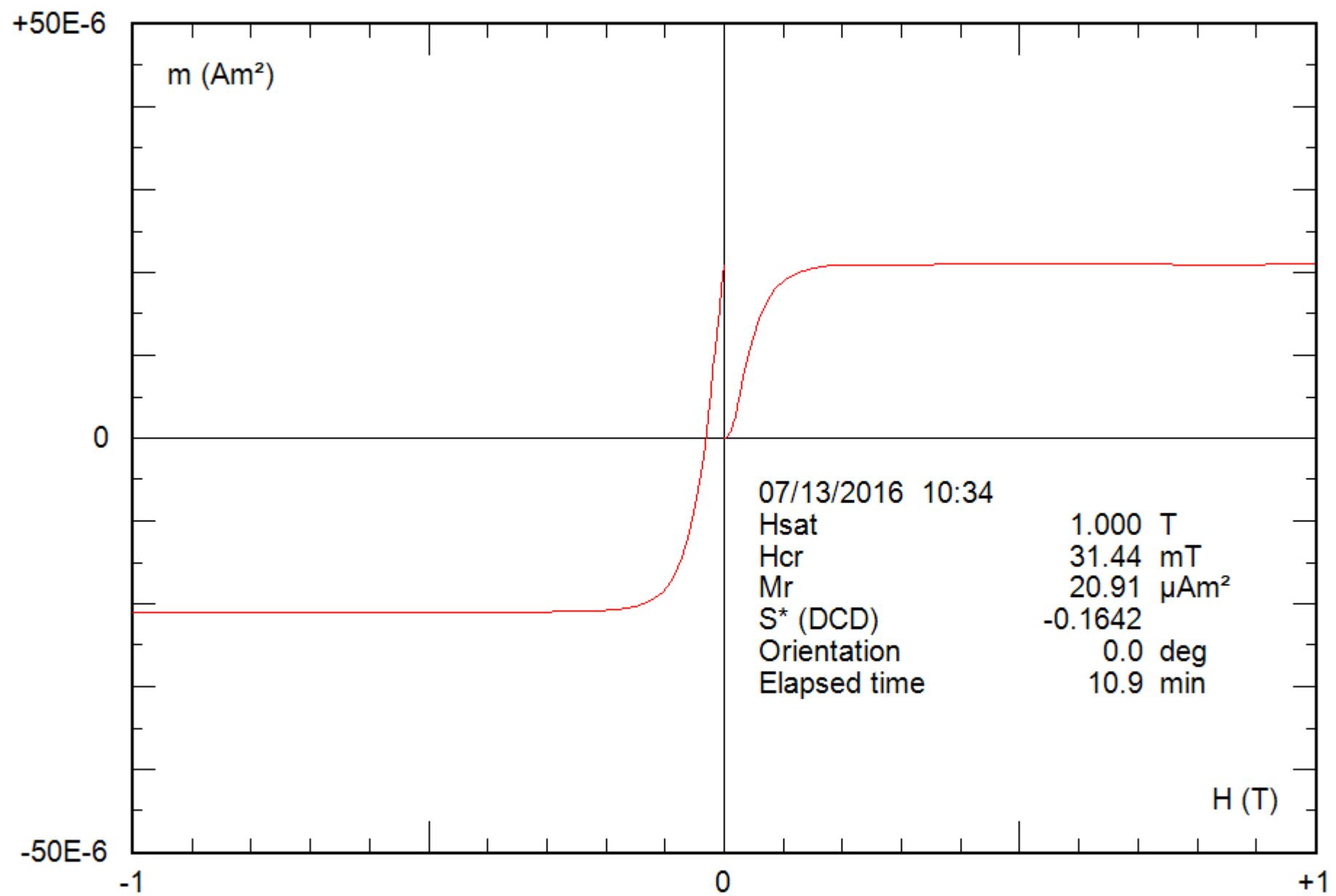
0.25 g

File: 20-04a\_hyst



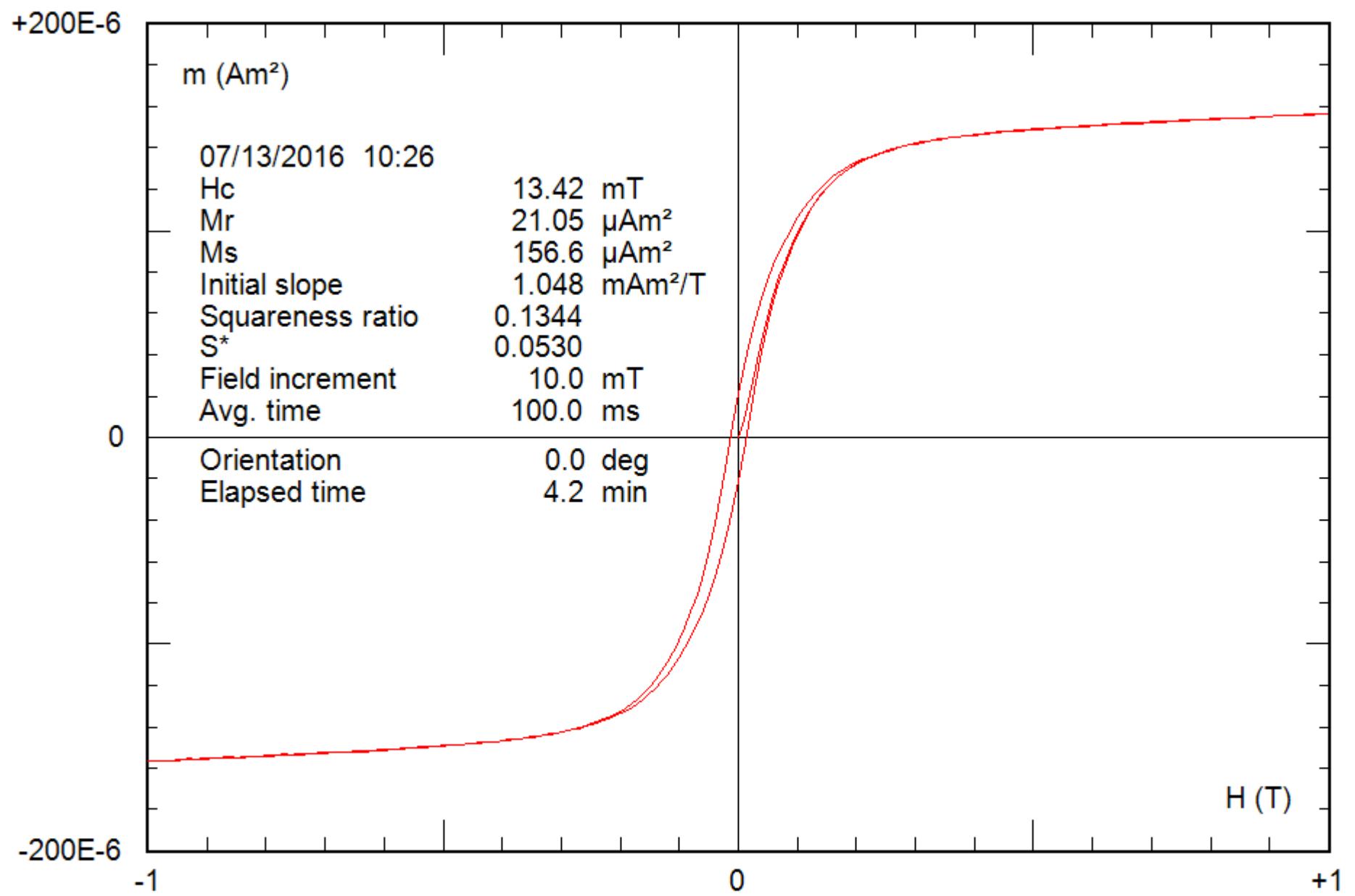
0.25 g

File: 20-04a\_hyst\_corr



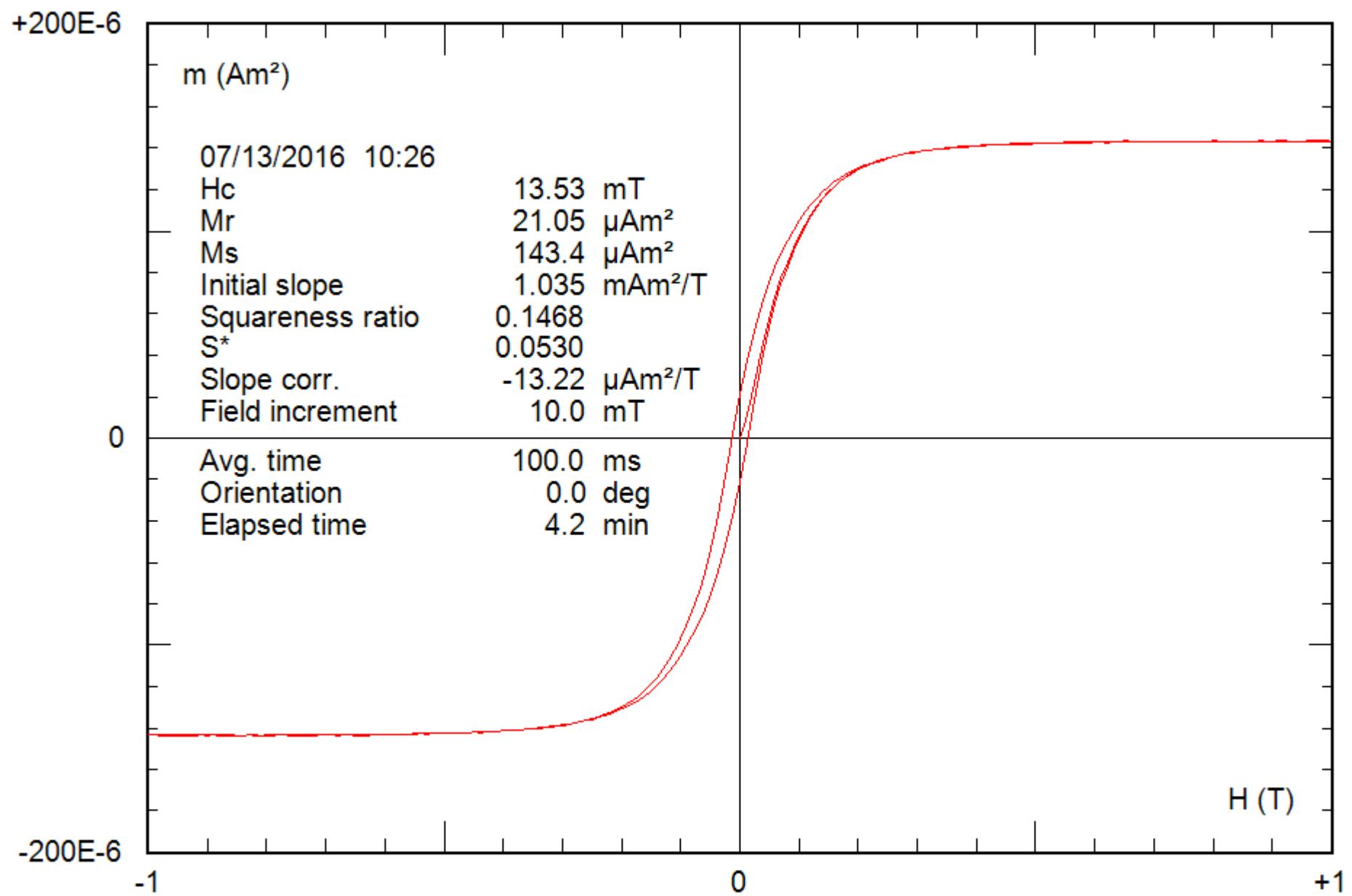
0.21 g

File: 21-04a\_bcr



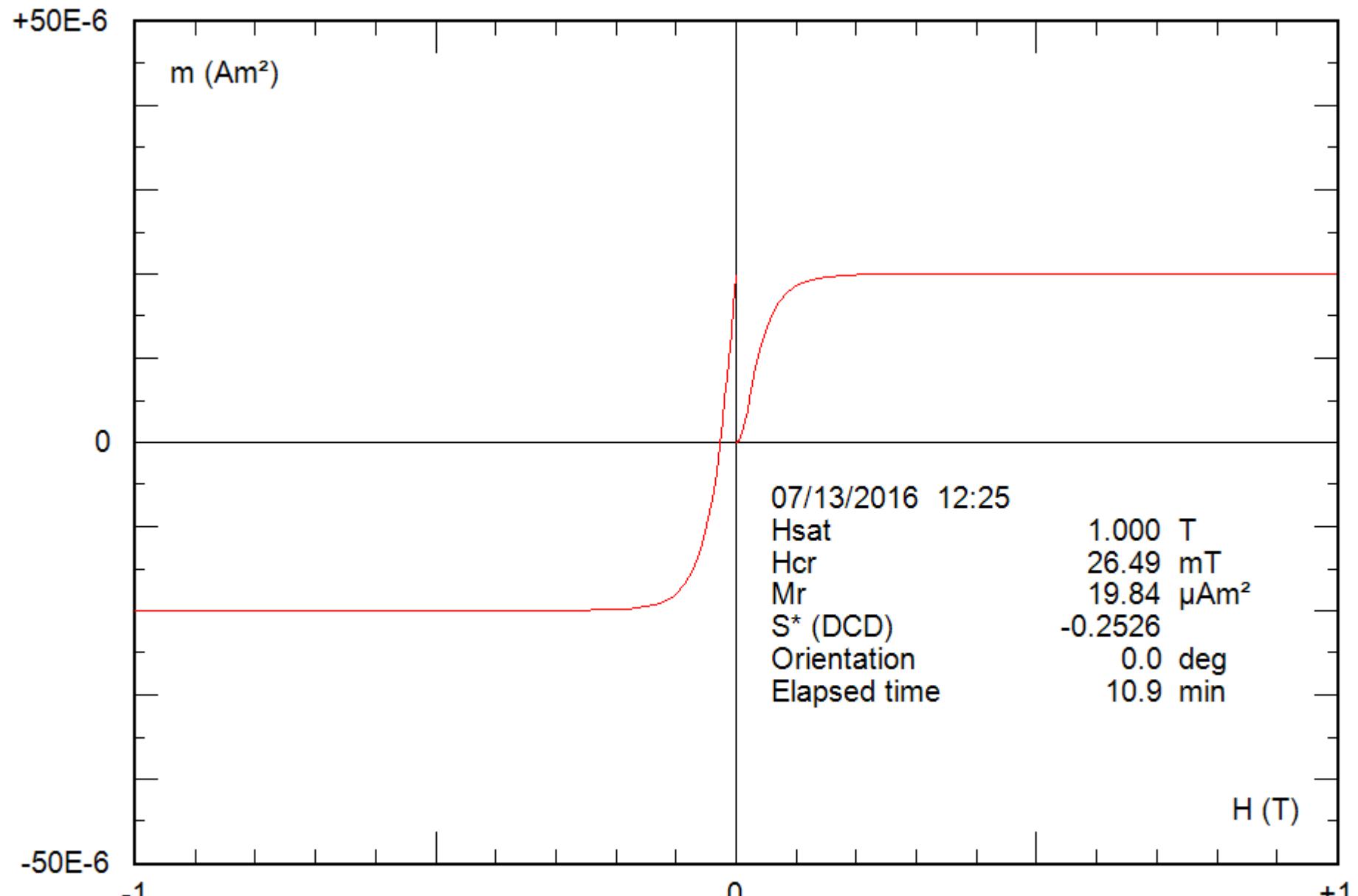
0.21 g

File: 21-04a\_hyst



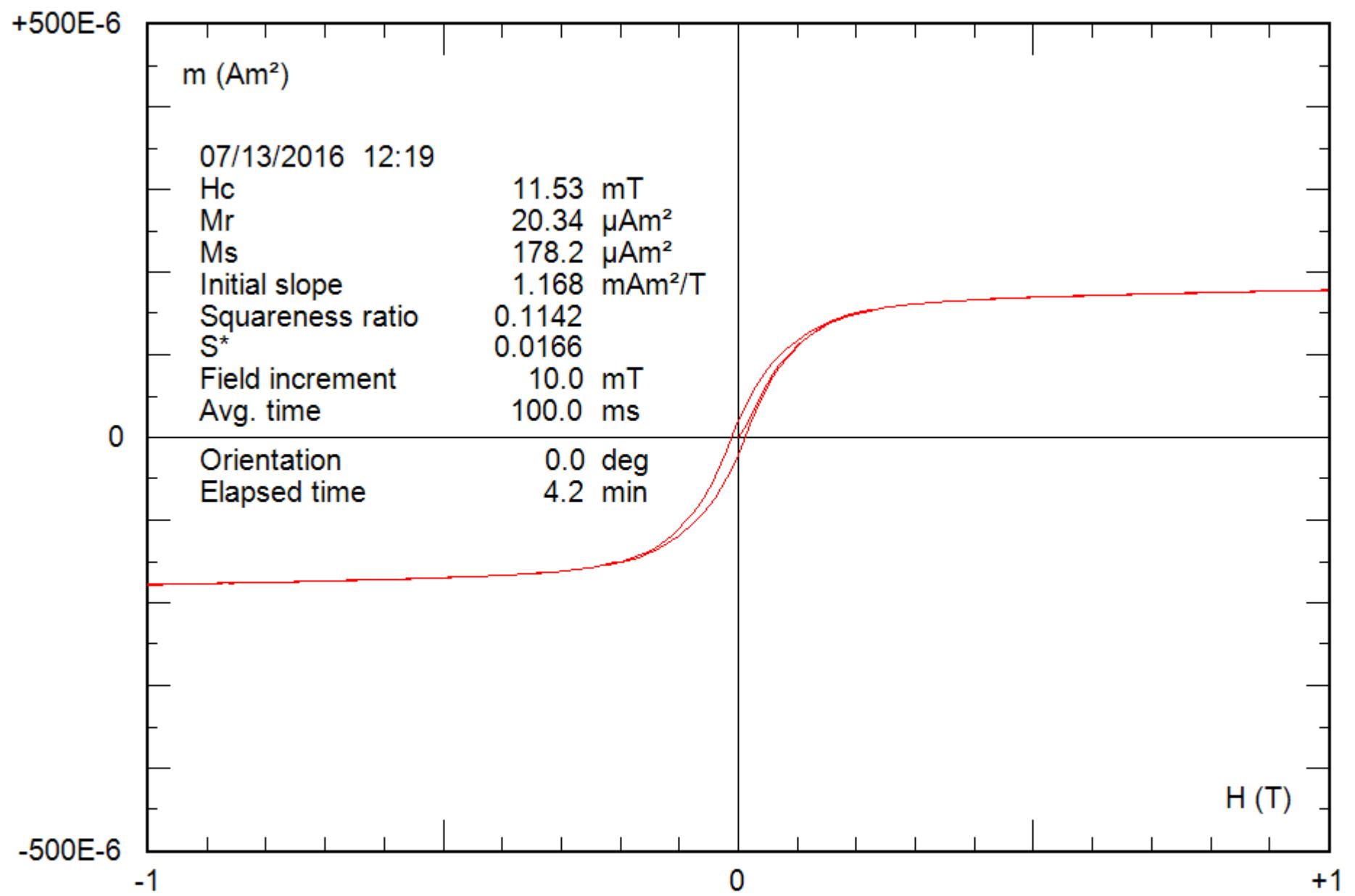
0.21 g

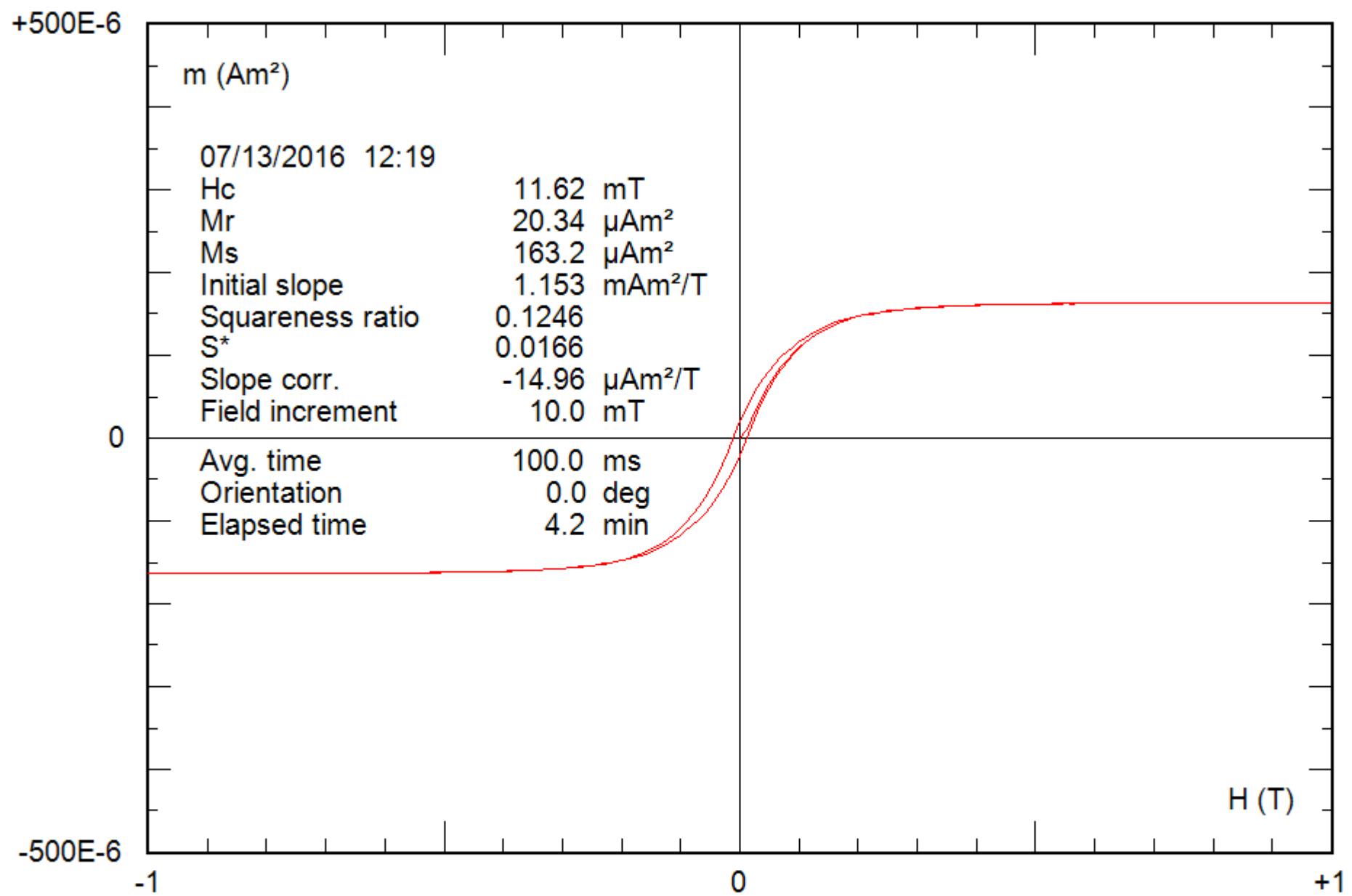
File: 21-04a\_hyst\_corr



0.24 g

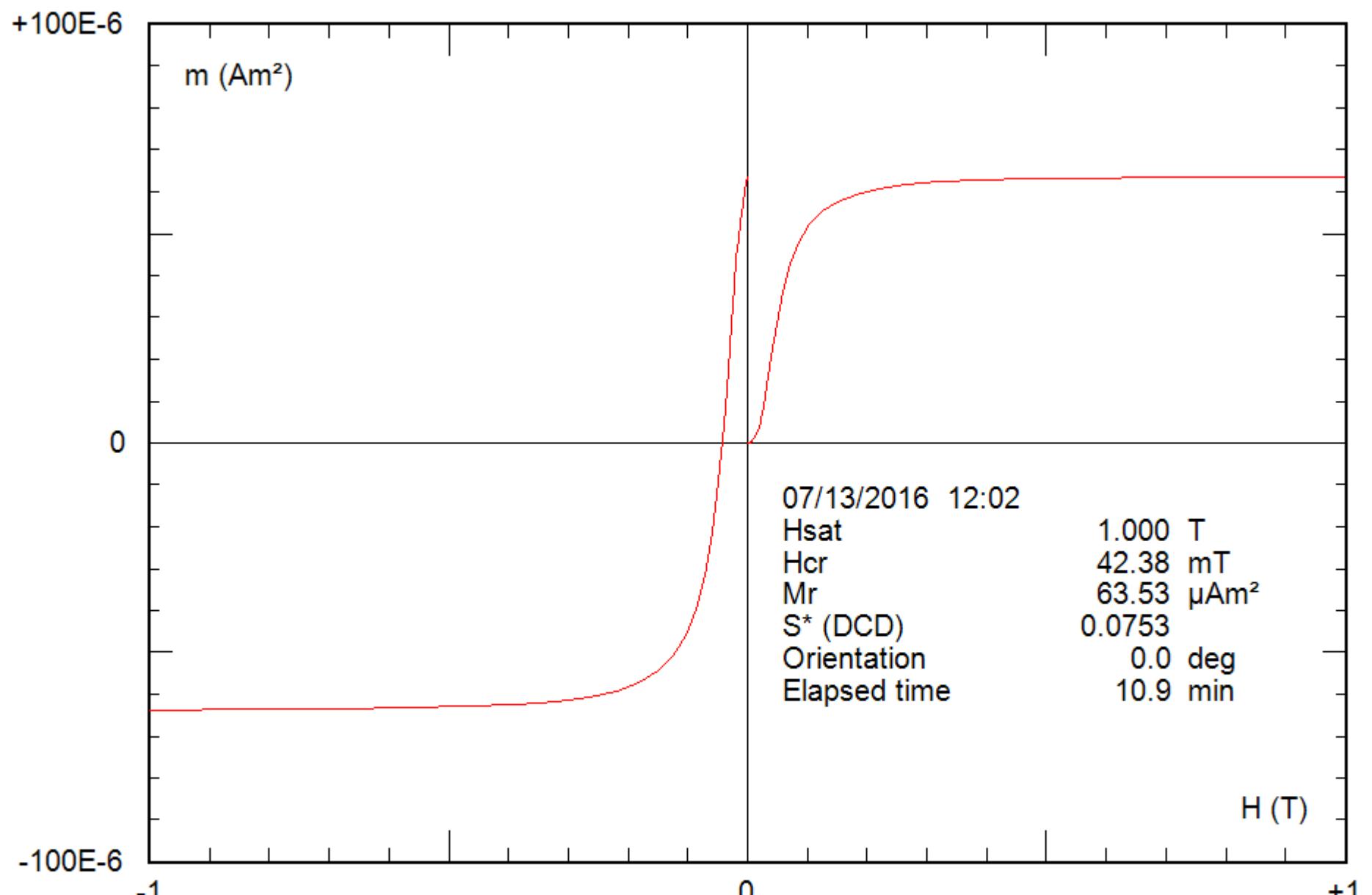
File: 22-04a\_bis bcr





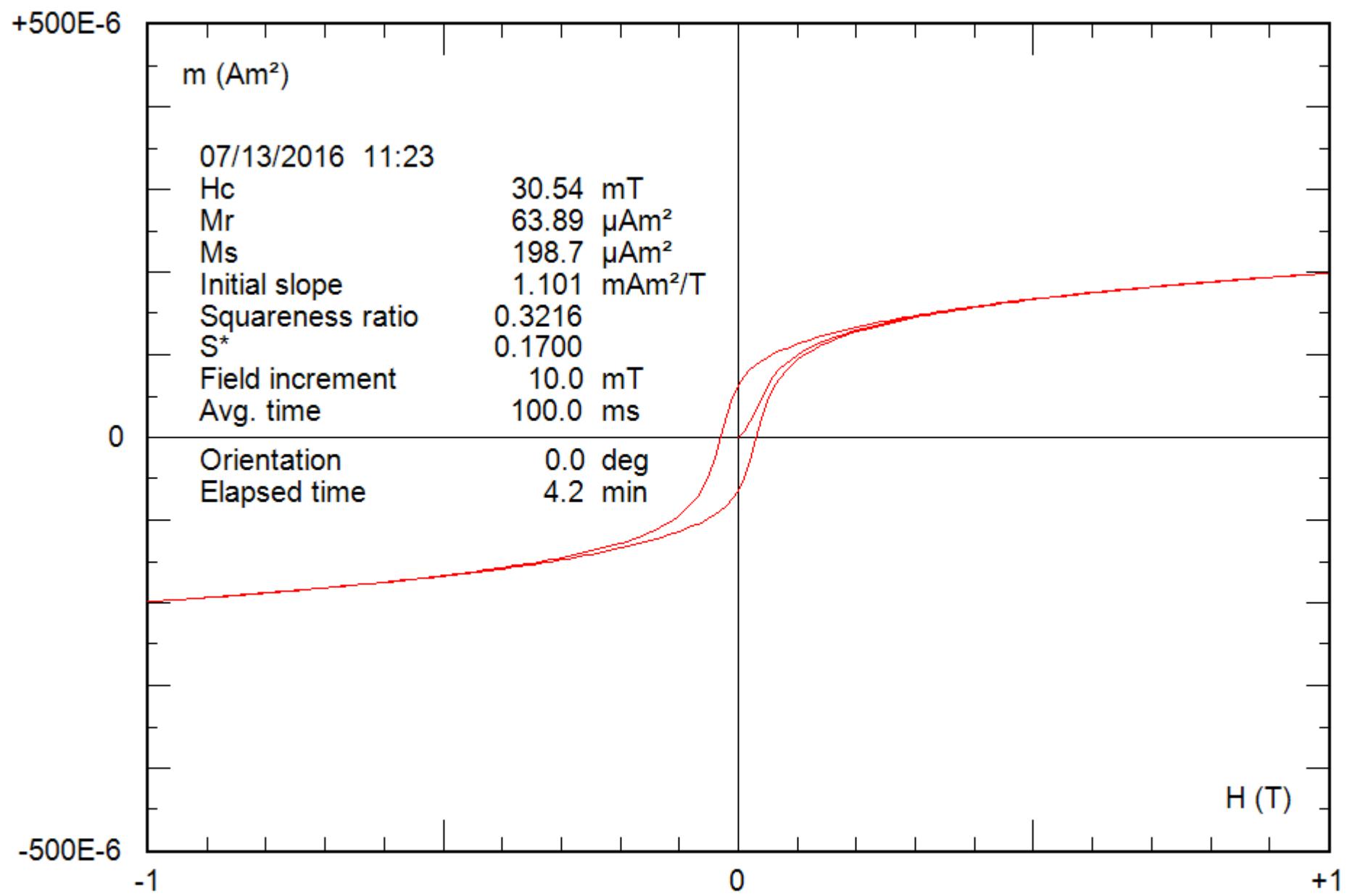
0.24 g

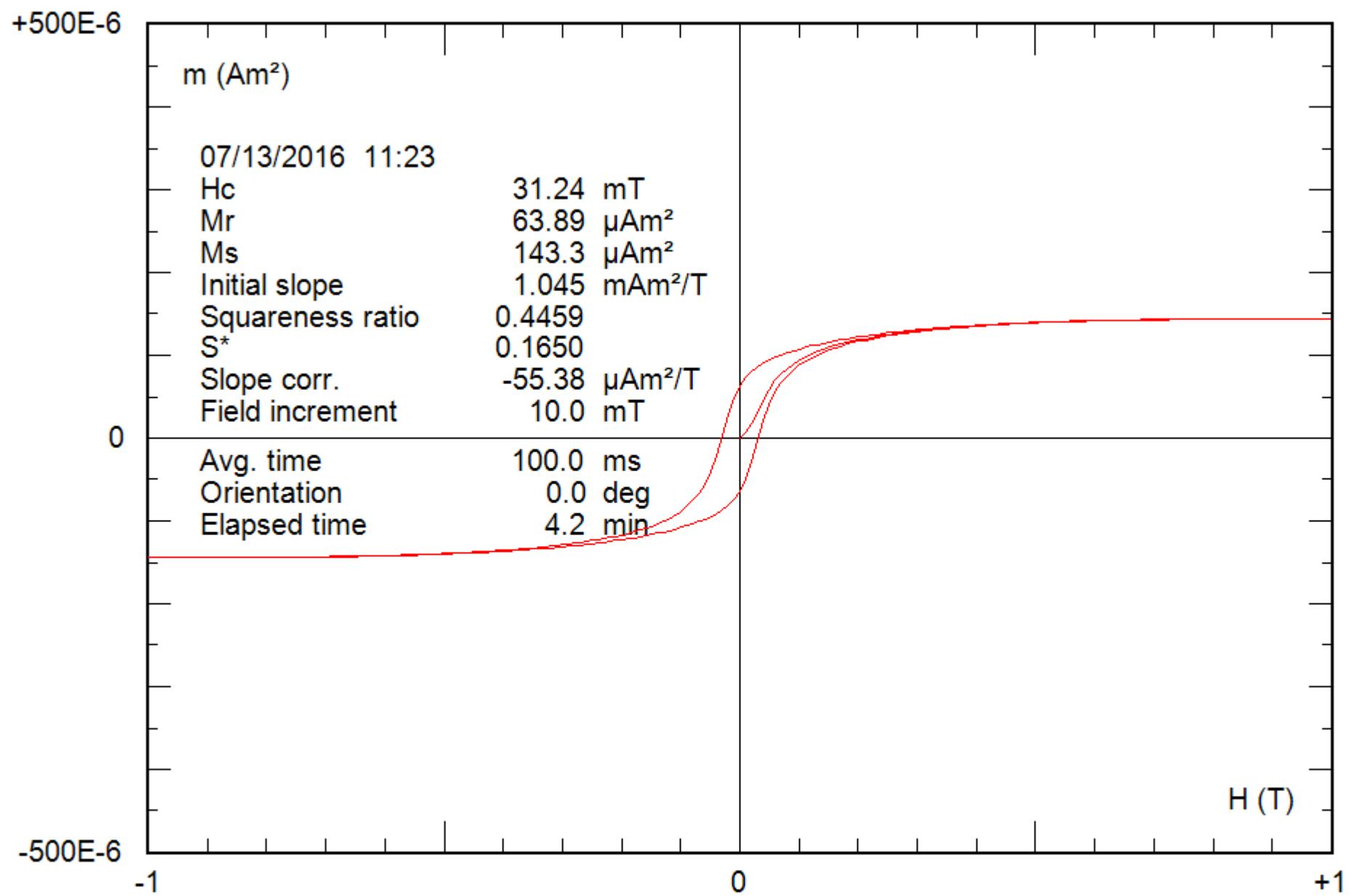
File: 22-04a\_bis hyst\_corr



0.26 g

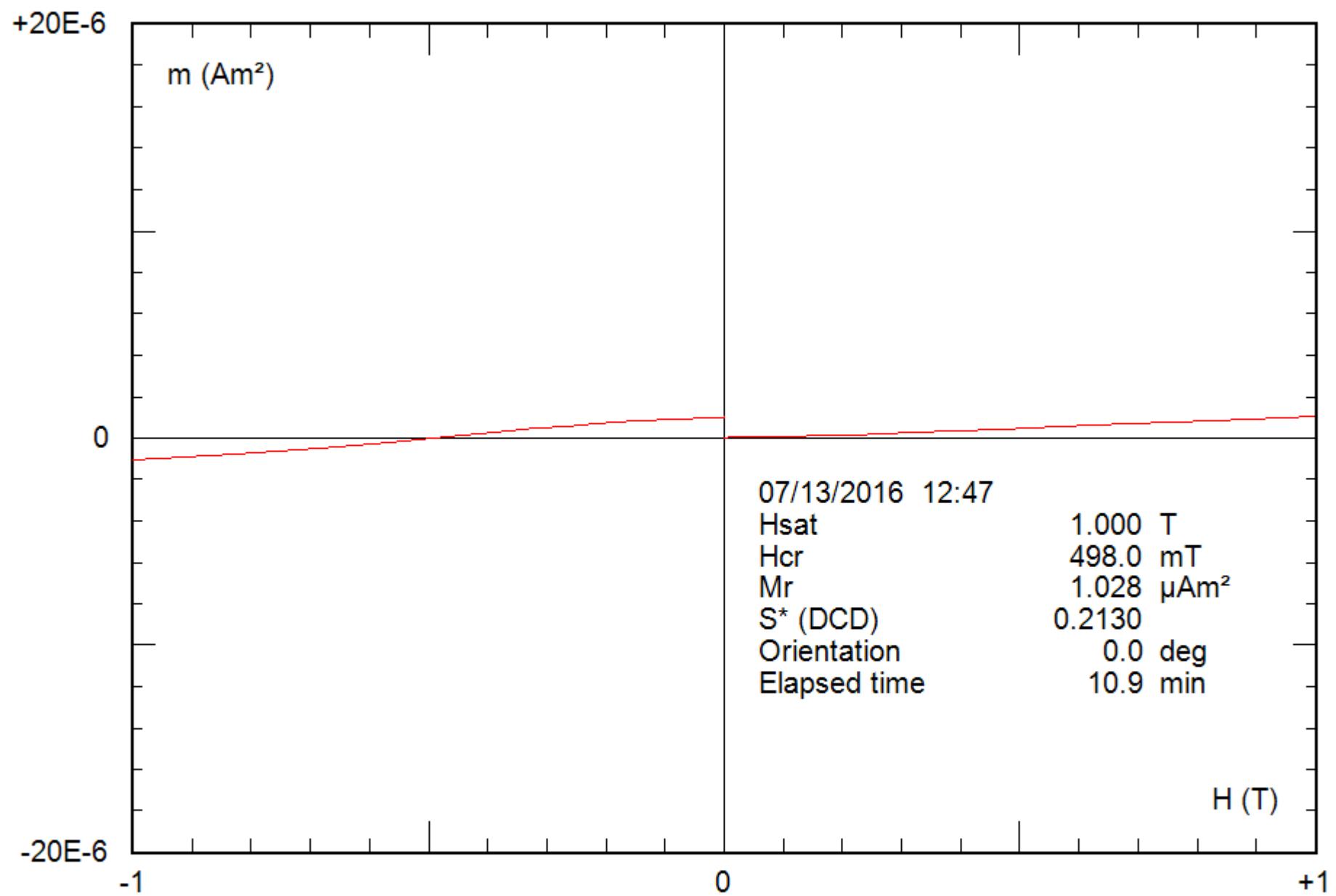
File: 23-04a\_bcr





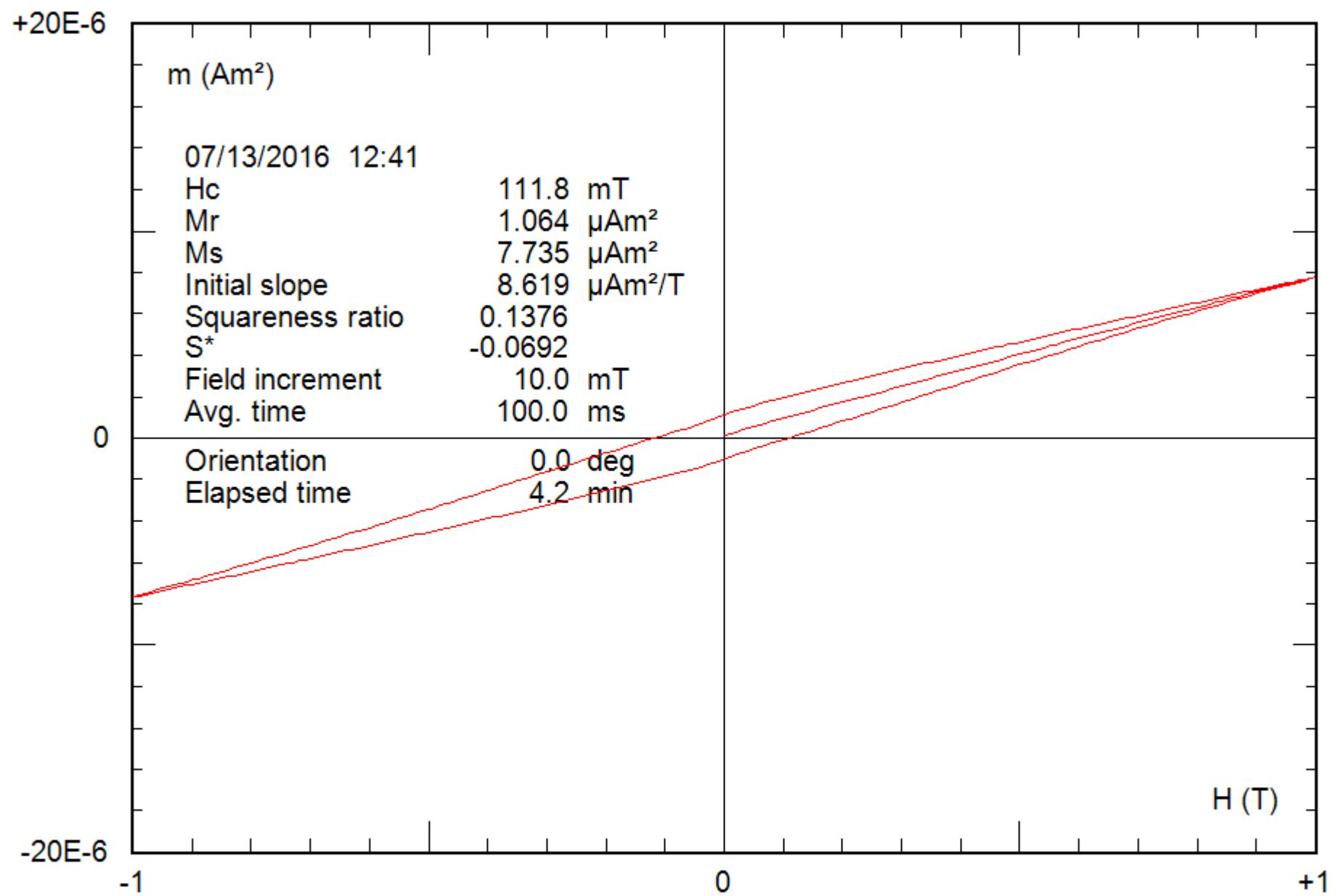
0.26 g

File: 23-04a\_hyst\_corr



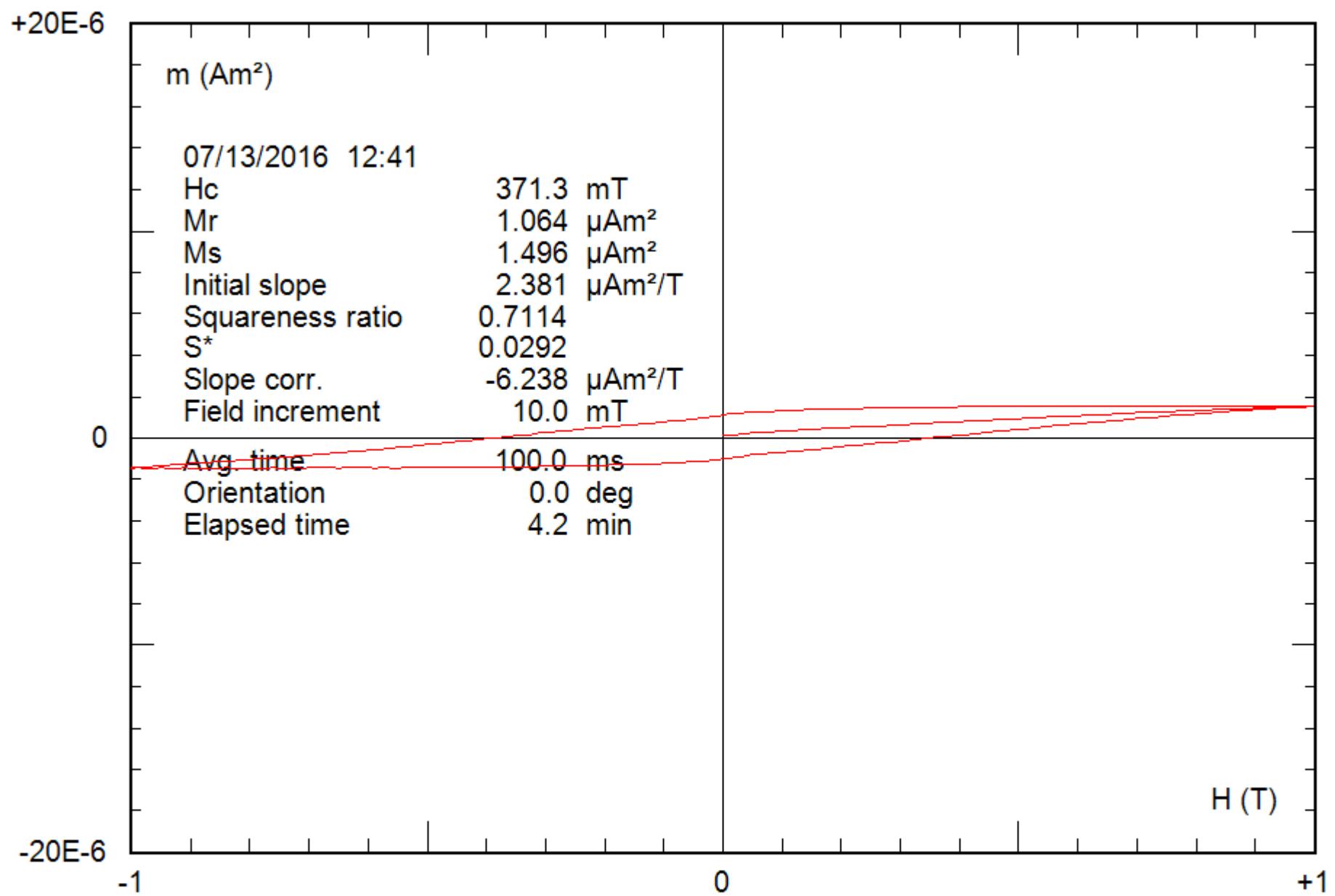
0.24 g

File: 24-08b\_bcr



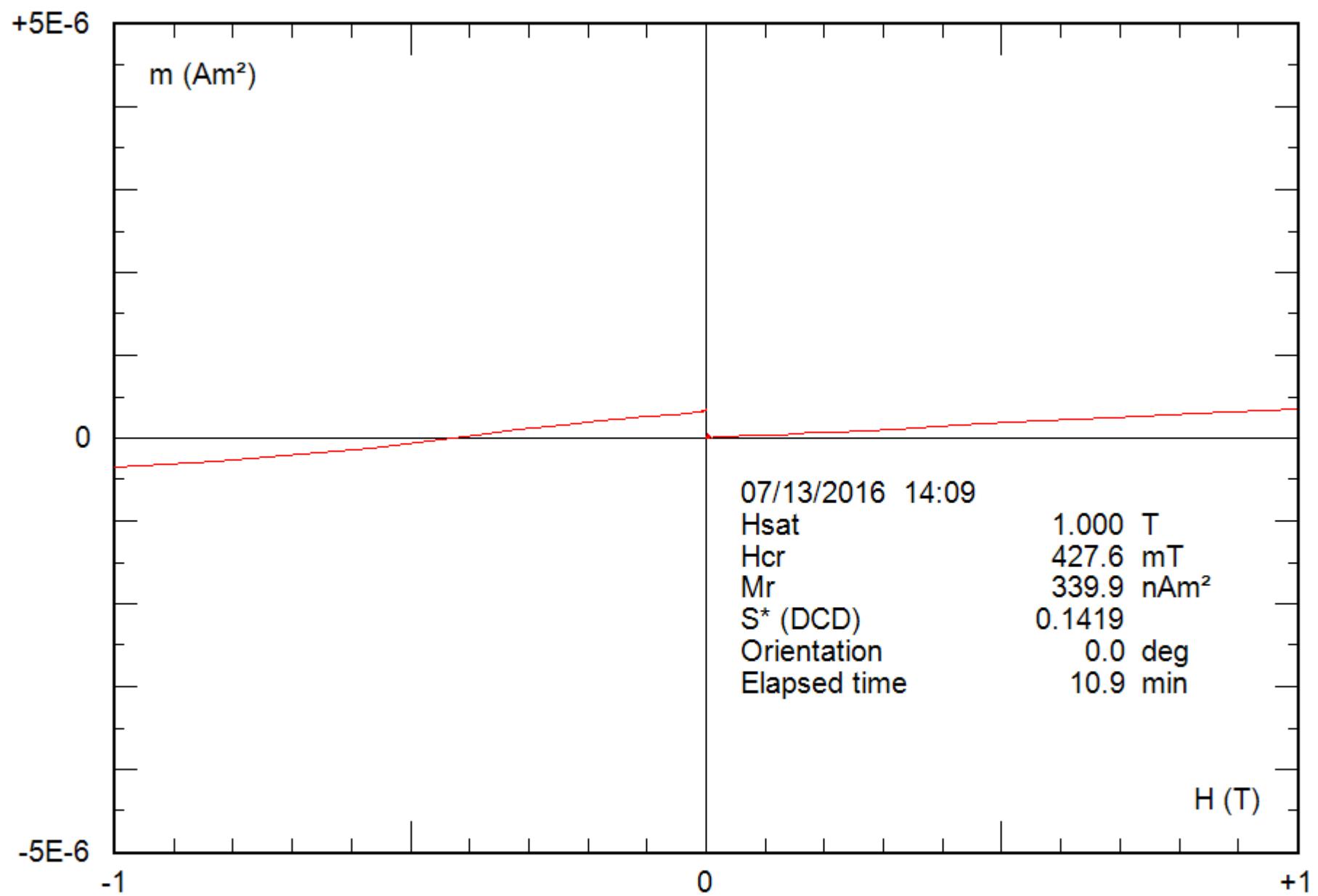
0.24 g

File: 24-08b\_hyst



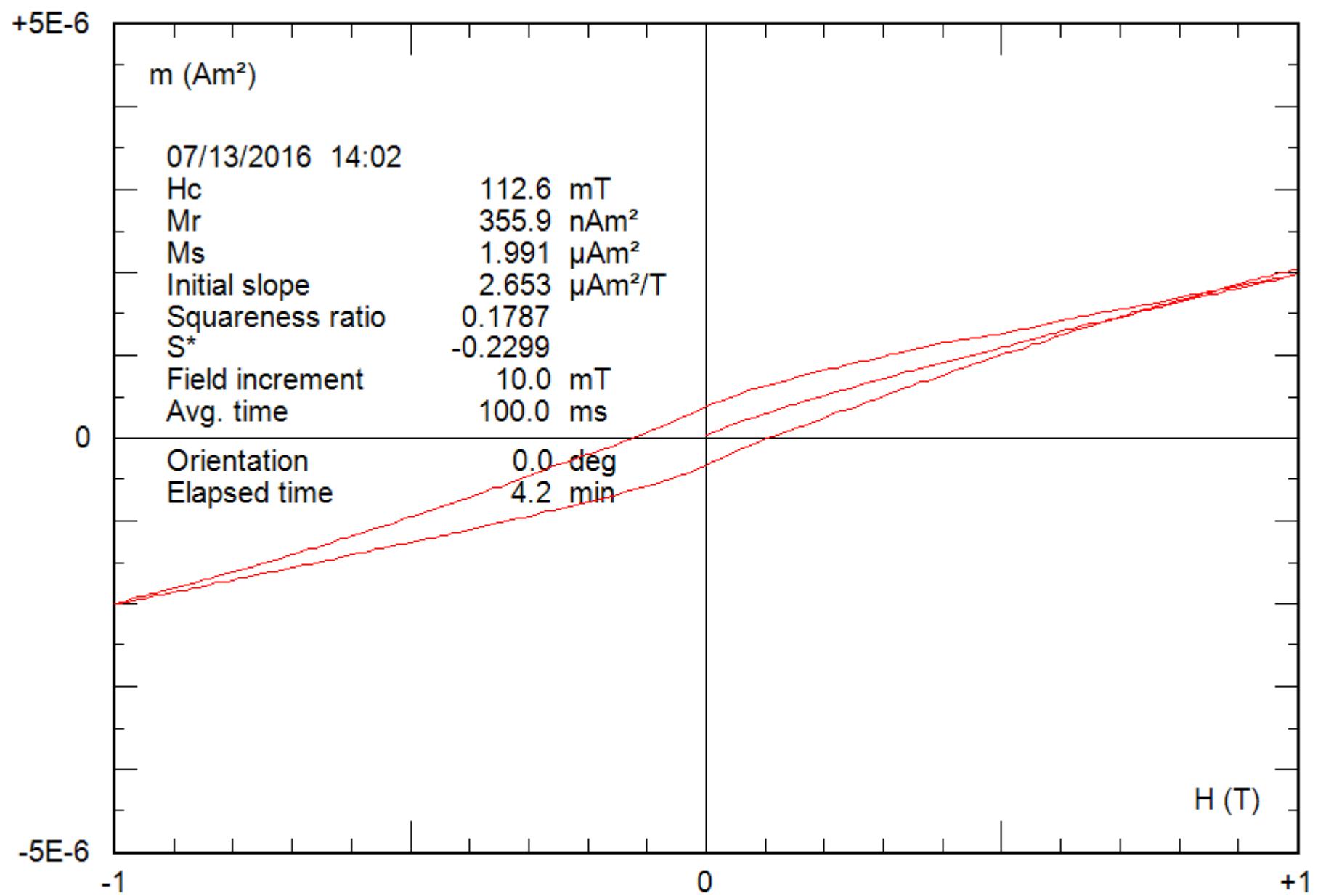
0.24 g

File: 24-08b\_hyst\_corr



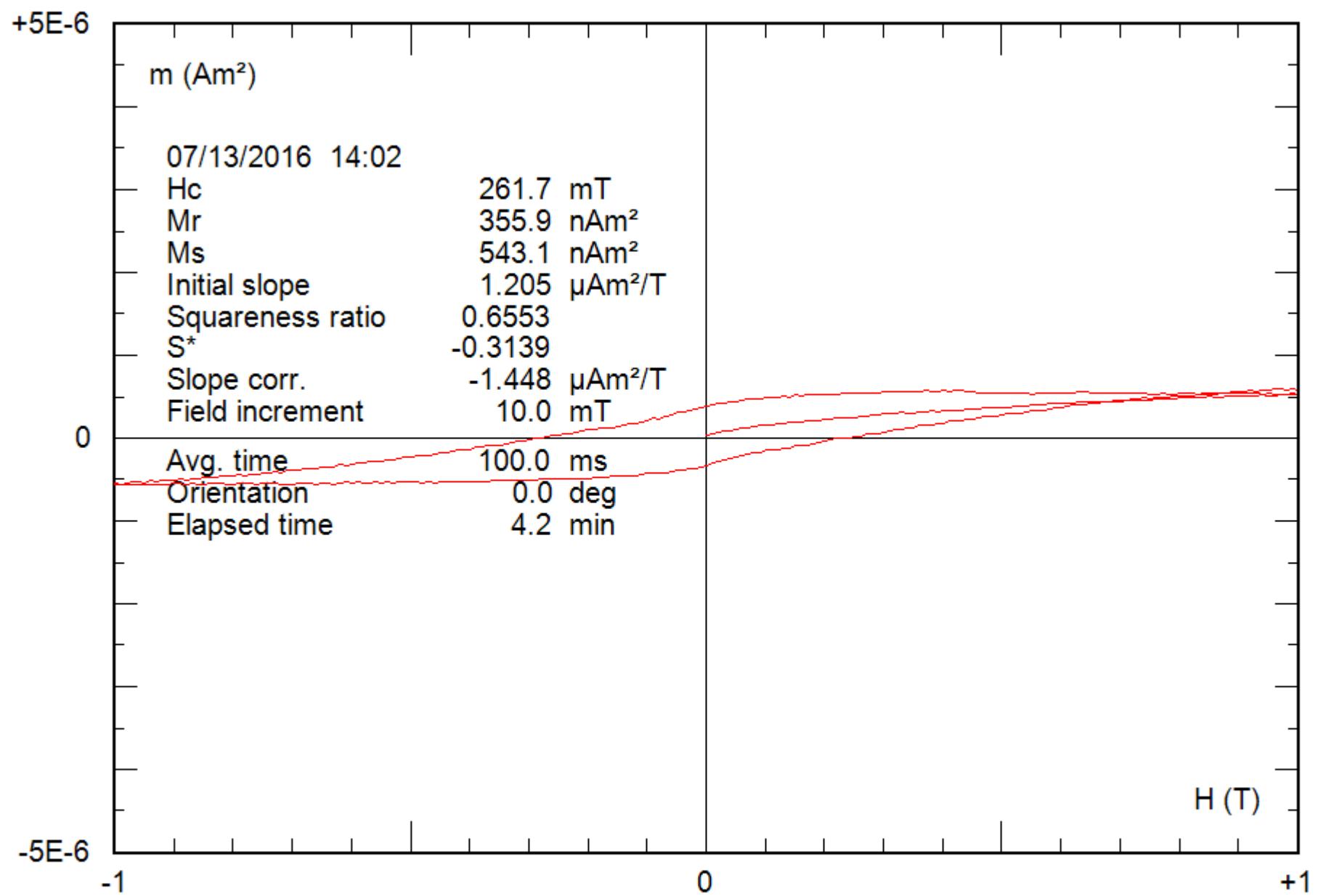
0.26 g

File: 25-04c\_bcr



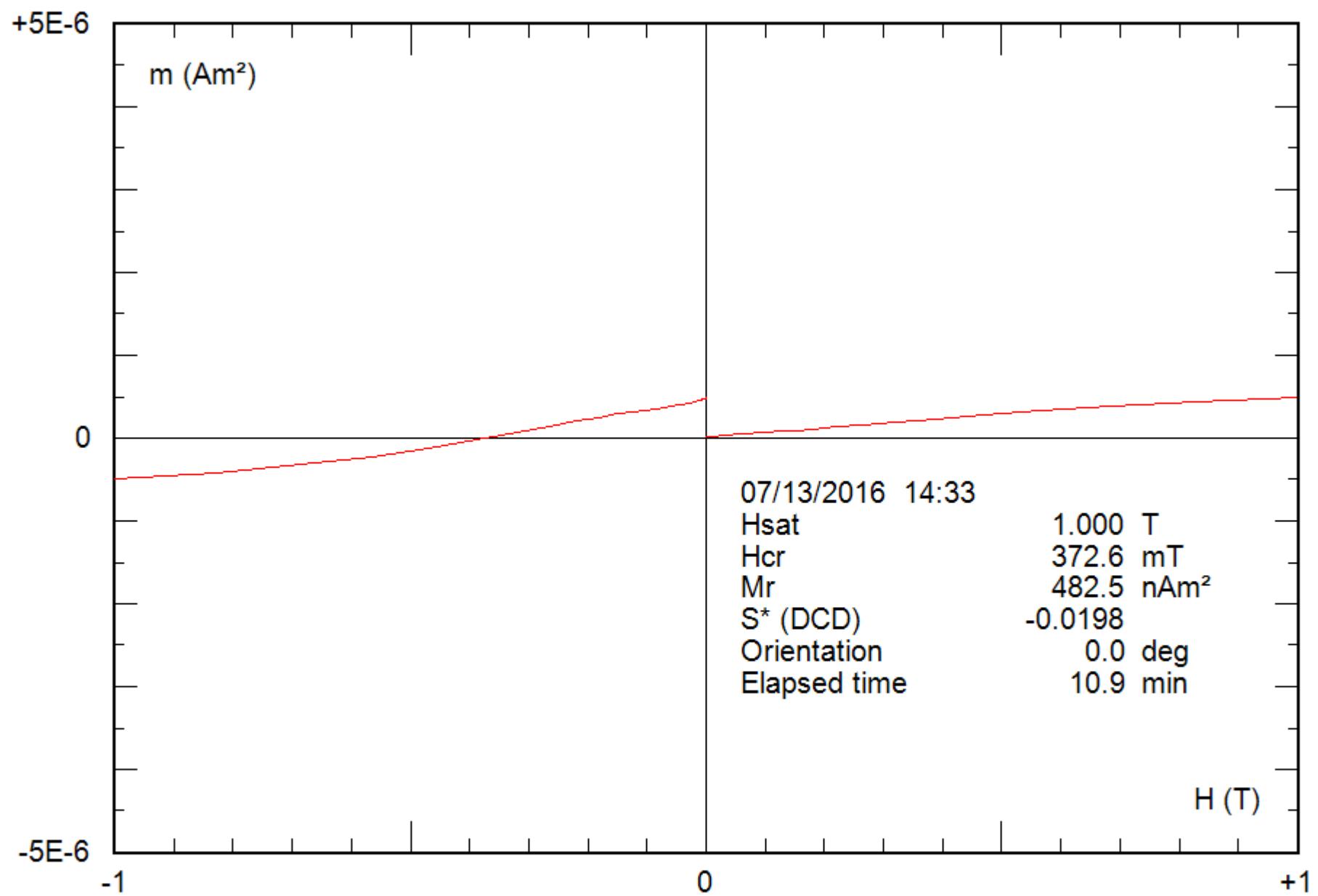
0.26 g

File: 25-04c\_hyst



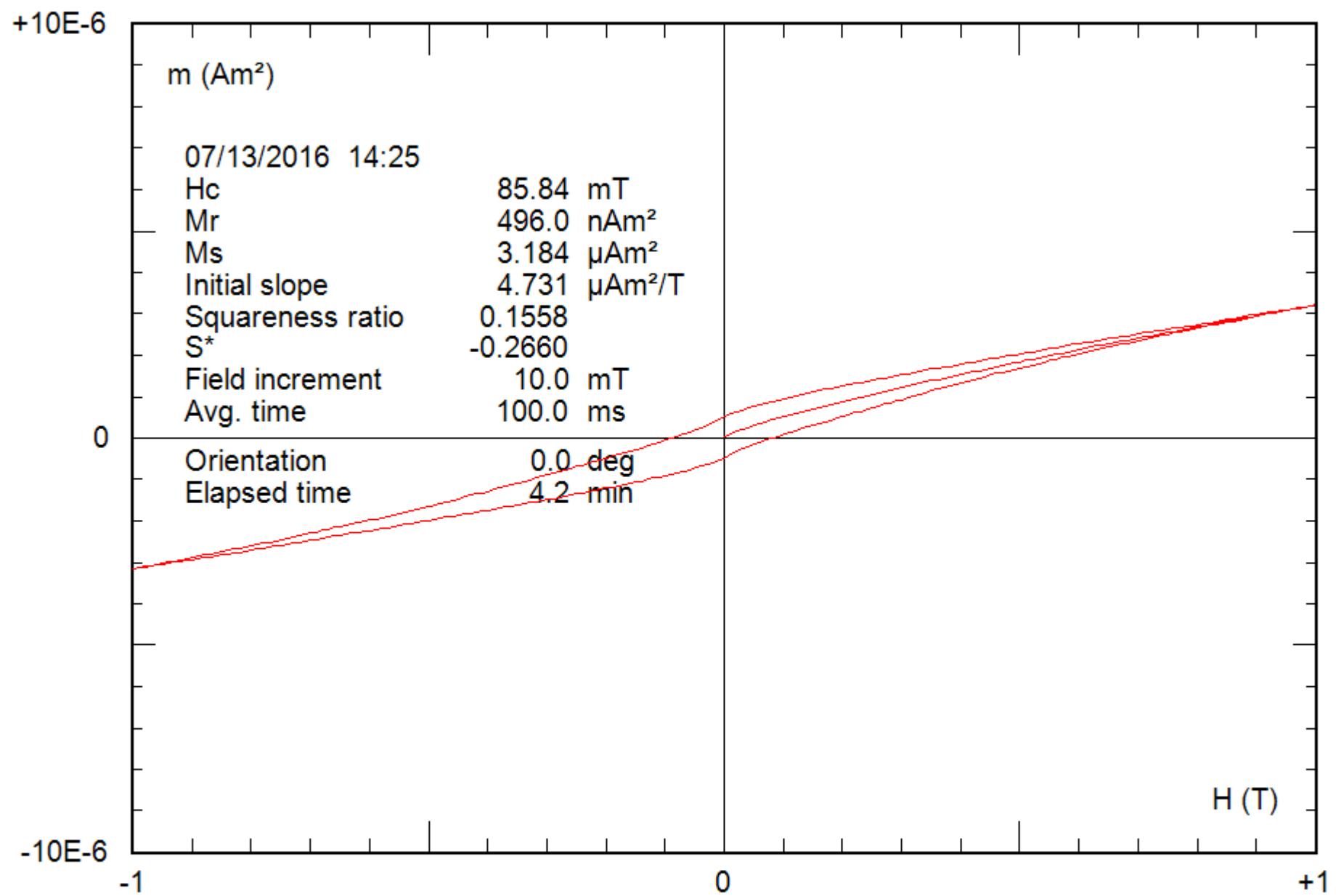
0.26 g

File: 25-04c\_hyst\_corr



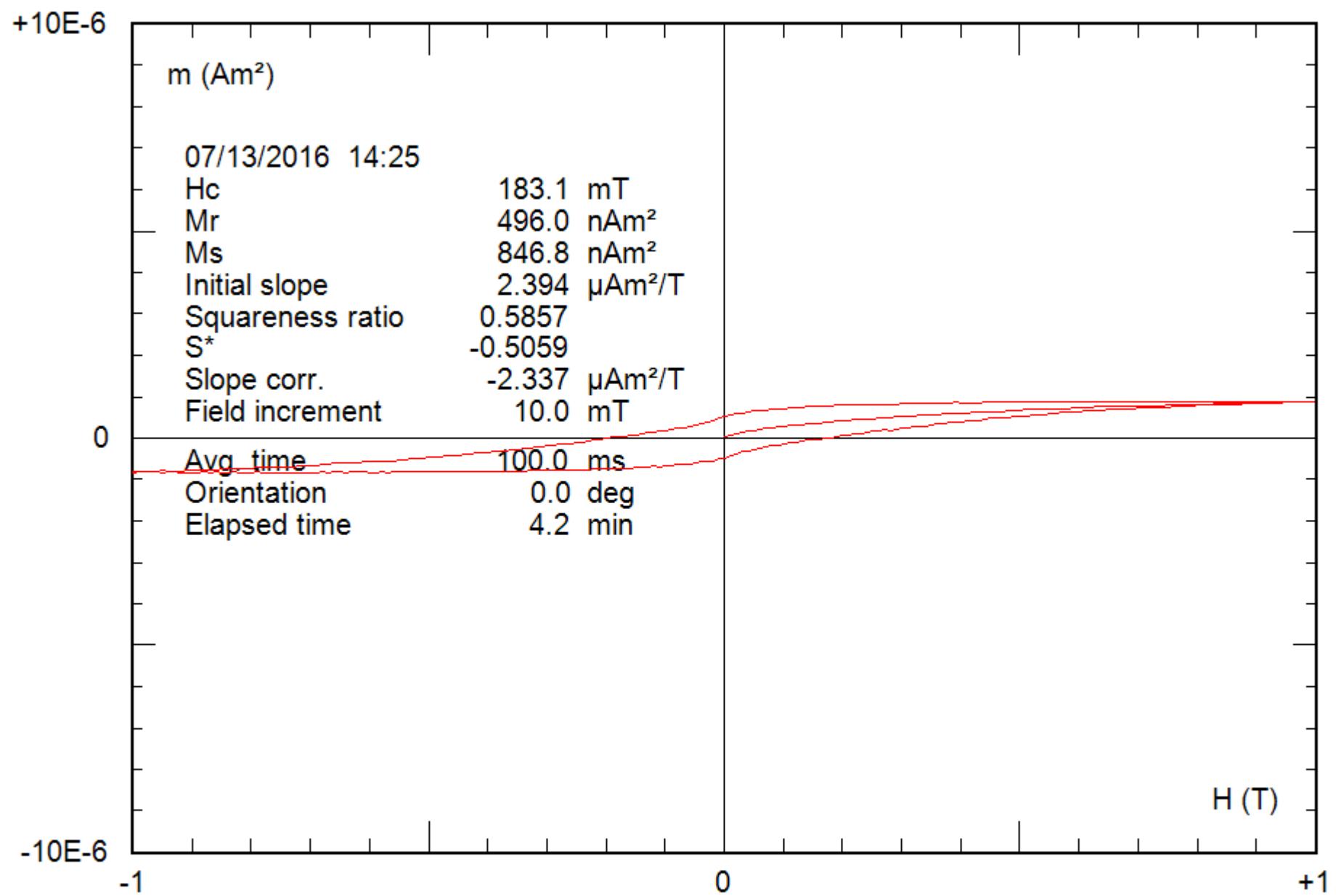
0.25 g

File: 26-04c\_bcr



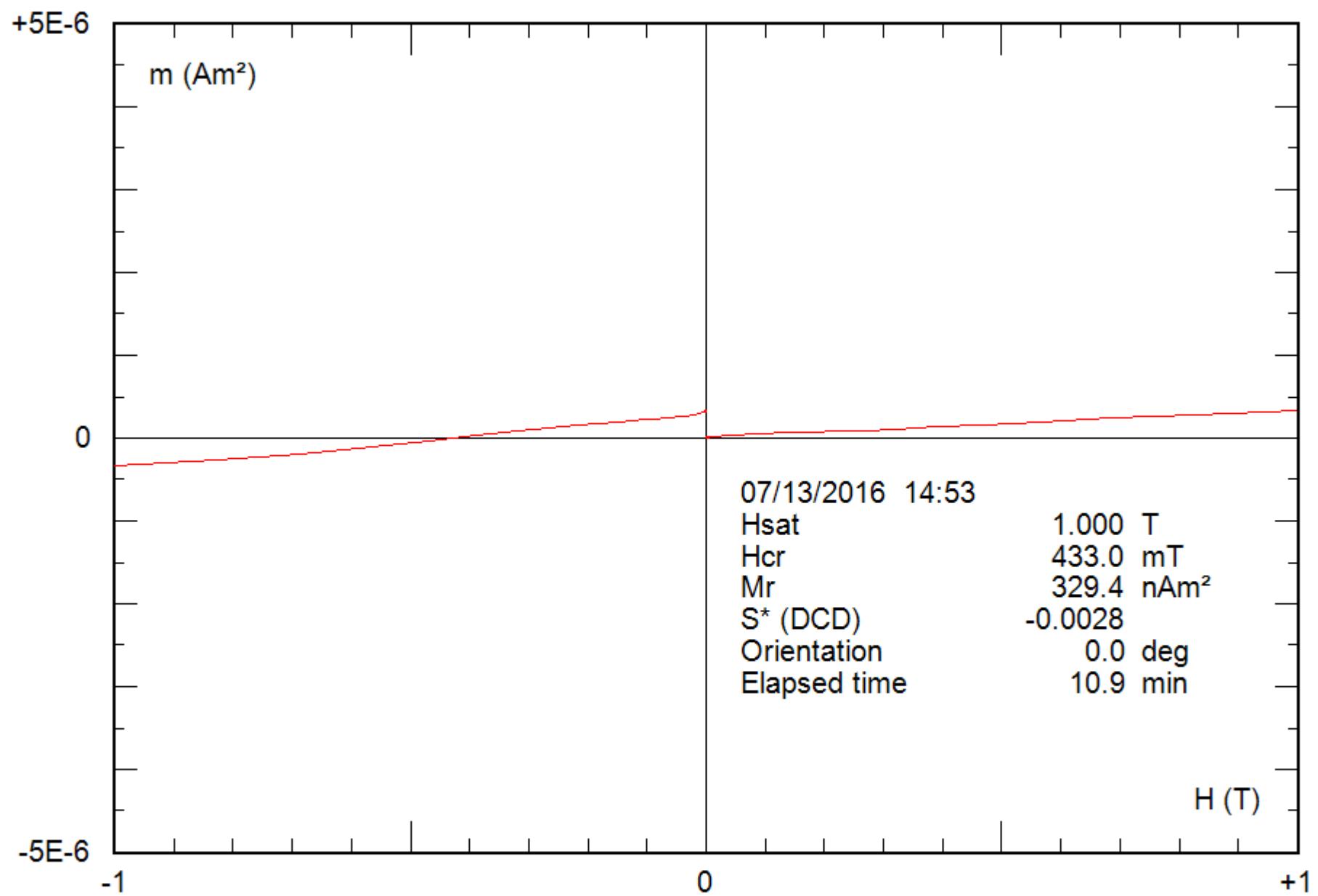
0.25 g

File: 26-04c\_hyst



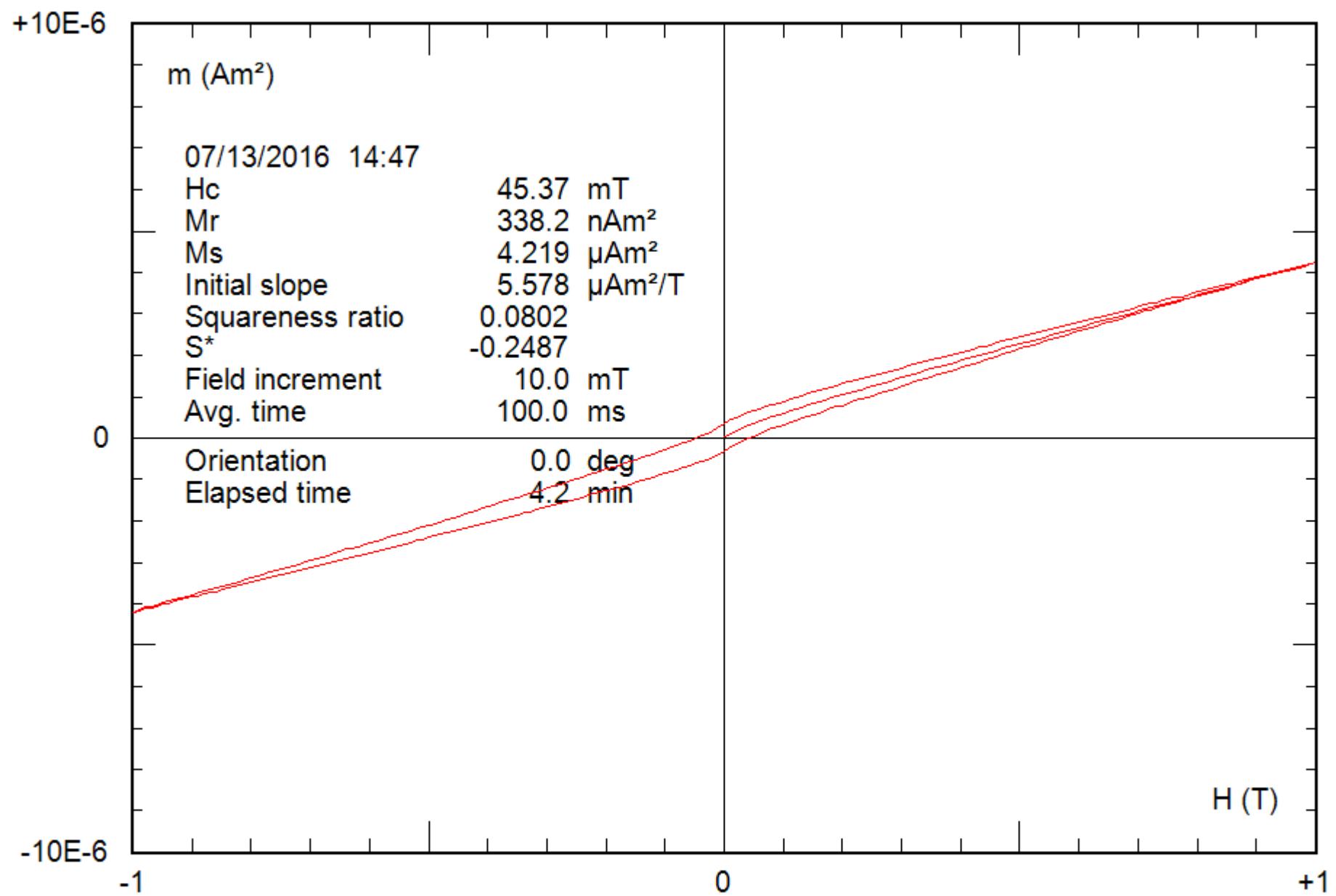
0.25 g

File: 26-04c\_hyst\_corr



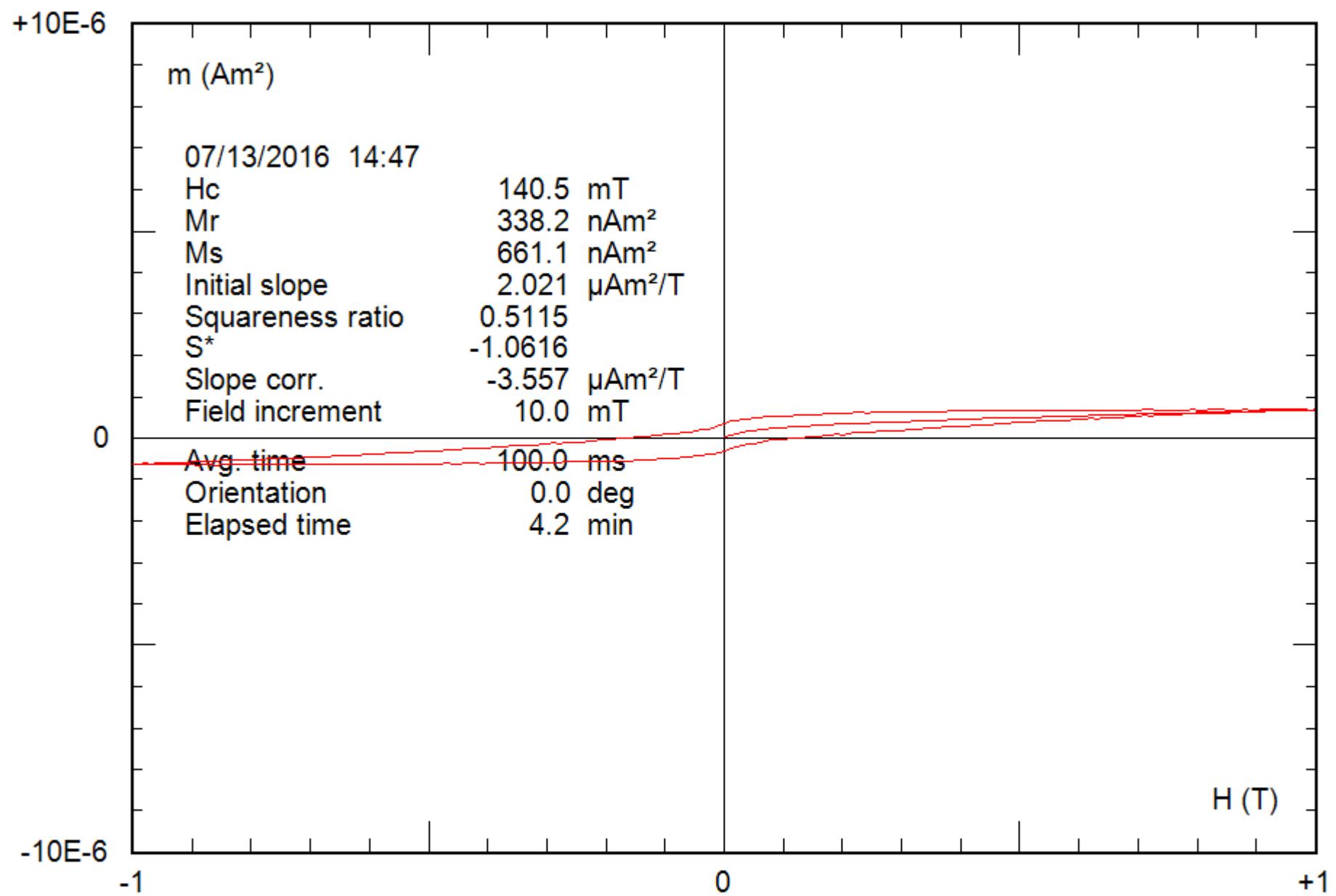
0.25 g

File: 27-08b\_bcr



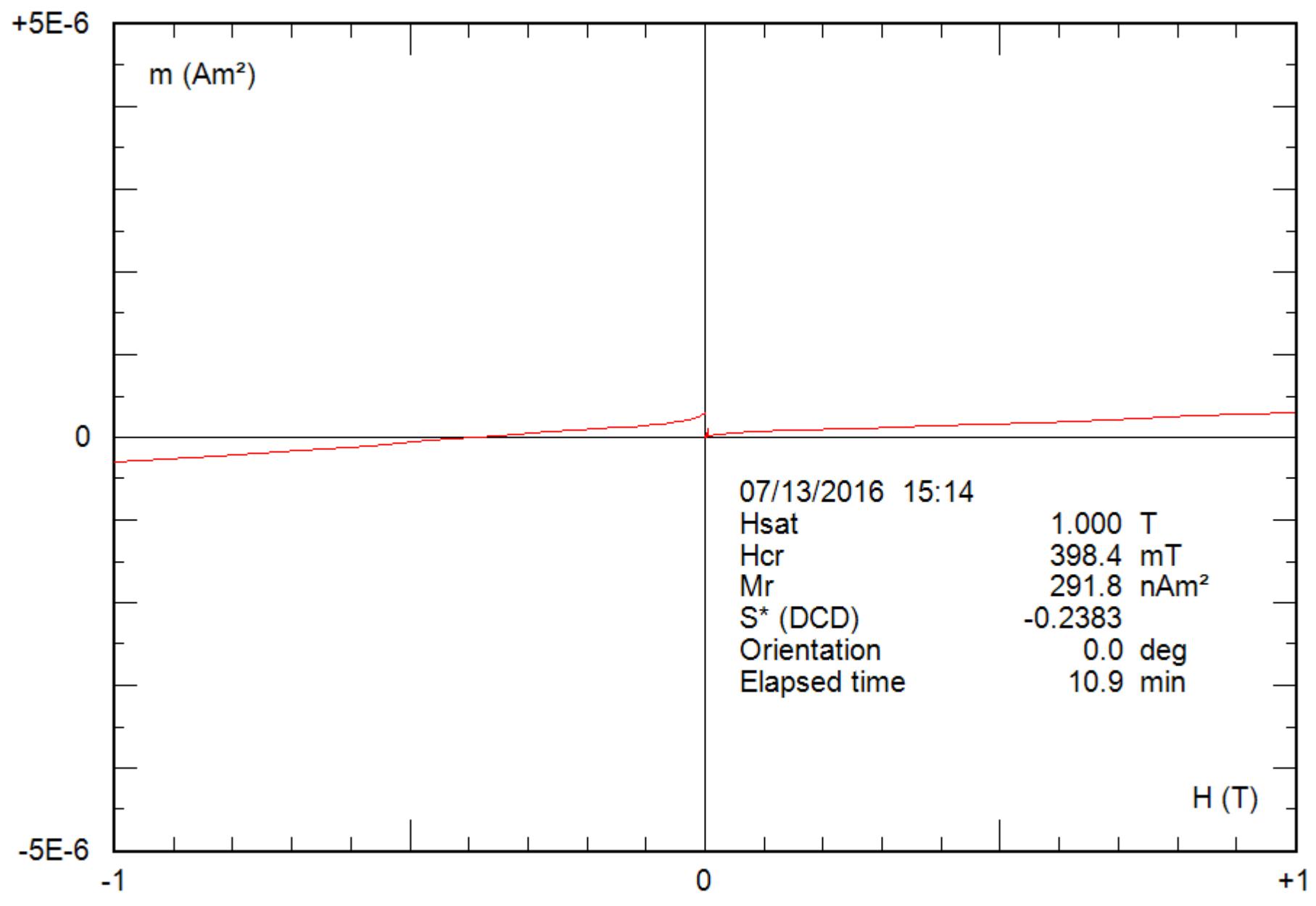
0.25 g

File: 27-08b\_hyst



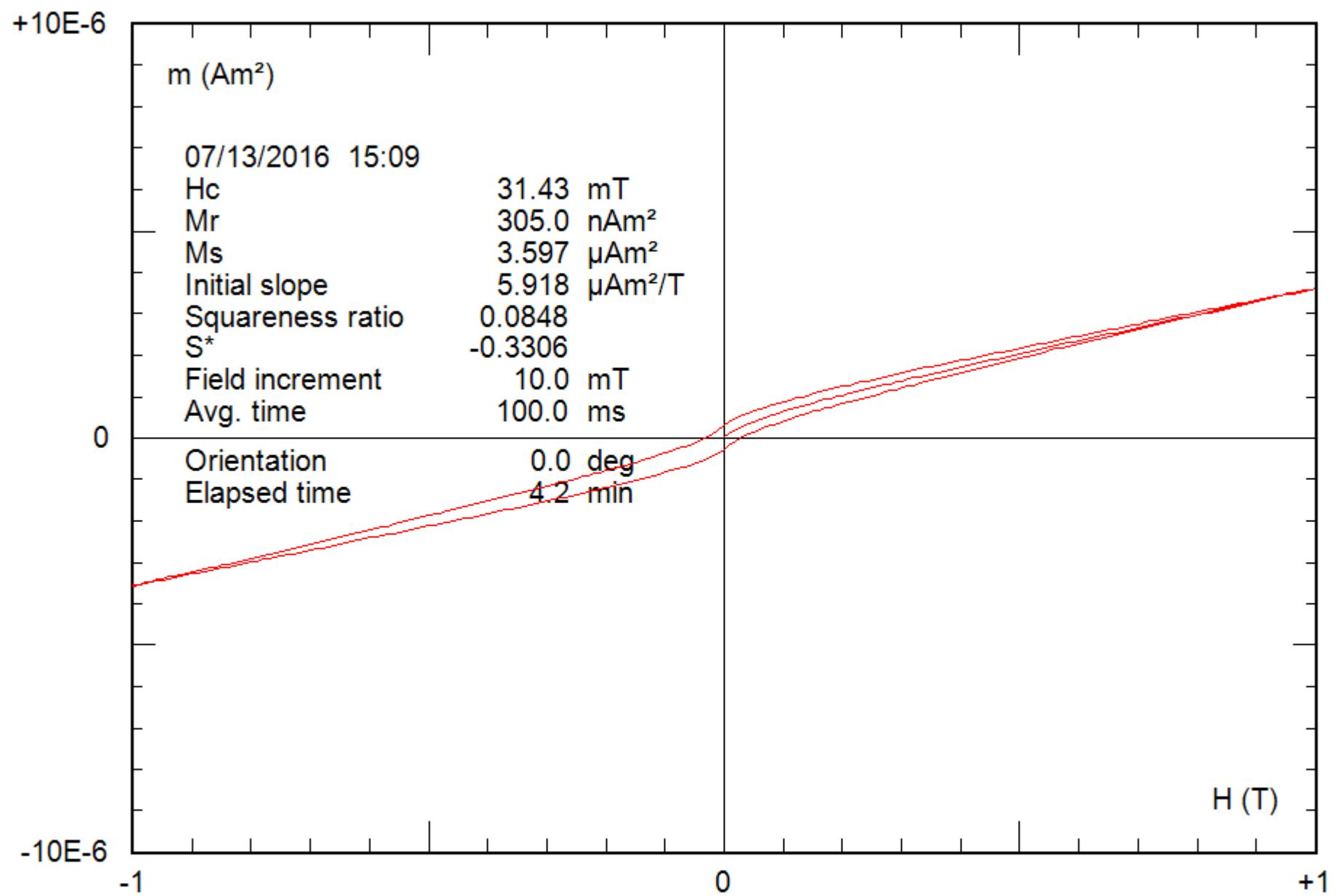
0.25 g

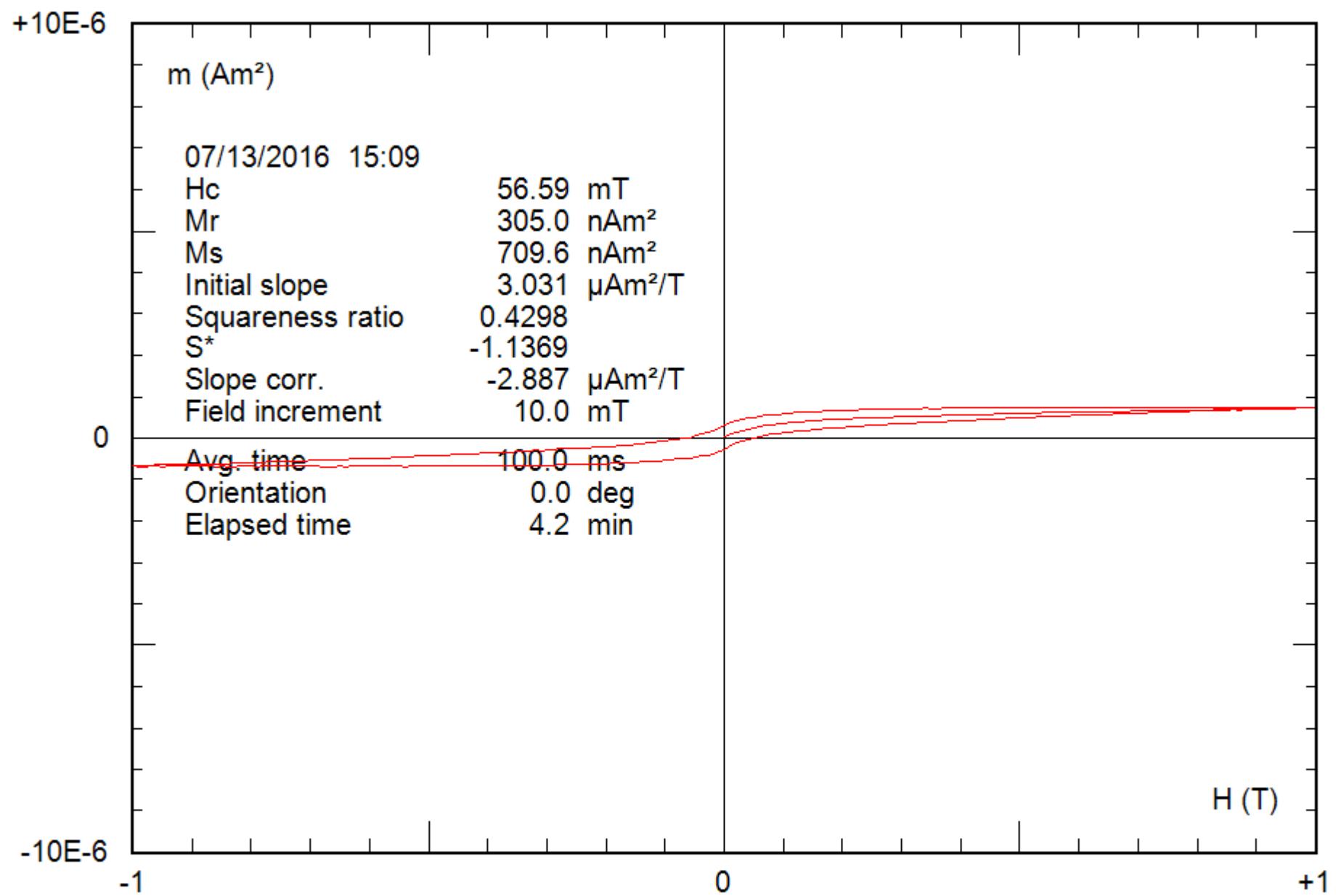
File: 27-08b\_hyst\_corr



0.20 g

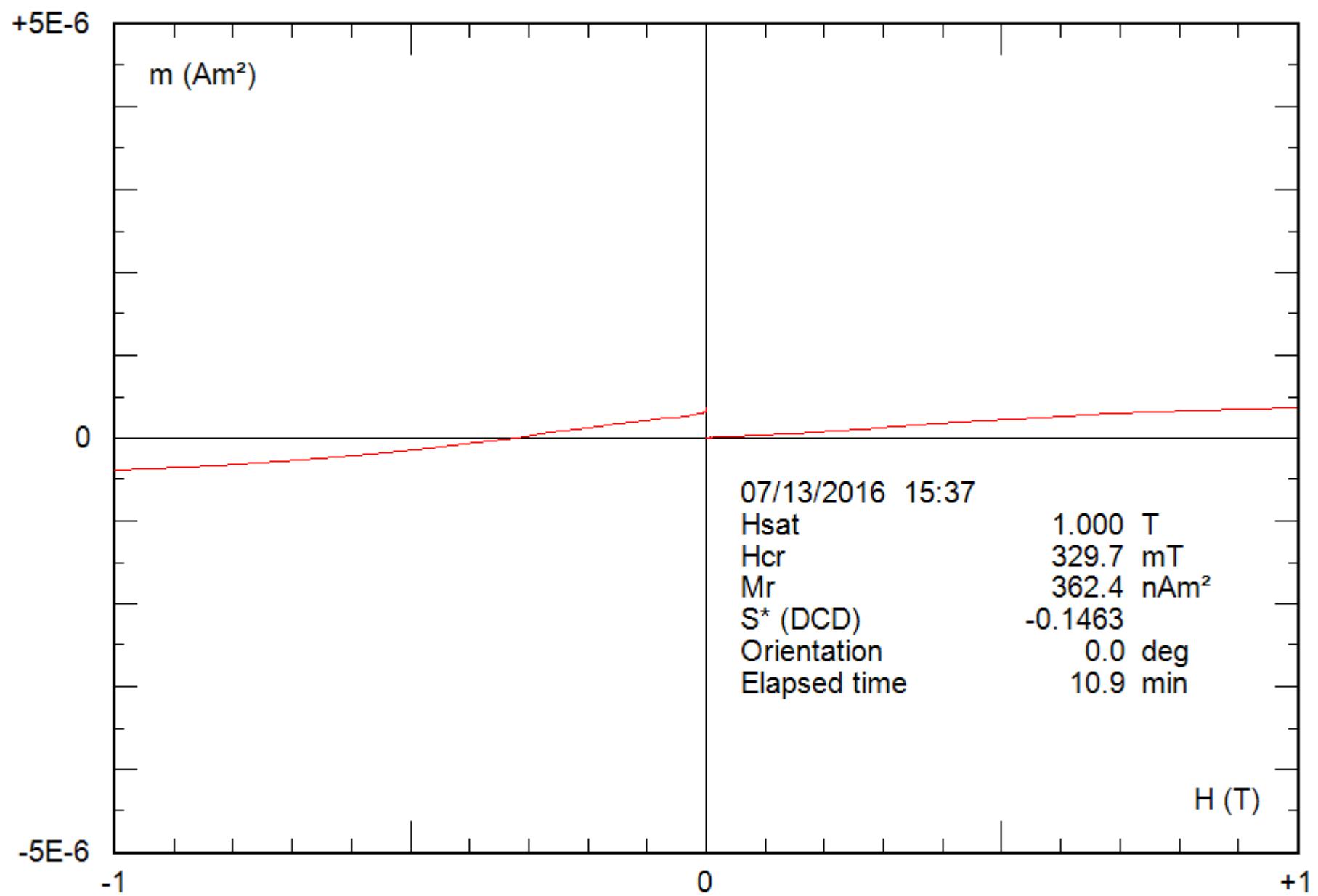
File: 28-05a\_bcr





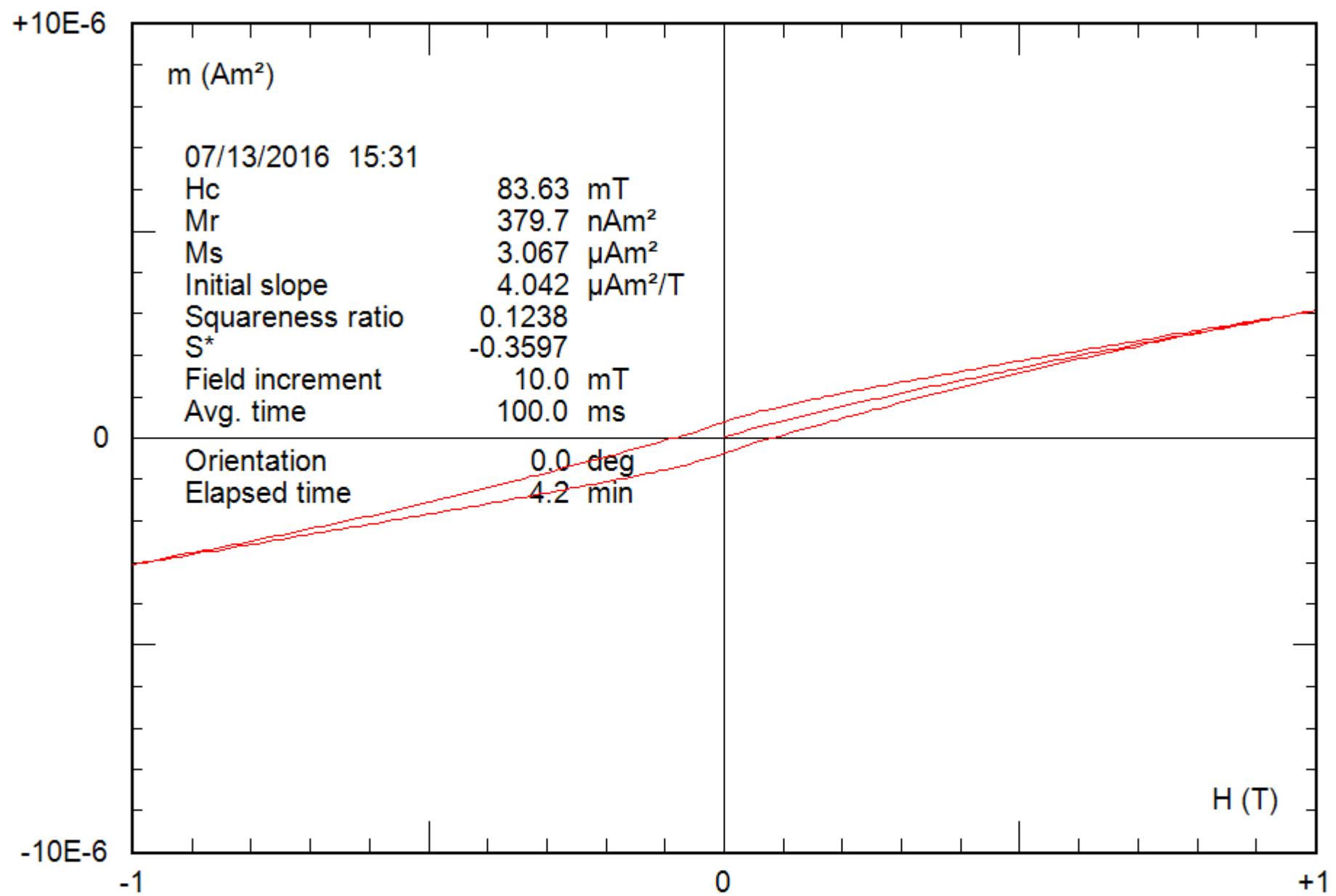
0.20 g

File: 28-05a\_hyst\_corr



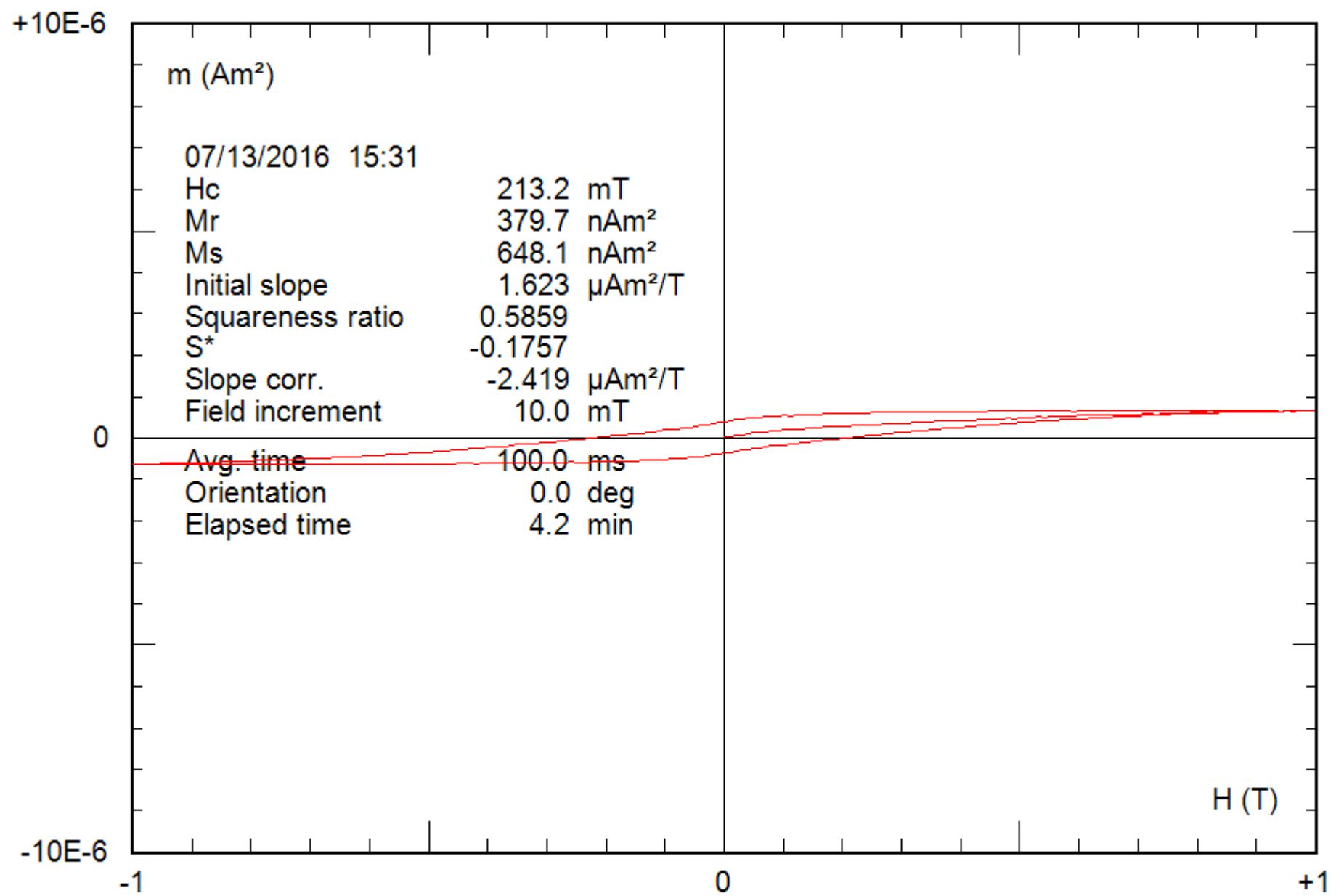
0.27 g

File: 29-02c\_bcr



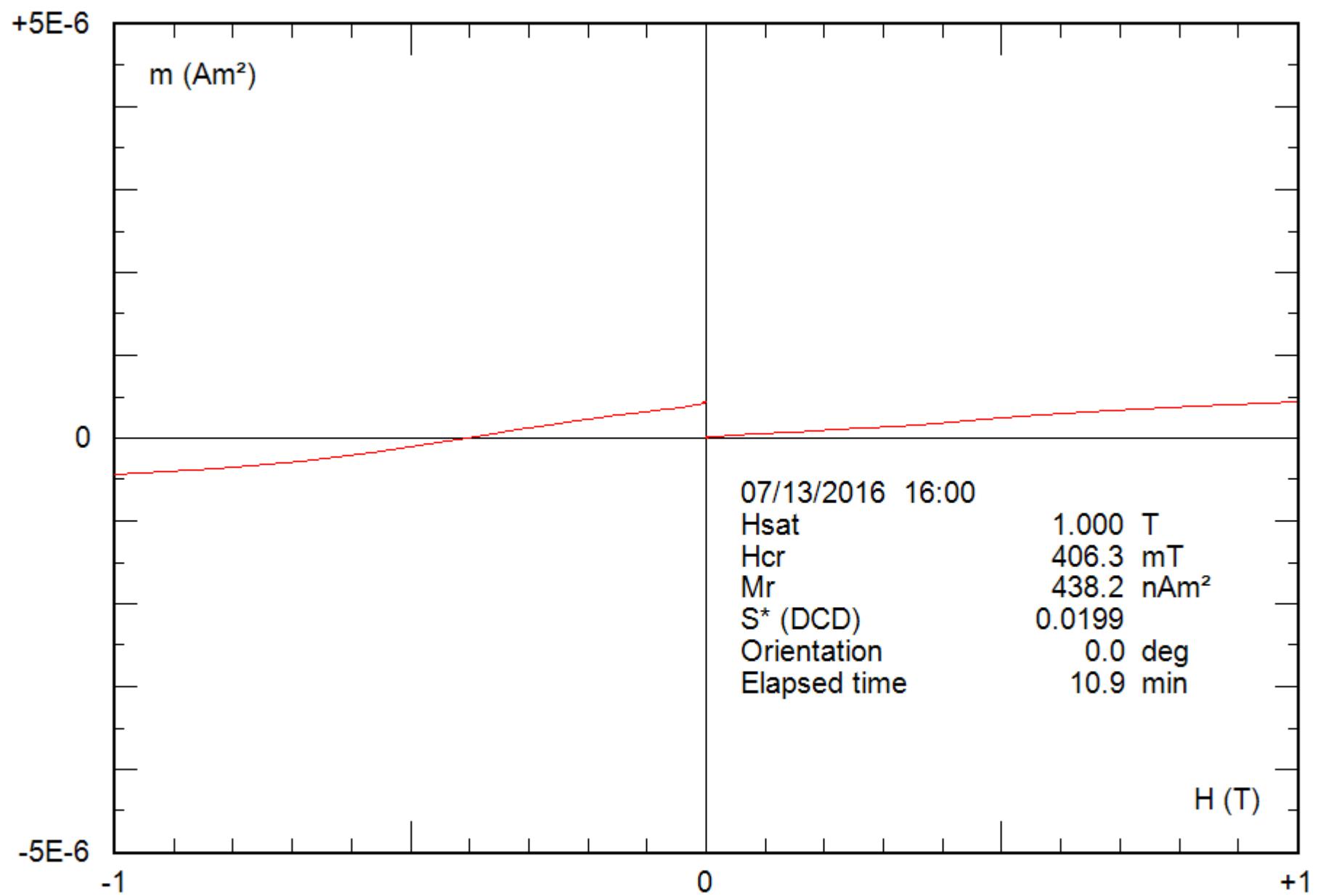
0.27 g

File: 29-02c\_hyst



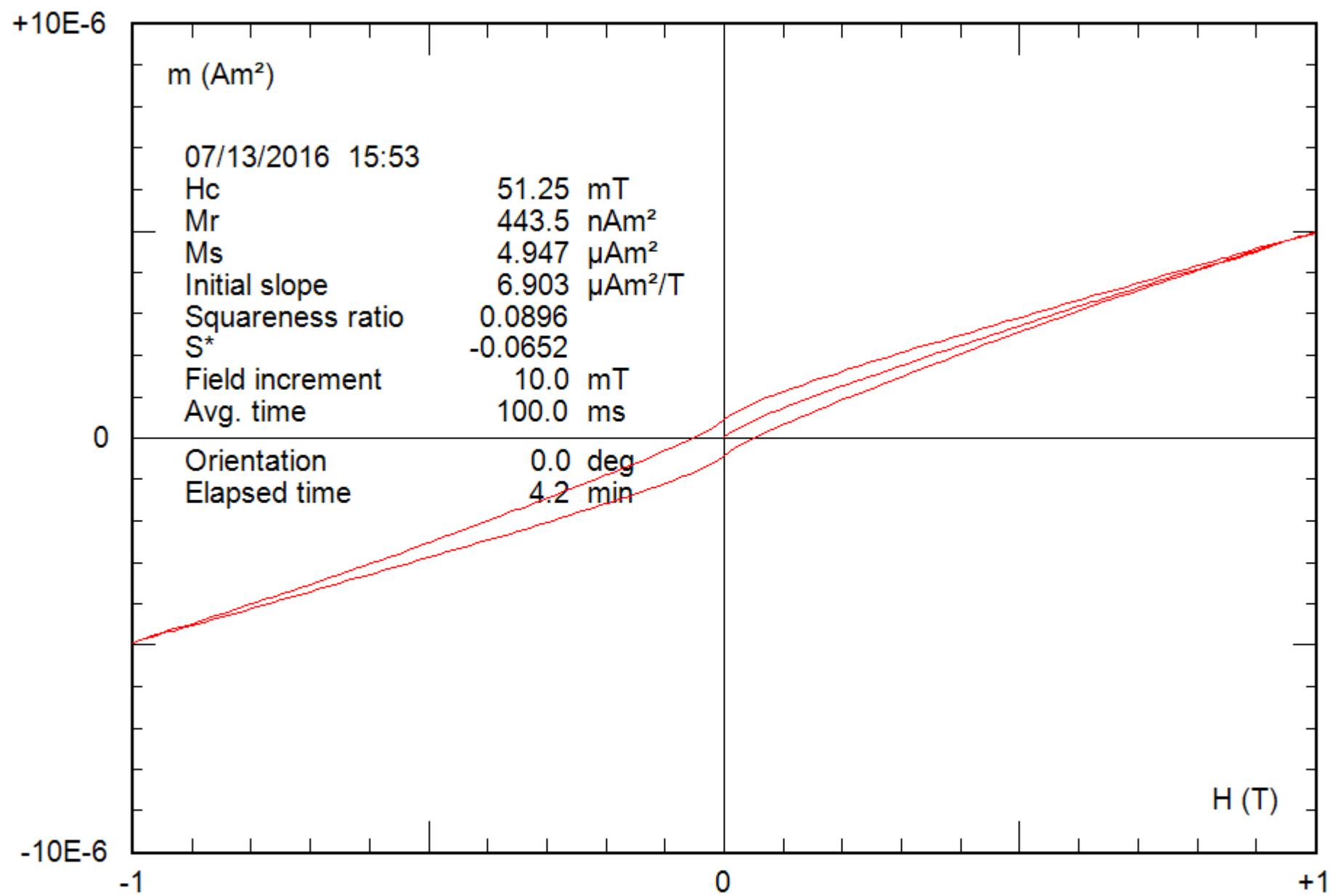
0.27 g

File: 29-02c\_hyst\_corr



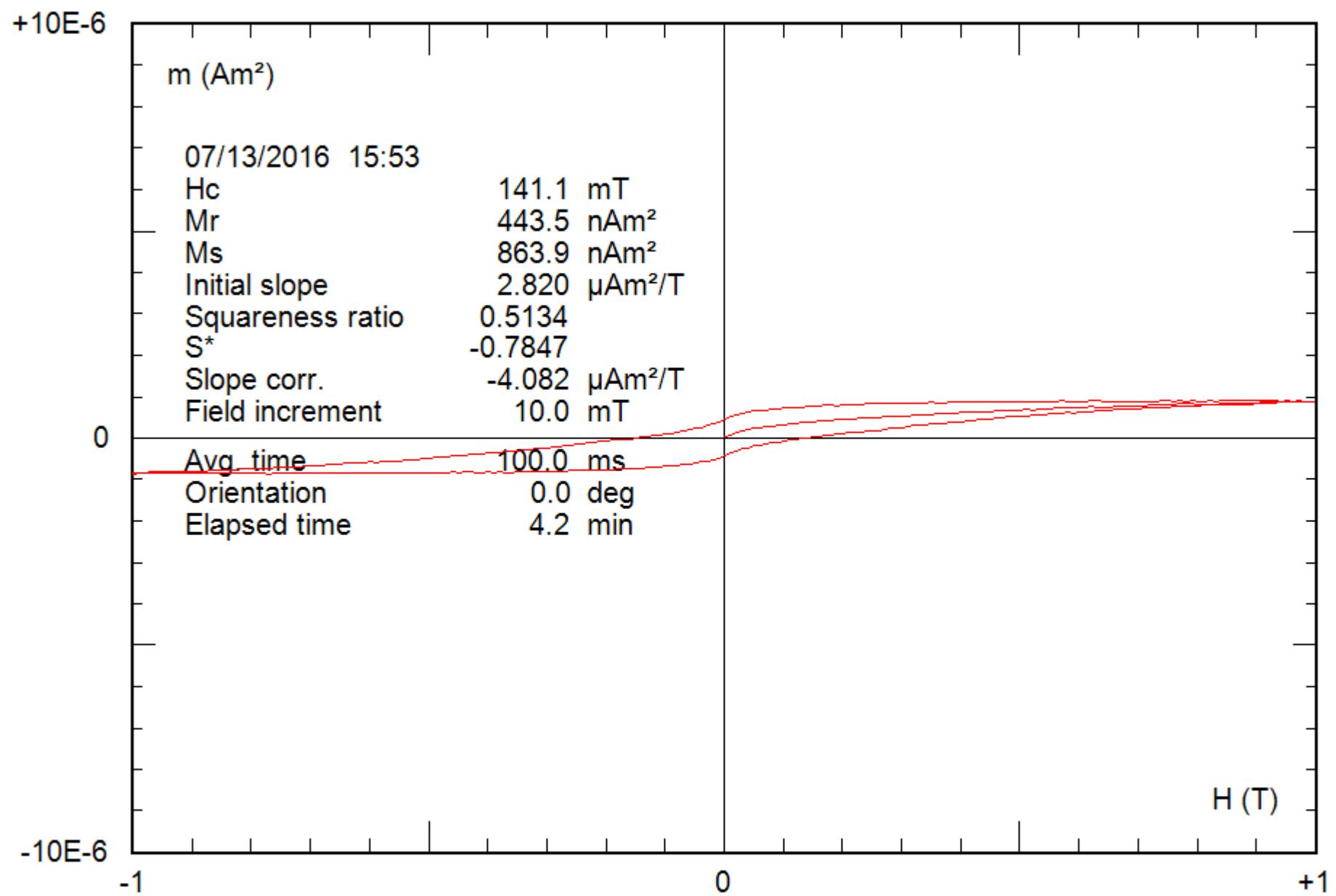
0.25 g

File: 30-10b\_bcr



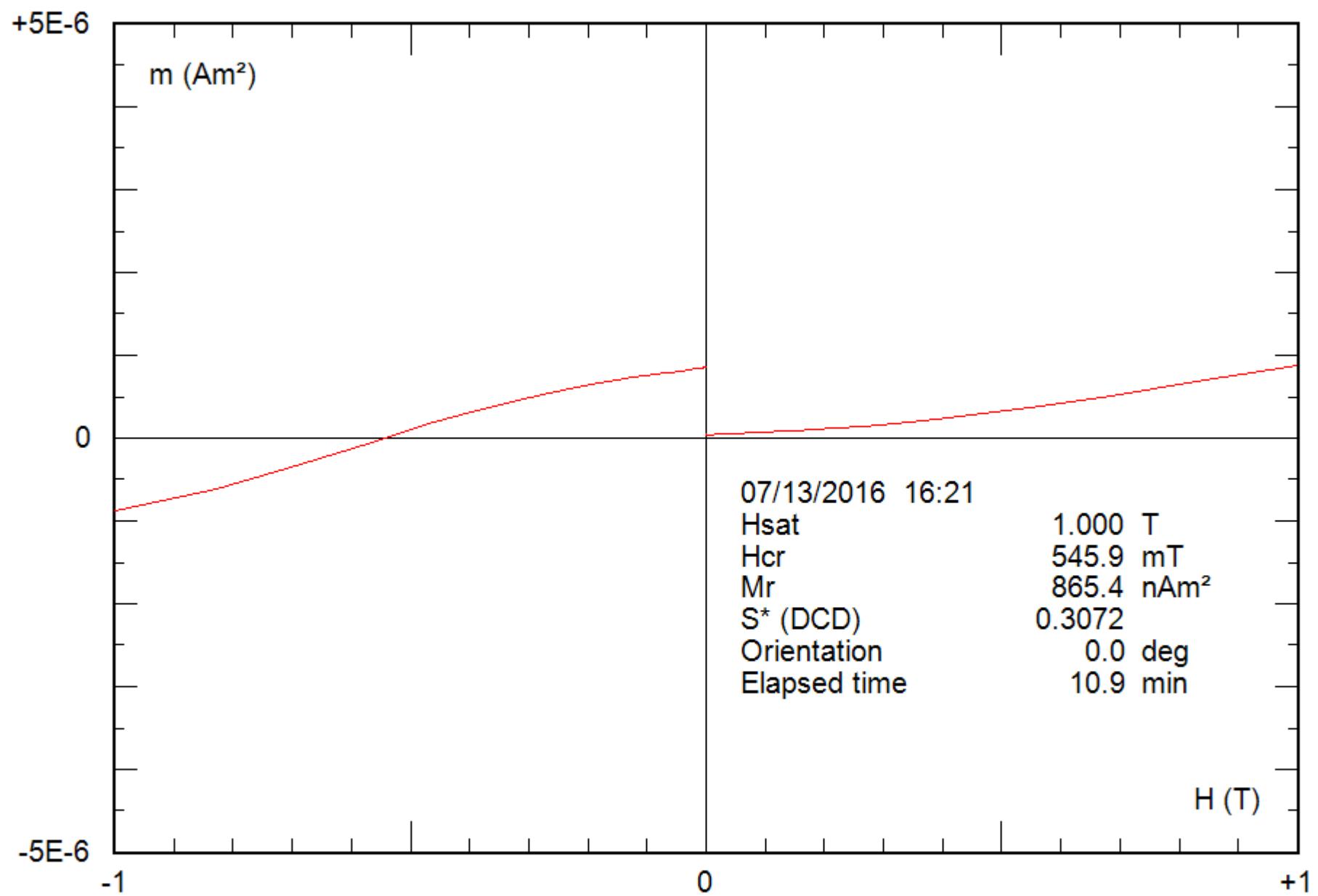
0.25 g

File: 30-10b\_hyst



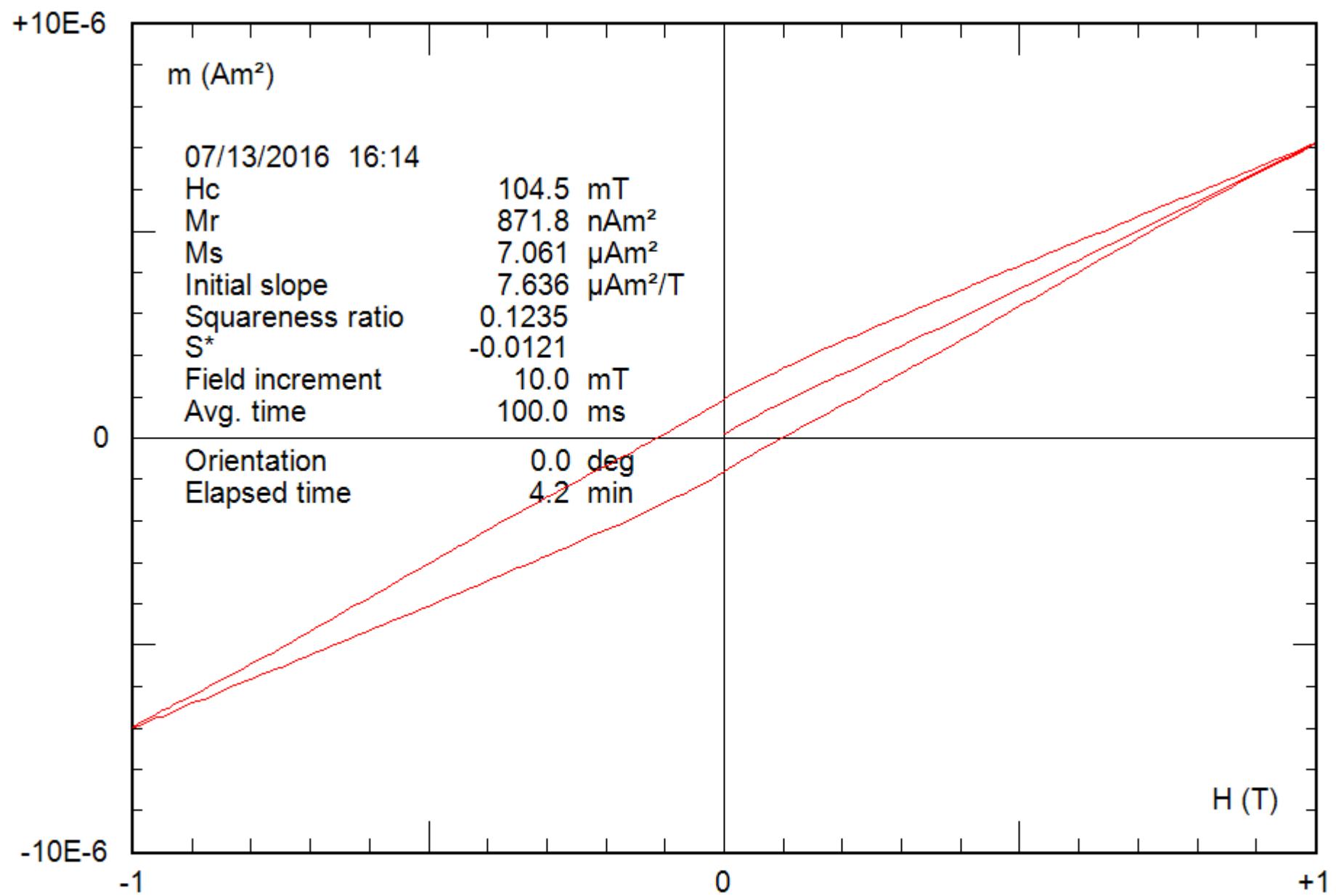
0.25 g

File: 30-10b\_hyst\_corr



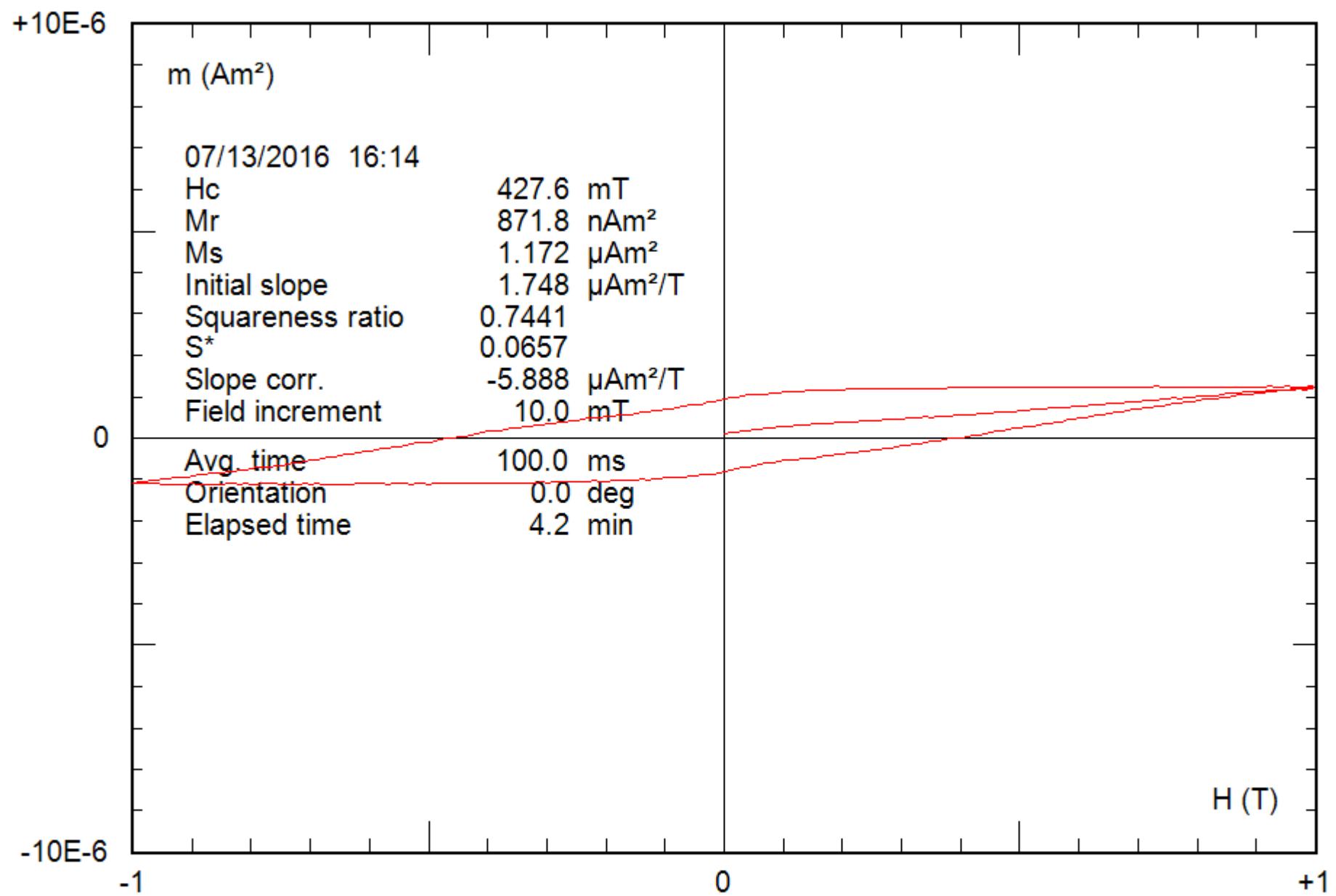
0.19 g

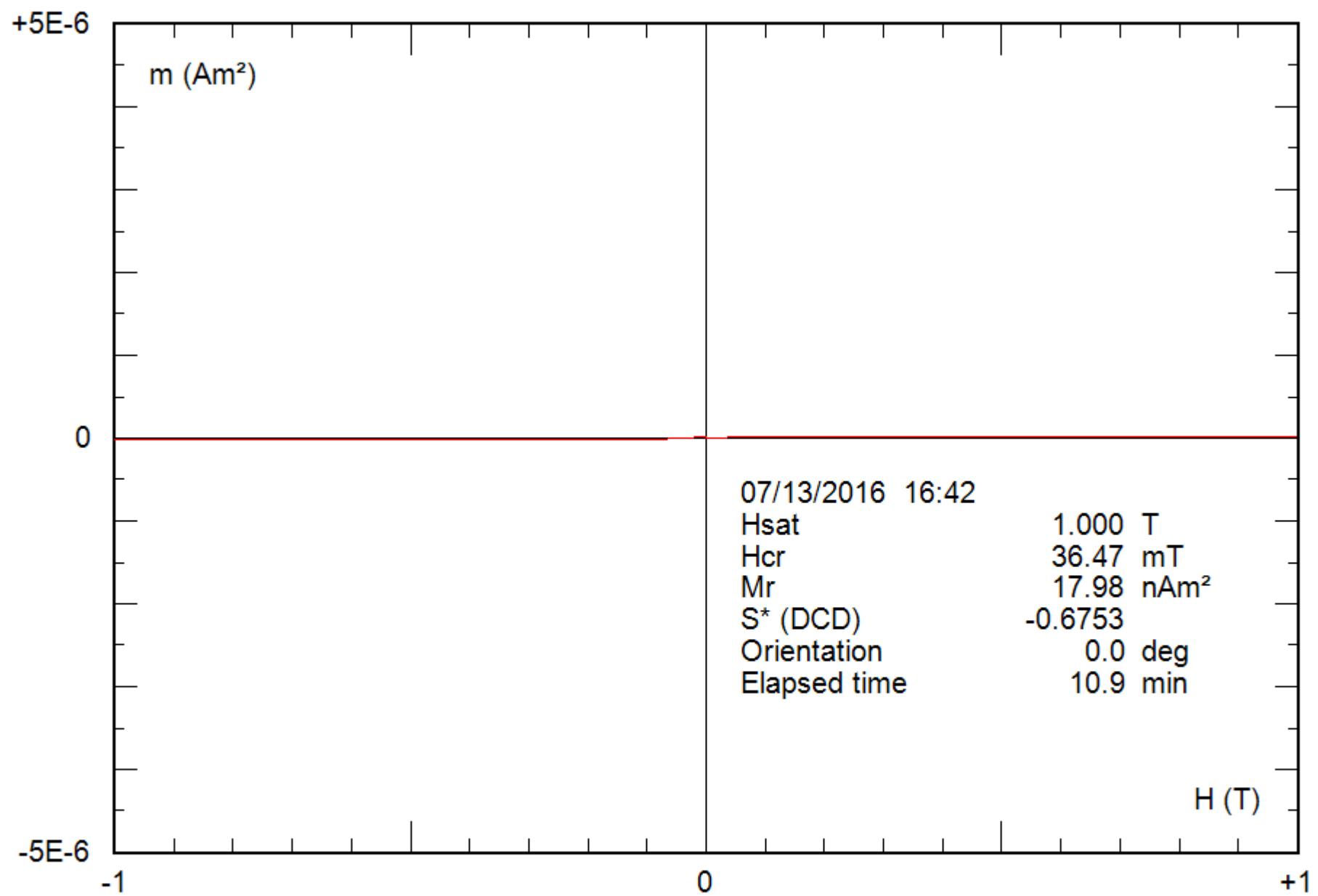
File: 31-06c\_bcr



0.19 g

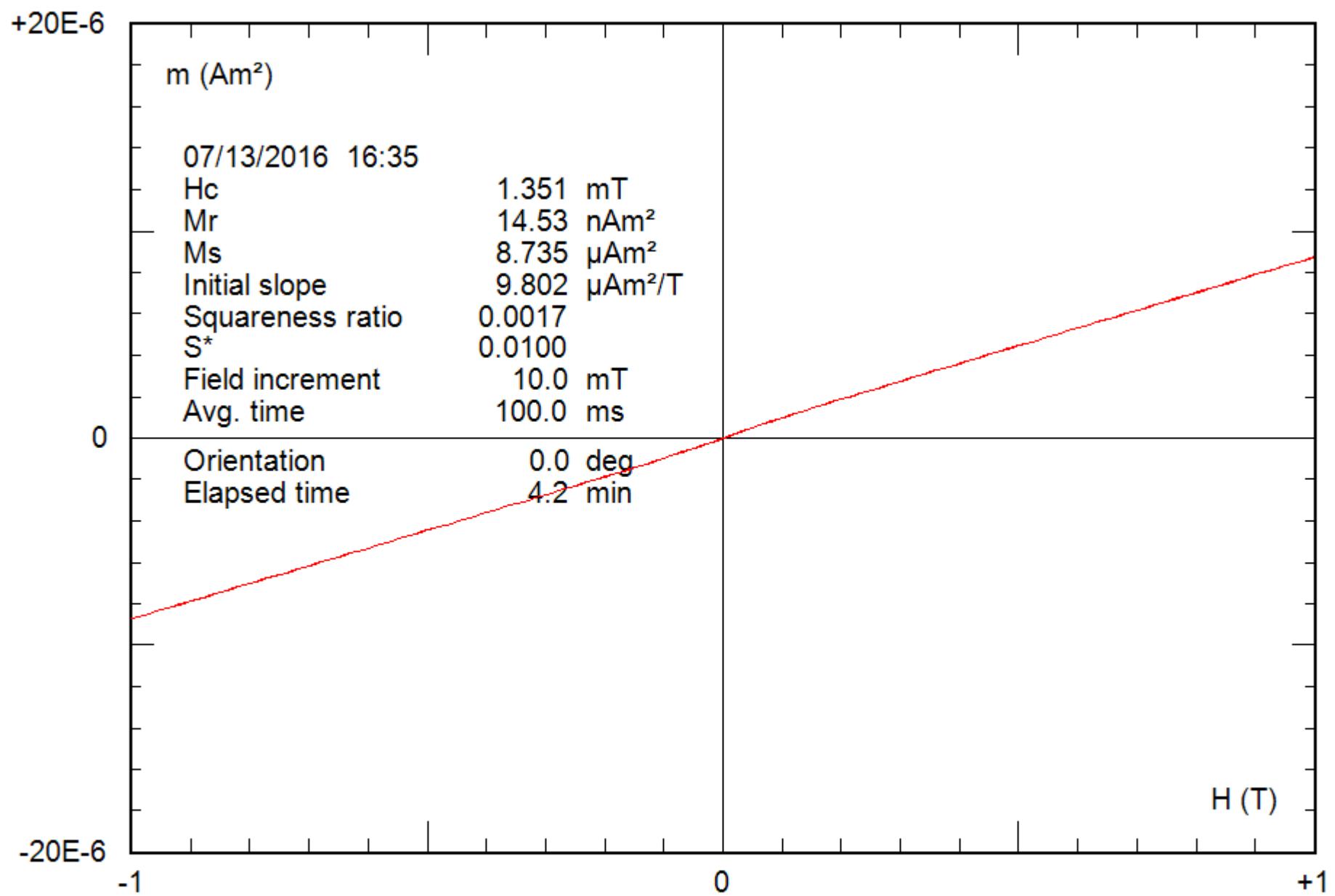
File: 31-06c\_hyst





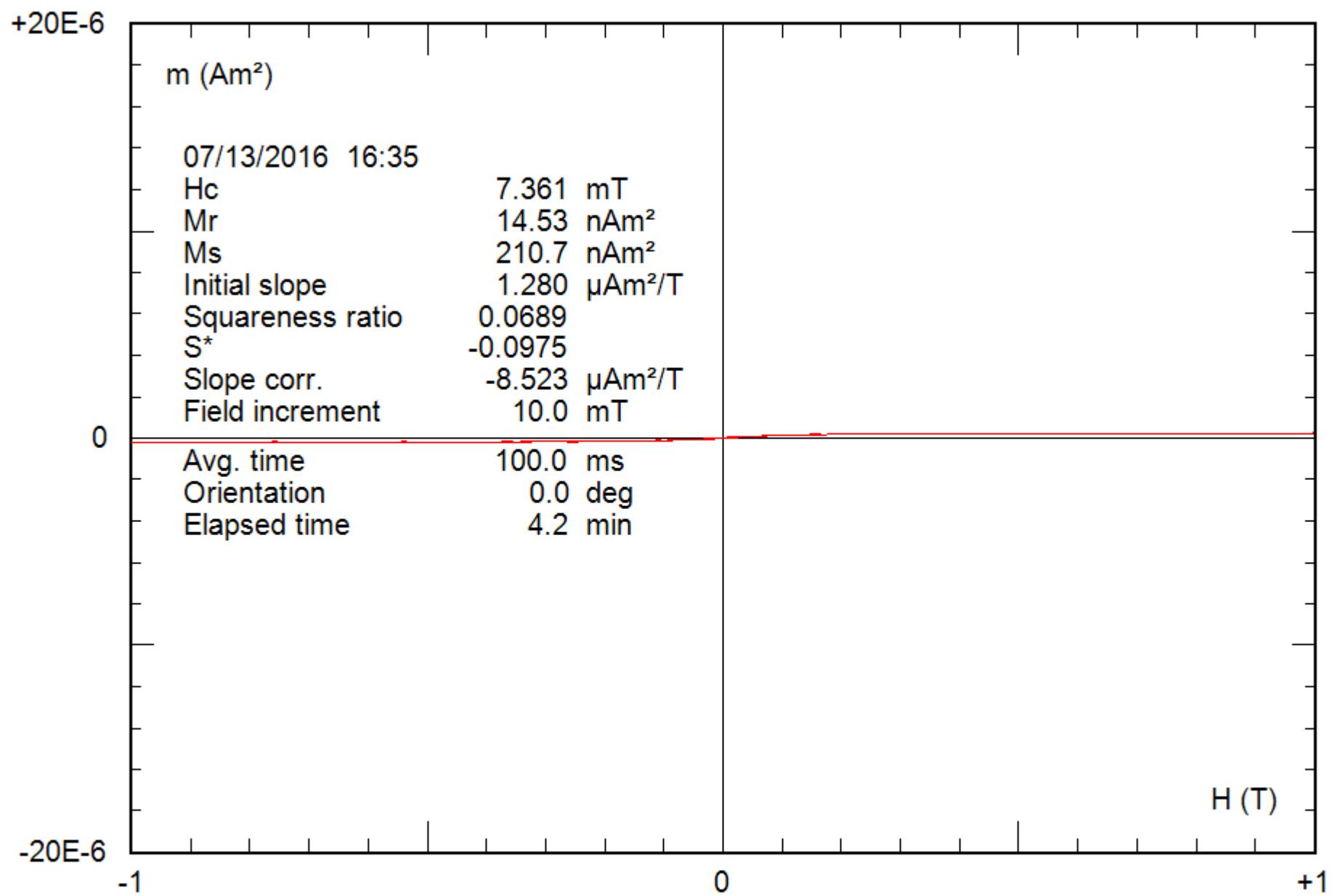
0.22 g

File: 32-01c\_bcr



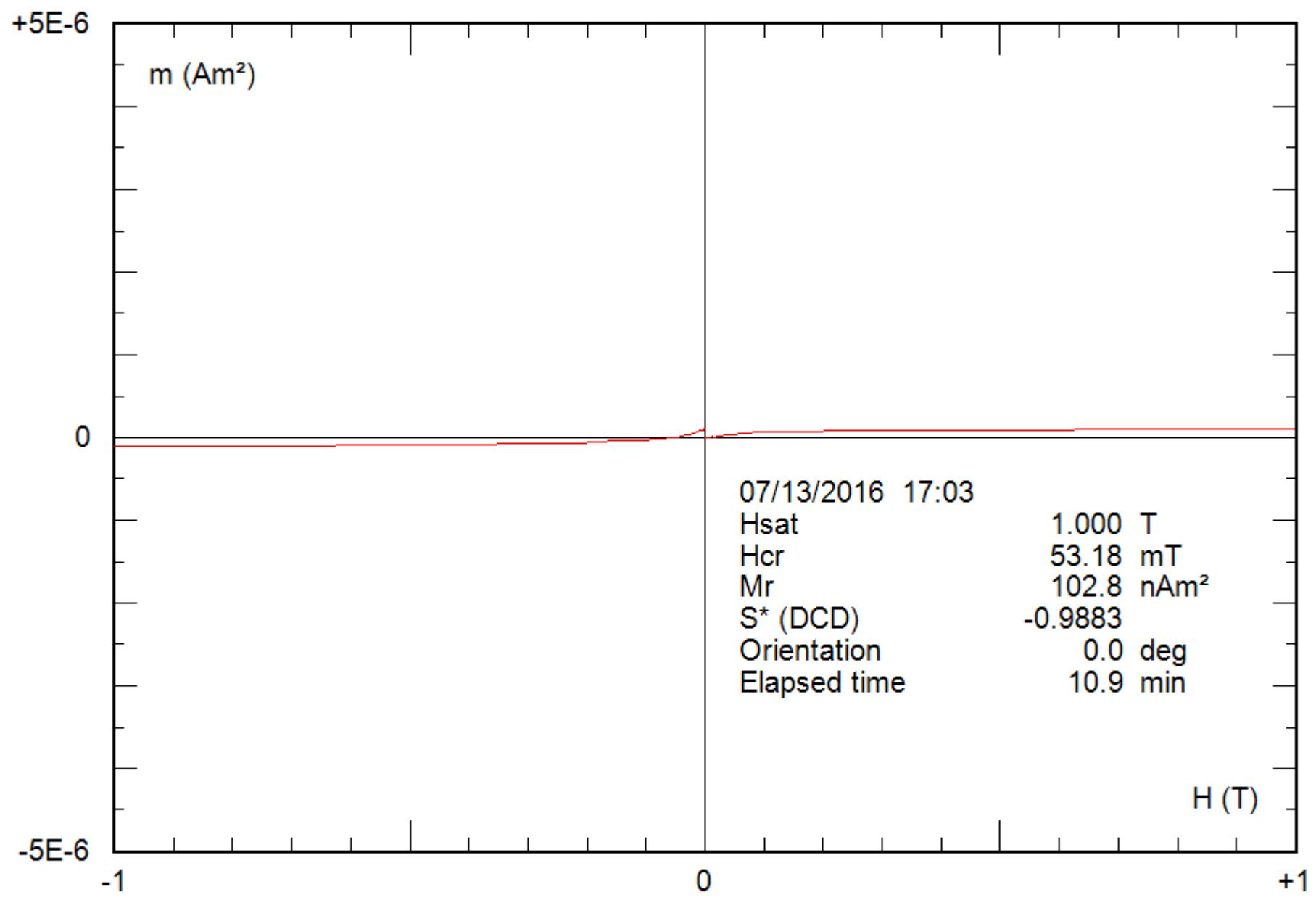
0.22 g

File: 32-01c\_hyst



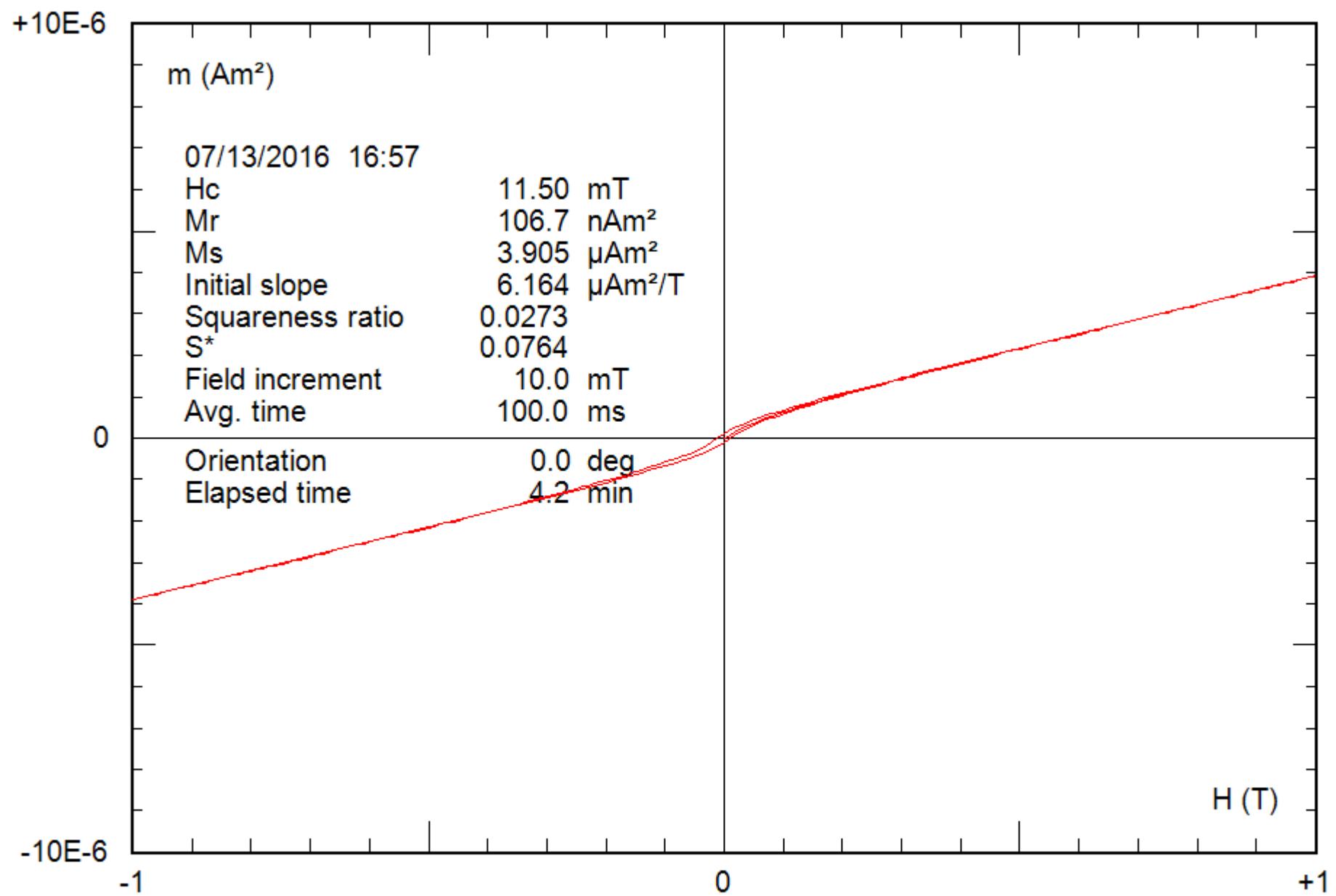
0.22 g

File: 32-01c\_hyst\_corr



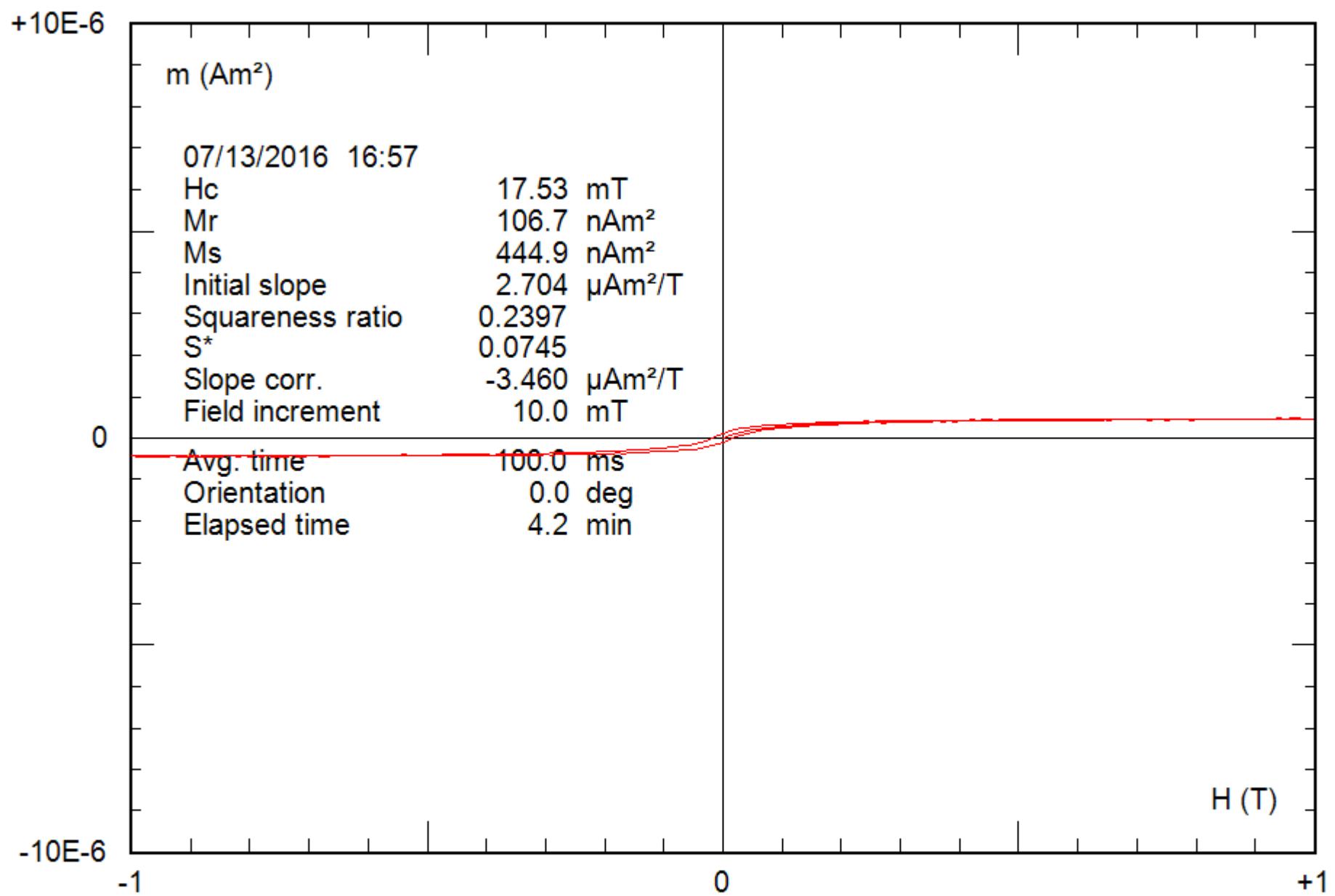
0.21 g

File: 33-04c\_bcr



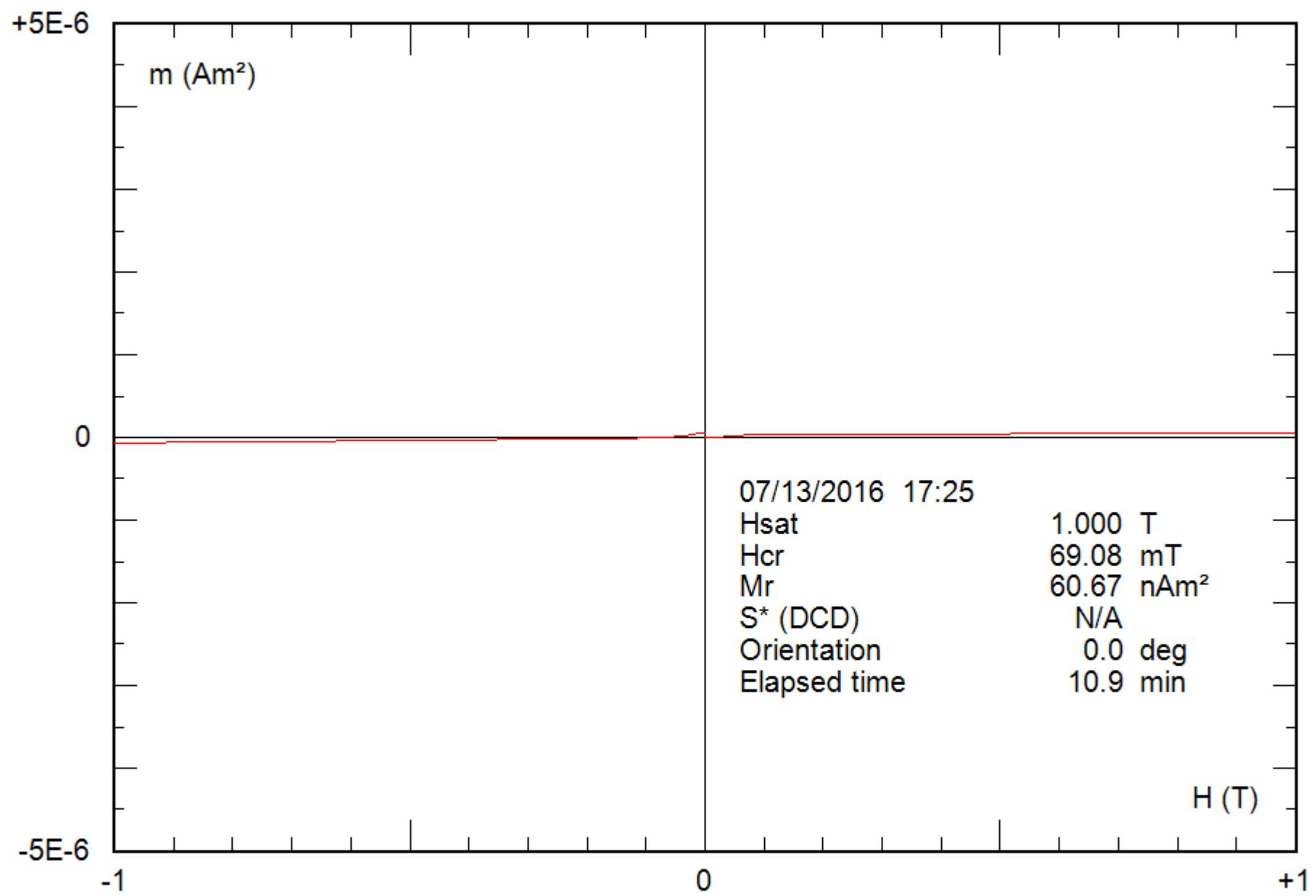
0.21 g

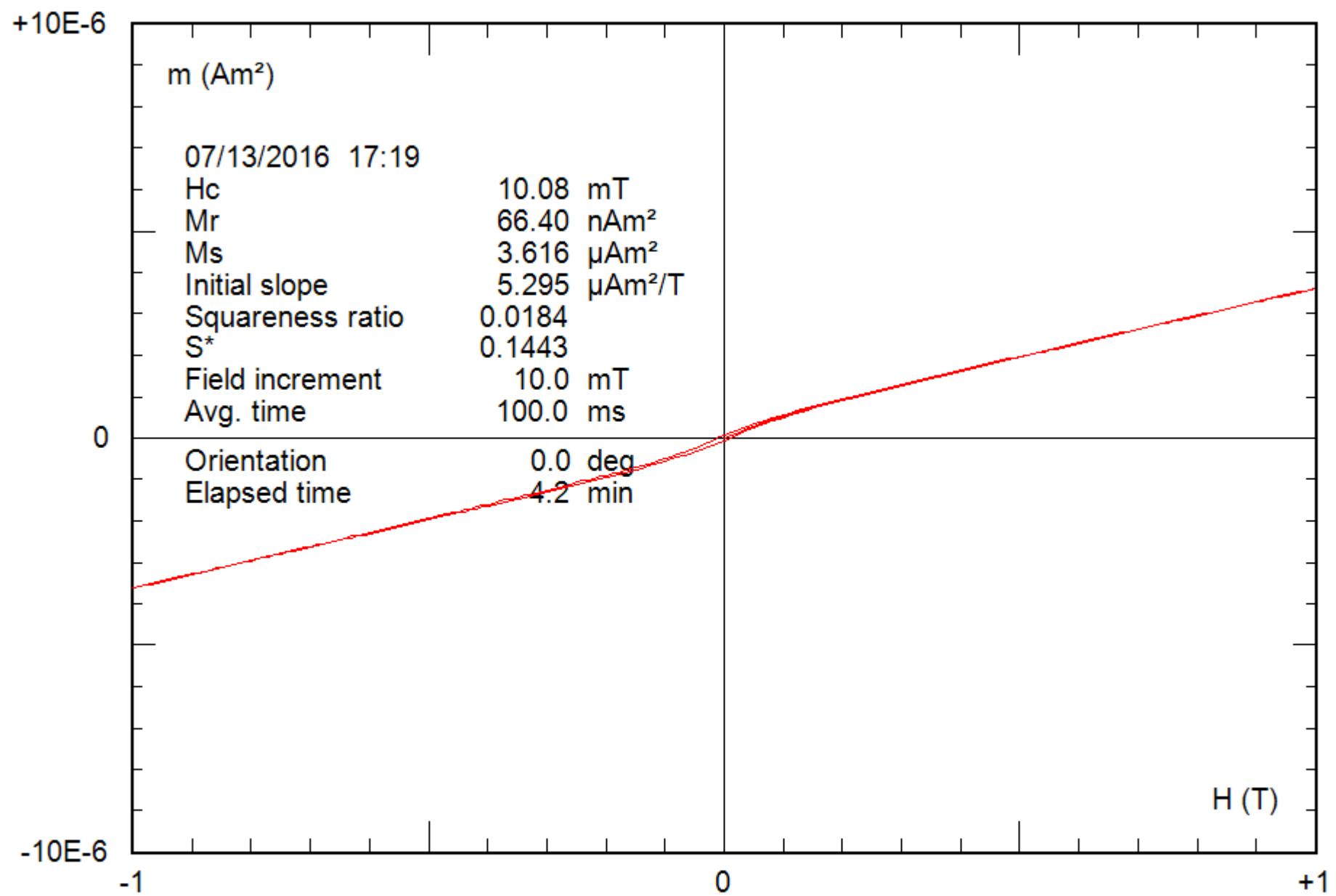
File: 33-04c\_hyst



0.21 g

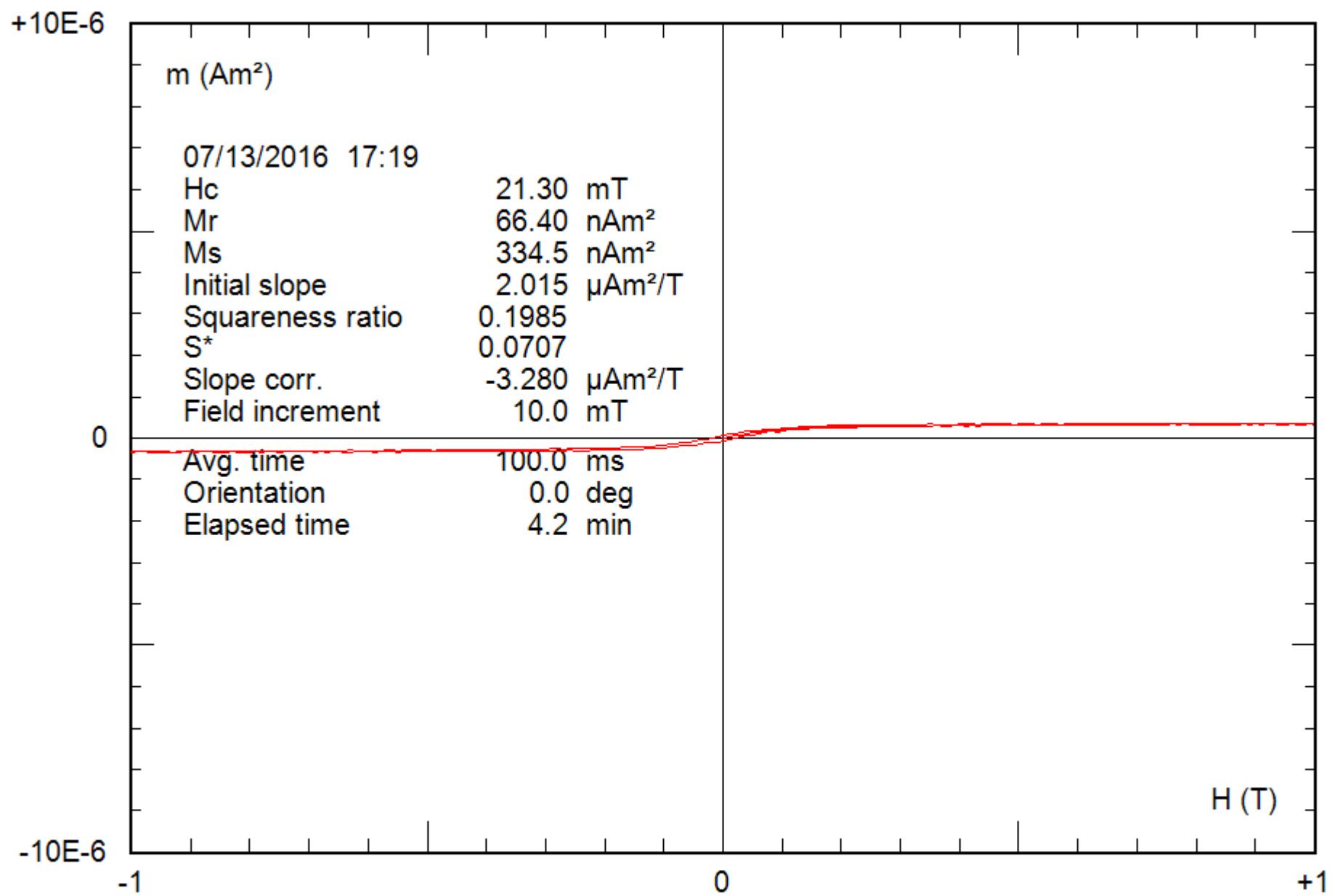
File: 33-04c\_hyst\_corr





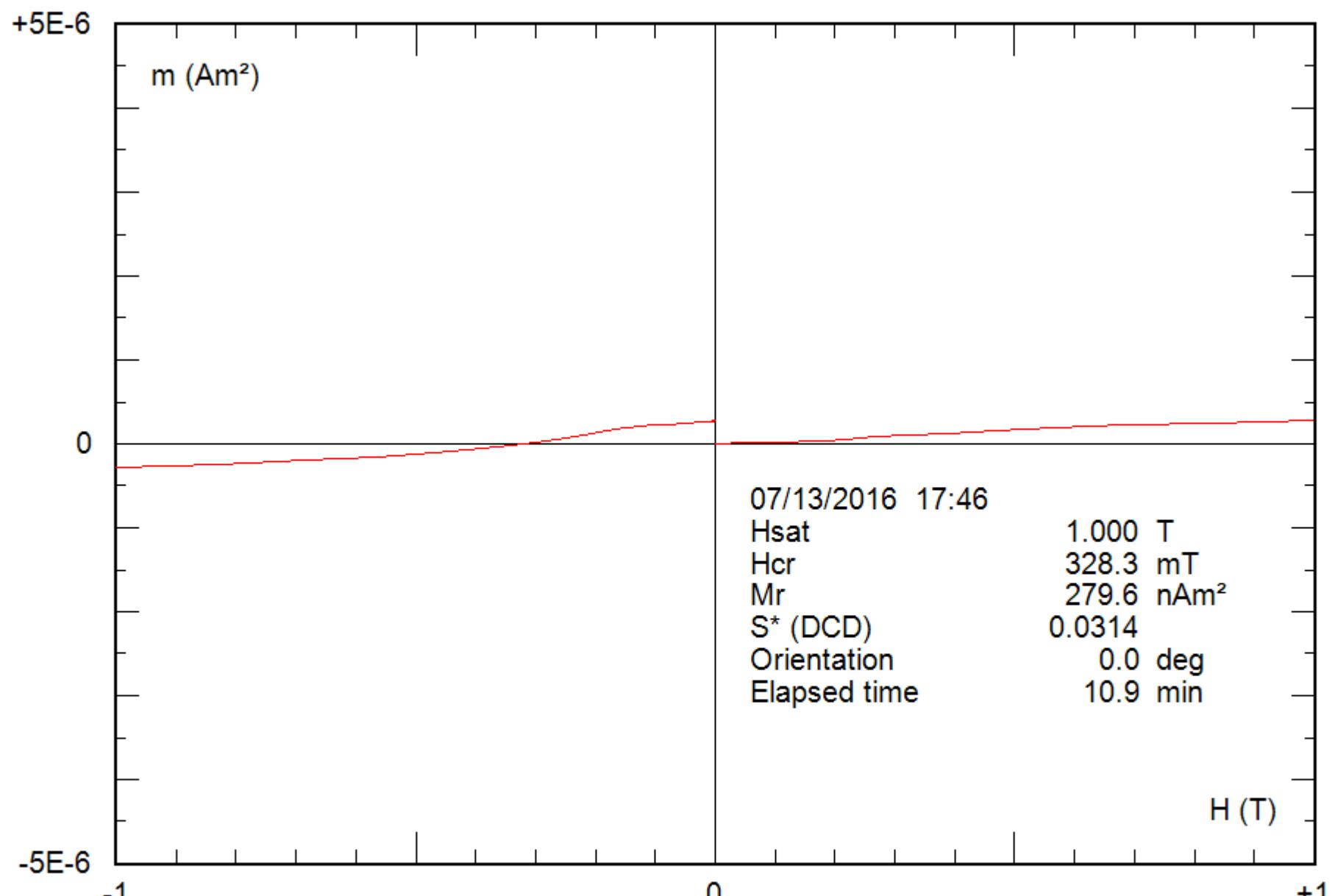
0.23 g

File: 34-03b\_hyst



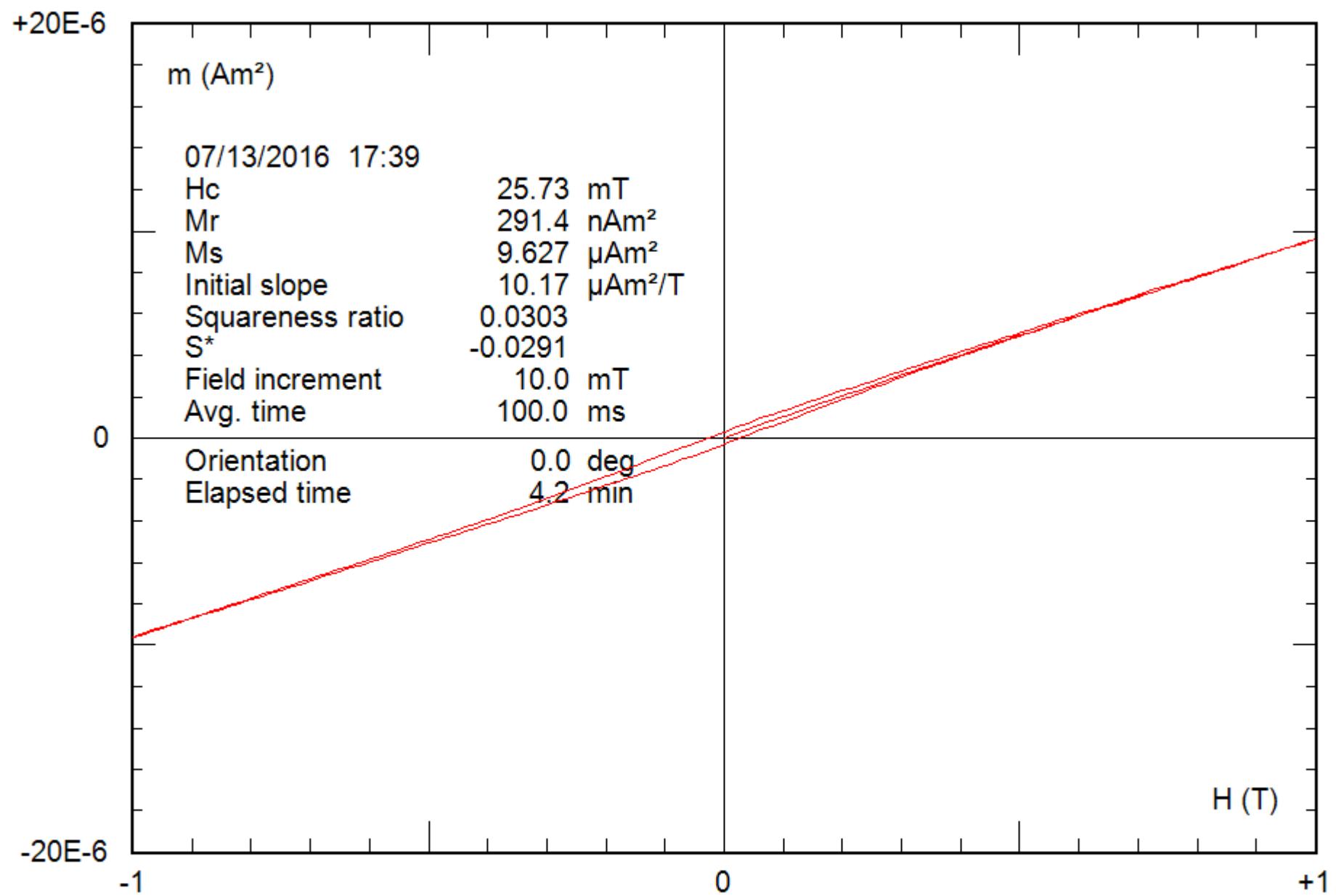
0.23 g

File: 34-03b\_hyst\_corr



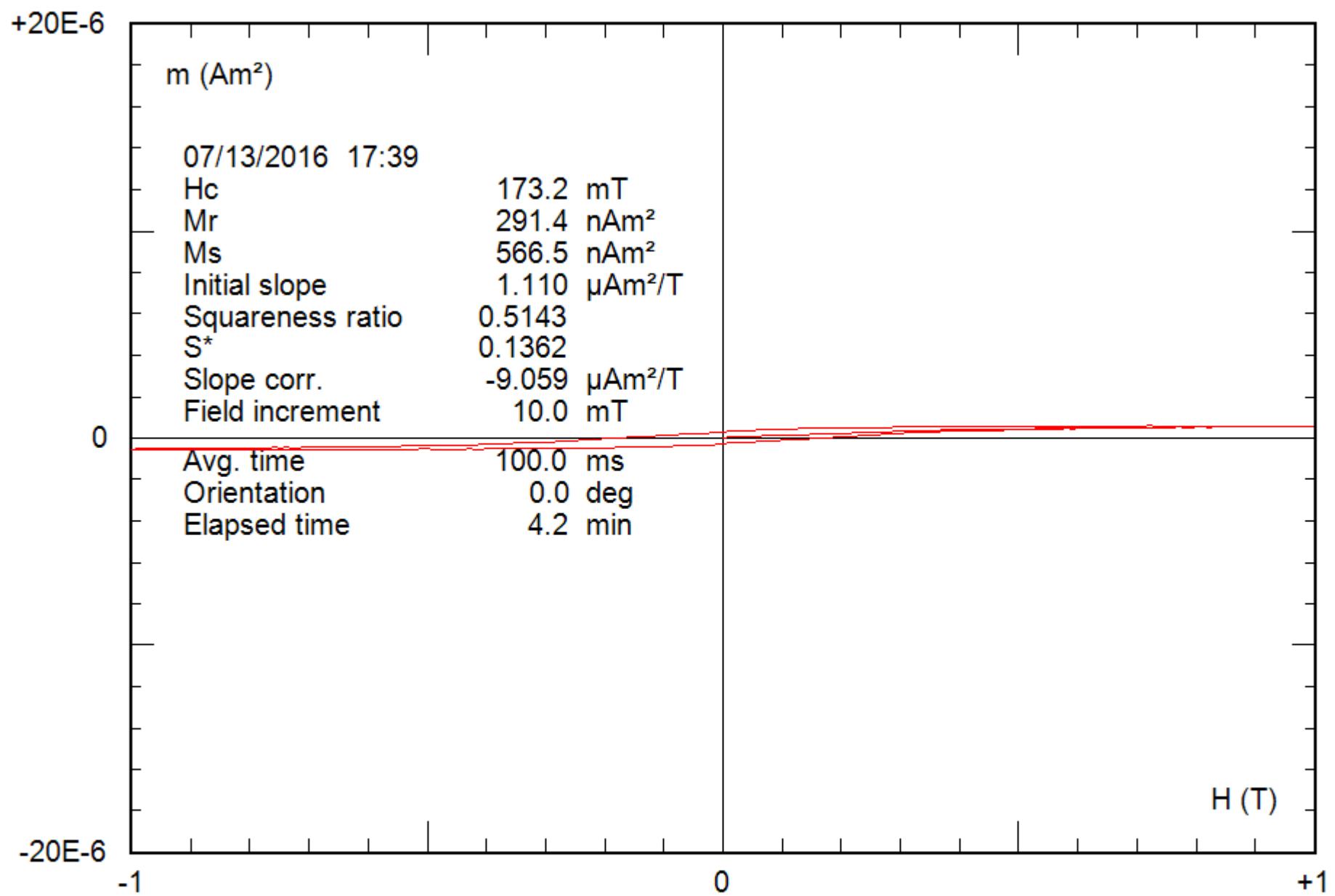
0.20 g

File: 35-02b\_bcr



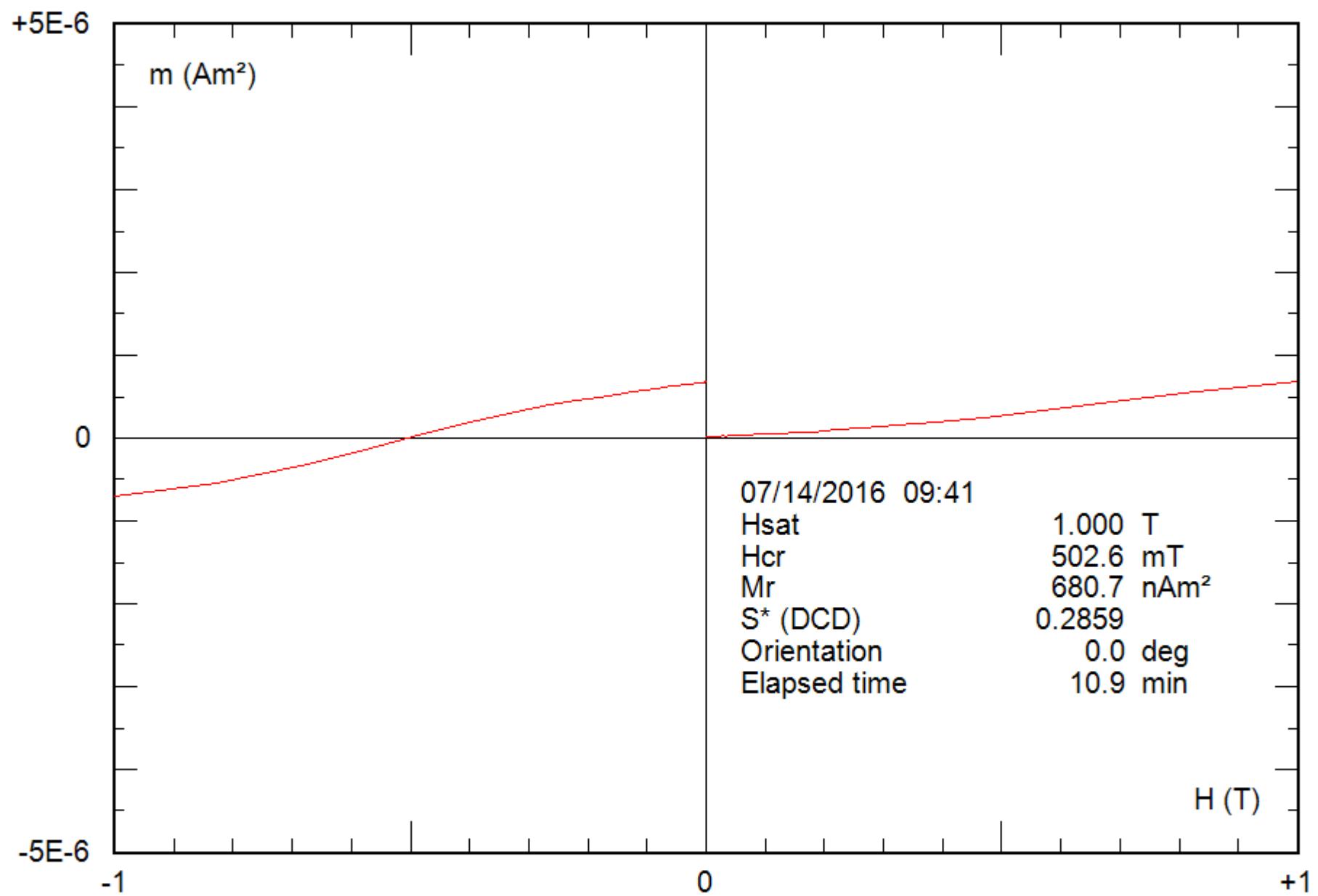
0.20 g

File: 35-02b\_hyst



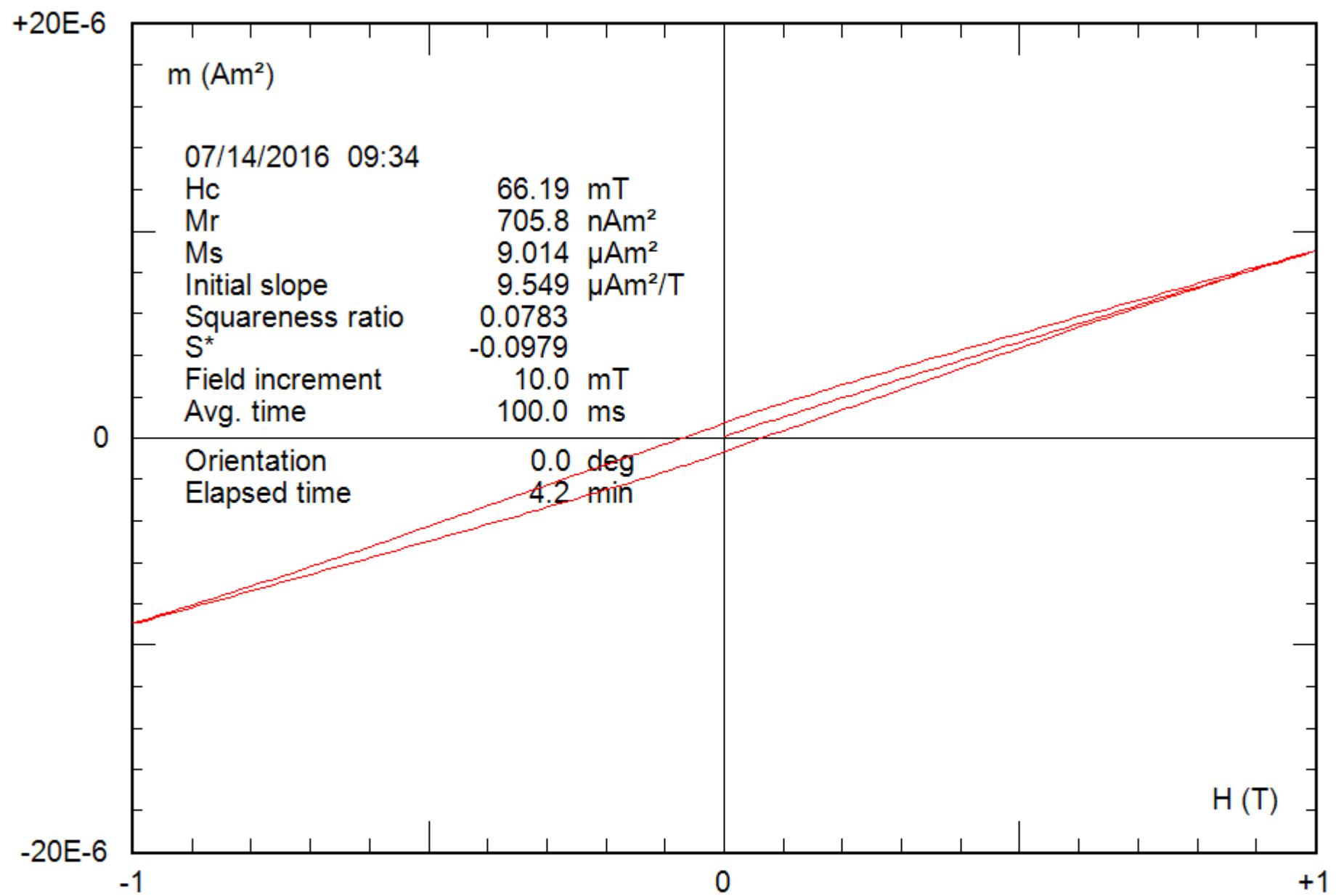
0.20 g

File: 35-02b\_hyst\_corr



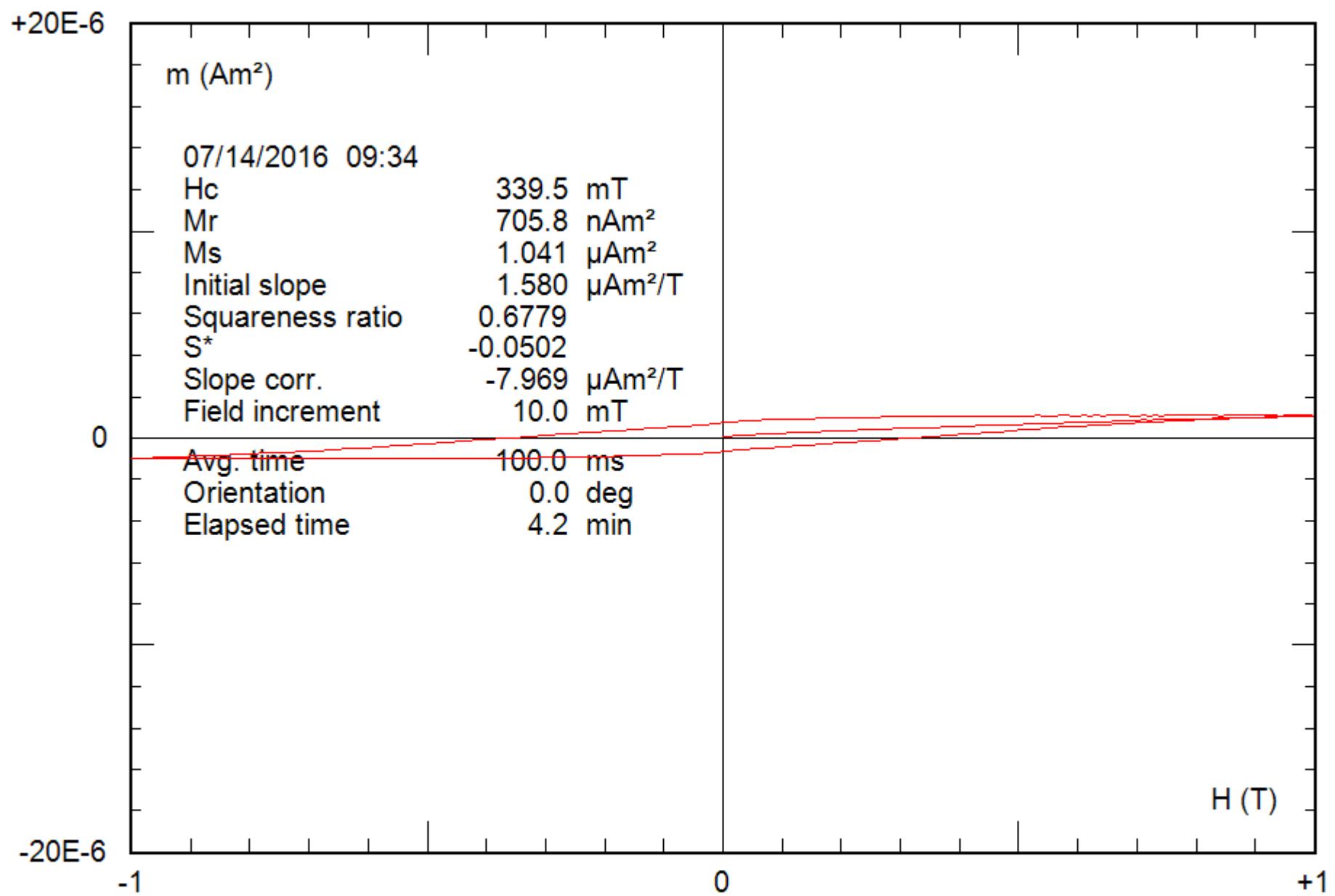
0.27 g

File: 36-07c\_bcr



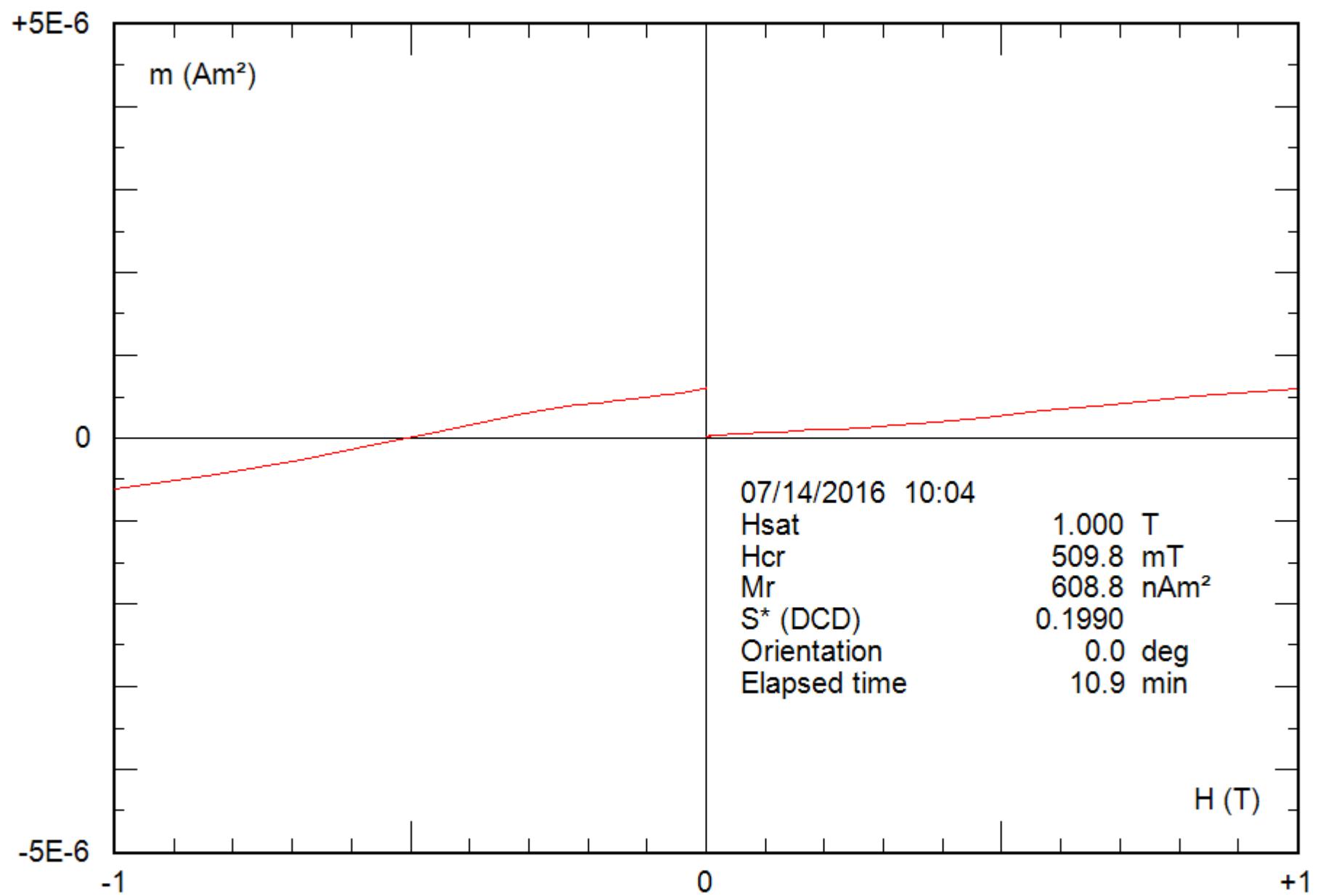
0.27 g

File: 36-07c\_hyst



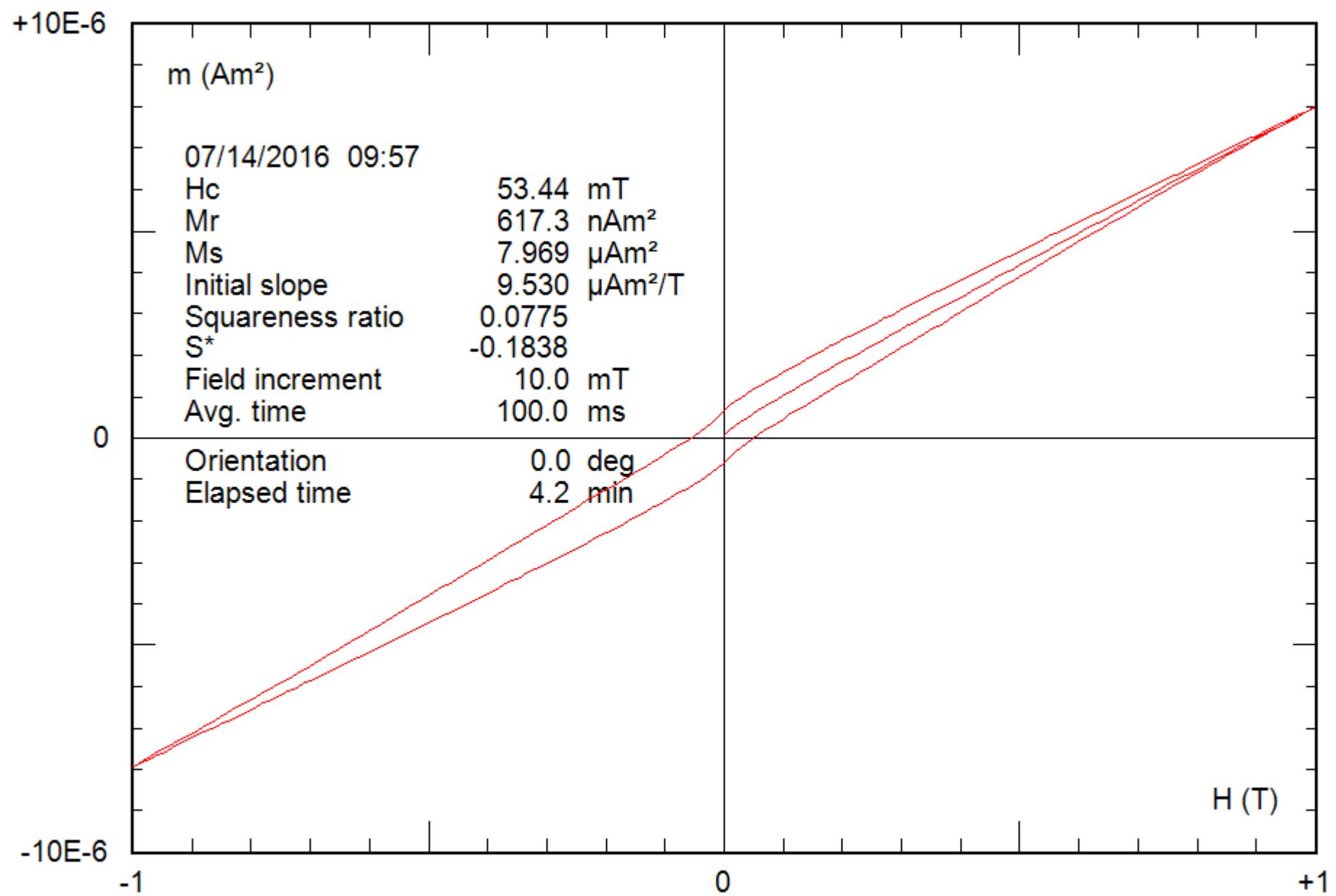
0.27 g

File: 36-07c\_hyst\_corr



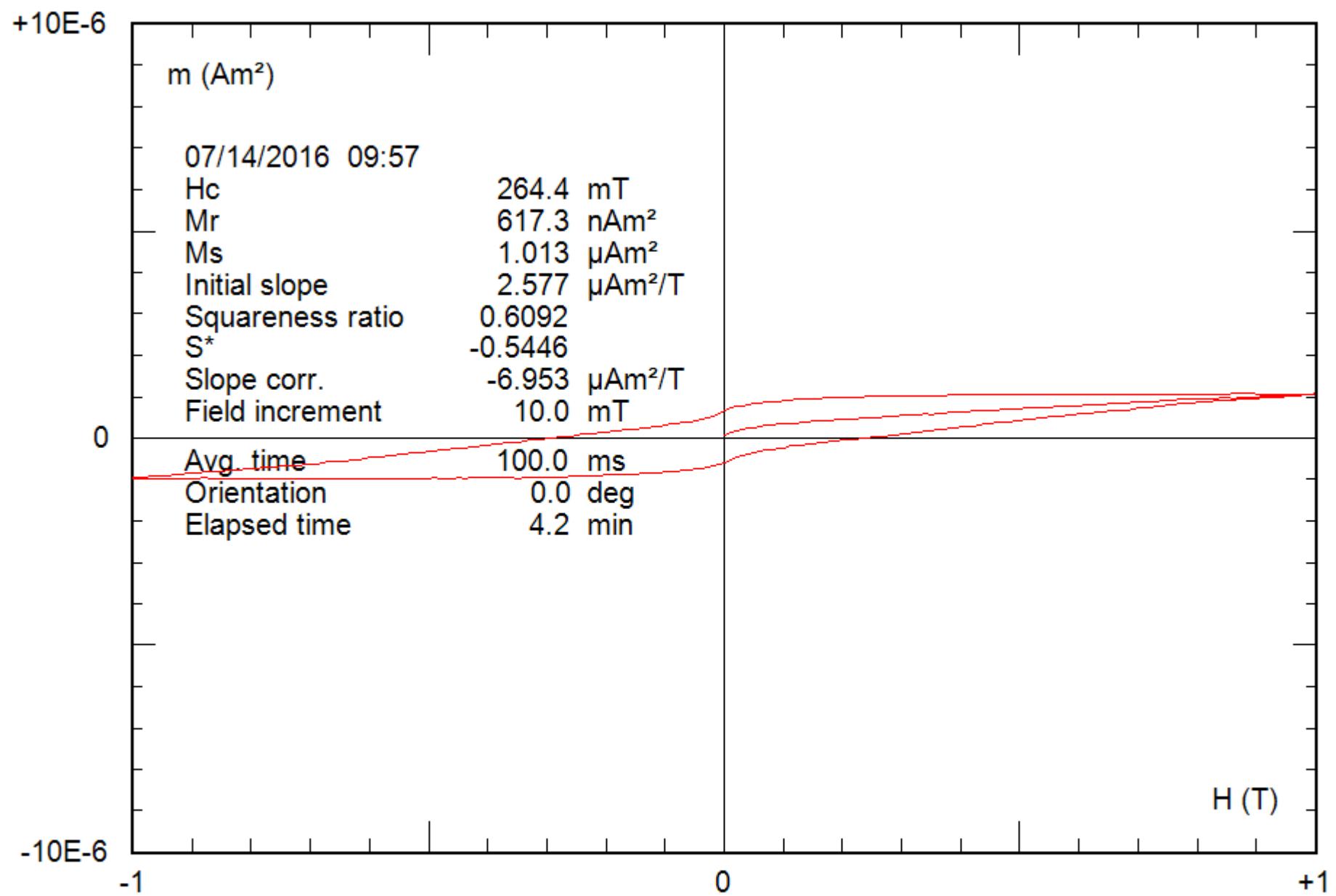
0.29 g

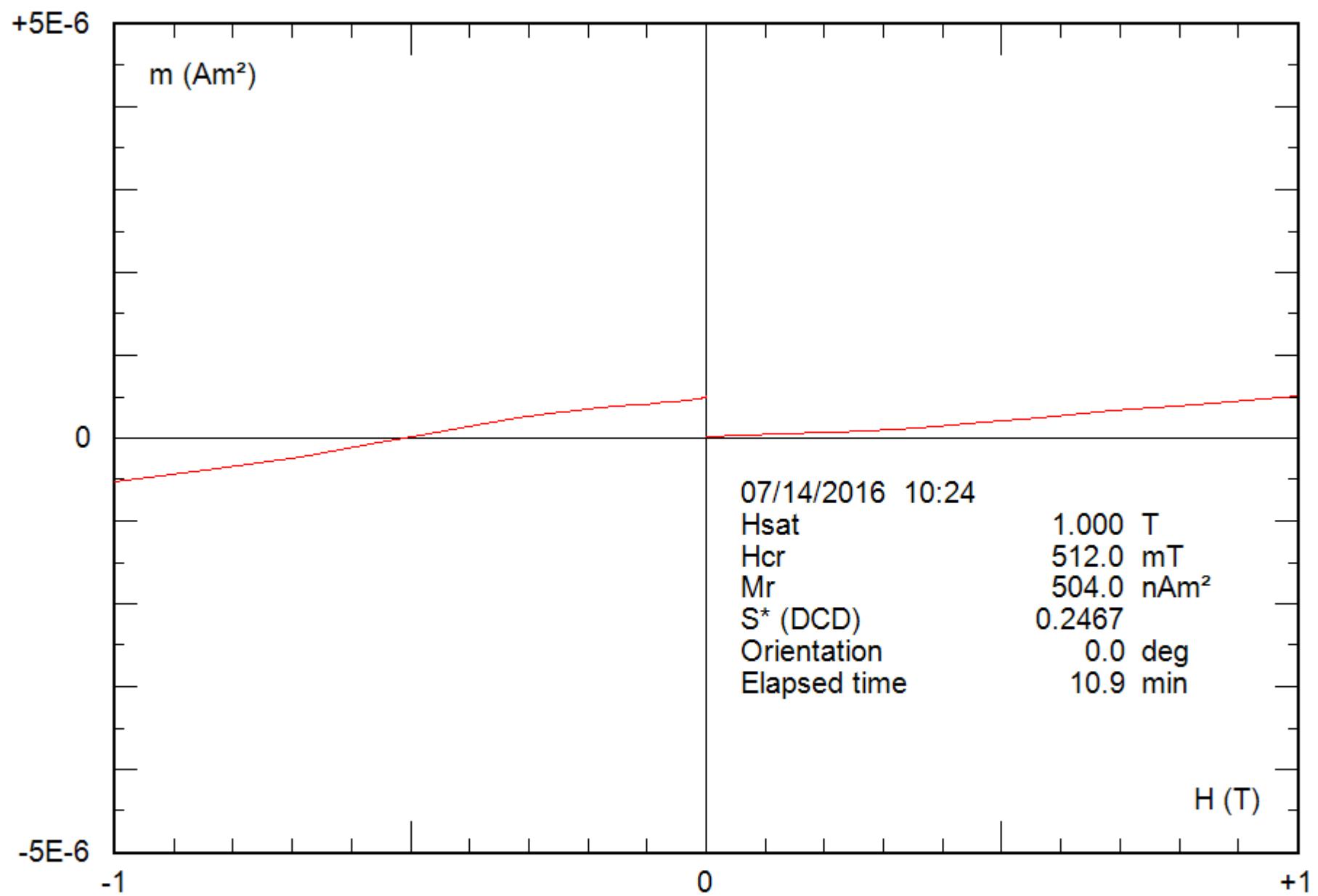
File: 37-03c\_bcr



0.29 g

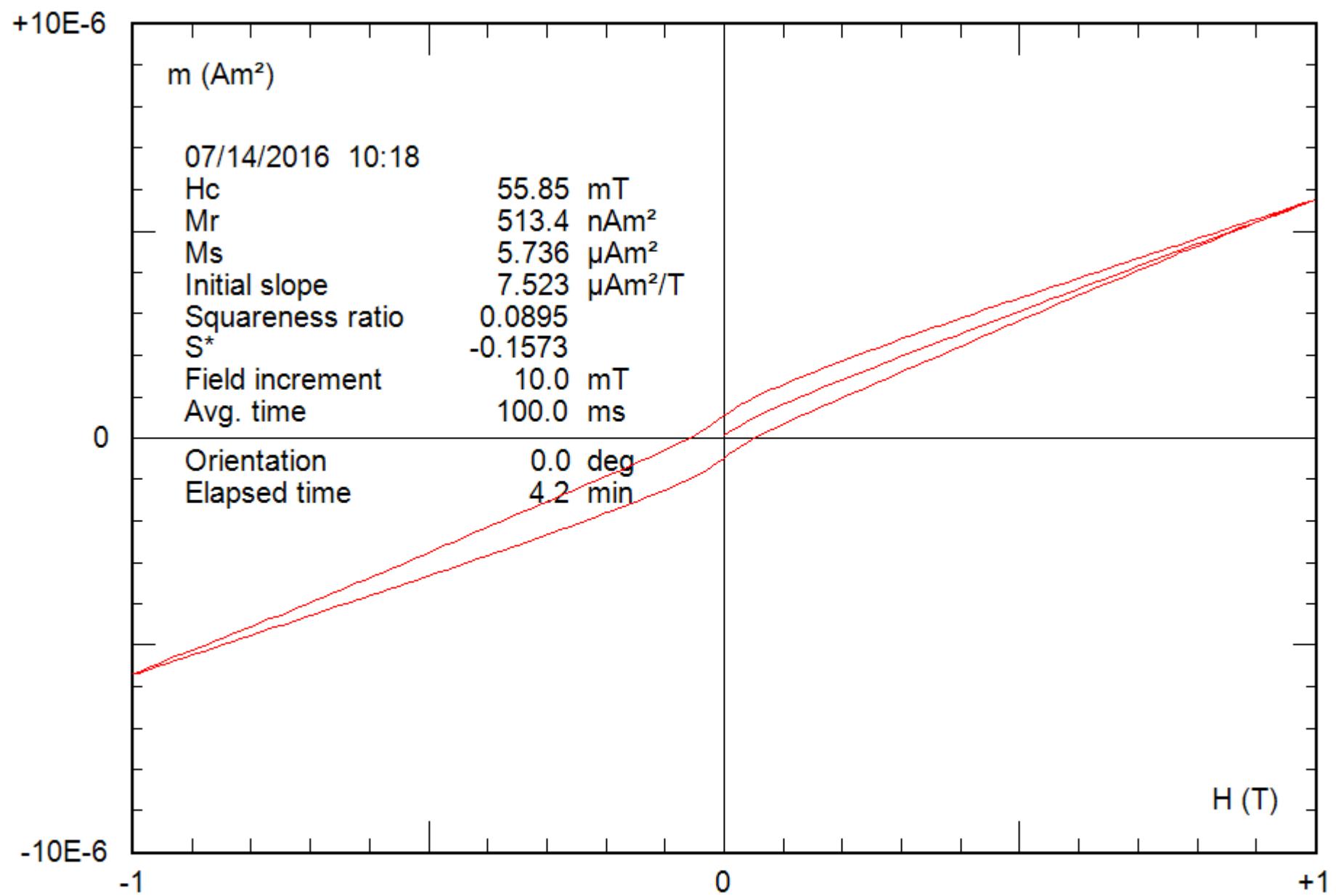
File: 37-03c\_hyst





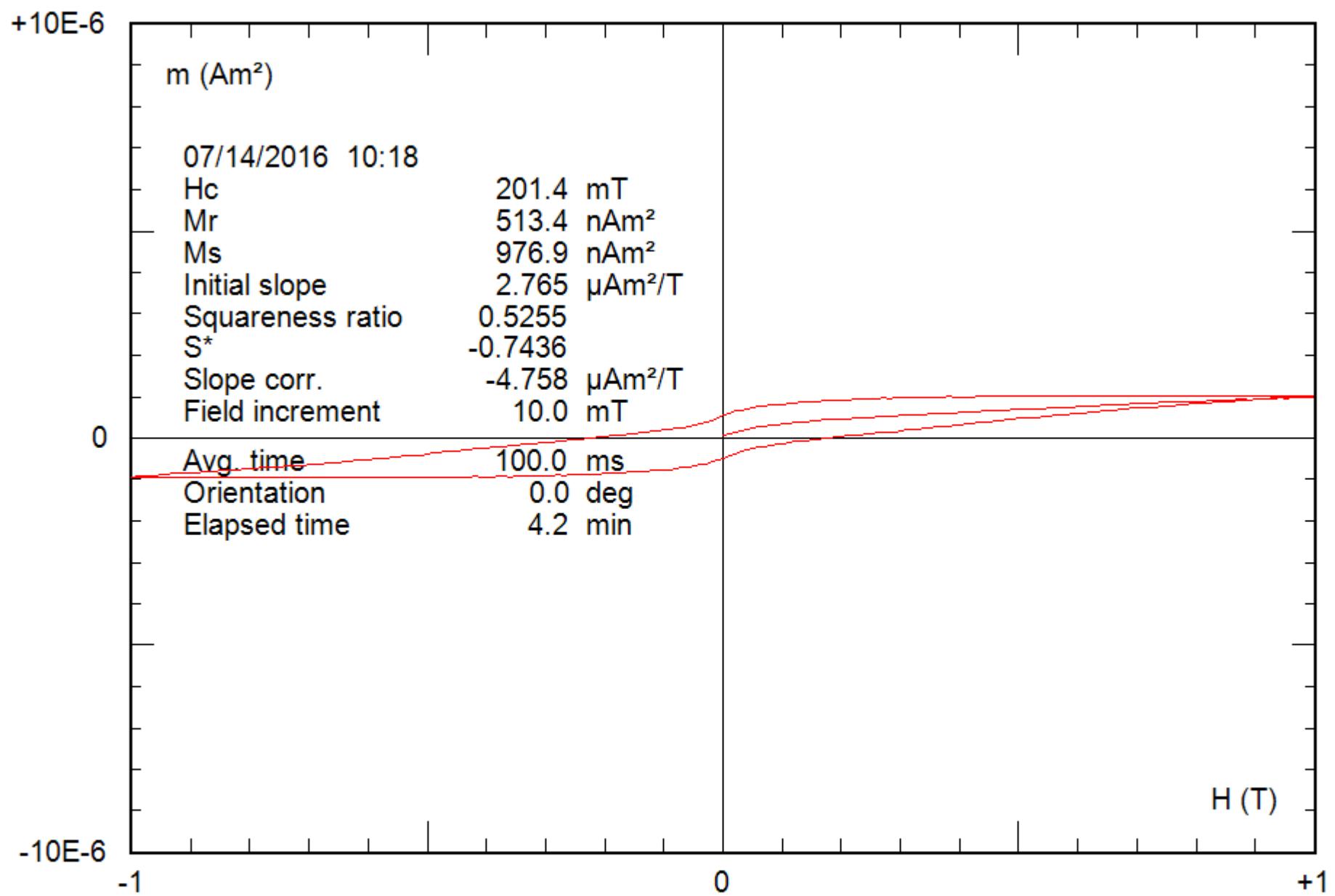
0.26 g

File: 38-04c\_bcr



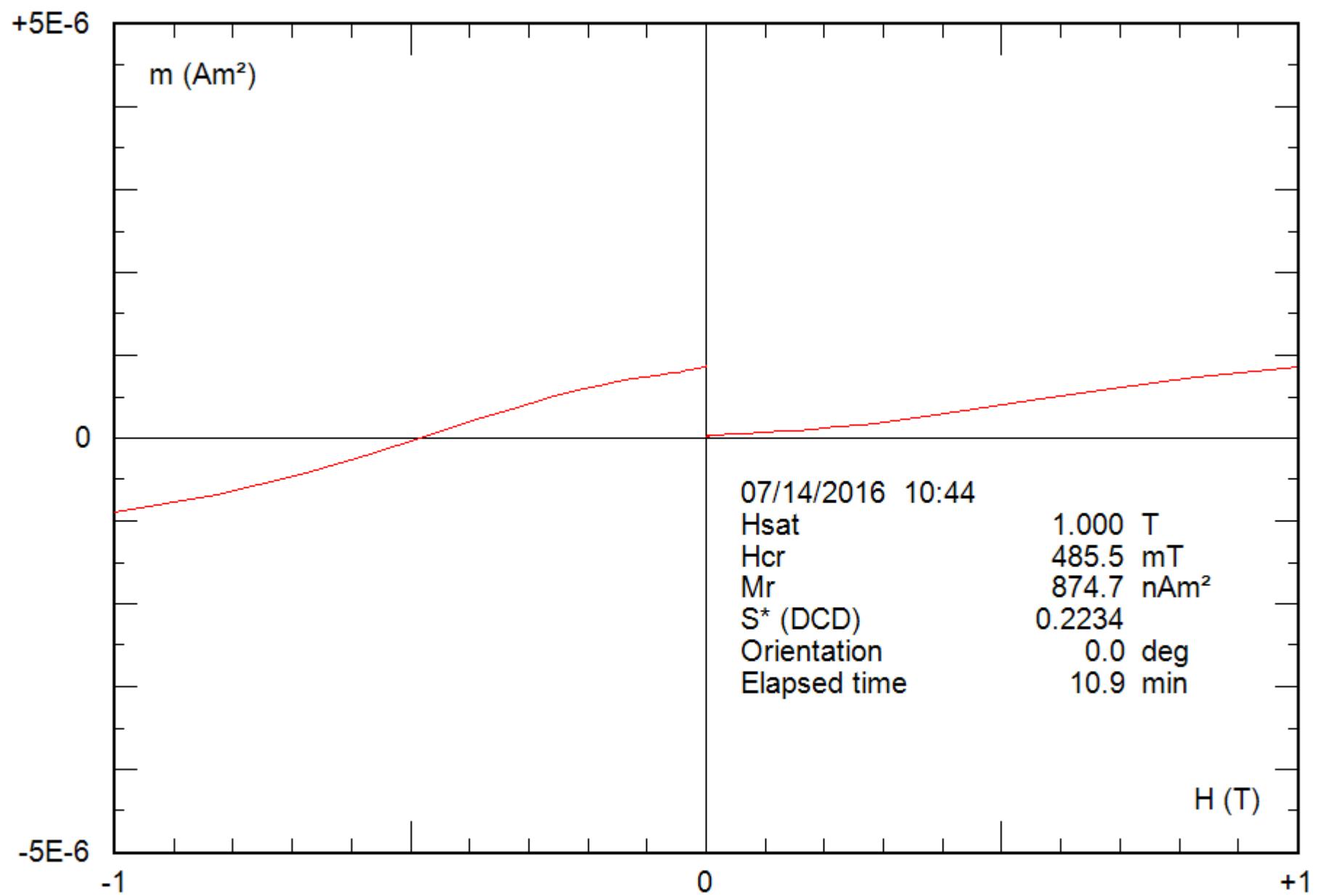
0.26 g

File: 38-04c\_hyst



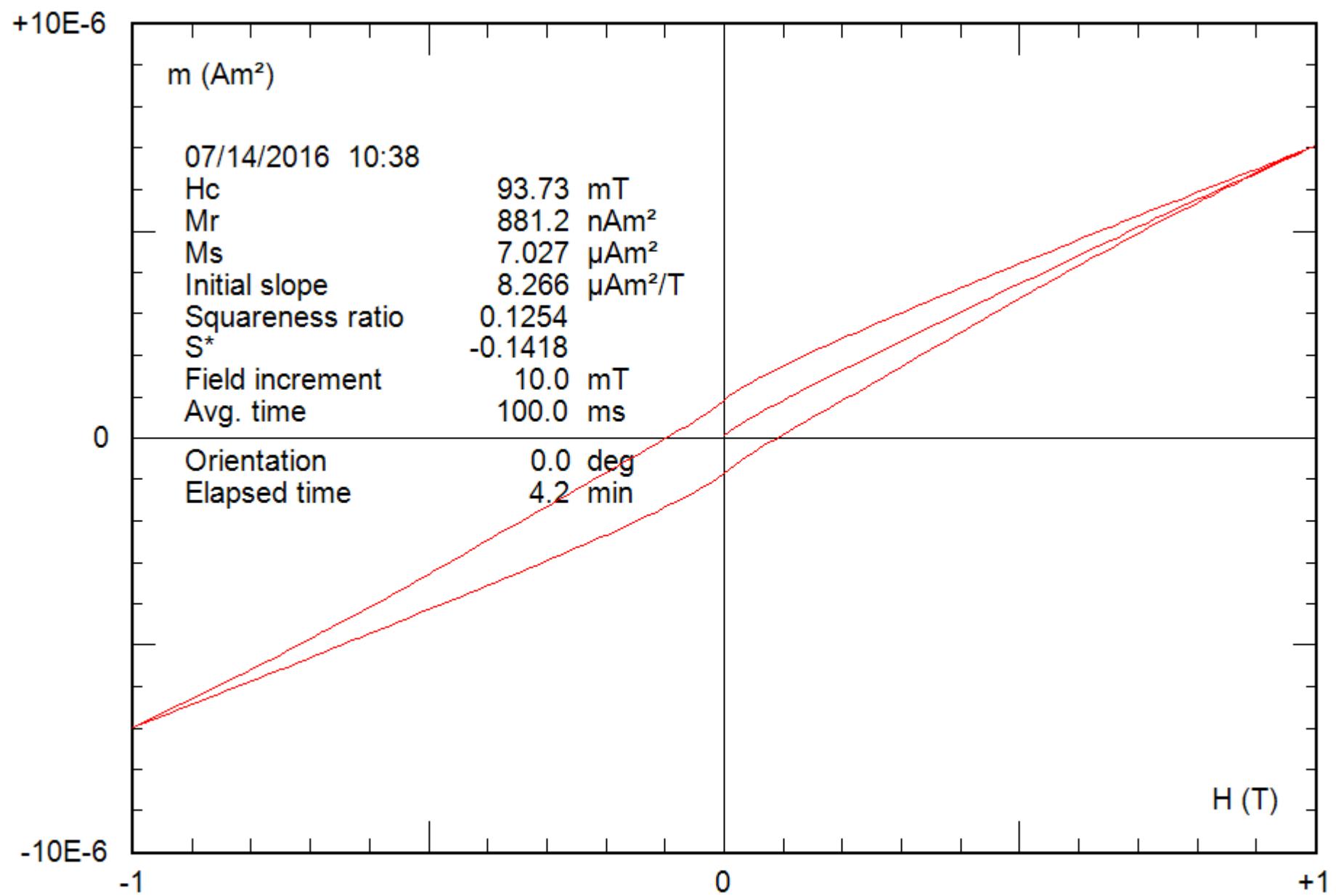
0.26 g

File: 38-04c\_hyst\_corr



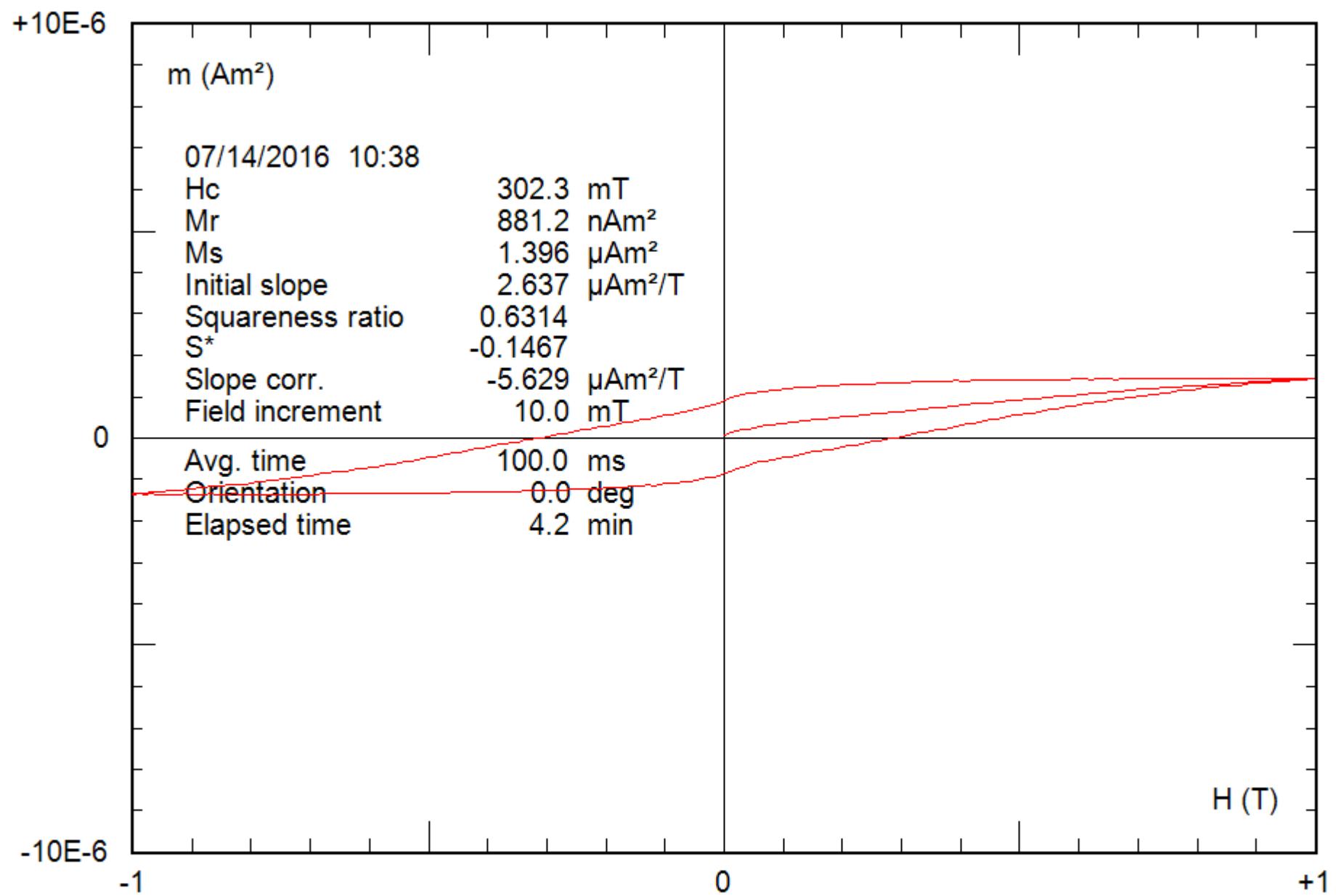
0.24 g

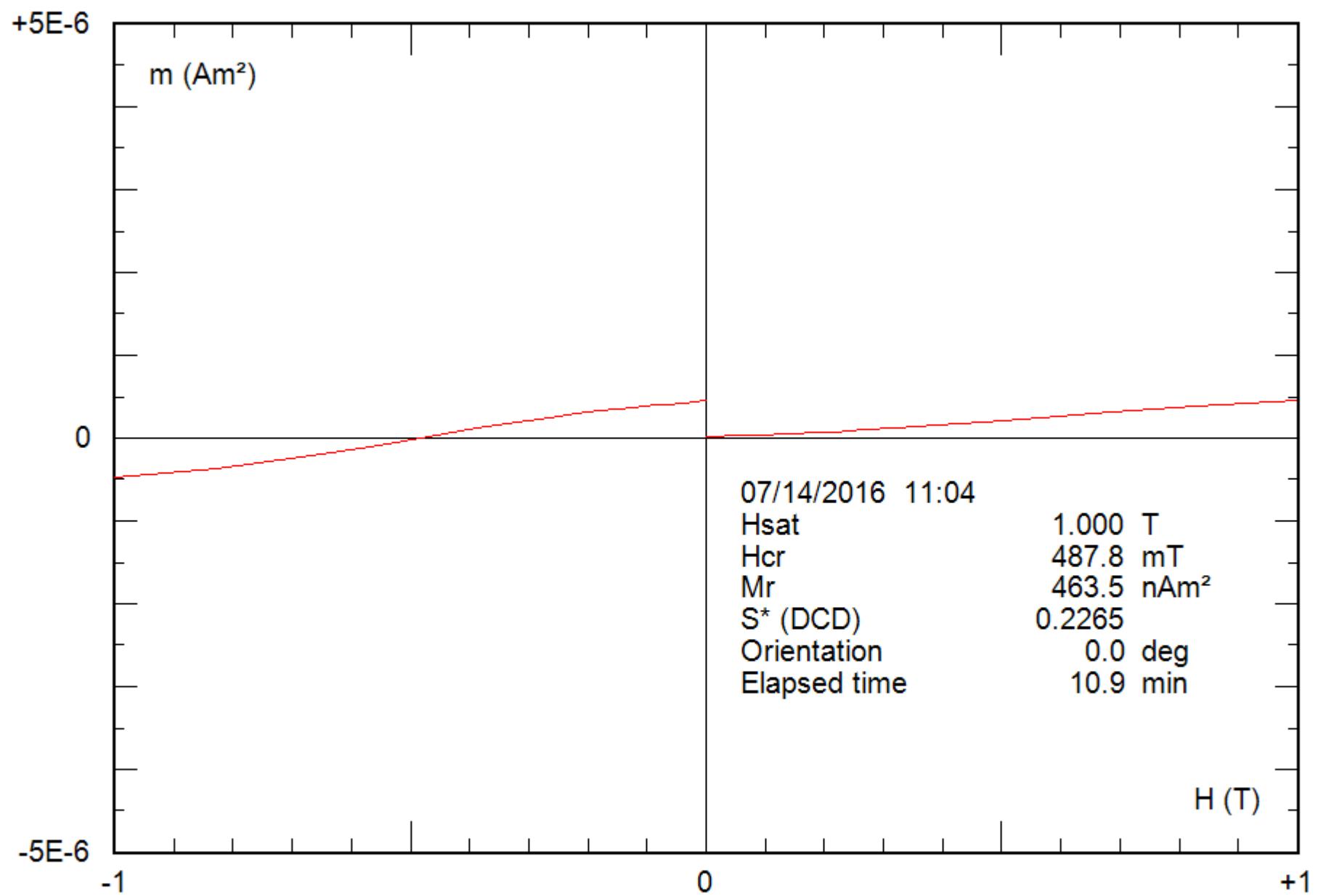
File: 39-04b\_bcr



0.24 g

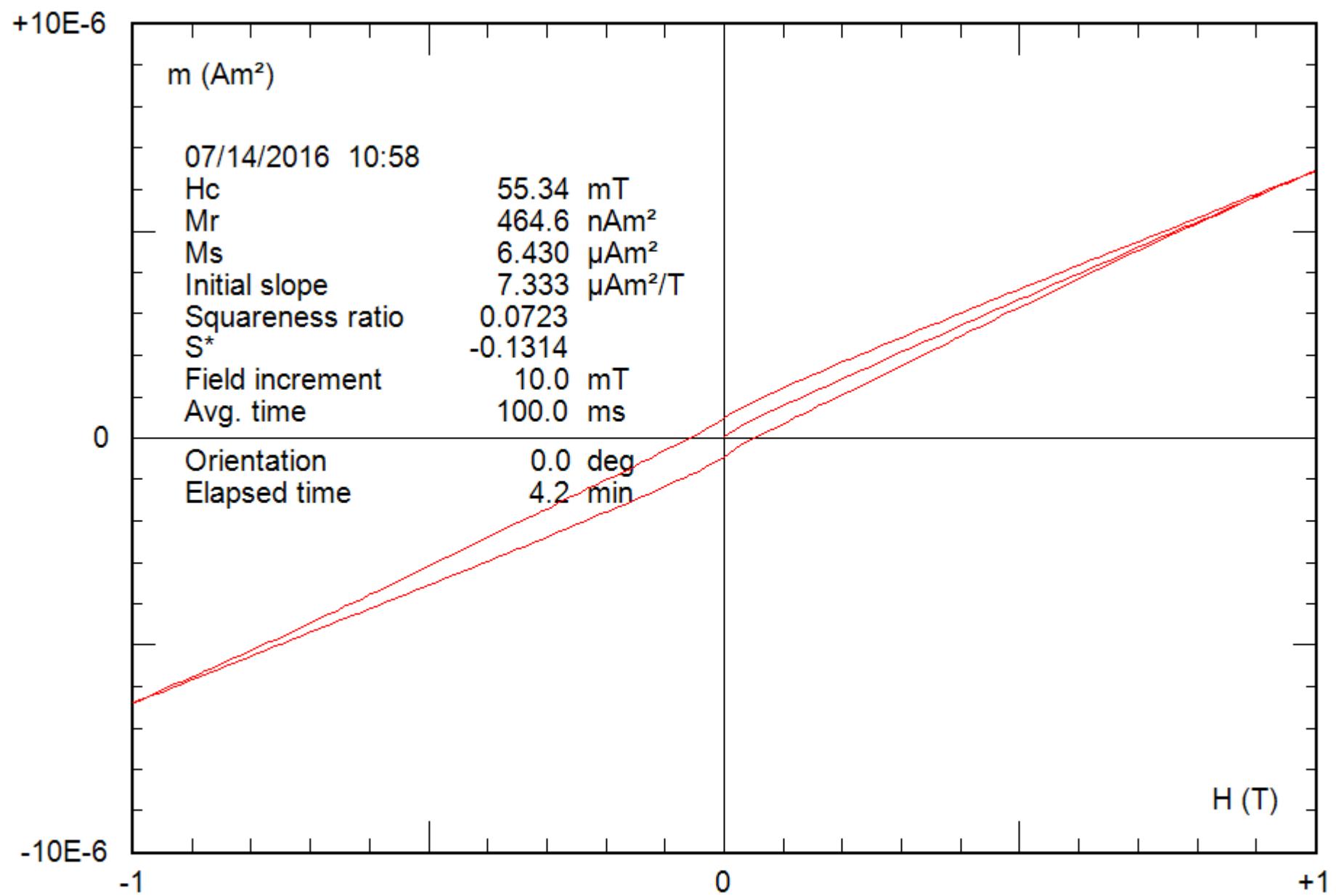
File: 39-04b\_hyst



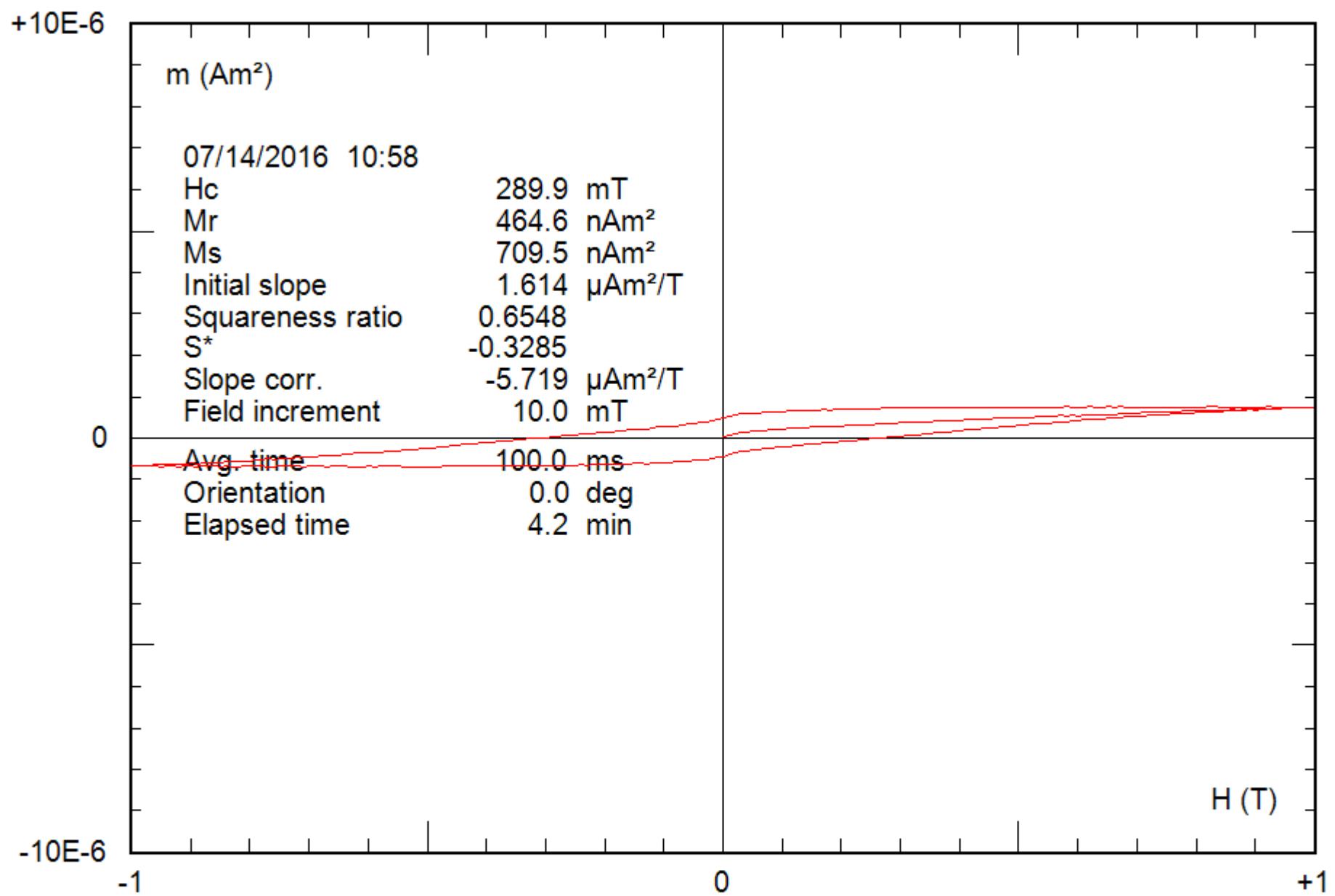


0.22 g

File: 40-05c\_bcr

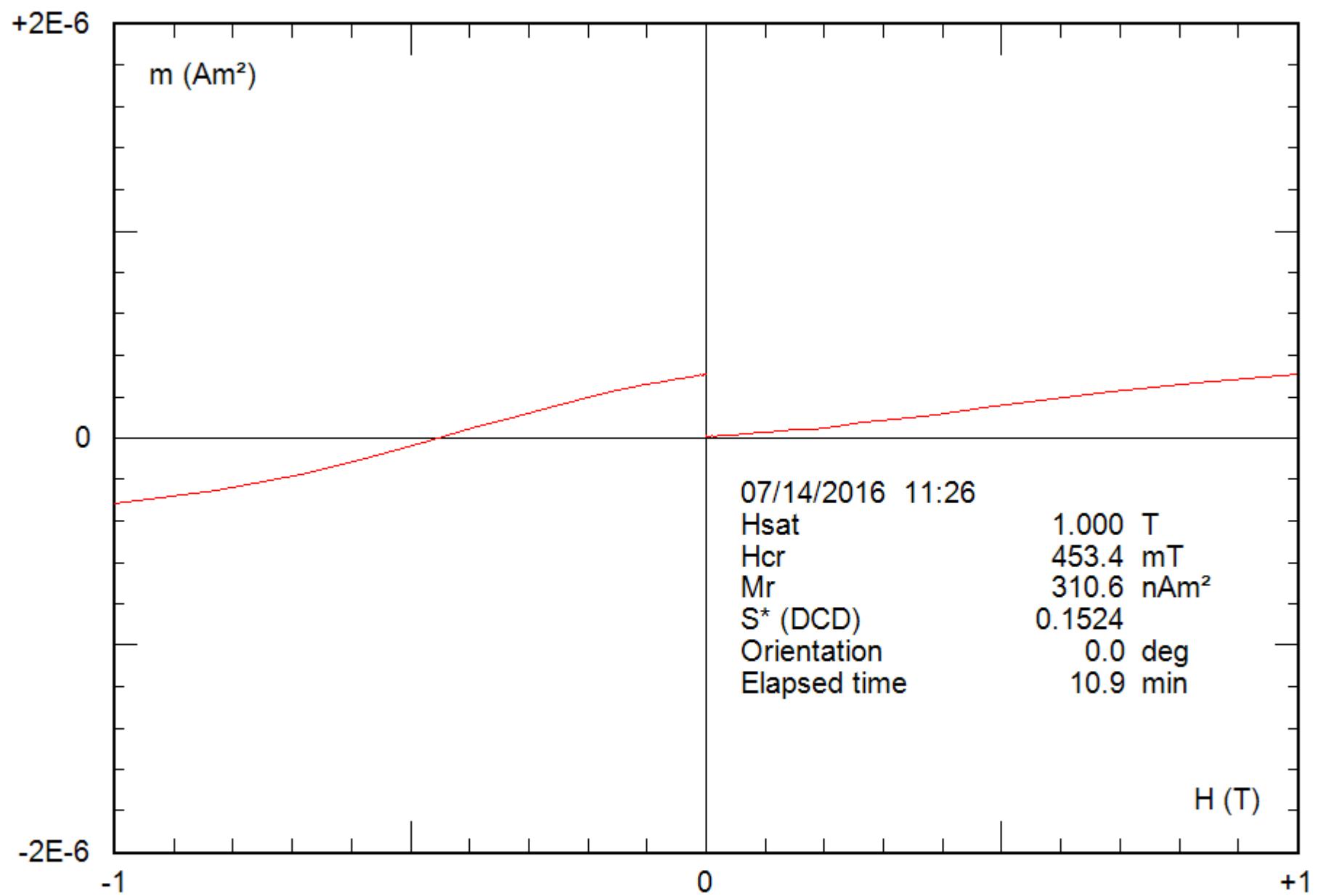


0.22 g  
File: 40-05c\_hyst



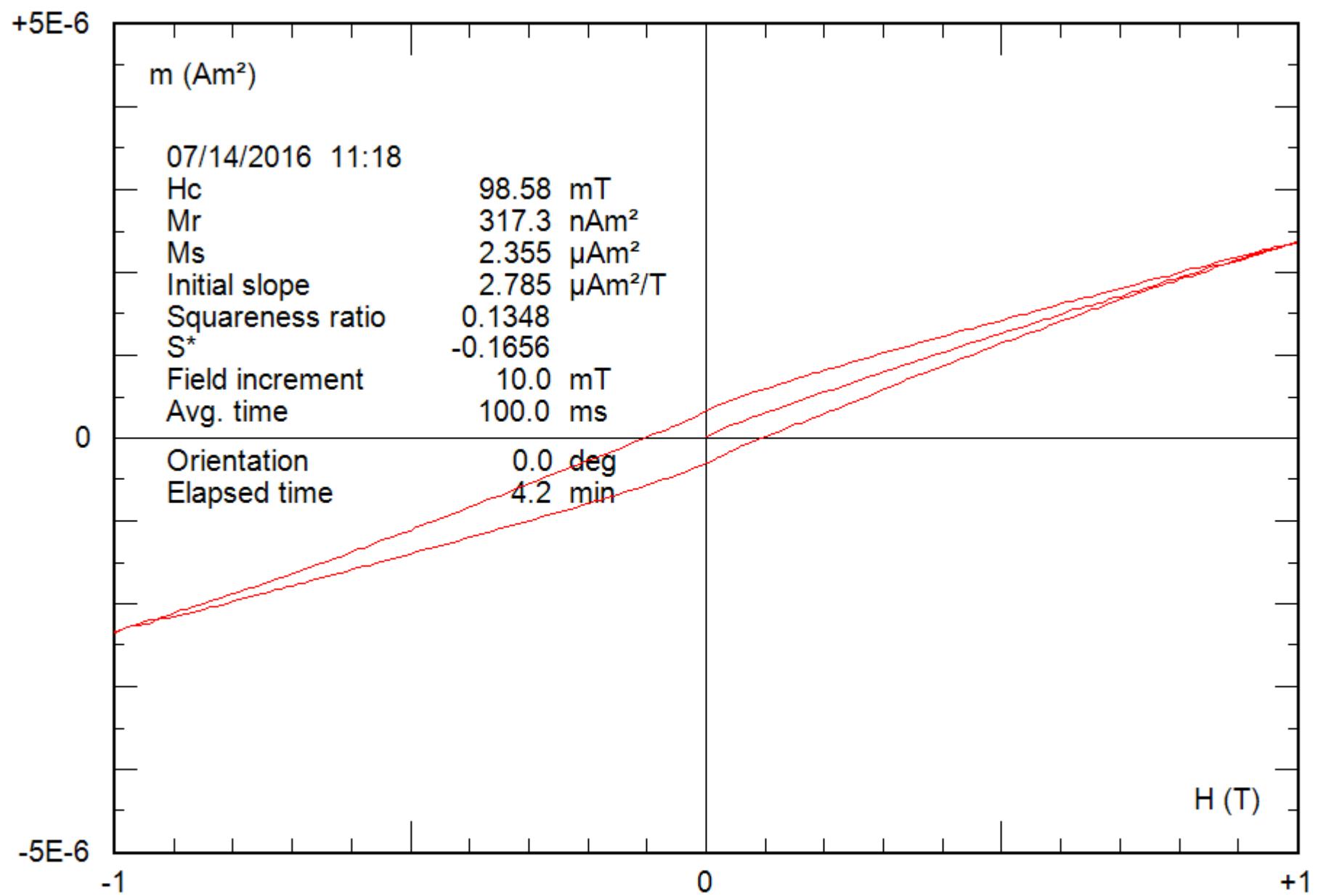
0.22 g

File: 40-05c\_hyst\_corr



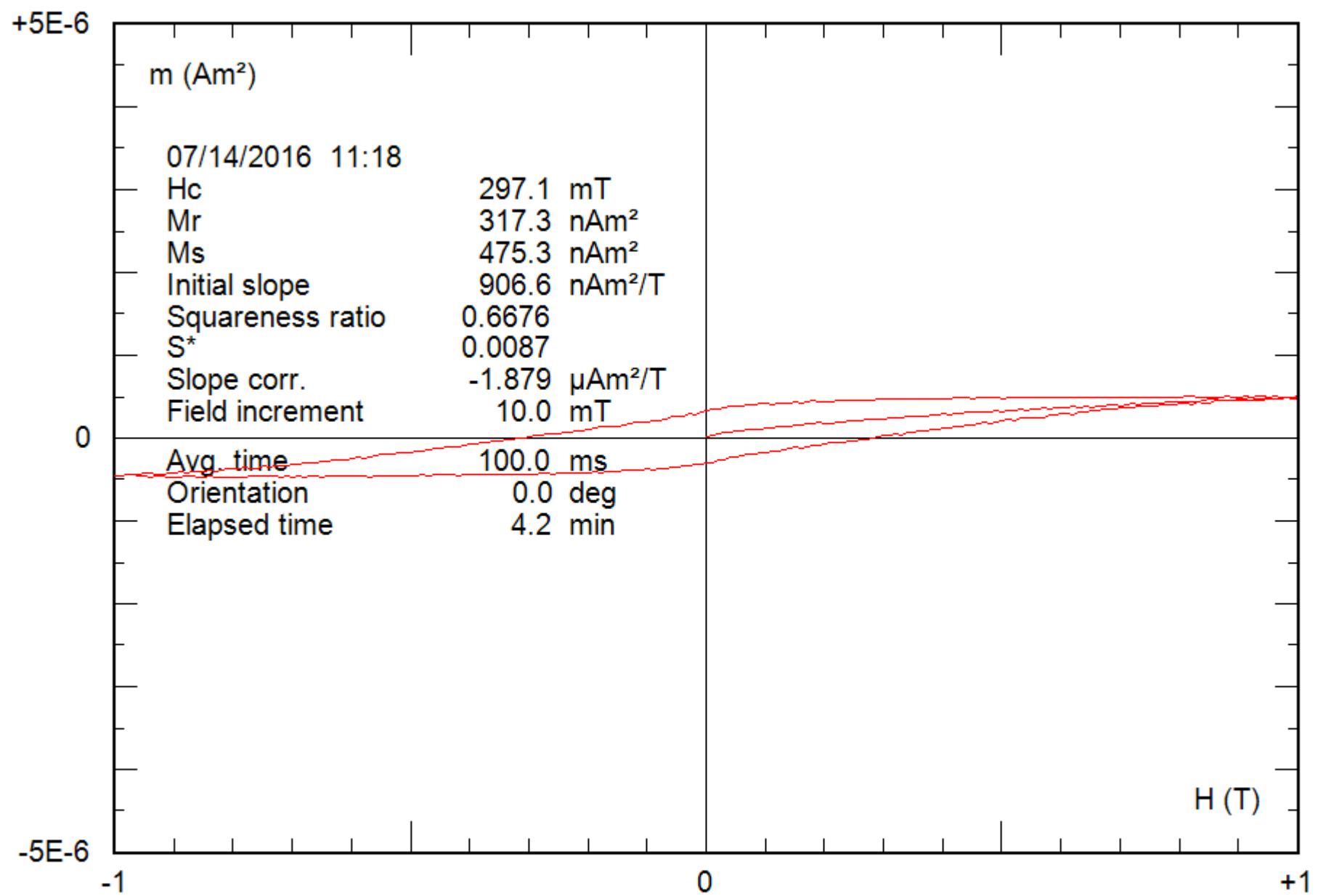
0.26 g

File: 41-04c\_bcr



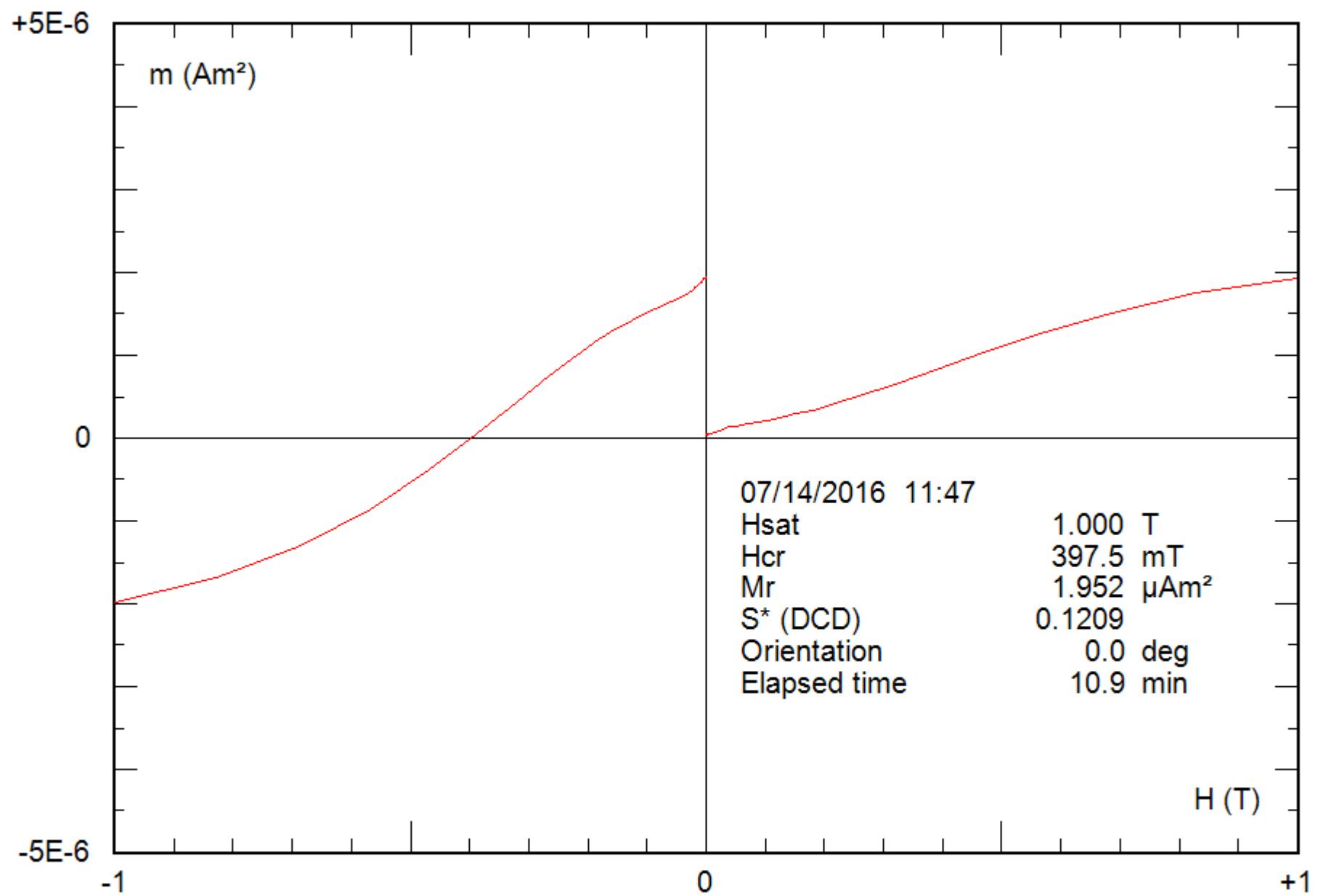
0.26 g

File: 41-04c\_hyst



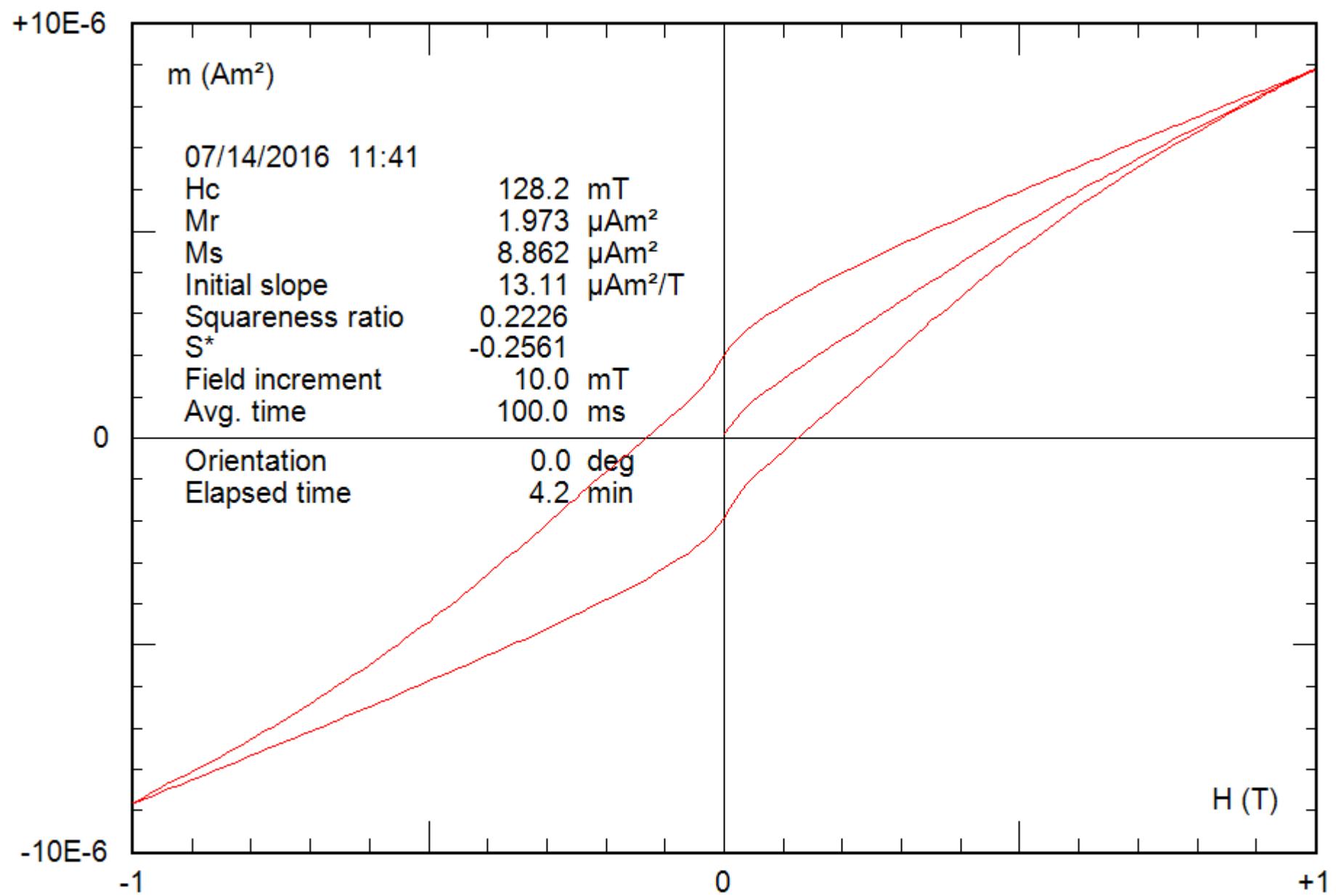
0.26 g

File: 41-04c\_hyst\_corr

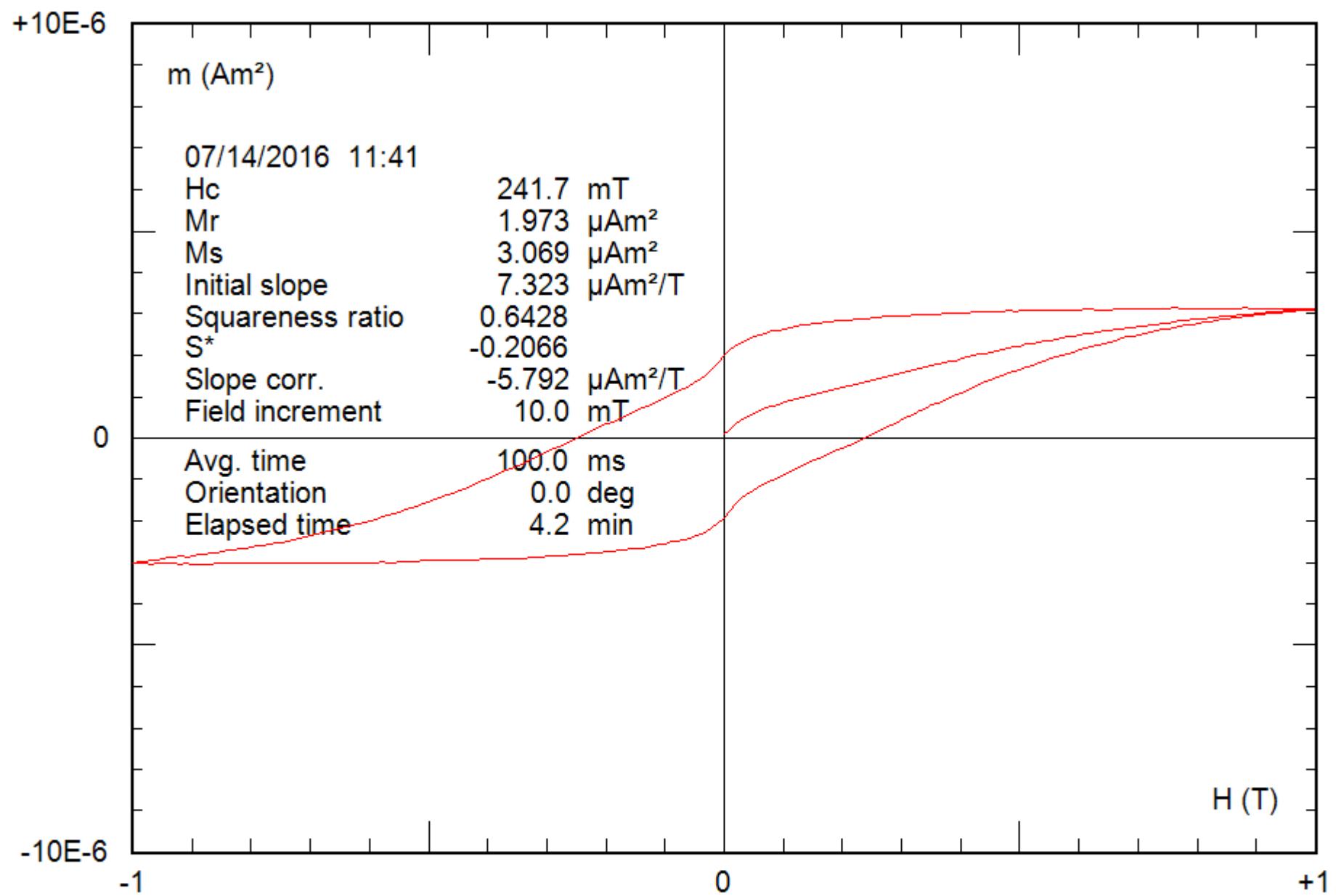


0.27 g

File: 42-03b\_bcr

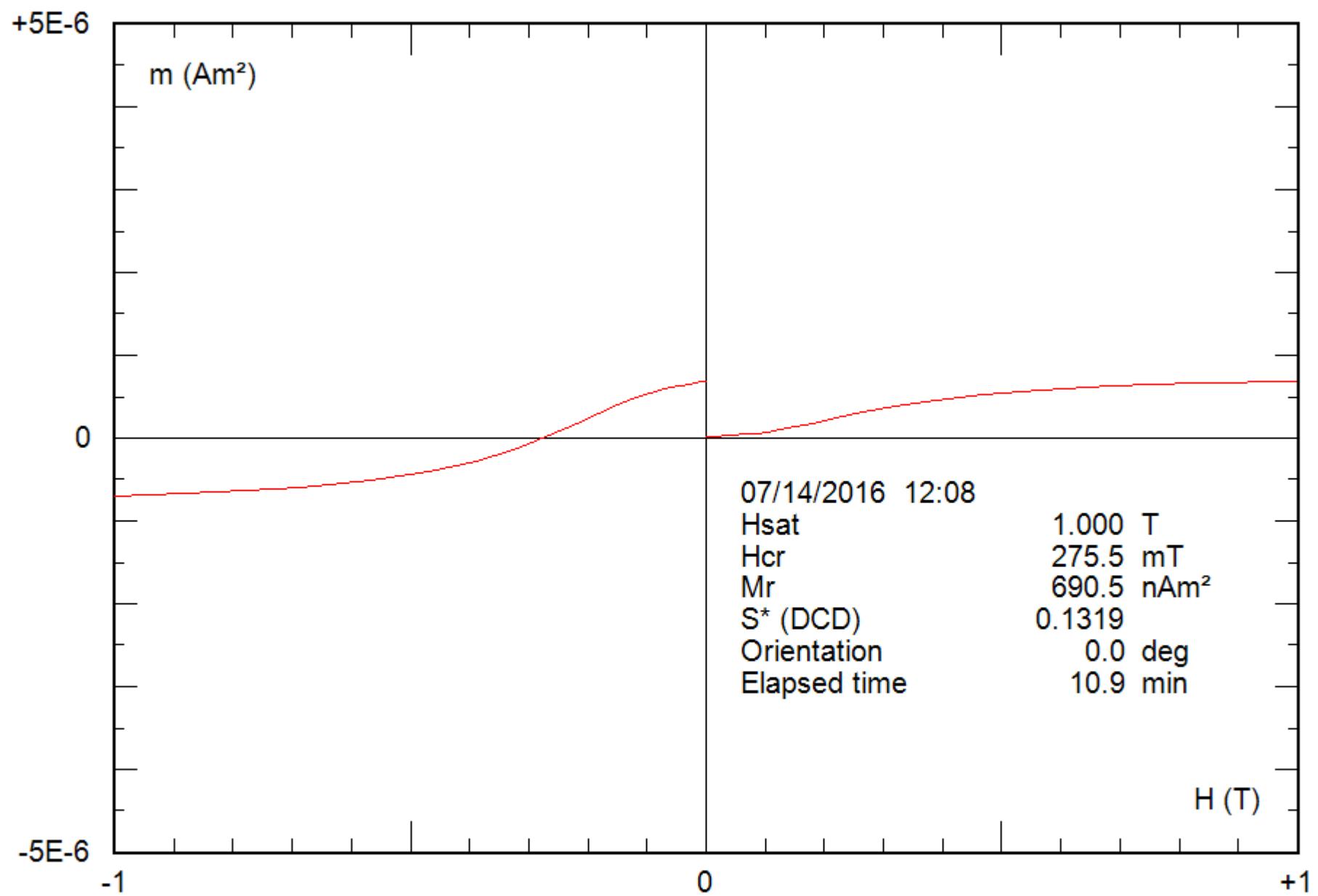


0.27 g  
File: 42-03b\_hyst



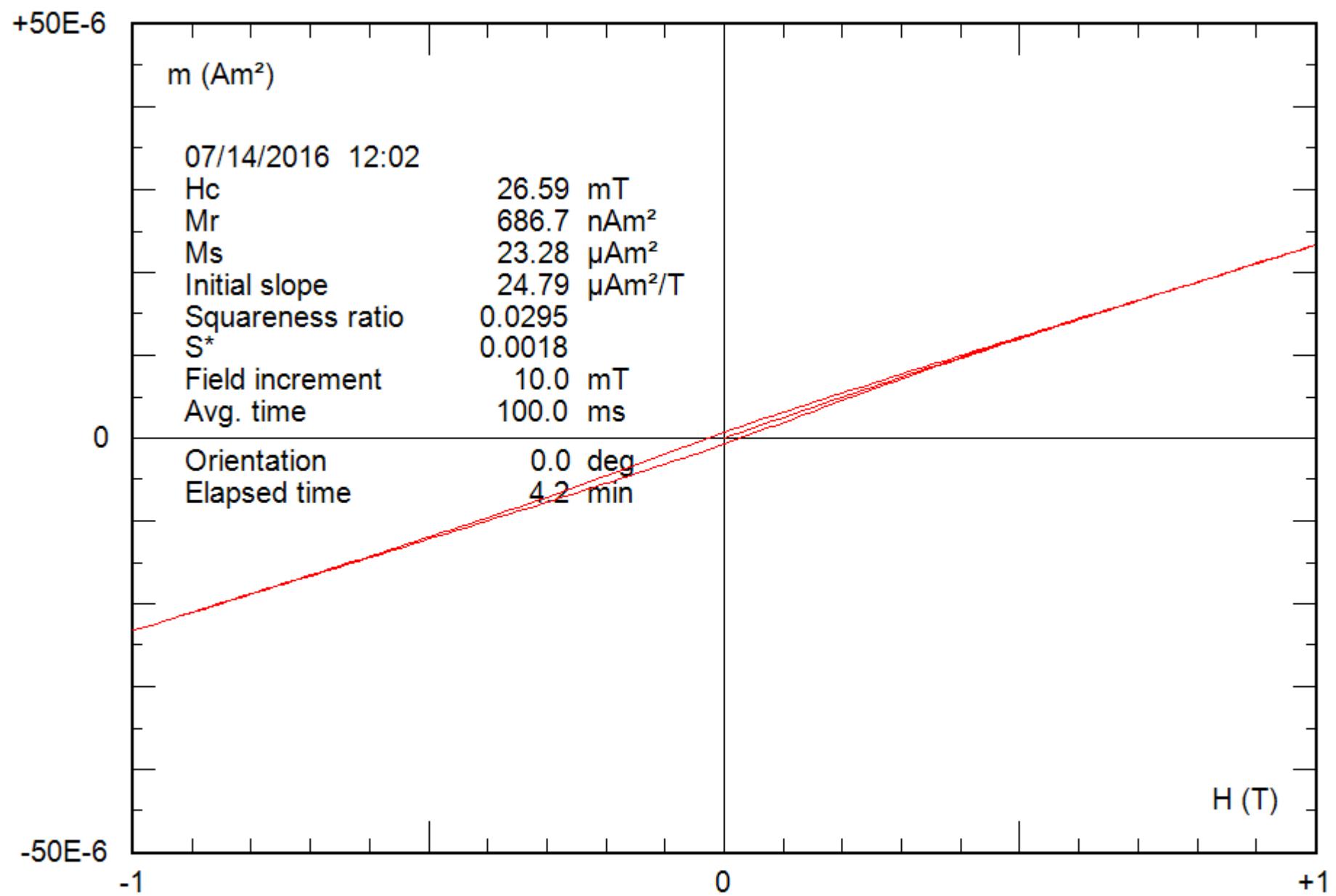
0.27 g

File: 42-03b\_hyst\_corr



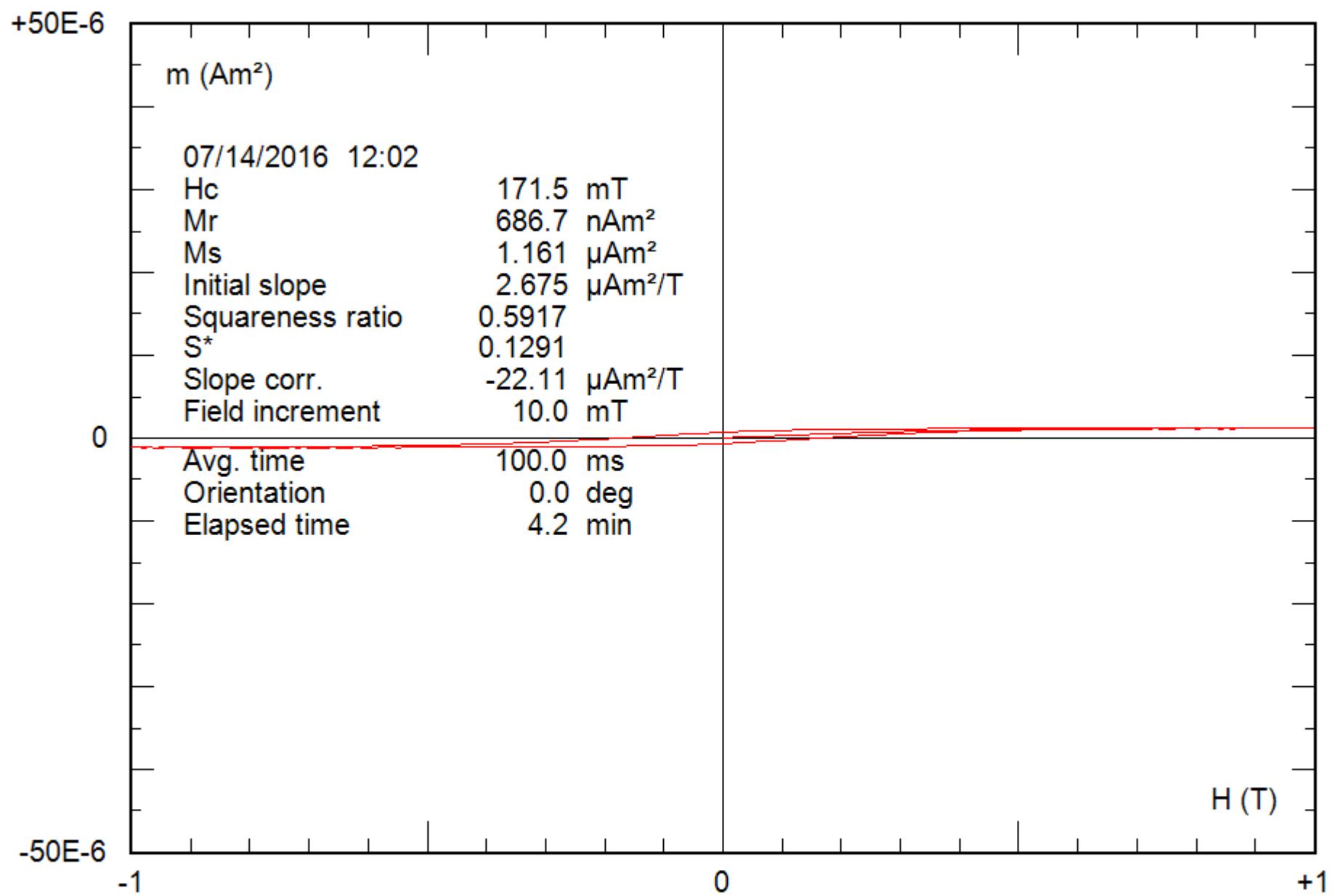
0.24 g

File: 43-05c\_bcr



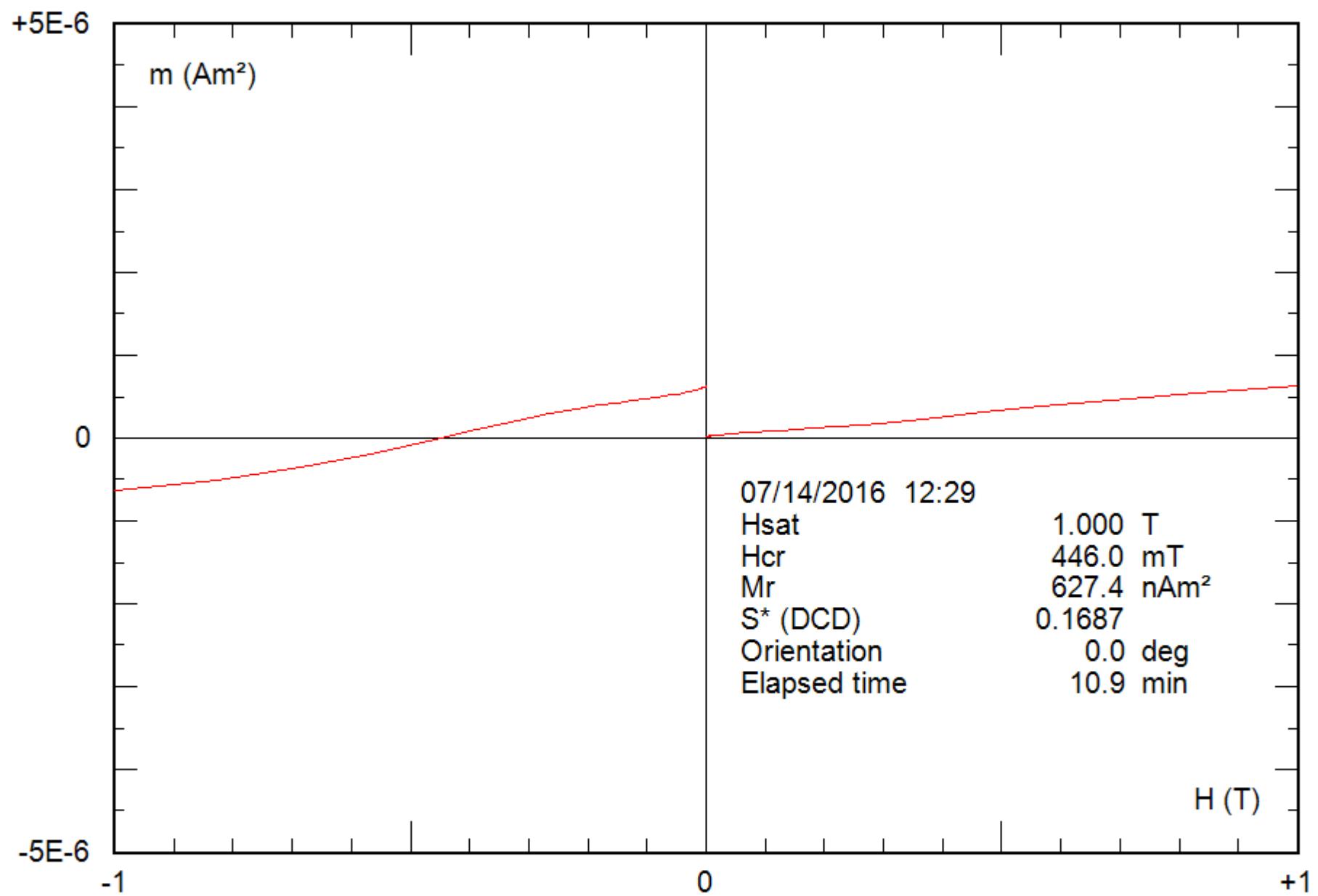
0.24 g

File: 43-05c\_hyst



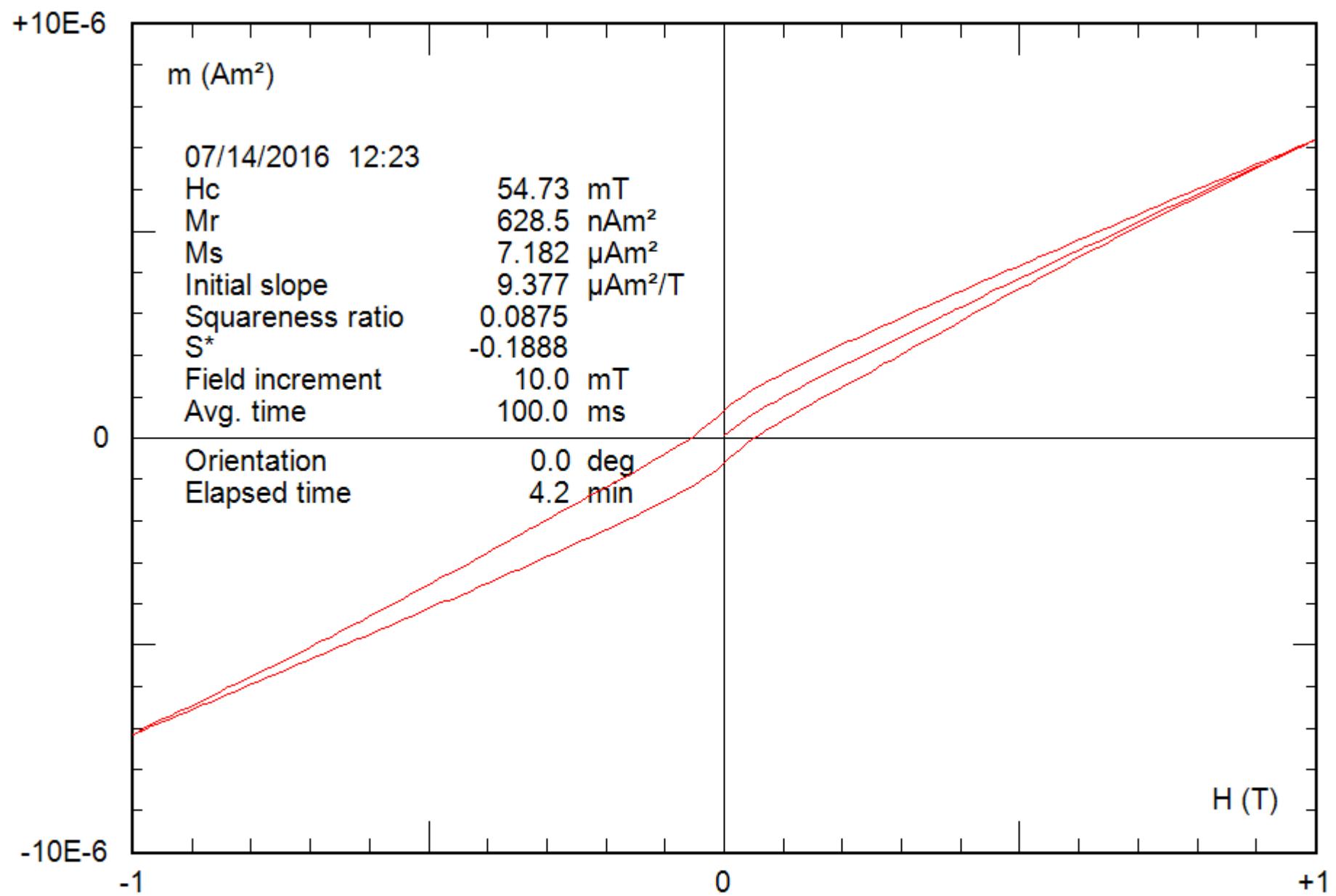
0.24 g

File: 43-05c\_hyst\_corr



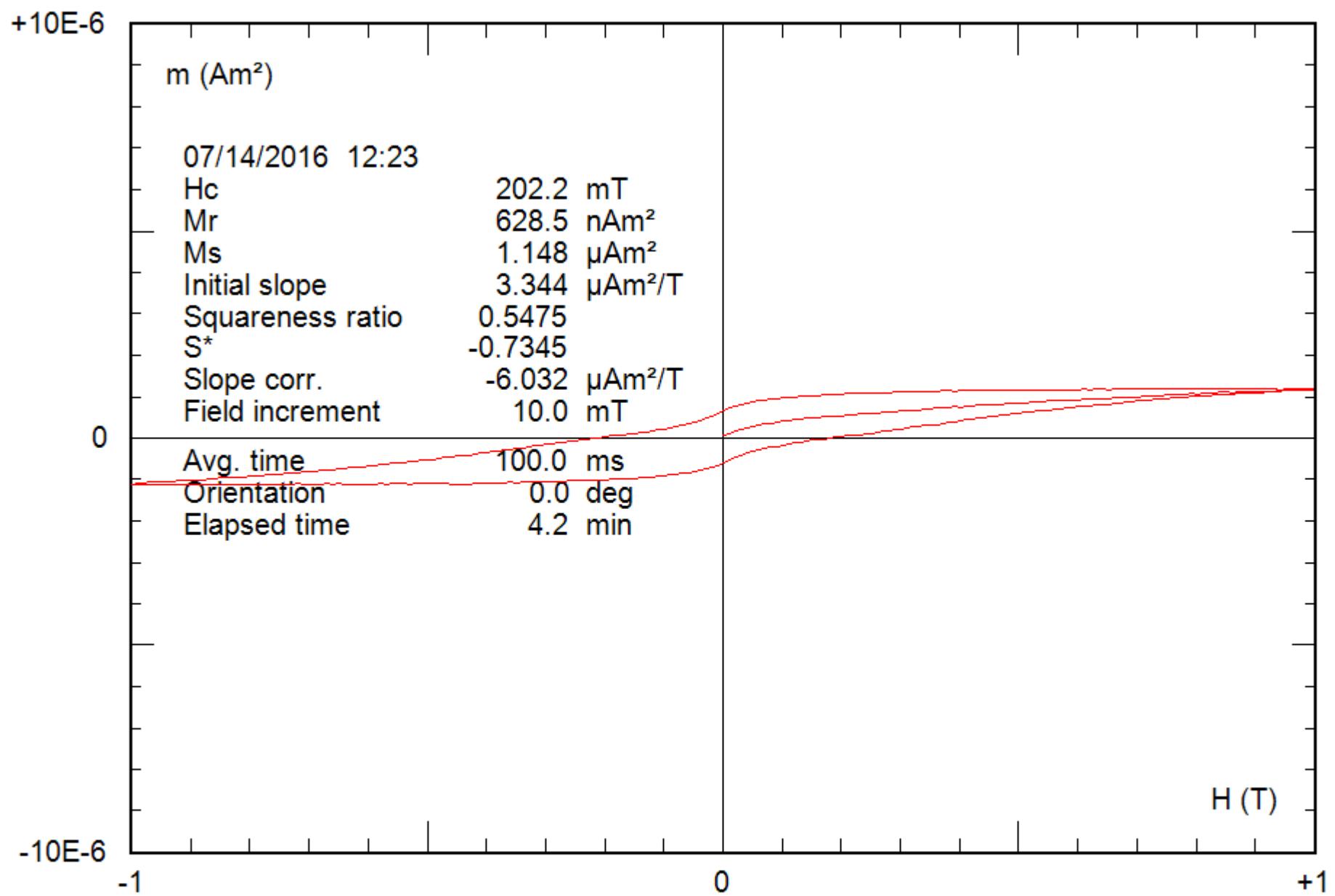
0.24 g

File: 44-10c\_bcr



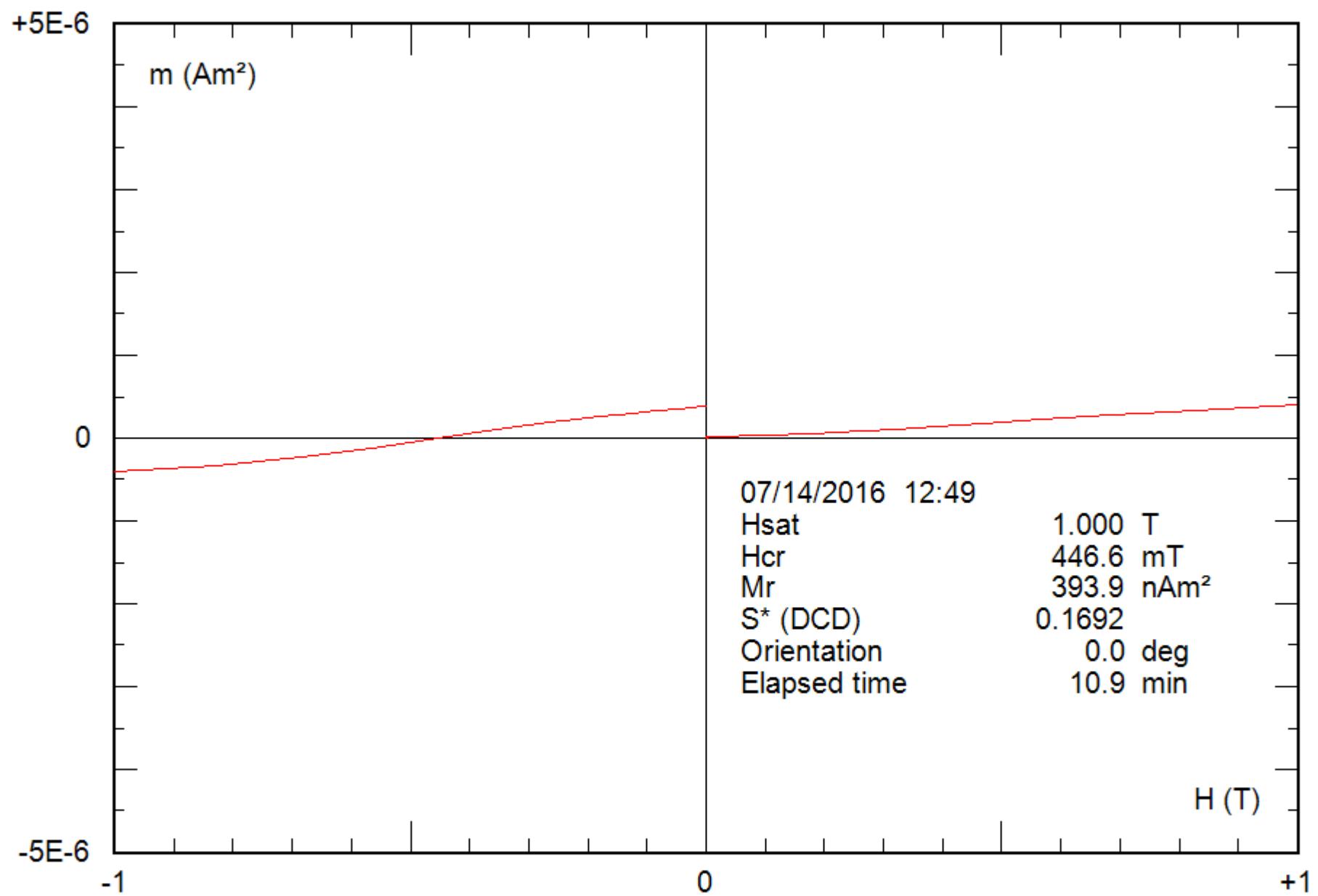
0.24 g

File: 44-10c\_hyst



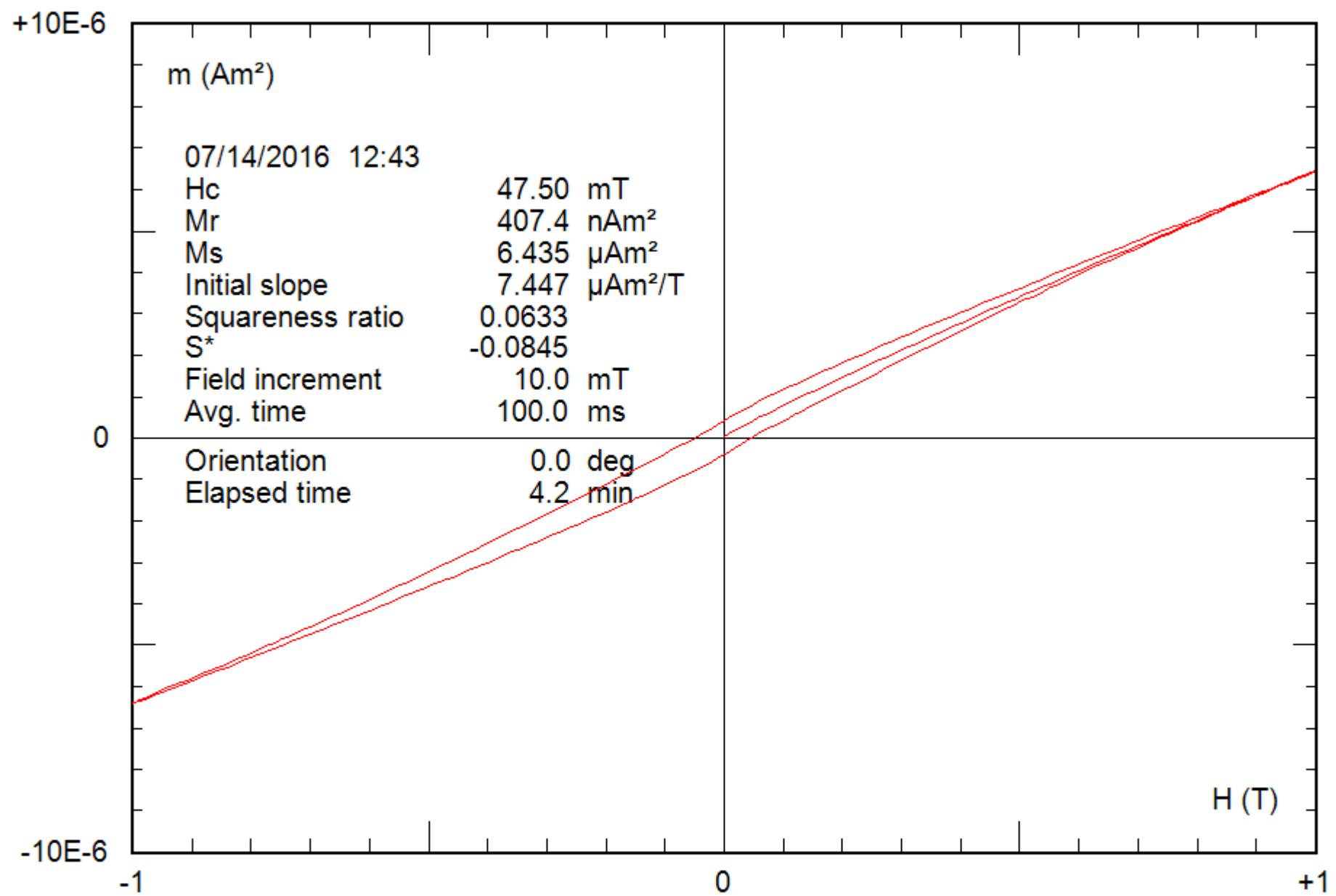
0.24 g

File: 44-10c\_hyst\_corr



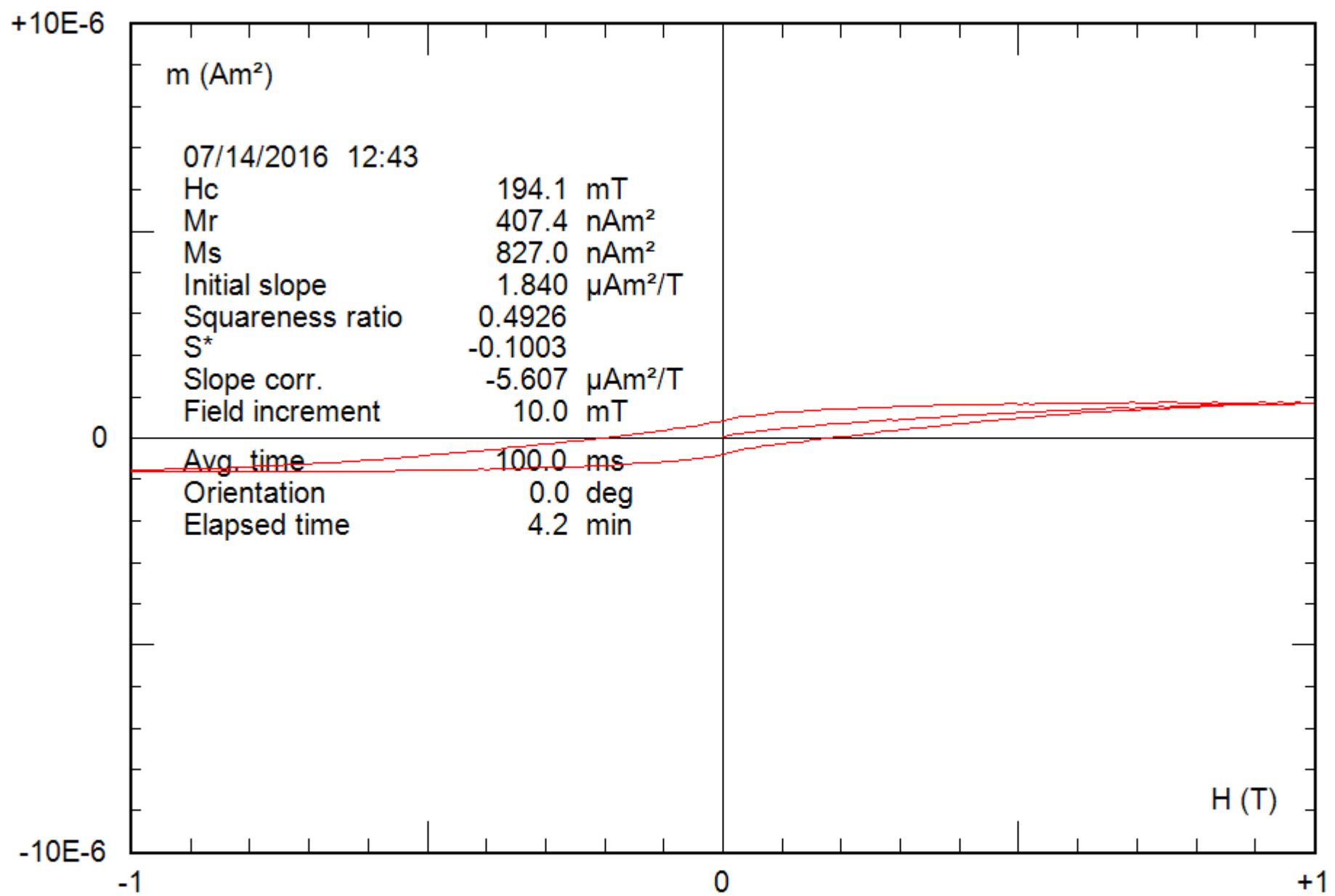
0.22 g

File: 45-03c\_bcr



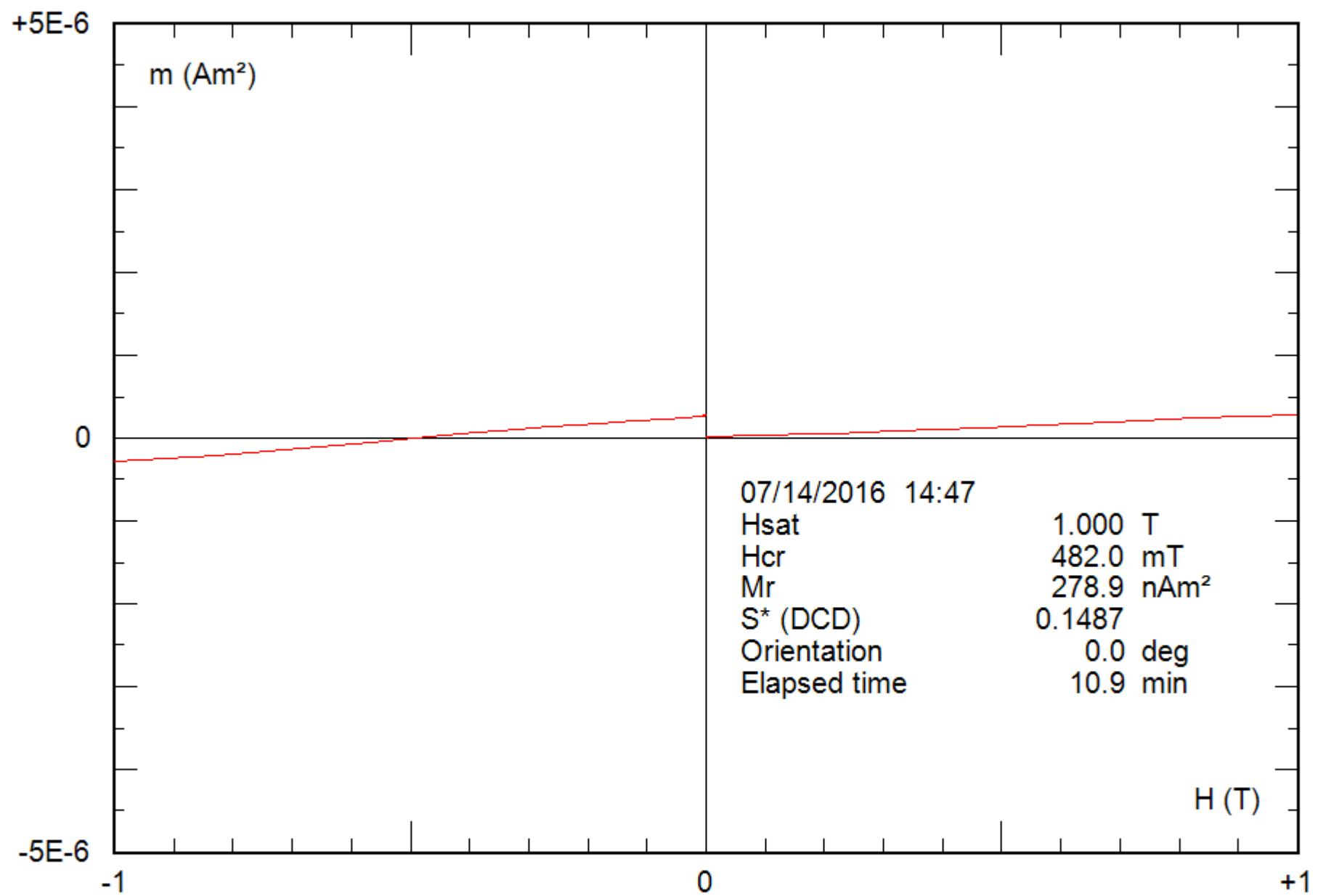
0.22 g

File: 45-03c\_hyst



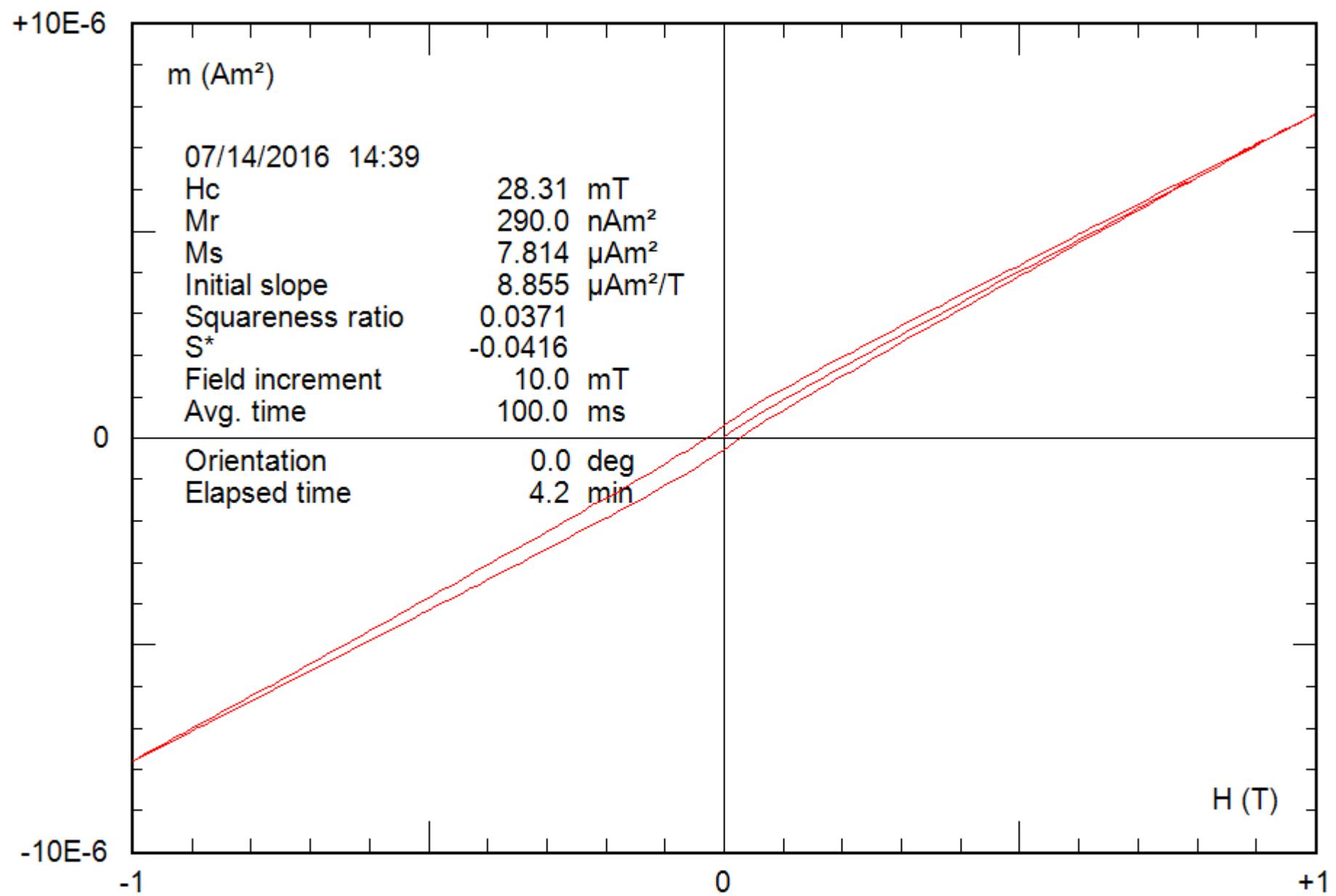
0.22 g

File: 45-03c\_hyst\_corr



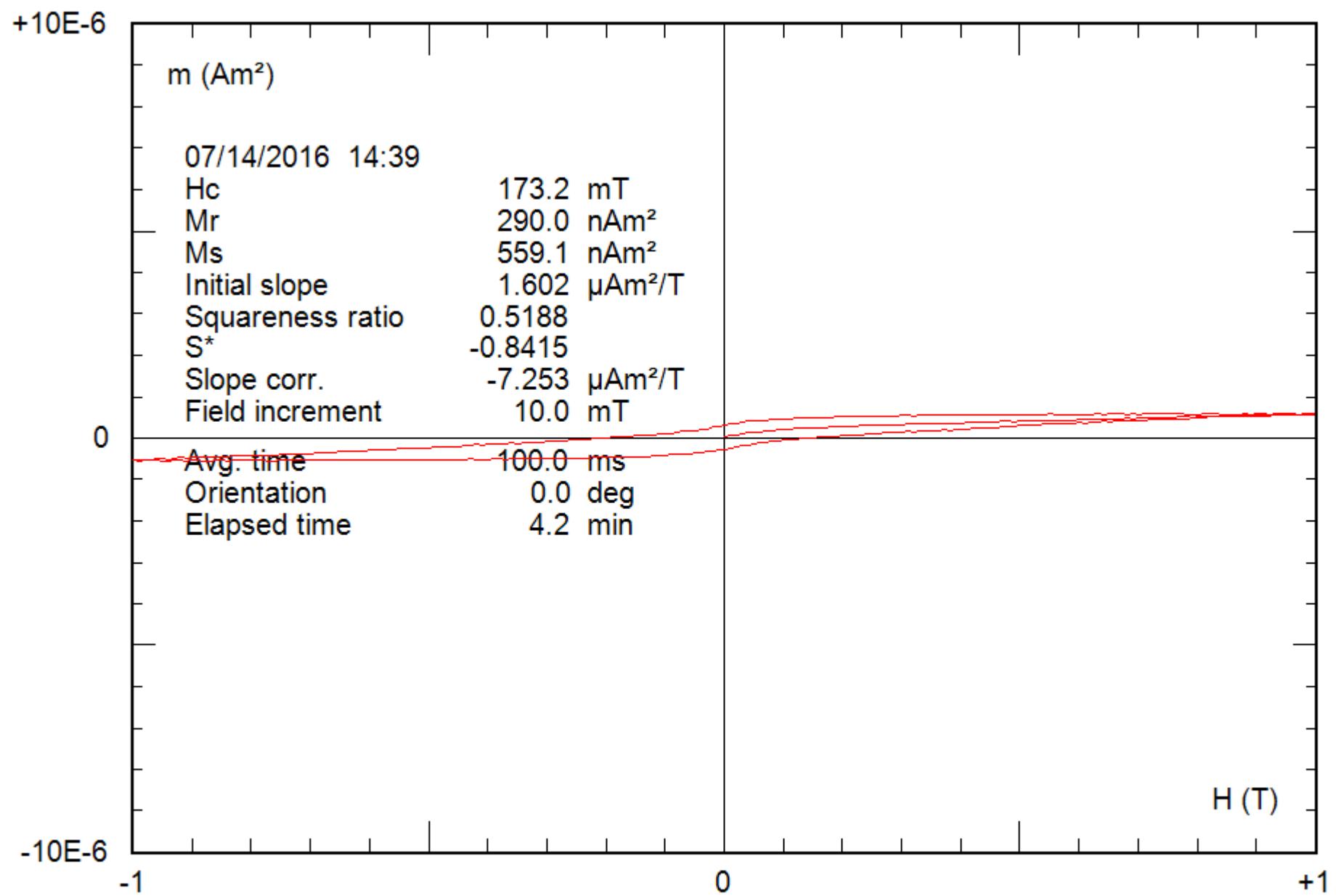
0.25 g

File: 46-09c\_bcr



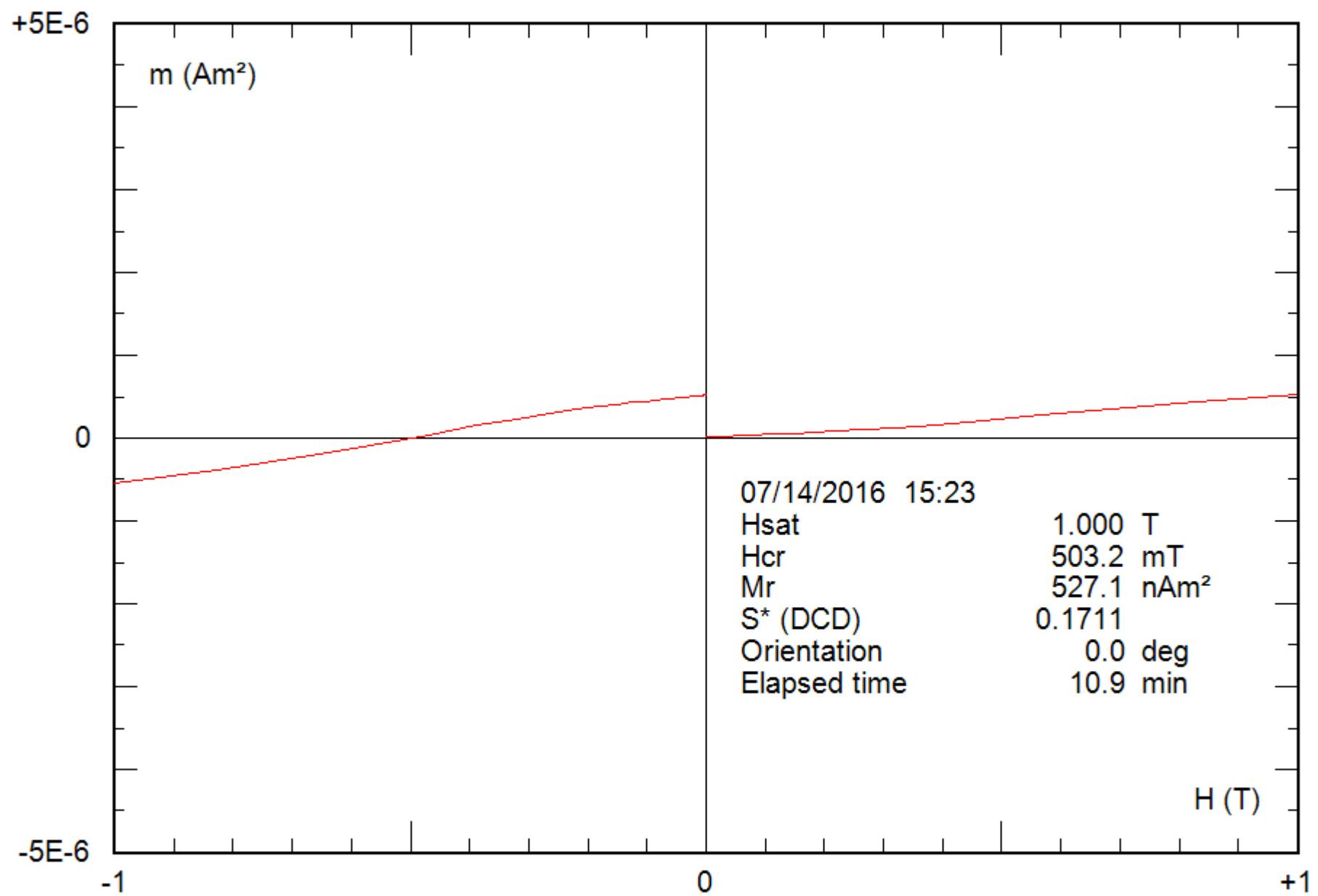
0.25 g

File: 46-09c\_hyst



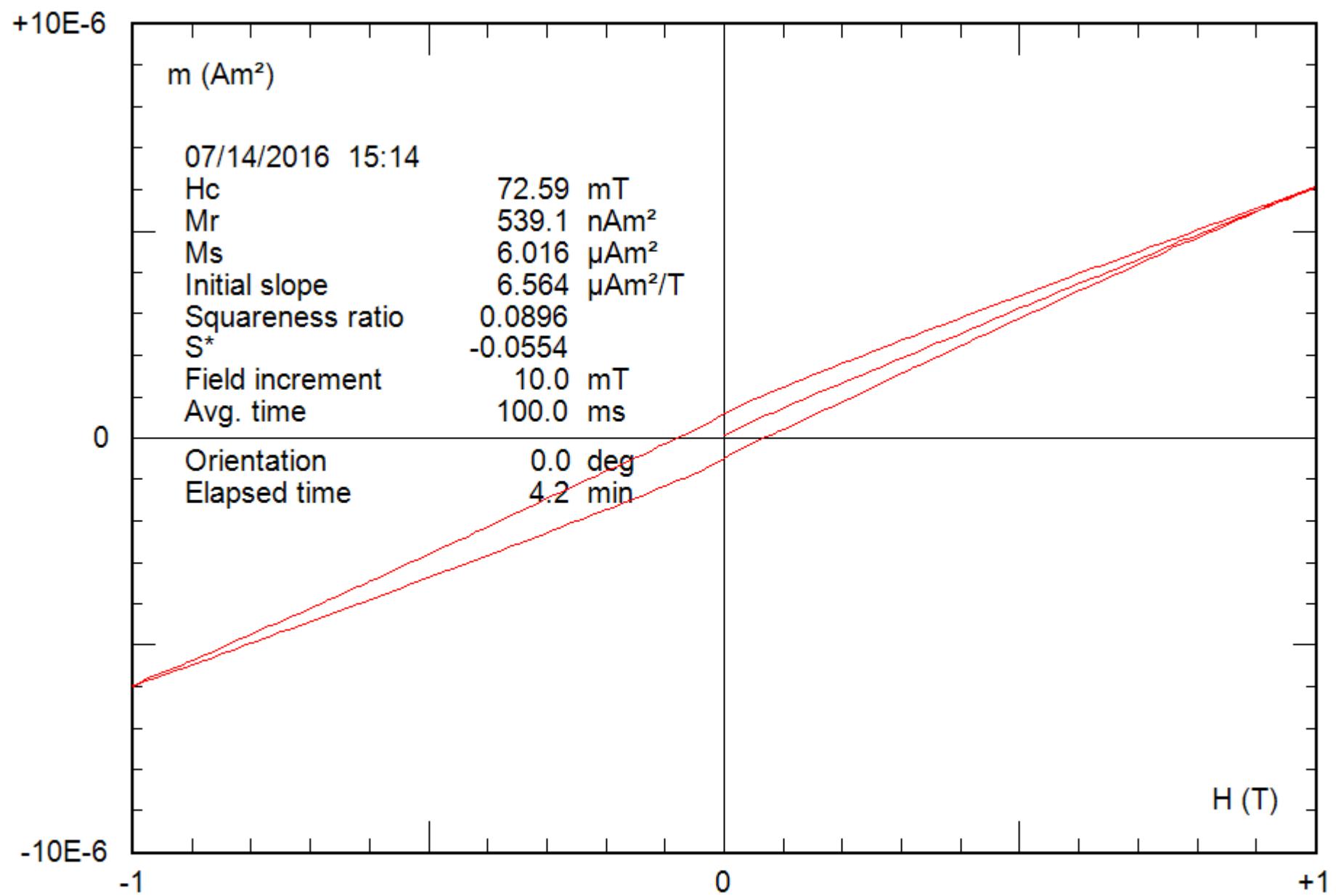
0.25 g

File: 46-09c\_hyst\_corr



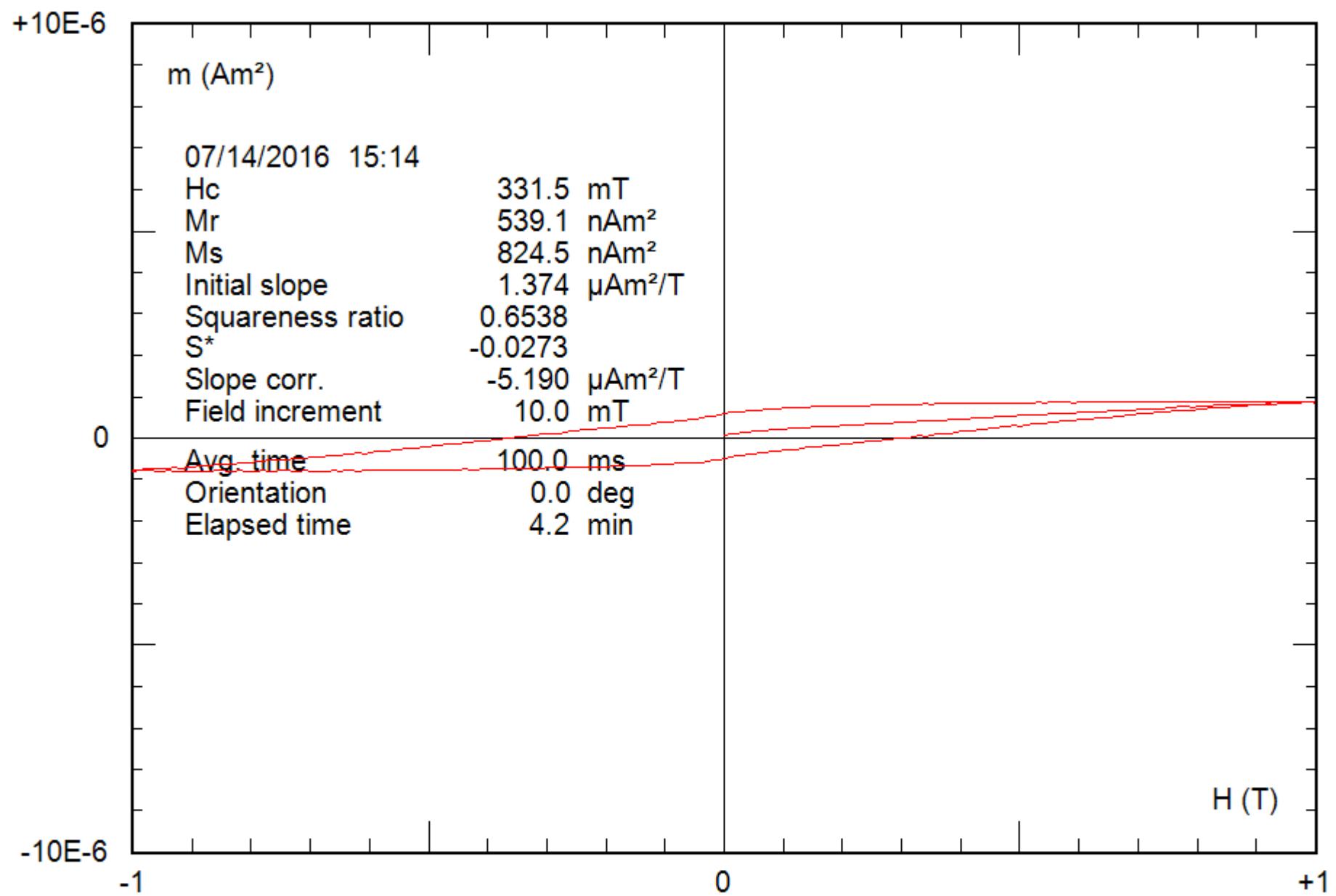
0.23 g

File: 47-08b\_bcr



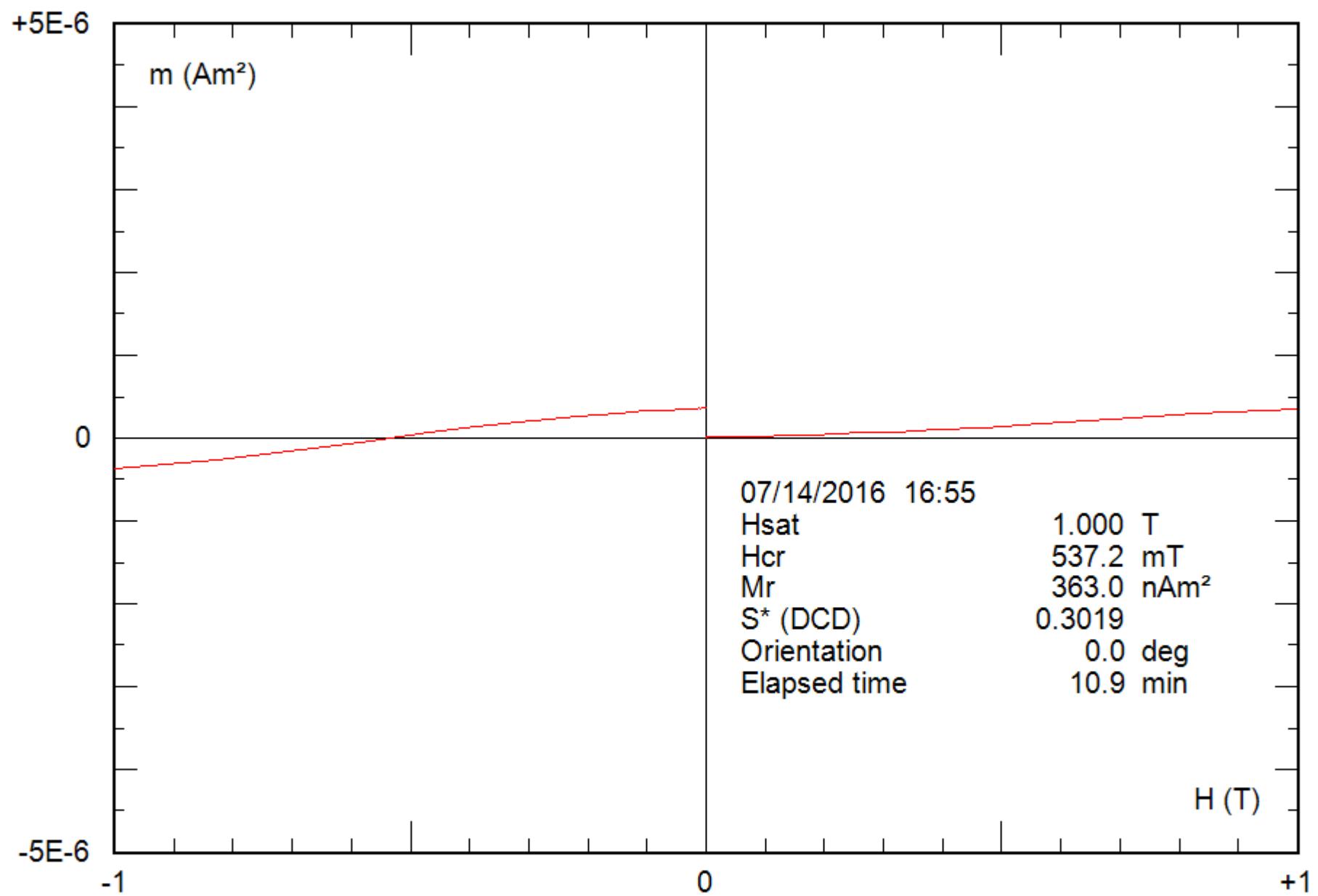
0.23 g

File: 47-08b\_hyst



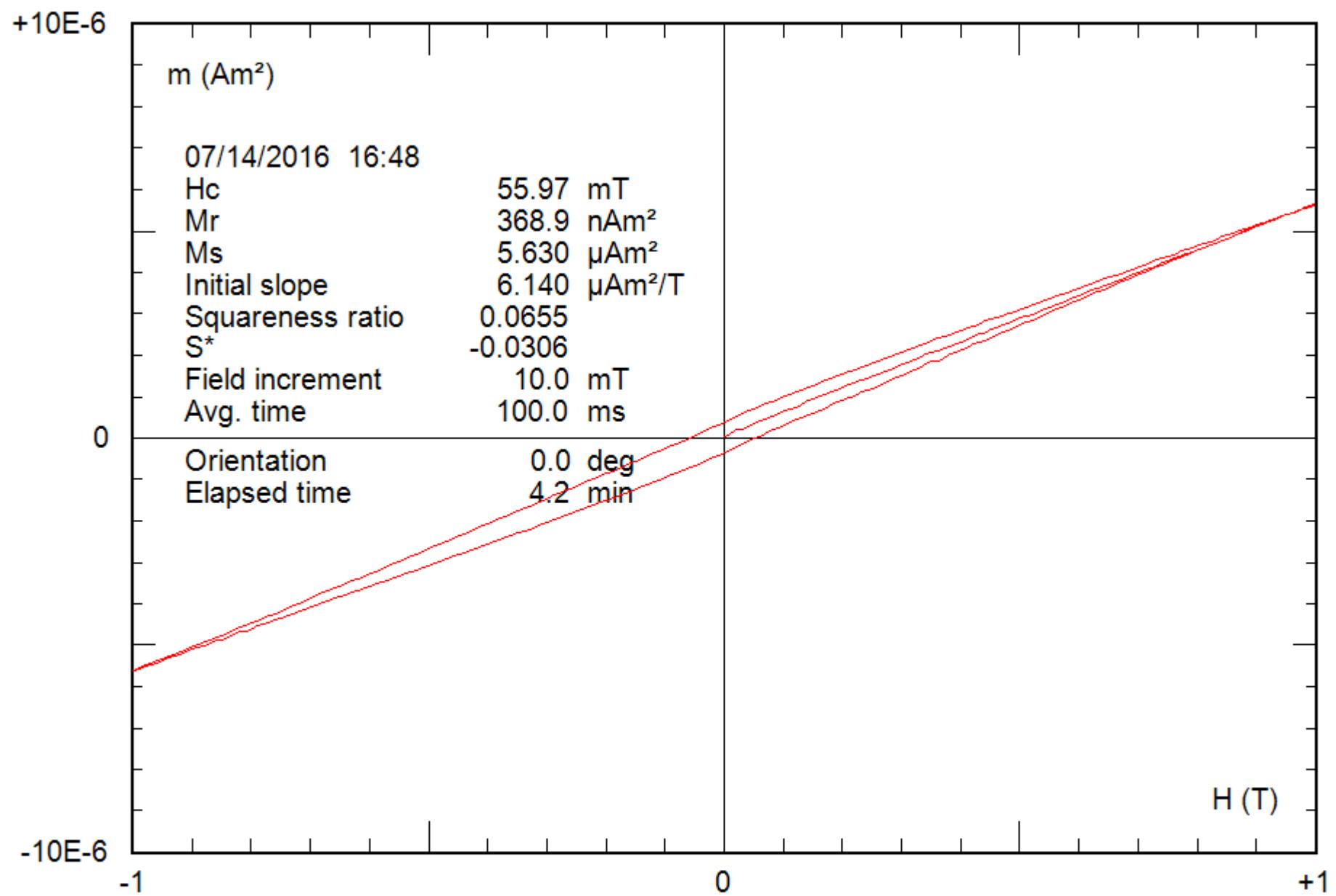
0.23 g

File: 47-08b\_hyst\_corr



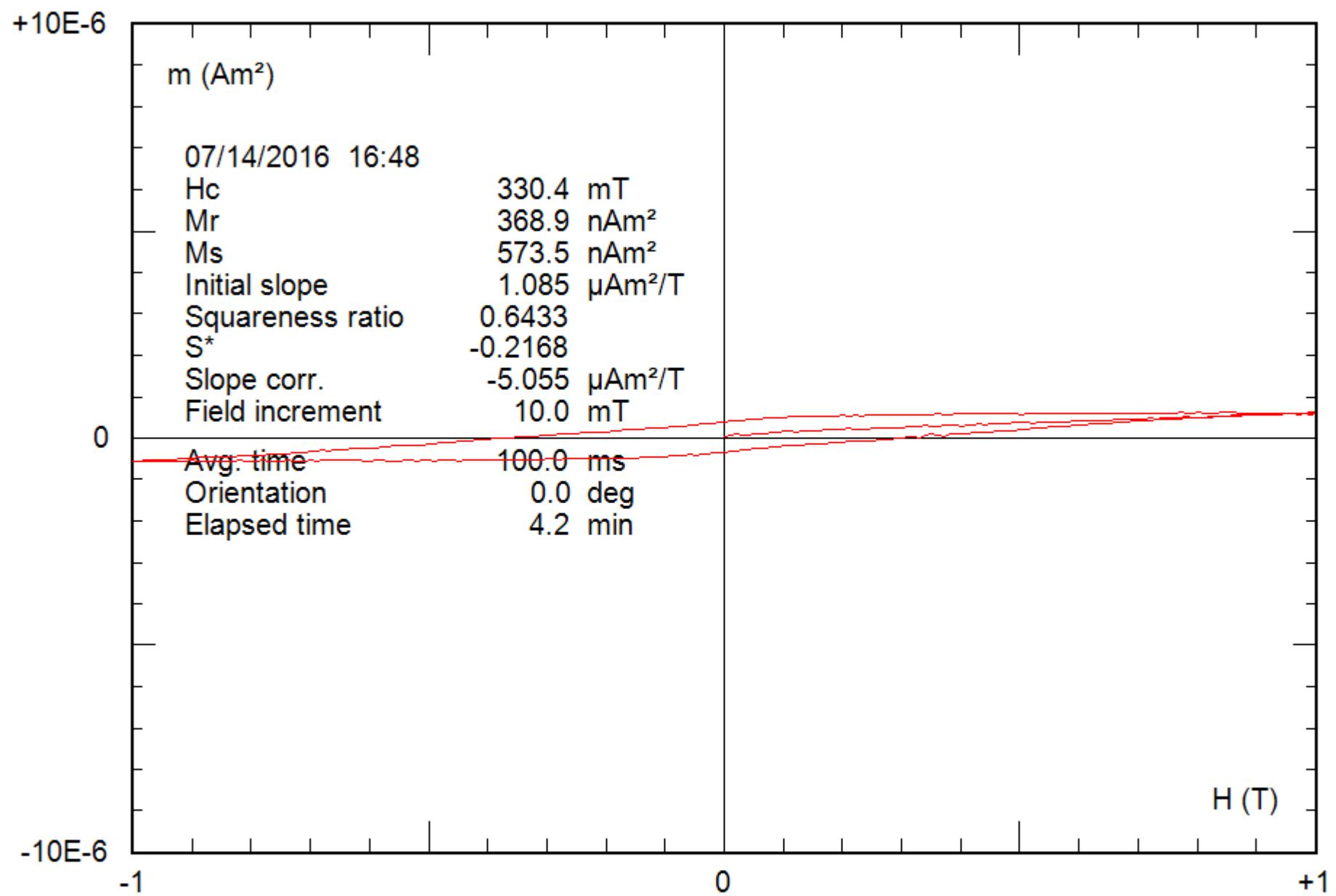
0.25 g

File: 48-05c\_bcr



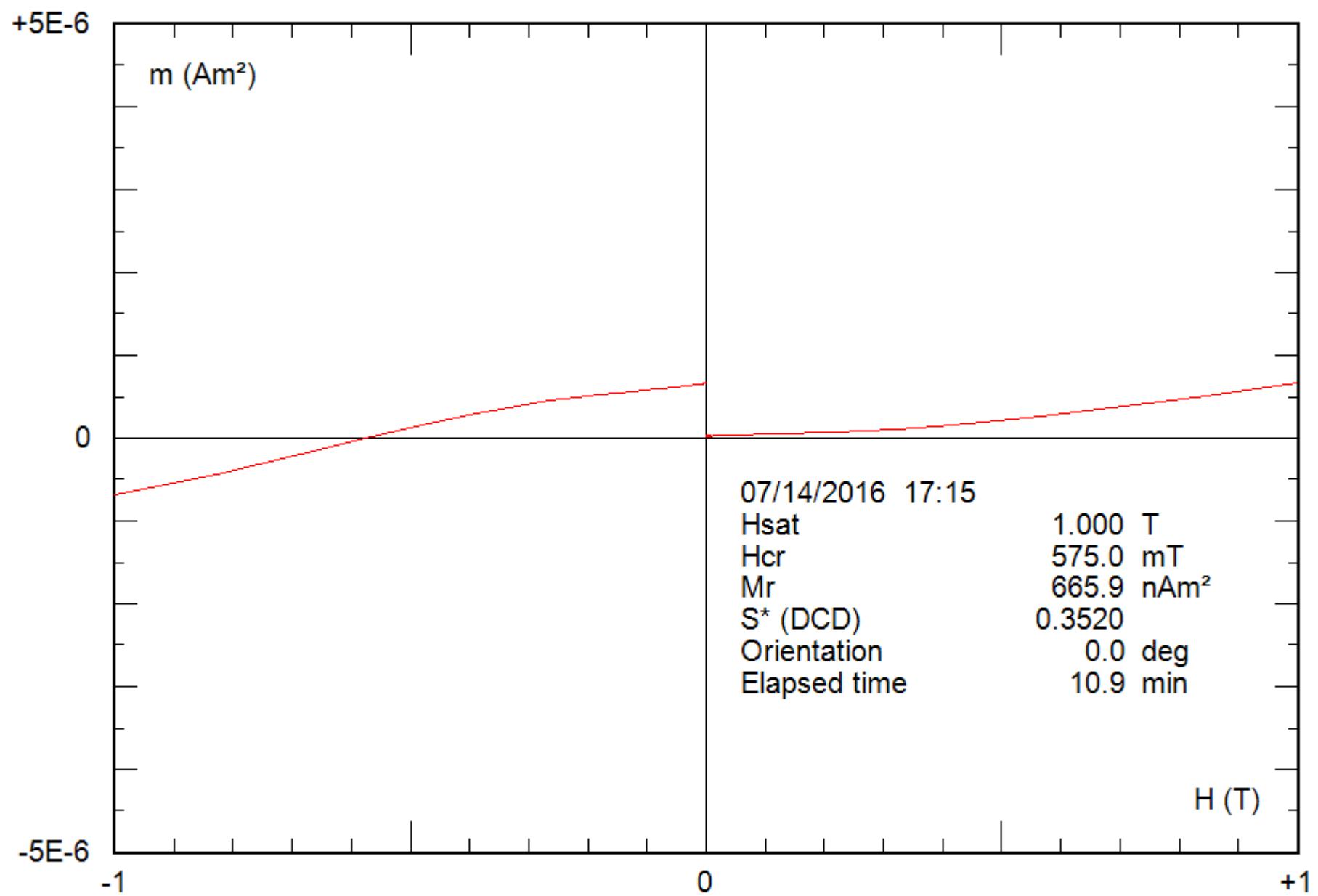
0.25 g

File: 48-05c\_hyst



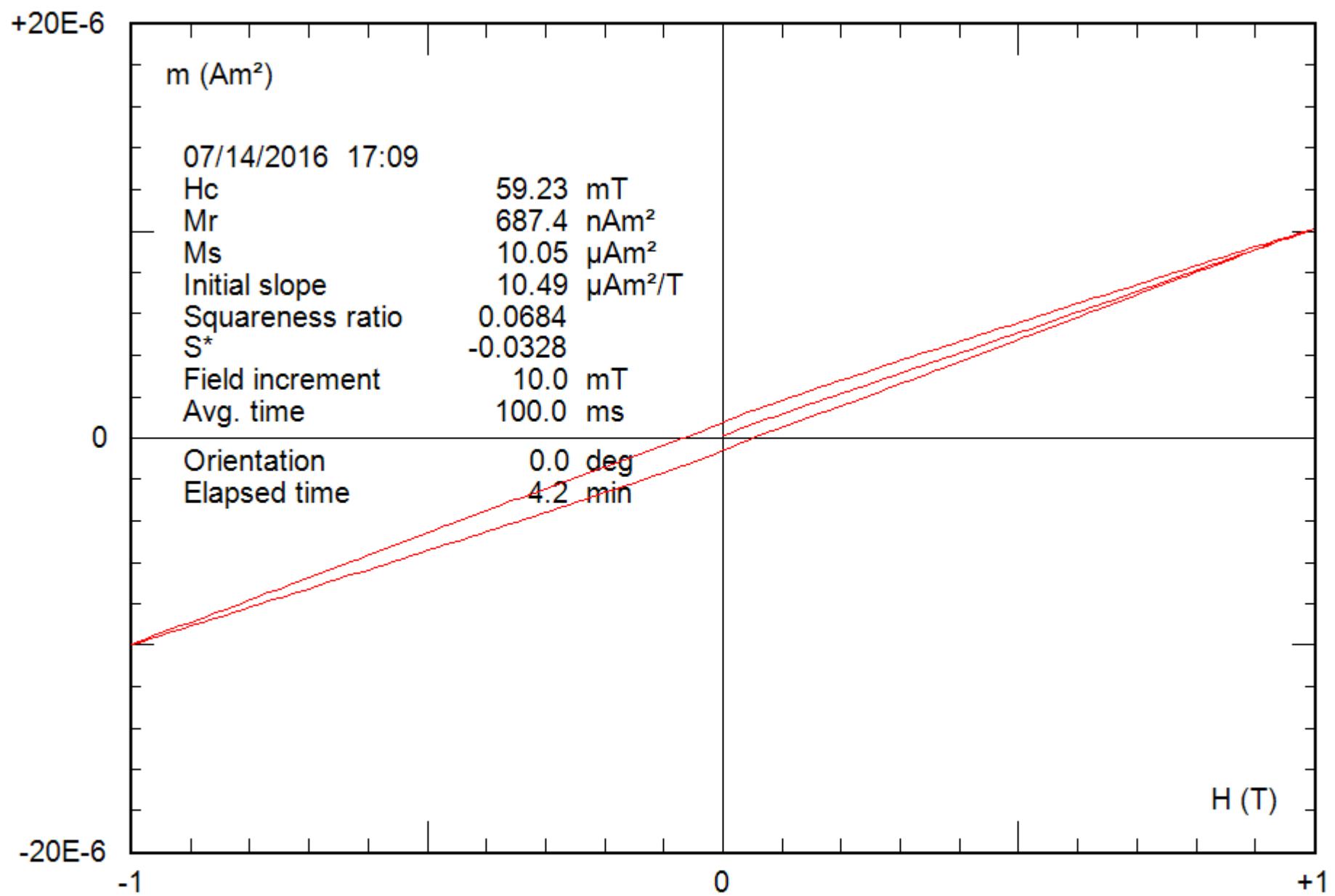
0.25 g

File: 48-05c\_hyst\_corr



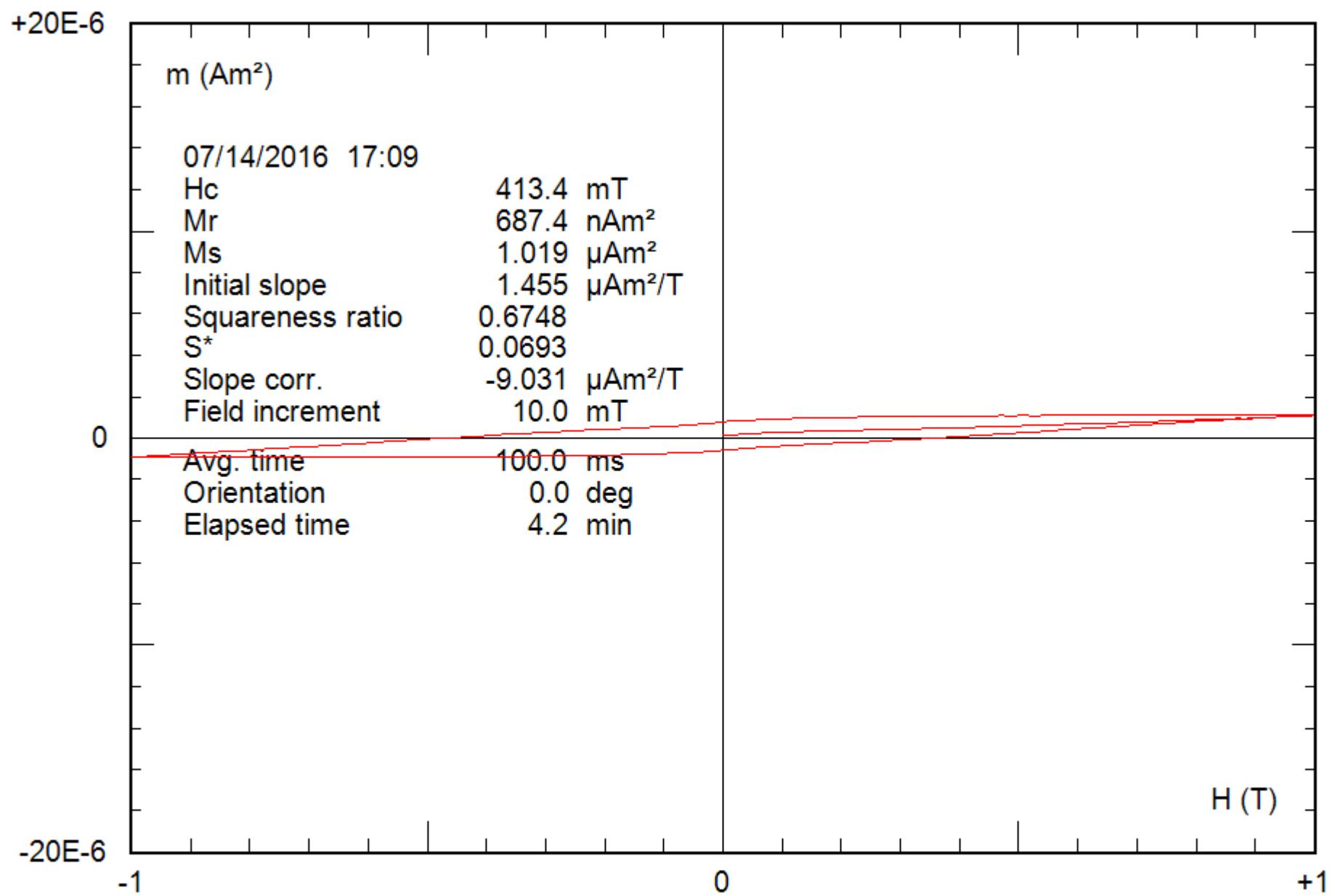
0.28 g

File: 49-04c\_bcr



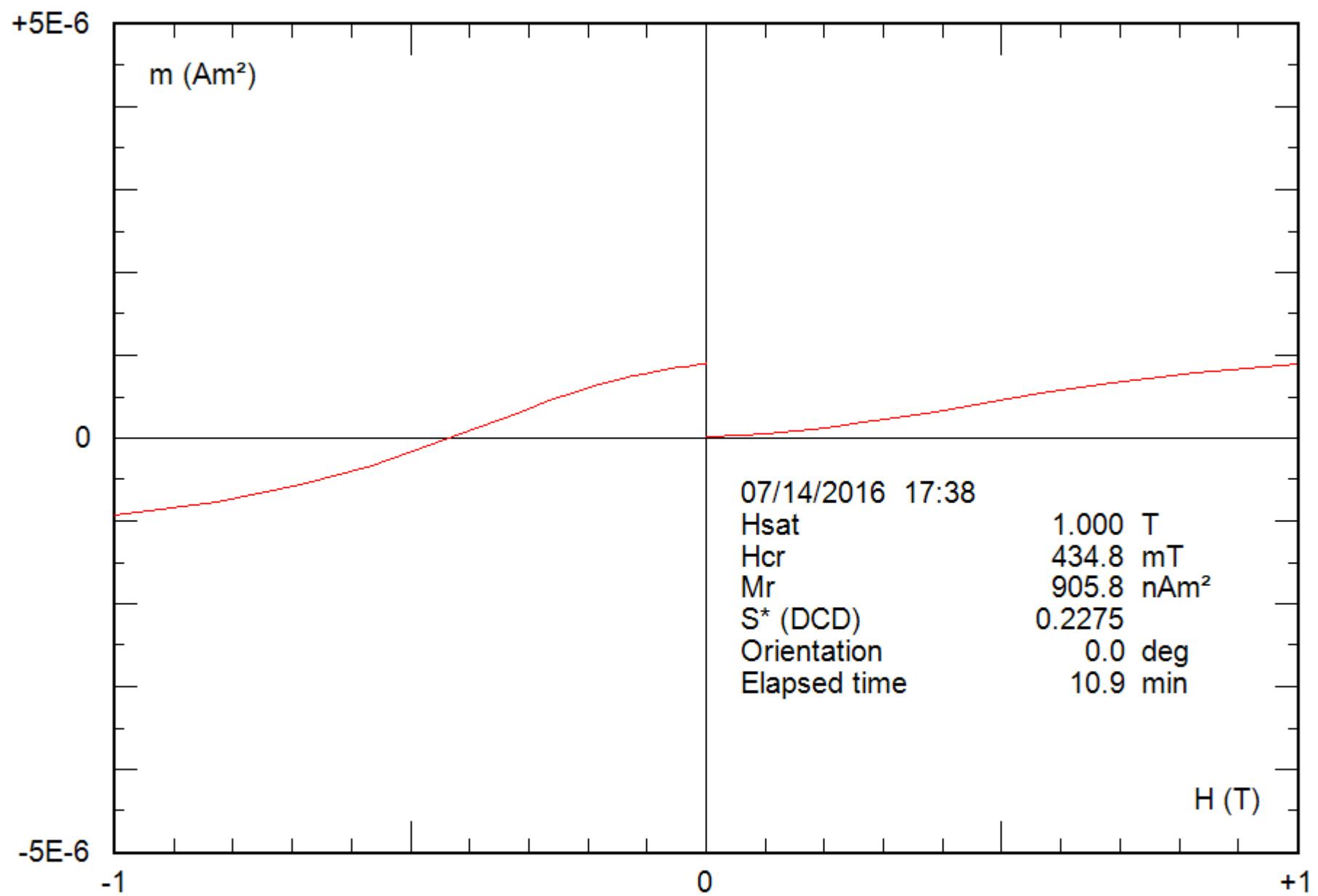
0.28 g

File: 49-04c\_hyst



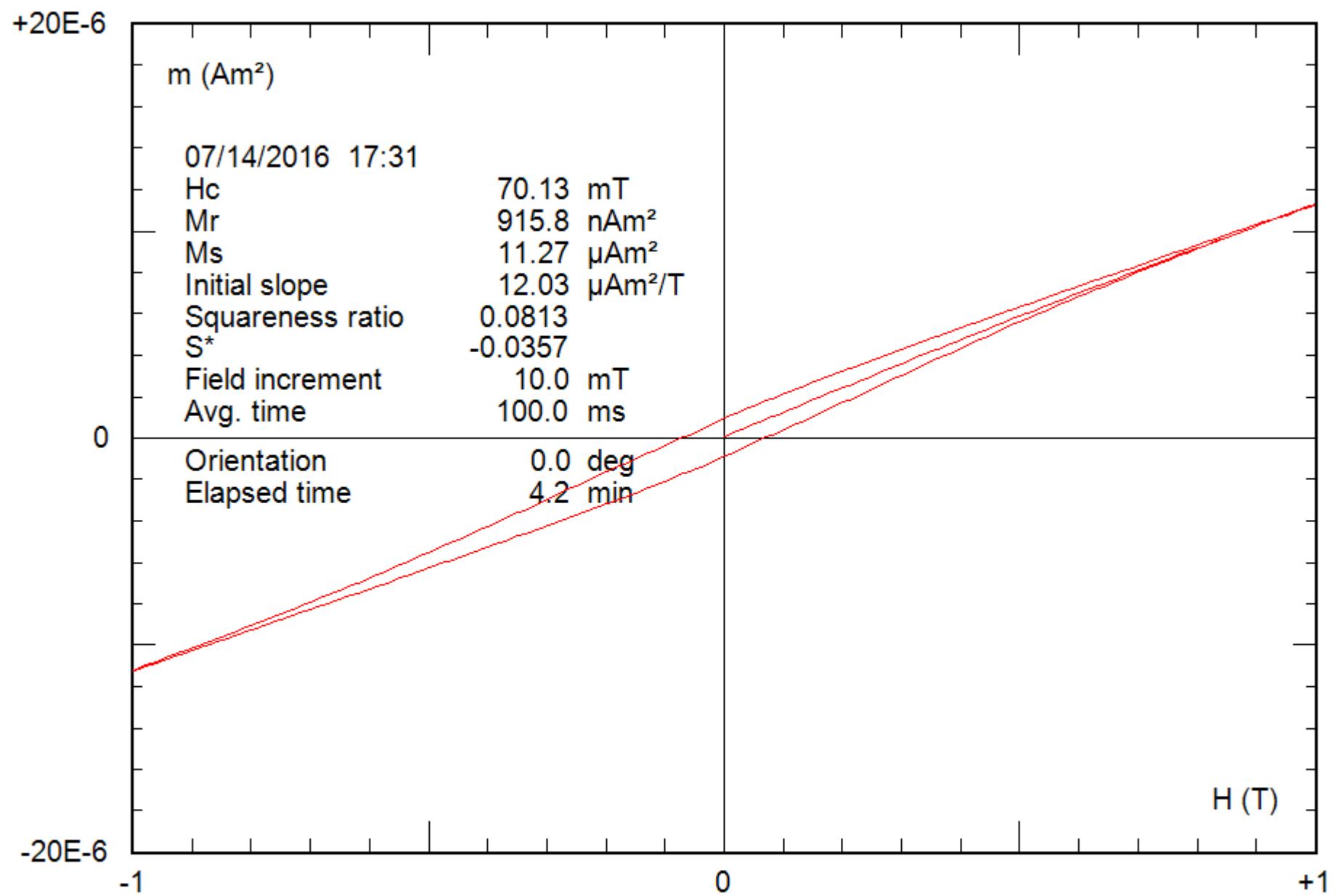
0.28 g

File: 49-04c\_hyst\_corr



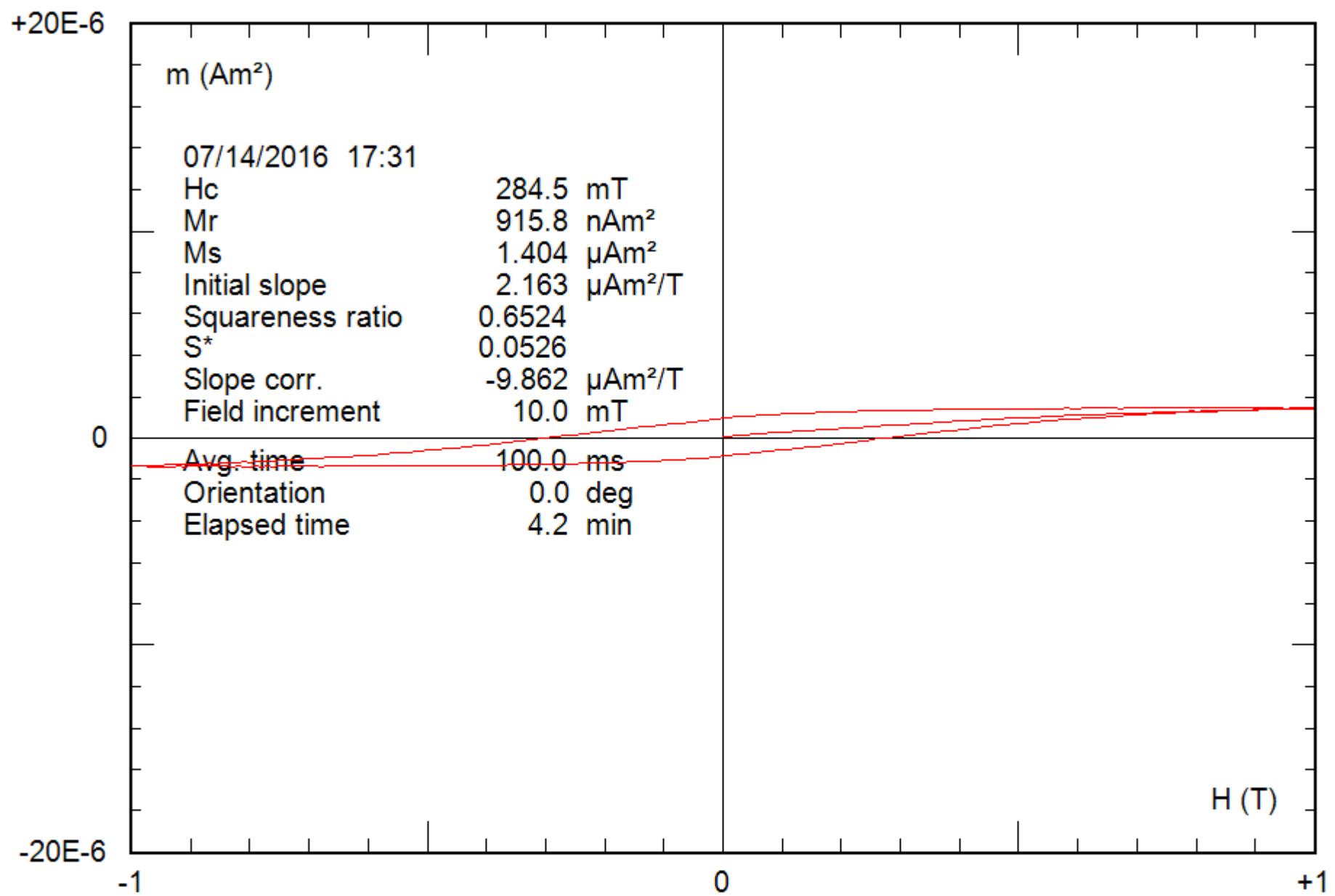
0.22 g

File: 50-10c\_bcr



0.22 g

File: 50-10c\_hyst



0.22 g

File: 50-10c\_hyst\_corr